

Interactive comment on “ARIOS: An acidification ocean database for the Iberian Upwelling Ecosystem (1976–2018)”

by Xosé Antonio Padin et al.

Anonymous Referee #1

Received and published: 9 May 2020

The manuscript is scientifically sound. With regard to the presentation, the paper is easily readable although some sentences are unclear, and there are a number of grammatical errors and typos in the text. If the authors have their text double-checked, I only have some small remarks, which need to be dealt with. Therefore, overall, I believe that the work is well suited for publication in ESSD.

Specific comments:

Title should be modified. I would suggest something like ARIOS: a database for ocean acidification assessment in the Iberian Upwelling System.

The title has been changed to the following:

ARIOS: a database for ocean acidification assessment in the Iberian Upwelling System (1976 - 2018).

Introduction

Line 41: change “fix” to “withdraw”

The suggestion has been included in the new version of the manuscript.

Lines 42-45: please rephrase “In any case, the rapid increase of CO₂ in the atmosphere decreases the ocean’s pH”. I would suggest: The gradual absorption of atmospheric CO₂ by the oceans decreases seawater pH, causing ocean acidification, which conditions the buffering capacity of seawater and in turn the exchange of CO₂ between the ocean and the atmosphere.

The suggestion has been included in the new version of the manuscript

Lines 52-53: Please rephrase. “observe and gather data about pH and other parameters of the marine carbon system to produce global and regional data products in order to help sustainably manage the ocean’s resources. I would suggest: to conduct accurate measurements of pH and ancillary parameters and provide data products for a sustainable management of marine resources.

The suggestion has been included in the new version of the manuscript

Line 55: change of marine ecosystems to for marine ecosystems

The suggestion has been included in the new version of the manuscript

Lines 63-65: I know what you mean but I would recommend to rephrase the sentence.

The sentence "However, the great spatial and temporal variability has been widely spaced and intermittent over time, preventing a complete view from being obtained of ocean acidification in the upwelling system" has been rewritten. The new sentence is: "However, the high physical/chemical variability in short temporal and spatial scales of upwelling systems and the lack of regular sampling in these waters prevents a complete picture of the acidification of these ecosystems."

Line 67-72: remove the s at the end of effects for grammar consistency. I would also move this paragraph after line 53 for coherence with the text.

The following paragraph has been moved after line 53.

"The effect of ocean acidification on marine ecosystems has stimulated impetus in the

international community for gathering high quality time-series measurements of the marine inorganic carbon system (Hofmann et al., 2011; Andersson and MacKenzie, 2012; McElhany and Busch, 2013; Takeshita et al., 2015; Wahl et al., 2016) and for predicting the future evolution of the pH caused by climate change."

Lines 72-75: I would modify the entire paragraph as: In the Iberian Upwelling System, accurate measurements of carbon system parameters commenced more than 30 years ago.

The first sentence of the new paragraph is the following "In the Iberian Upwelling System, the researchers of the Instituto de Investigaciones Mariñas (IIM-CSIC) since 1976 commenced accurate measurements of marine inorganic carbon system and associated parameters. As a result, a collection of pH observations and ancillary biogeochemical information along the Galicia coast (40°N and 45°N, 11°W) has been gathered under the framework of different projects over the past 40 years. The current database, hereinafter called ARIOS (Acidification in the rias and the Iberian continental shelf) database, holds biogeochemical information from 3,357 oceanographic stations, giving 17,653 discrete samples."

Lines 75-80: which changes? Results have not been presented yet. I would therefore, continue the statement as: Researchers of the Instituto de Investigaciones Mariñas (IIM-CSIC) have been collecting pH and biogeochemical data along the Galicia coast (40_N and 45_N, 11_W) under the framework of different projects. This has allowed to generate a database, ARIOS (Acidification in the rias and the Iberian continental shelf), containing 17,653 discrete records gathered in 3,357 sampling sites.

In response to the previous two comments, the paragraph has been rewritten in order to include the suggestions: "In the Iberian Upwelling System, the researchers of the Instituto de Investigaciones Mariñas (IIM-CSIC) since 1976 commenced accurate measurements of marine inorganic carbon system and associated parameters. As a result, a collection of pH observations and ancillary biogeochemical information along the Galicia coast (40°N and 45°N, 11°W) has been gathered under the framework of different projects over the past 40 years."

Data provenance: I would remove provenance

Line 86: I would replace Region by Data Coverage

Region has been replaced by Data spatial coverage

Line 115: change In addition to for Besides

The suggestion has been included in the new version of the manuscript

Line 119: delete the article before acidification and it would be convenient to specify the exact region/s where the mentioned acidification rate was estimated

The suggestion has been included and the region was mentioned as "the acidification in the first 700 metres for the geographical area from the Iberian Peninsula to the 20° W meridian and from 36°N to 43°N has also been observed at a rate of -0.0164 pH units per decade"

Data sources:

In general, I very much appreciated the comprehensive explanation of the projects that provided data for the dataset. However, considering that explanatory information of the cruises is given in Table 1 and each individual project is associated to a database included in a public repository, I do not find section 3.2 essential for the manuscript, as all those details can be mentioned (and possible are) in the repository. The authors might re-consider to shorten this section by keeping the first paragraph and refer subsequent info to Table 1.

We understand that detailed information about the methods and materials of the oceanographic cruises that are part of the ARIOS database is mandatory according to the requirements of ESSD. Therefore, we prefer not to modify this section of the manuscript.

Methods:

Lines 337-338: Please rephrase. I would suggest: Except for the Galicia cruises (Table 1), in which nutrient samples were analysed on board, samples were kept in the dark and cold (4_C) after collection for further analyses in the shore based laboratory.

The suggestion has been included in the new version of the manuscript

Line 341: change is to was.

The suggestion has been included in the new version of the manuscript

Line 351: Same as above

The suggestion has been included in the new version of the manuscript

Line 371: remove the article before Table 1

The suggestion has been included in the new version of the manuscript

Line 374: You possibly mean Table 1 instead?

The suggestion has been included in the new version of the manuscript

Line 377: by the high variability present in a system characterized by an intense biological activity

The suggestion has been included in the new version of the manuscript

Line 384: This section should be moved and either merged with 3.1 or placed right below it for the sake of consistency and for a better introduction of the sampling region.

This section has been moved just below section 3.1. The other sections of the Material and Methods have been renumbered because of this change.

Results:

Line 424: vertical profile of what? Please indicate.

These sentences had been rewritten. The new text is: "The vertical profile of the temperature, salinity, pH_T, NO₃⁻ and oxygen concentration in the ocean region between 41°N and 43°N was estimated for each oceanographic station as the mean value of the depth ranges described in Figure 2b"

Lines 420-445: why do not you show the standard deviations for T and S for all the depth ranges as you do for the water column comprised between 500m and 1100m?

The standard deviations were not included in Figure 3 simply to facilitate the clarity of the graphical representation of the vertical biogeochemical profiles. For the same reason, they were also not included in the monthly distribution of the biogeochemical variables shown in Figure 4. In any case, the mean values with their corresponding standard deviation values (mean±standard deviation) represented in Figure 3 and Figure 4 are shown in two Tables at the end of the reply to the referee. This information could be included as supplementary material to the manuscript if it deems this appropriate by the Editor.

Line 443: add : : down to 1100m

The suggestion has been included in the new version of the manuscript

Line 447: change distribution to profile

The suggestion has been included in the new version of the manuscript

Line 448: replace at this depth by within this depth range

The suggestion has been included in the new version of the manuscript

Lines 452-454: Speculative as it is not demonstrated or shown in the graph. Therefore, I would just say: The highest pH values could be attributed to the biological CO₂ drawdown by phytoplankton activity, which brought the pH to a peak value of 8.13 at 40 metres deep during the spring bloom.

The suggestion has been included in the new version of the manuscript.

Lines 454-457. Same as above. I would suggest to rephrase the paragraph as it is also confusing. Below 100 metres, respiration of organic matter possibly was responsible of lowering pH: : .but anyhow the text s counterintuitive To me, pH values between 200 and 500 m depth seem to be lower than those from 500 m down to 1100m, which were also constant and similar within the entire depth range regardless of the season.

The commented lines: "Underneath this intense photosynthesis activity between the surface and 100 metres, the respiration of organic matter took the pH to lower values than those measured in winter between 200 to 500 metres, a depth at which the spring and winter values were practically equal." were rewritten. This was the new text: "The higher pH values could be attributed to the biological reduction of CO₂ by phytoplankton activity, which brought the pH to a maximum value of 8.13 to 40 meters during the spring bloom. After the intense photosynthetic activity observed in surface waters during spring and summer, pH values reached minimum values in the first 200 meters of depth during autumn due to respiration of organic matter. However, it was at a depth of 500 metres that the minimum pH values were measured in all seasons where is found the subpolar Eastern North Atlantic Central Water proceeding from the northeastern cyclonic gyre (Harvey, 1982; Ríos et al., 1992)."

Lines 458-462: Please rewrite: I would propose: The influence of phytoplankton growth on biogeochemistry during spring can be also evidenced by the oxygen concentration pattern during this season. In the upper layer (depth range?) spring oxygen levels exceeded those in winter, whereas a decrease in oxygen concentration was found from 300 m depth down to 1000 m, possible due to enhanced respiration from cascading organic matter. It would be helpful to add in this section a table with averaged concentrations and SD of each parameter within the different depth ranges and for each season.

The suggestion has been included in the new version of the manuscript. The new lines are the following: The influence of phytoplankton growth on biogeochemistry during spring can be also evidenced by the oxygen concentration pattern (Fig. 3e). In the upper layer above 250 metres depth, spring oxygen levels exceeded those in winter, whereas a decrease in oxygen concentration was found from this depth down to 1000 metres, possible due to enhanced respiration from cascading organic matter.

Line 473: seasonal cycle of what? Please specify. I would recommend to rewrite the whole paragraph, as in Fig 4 what you actually show is the seasonal cycle of different biogeochemical parameters in surface waters of 5 regions and not the five regions themselves, as it can be deduced from the text the way it is right now.

The suggestion has been included in the new version of the manuscript. The new lines are the following: "The seasonal cycle of the biogeochemical properties (temperature, salinity, pH, oxygen concentration, nitrate concentration and chlorophyll) in the surface waters (0 to 5 metres) of five geographical boxes was estimated as a monthly average previously filtering

values outside of two standard deviations of the mean. Five regions that were located as a longitudinal transect between the inner Ría de Vigo and the ocean zone are shown in Fig. 4."

Line 506: replace seasonable by seasonal.. you could also rewrite the following sentence as: with maximum and minimum pH values in spring and autumn, respectively, and in all regions (Fig. 4c).
The suggestion has been included in the new version of the manuscript

Lines 541-549: considering change to past tense for consistency with the rest of the paragraph. Moreover, a reference could be well added at the end of the paragraph to reinforce your statement regarding the relevance of benthic and vertical fluxes in the Ria.

The suggestion has been included in the new version of the manuscript. The new lines are the following: "The nutrient concentration during spring and summer was only detectable in the newly upwelled waters that can show values up to $6 \mu\text{mol L}^{-1}$ (Fraga, 1981; Castro et al., 1994). During the cessation of the upwelling season in September and October, the chlorophyll concentration (Fig. 5f) increased again, sustained by nutrients that entered from deeper waters through vertical mixing. It should be noted that there was a coincidence of high chlorophyll in the water column and low oxygen concentration in the inner *Ría de Vigo* from May to November, indicating the potential importance of benthic fluxes and vertical fluxes (reference).

Line 551: please add a "s" to trend

The suggestion has been included in the new version of the manuscript

Line 552: long term trends of what? Please specify. I assume the temporal trends are estimated over parameters and they do not refer to surface waters themselves. Therefore, it needs to be re-written.

The text has been re-written. "The long-term trends of the biogeochemical properties in these surface waters were estimated to be the interannual linear rate of the deseasonalized time series, previously removing the monthly means in these regions and assuming a null spatial variability"

Line 562: any suggestion why a warming trend is not found as it was previously reported?

The fact that we did not find a statistically significant warming trend as would be expected on the basis of the overall ocean behavior is mainly due to the fact that the warming trend was estimated for the surface waters of the study area. The first meters of the water column in this coastal zone are under the influence of important oceanographic phenomena that affect temperature such as coastal upwelling pulses, the presence of surface currents or river inputs. The different temporal variability of these processes together with the remarkable spatial variability of the study area prevents the characterization of the warming in a statistically significant way from measurements with the irregular frequency of the ARIOS database. Therefore, information from other temperature data sources and a more detailed statistical analysis should be considered for this specific purpose.

Line 563: consequence of climate change: : I would add "in marine ecosystems".. and move the reference by Caldeira and Wicket 2003 after ocean acidification.

The suggestion has been included in the new version of the manuscript

Line 565: What do you mean by pH number? Value? Number of measurements?

The text has been re-written. "with a greater decline in pH number towards the coast" has been changed by "with a greater decrease in the long-term trend of pH towards the coast

Line 566: please replace : : : was about triple the change of : : by: : : .. was three fold higher than the

trend observed in the open ocean zone, equivalent to $0.0012 \pm 0.0002 \text{ yr}^{-1}$

The suggestion has been included in the new version of the manuscript

Line 569-575: I would rephrase the paragraph as: These pH decrease rates found in both coastal and open ocean regions of the Iberian Upwelling System lie within the range of other acidification rates estimated in different sites of the North Atlantic Ocean (Lauvset and Gruber, 2014; Bates et al., 2014), being also coherent with the mean rates calculated for the global ocean and for the Eastern North Atlantic and equal to -0.018 and $-0.0164 \text{ decade}^{-1}$, respectively (Lauvset et al., 2015; Rios et al 2001)

The suggestion has been included in the new version of the manuscript

Line 575: just indicate: Salinity exhibited an increasing long-term trend (value?) that was dependent on the distance to the mouth of the Ria (de Vigo? All of them?). I do not see the salinity trend indicated anywhere.

The suggestion has been included in the new version of the manuscript

Line 581: Change So to Therefore.

The suggestion has been included in the new version of the manuscript

Line 586: add a “s” to nutrient and remove the previous article

The suggestion has been included in the new version of the manuscript

Line 587: same as above: delete the article before nitrate. But anyhow, it is not clear if the trends in nutrients level come from the previous study by Doval et al (2016) or are the result of your analysis. Please clarify.

The rates had been estimated from ARIOS database. This paragraph has been rewritten in order to clarify the meaning. The new text is the following: " The long-term trend of the concentrations of nutrients in the inner Ría de Vigo that had been previously reported for the period 2001-2011 by Doval et al. (2016) showed a significant increase in nitrate, phosphate and ammonium concentrations of $0.0559 \pm 0.0158 \mu\text{mol kg}^{-1} \text{ yr}^{-1}$, $0.0076 \pm 0.0016 \mu\text{mol kg}^{-1} \text{ yr}^{-1}$ and $0.0560 \pm 0.0011 \mu\text{mol kg}^{-1} \text{ yr}^{-1}$ respectively."

Lines 590-596: do you mean that your AOU temporal trend coincides with the deoxygenation rate calculated previously by Doval? It is not clear enough in the text.

Long-term trends in oxygen concentration and AOU were estimated from the ARIOS database. The paragraph was rewritten to clarify this information. This is the new text: "This fertilization on a long-term scale estimated from ARIOS database in the surface waters of the inner ria was observed in parallel to the deoxygenation of $-0.7 \pm 0.2 \mu\text{mol kg}^{-1} \text{ yr}^{-1}$. The apparent oxygen utilisation (AOU), calculated using the concentration of oxygen at saturation calculated according to Benson and Krause (1984), underwent a long-term change of $0.7 \pm 0.2 \mu\text{mol kg}^{-1} \text{ yr}^{-1}$ equal to the observed in the measurements of oxygen concentration. This coincidence may indicate that the long-term reduction of oxygen is due to the changes in the biological consumption rates, in the rates of the waters ventilation or even in sediment-water interactions rather than due to the effect of temperature and salinity on oxygen saturation."

Lines 598-604: Speculative. Please support with references

This paragraph was intended to provide an integrative hypothesis on observed large-scale trends. The text has been changed in the manuscript to the following paragraph: "These findings found in the shallower waters of the *Ría de Vigo* allow us to hypothesize that the long-term increase in salinity would produce an increasingly weak vertical salinity gradient in the

water column that would favour the vertical fluxes between the bottom and surface waters. Therefore the observed changes of oxygen and remineralized nutrient inputs in the surface waters could be due to an increasing footprint of benthic respiration, that has a major importance in the net ecosystem metabolism of this coastal region (Alonso-Pérez et al., 2015). This hypothesis would also explain the intense acidification in the inner waters in spite of growing alkalinity buffering."

Line 614: correlation coefficient should be indicated even though it is contained in the Figure. **The coefficient of determination of 0.52 between pH and AOU was included in the sentence as follows: "The relationship between pH and AOU (Fig. 5b) showed an inverse linear correlation of $-399 \pm 5 \mu\text{mol kg}^{-1}$ and a coefficient of determination (r-squared) of 0.52."**

Line 606-619: To me, the entire paragraph is the highlight of the paper, as it evidences the relevance of the dataset and gives insight on the processes responsible for the mean decreasing pH trend found in the area. In my opinion, this finding gets somehow diluted between the other results when it should be emphasized by the authors.

Line 623: photosynthesis of organic matter???

The sentence has been corrected. The new text is: "... the oxygen concentration in addition to the remineralization of the organic matter and the photosynthesis is conditioned by changes..."

Line 627: I would finish as: Hence, the analysis performed over the database presented here confirms that the future evolution of ocean acidification in this productive region is likely to depend on both the potential CO₂ increase in the atmosphere and other long-term changes (of natural and/or anthropogenic origin) affecting the seawater's carbonate system.

Following the suggestion, the new manuscript finish as follows: "Therefore, the long-term drop in seawater pH measurements estimated from the ARIOS database presented here confirms that the future evolution of ocean acidification in this productive region is likely to depend on both the potential CO₂ increase in the atmosphere and other long-term changes (of natural and/or anthropogenic origin) affecting the seawater's carbonate system."

Interactive comment on “ARIOS: An acidification ocean database for the Iberian Upwelling Ecosystem (1976–2018)” by Xosé Antonio Padin et al.

Michele Giani (Referee)

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Received and published: 5 June 2020

The data set surely will be relevant also for the future understanding of the interactive effects on acidification in the Iberian Upwelling System of coastal processes and global changes. However there is a need of a careful revision of the data set and of improvements in the ms. In the abstract the author give acidification rates ranging from -0.0016 to -0.0032 pH units/yr whereas in the ms (L. 566) they give a -0.0039 pH units/yr for the inner waters. This discrepancy should be resolved. **The discrepancy has been corrected. The correct acidification rate is -0.0039 pH units yr⁻¹.**

As the estimated acidification rate is higher than the average ocean acidification it would be important to discuss the potential effects of the gaps in the times series some spanning also 7 consecutive years.

It would be relevant to compare the trends on periods without long gaps of data, which could strongly affect the slope of the trend.

Following the reviewer's recommendation, the long-term trend in acidification was partially assessed in those periods best analysed, namely 1981-1998 and 2001-2009. The main conclusion of the results obtained was the loss of statistical significance of acidification rates. In fact, every interannual changes of deseasonalized pH time serie during the period 1981-1998 showed p-values > 0.05 while only the surface waters of the ocean and the continental shelf with pH trends of -0.0040 ± 0.0006 yr⁻¹ and -0.0140 ± 0.0017 yr⁻¹, respectively, were significant. In any case, we understand that the inclusion of the data shown in Table 2 beyond an intense study on the evolution of acidification in an upwelling system want to highlight the fact that direct observations in the Iberian Upwelling System over the last 40 years indicate that pH is decreasing.

A comparison with other articles reporting ranges for coastal acidification trends could be interesting for improving the discussion of results.

Other measures of acidification rates in near-shore areas such as the well-known ESTOC or CARIACO stations show year-on-year trends of -0.0018 ± 0.0001 yr⁻¹ and -0.0025 ± 0.0004 yr⁻¹, respectively. Other acidification rates like the one found at the DYFAMED station in the Mediterranean Sea was -0.0028 ± 0.0003 yr⁻¹ (<https://hal.sorbonne-universite.fr/hal-01534516/document>). In general these rates of change are in the range of the information obtained from the ARIOS database as well as the acidification rates used as reference by Lauvset et al. 2015.

Coastal acidification includes local changes in water chemistry from changes in temperature or salinity, high nutrient inputs or inputs from freshwater rivers, or excess nutrient runoff (e.g., nitrogen and organic carbon). An ecosystem's ability to cope with acidification is influenced by the number of local stresses it faces. Some ecosystems may be more resilient to ocean acidification by minimizing biogeochemical changes. Because of these characteristics, the impact of acidification between coastal areas is difficult to compare, and extensive work is needed to analyse the similarities and differences between coastal areas.

In the section “Cruises in the 2000s and recent years” (L. 261- 265), the information about each cruise is given but not always the months and years are given. I strongly suggest to provide similar information for each cruise or to refer to a more specific table where the time span of each cruise is given. In particular, regarding the last ARIOS project it is not clear in which months was carried out. **The number of days between the start and the end of the sampling period of each project was included in the Table 1.**

It would be important if the authors could be the precision for the temperature and salinity measurements in the period 1976-1984.

The information has been included in the new version of the manuscript. Namely, the precision of the temperature and salinity measurements in that period was 0.02°C and 0.005 respectively.

For chlorophyll measurements, as different filters were used, could the authors provide an estimation of the pore size given and of potential effects of the change. The indication of the volume filtered (range) could be also important if available.

The 6 cm Schleicher and Scholl and Whatman GF/F 2.5 cm filters are made of glass fibre and have a similar nominal pore of 0.7 micron. The Schleicher and Scholl filters were used for chlorophyll measurements with spectrophotometers that needed a larger sample volume because of their lower detection sensitivity while the Whatman filters always for fluoremetric measurements.

Regarding the adopted Quality control procedure (L.370-L.382) it would be useful if the authors could provide a synthetic information on the first and second level of the quality control cited in this section.

Regarding the presented ARIOS data set there are some corrections to the data that the authors should consider as there are many negative concentrations for nitrites (n=4), nitrates (n=16), ammonia (n=13), and chlorophyll a (n=2).

The values of any parameter lower than the precision of those measurements as well as the negative measurements were replaced by zero and their corresponding flag by 6.

There are concentrations for nutrients and chlorophyll a in the range of 10^{-3} to $<10^{-7}$ that should be correctly reported, presumably, as less than the detection limits given in the methods, and properly flagged with QF = 6.

The values of any parameter lower than the precision of those measurements as well as the negative measurements were replaced by zero and their corresponding flag by 6.

For all nutrients there are many values equal to 0 with QF= 2, these values presumably are below the detection limits and should be flagged with QF=6.

The values of any parameter lower than the precision of those measurements as well as the negative measurements were replaced by zero and their corresponding flag by 6.

There are three in situ pH values in the range 7-7.6 that should be checked to evaluate if they can be considered reliable or doubtful.

These 3 measurements of pH were flagged = 3

Below some minor comments are given:

L.52 I suggest to correct as follows: to help a sustainable management of the : : :

The suggestion has been included in the new version of the manuscript

L. 290 I think that “as well“ should be omitted and I suggest to substitute “to create” with “to record”.

The suggestion has been included in the new version of the manuscript

L. 306 “pH value to do so”: unclear.

The sentence was rewritten as follows: "the seawater pH measurements were determined with a spectrophotometric method following Clayton and Byrne (1993), subsequently adding 0.0047 to the pH value according to DelValls and Dickson (1998)."

L325-327 it is unclear if for the titration the HCl concentration was 0.1 or 0.13 M.

0.1M

L.338, L.341 change the conjugation of verb to the past.

The suggestion has been included in the new version of the manuscript

L.351 “Cl₂Mn” should be written as “MnCl₂”.

The suggestion has been included in the new version of the manuscript

L.352 change the conjugation of verb to the past.

The suggestion has been included in the new version of the manuscript

L. 382 check the year in the reference list is 2010.

The correct year is 2010. The manuscript has been corrected.

L. 506 “large seasonable variability“ change with “ large seasonal variability”.

The suggestion has been included in the new version of the manuscript

L. 523 “and so it would” I suggest changing as: “and therefore it would”.

The suggestion has been included in the new version of the manuscript

I suggest changing the yellow colour in Figure 1, as on the printed version is not clearly visible.

The suggestion has been included in the new version of the manuscript. The new line is a green line.

To enhance the readability of Figure 2, I suggest to enlarge them or to split the figure in two.

In Figures 3 and 4, for Salinity, I suggest omitting “psu” as it is not a real measurement unit, but a conductivity ratio. I suggest to indicate the pH is on the total scale similarly to figure 5.

The pH on the total scale was indicated as pH_T in every figure. The psu is mentioned to indicate salinity is reported in practical salinity scale. The new Figures 3 and 4 are attached to the reply to the referee.

There is the need to correct the units of oxygen in micromole kg⁻¹ in figure 4.

The units of oxygen has been corrected.

The subscript of pH_T in the Figure 5 is not well readable and should be explained in the caption.

I suggest, all over the figures, to indicate the pH as pHT for clarifying that is expressed on the total scale, moreover, to increase the readability, I suggest to enlarge or split the two graphs.

TABLE 1. According to the data set, the ARIOS cruises were carried out during different months of 2018 and not in one month of 2017. As some cruises/projects could span over more month perhaps it would be better to provide the period of the study instead of a single date.

The number of days (#d) between the start and the end of sampling period of each project was included in the Table 1 that is attached at the end of the reply to the referee.

1 **ARIOS: a database for ocean acidification assessment in the Iberian Upwelling**
2 **System (1976 - 2018).**

3
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7
8 **1. Abstract**

9 A data product of 17,653 discrete samples from 3,343 oceanographic stations
10 combining measurements of pH, alkalinity and other biogeochemical parameters off the
11 North-western Iberian Peninsula from June 1976 to September 2018 is presented in this
12 study. The oceanography cruises funded by 24 projects were primarily carried out in the
13 *Ría de Vigo* coastal inlet, but also in an area ranging from the Bay of Biscay to the
14 Portuguese coast. The robust seasonal cycles and long-term trends were only calculated
15 along a longitudinal section, gathering data from the coastal and oceanic zone of the
16 Iberian Upwelling System. The pH in the surface waters of these separated regions,
17 which were highly variable due to intense photosynthesis and the remineralization of
18 organic matter, showed an interannual acidification ranging from -0.0016 yr⁻¹ to -0.0039
19 yr⁻¹ that grew towards the coastline. This result is obtained despite the buffering
20 capacity increasing in the coastal waters further inland as shown by the increase in
21 alkalinity by 1.1±0.7 μmol kg⁻¹ yr⁻¹ and 2.6±1.0 μmol kg⁻¹ yr⁻¹ in the inner and outer *Ría*
22 *de Vigo* respectively, driven by interannual changes in the surface salinity of
23 0.0193±0.0056 psu yr⁻¹ and 0.0426±0.016 psu yr⁻¹ respectively. The loss of the vertical
24 salinity gradient in the long-term trend in the inner ria was consistent with other
25 significant biogeochemical changes such as a lower oxygen concentration and
26 fertilization of the surface waters. These findings seem to be related to a growing
27 footprint of sediment remineralization of organic matter in the surface layer of a more
28 homogeneous water column.

29 Data are available at: <http://dx.doi.org/10.20350/digitalCSIC/12498> (Pérez et al., 2020).

30
31 **2. Introduction**

32 CO₂ emissions of anthropogenic origin (fossil fuels, land use and cement
33 manufacturing) into the atmosphere are the main cause behind the warming of the Earth
34 due to the greenhouse effect (IPCC, 2013). Given the constant exchange of gases

TONI PADIN 25/6/2020 14:25

Eliminado: ARIOS: An acidification
ocean database for the Iberian Upwelling
Ecosystem (1976 - 2018) .

TONI PADIN 25/6/2020 11:15

Eliminado: 0032

39 through the air-sea interface, the oceanic reservoir plays a key role as a sink for about
40 31% of anthropogenic CO₂ emissions (Sabine et al., 2004), controlling the partial
41 pressure of carbon dioxide in the atmosphere and regulating global temperatures.

42

43 The CO₂ uptake by the oceans produces changes in the inorganic carbon system in spite
44 of being partially dampened by the seawater buffering capacity. This ability of seawater
45 to withdraw anthropogenic CO₂ becomes more limited as more CO₂ is absorbed, which
46 will make it difficult to stabilize atmospheric CO₂ in the future (Orr et al., 2009). The
47 gradual absorption of atmospheric CO₂ by the oceans decreases seawater pH, causing
48 ocean acidification, which conditions the buffering capacity of seawater and in turn the
49 exchange of CO₂ between the ocean and the atmosphere (Caldeira and Wickett, 2003;
50 Raven et al., 2005). This effect of CO₂ absorption, which is known as ocean
51 acidification, conditions the buffering capacity of seawater and to some extent the
52 exchange of CO₂ between the ocean and the atmosphere. The Intergovernmental
53 Oceanographic Commission of the United Nations identified the chemical change in
54 seawater brought about by ocean acidification as an indicator of a stressor on marine
55 ecosystems with a negative impact on socio-economic activities such as fishing and
56 shellfish farming. Hence, it was necessary for the oceanography community to observe
57 and gather data about pH and other parameters of the marine carbon system to conduct
58 accurate measurements of pH and ancillary parameters and provide data products to
59 help a sustainable management of the marine resources. The effect of ocean
60 acidification on marine ecosystems has stimulated impetus in the international
61 community for gathering high quality time-series measurements of the marine inorganic
62 carbon system (Hofmann et al., 2011; Andersson and MacKenzie, 2012; McElhany and
63 Busch, 2013; Takeshita et al., 2015; Wahl et al., 2016) and for predicting the future
64 evolution of the pH caused by climate change.

65

66 The threat for oceanic acidification of marine ecosystems is especially significant in
67 regions like coastal upwelling areas, which are more sensitive and appear to respond
68 faster to anthropogenic perturbations (Feely et al., 2008; Gruber et al., 2012; Lachkar,
69 2014; Hauri et al., 2013). These ecosystems are characteristic for their complex physical
70 and biogeochemical interactions and for sustaining enormous biological productivity
71 and productive fisheries (Pauly and Christensen, 1995; Hauri et al., 2009). The
72 photosynthetic activity in these regions is also an important mechanism for the seawater

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Eliminado: the rapid increase of CO₂ in the atmosphere decreases the ocean's pH

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Eliminado: produce global and regional data products in order to help sustainably manage the ocean's resources. .

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82 | CO₂ uptake, converting most of these areas into atmospheric CO₂ sinks (Pérez et al.,
83 | 1999; Cobo-Viveros et al., 2013). However, the high physical/chemical variability in
84 | short temporal and spatial scales of upwelling systems and the lack of regular sampling
85 | in these waters prevents a complete picture of the acidification of these ecosystems.

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Eliminado: However, the great spatial and temporal variability has been widely spaced and intermittent over time, preventing a complete view from being obtained of ocean acidification in the upwelling system.

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86 |
87 | In the Iberian Upwelling System, the researchers of the Instituto de Investigaciones
88 | Mariñas (IIM-CSIC) since 1976 commenced accurate measurements of marine
89 | inorganic carbon system and associated parameters. As a result, a collection of pH
90 | observations and ancillary biogeochemical information along the Galicia coast (40°N
91 | and 45°N, 11°W) has been gathered under the framework of different projects over the
92 | past 40 years. The current database, hereinafter called ARIOS (Acidification in the rias
93 | and the Iberian continental shelf) database, holds biogeochemical information from
94 | 3,357 oceanographic stations, giving 17,653 discrete samples. This unique collection is
95 | a starting point for evaluating the ocean acidification in the Iberian Upwelling System
96 | characterized by intense biogeochemical interactions as an observation-based analysis,
97 | or for use as inputs in a coupled physical-biogeochemical model to disentangle these
98 | interactions at the ecosystem level.

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Eliminado: The effects of ocean acidification on marine ecosystems has awoken the interest of the international community and has stimulated global impetus for gathering high quality time-series measurements of the marine inorganic carbon system (Hofmann et al., 2011; Andersson and MacKenzie, 2012; McElhany and Busch, 2013; Takeshita et al., 2015; Wahl et al., 2016) and for predicting the future evolution of the pH caused by climate change. Researchers at the *Instituto de Investigaciones Mariñas* (IIM-CSIC) in Vigo (Spain) have already started this task, measuring the marine inorganic carbon system and associated parameters on the Galicia coast in the northwest of the Iberian Peninsula, since 1976. These biogeochemical changes, including the pH variation, have been gathered via different projects with particular objectives between 40°N and 45°N, 11°W and the Galician coast over the past 40 years. The current database, hereinafter called ARIOS (Acidification in the rias and the Iberian continental shelf)

100 | 3. Data provenance

101 | 3.1. Data spatial coverage

102 | The main characteristic of the Galician coastline, located in the north-west of the
103 | Iberian Peninsula, is the *Rías Baixas*, four long coastal estuaries or rias (>2.5 km³)
104 | between 42°N and 43°N (Fig. 1). The water exchange between the *Rías Baixas* and open
105 | waters is drastically affected by the coastal wind pattern as part of the Canary Current
106 | Upwelling System (Wooster et al., 1976; Fraga 1981; Aristegui et al., 2004). Under the
107 | predominance of northeasterly winds (Blanton et al., 1984) during spring-summer, the
108 | surface offshore transport of surface waters leads to a rising cold, nutrient-rich, deep
109 | water mass called the Eastern North Atlantic Central Water (ENACW) (Ríos et al.,
110 | 1992). Under these conditions, the *Rías Baixas* act as an extension of the continental
111 | shelf (Rosón et al., 1995; Souto et al., 2003; Gilcoto et al., 2017), where upwelling
112 | filaments extending westward export primary production from the coast into the ocean
113 | (Álvarez-Salgado et al., 2001). In the opposite direction, the prevalence of northward
114 | winds (Blanton et al., 1984) moves the surface waters towards the coast, where they
115 | accumulate, sink and thus isolate the coast. This process, known as downwelling, is

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147 | typical during the autumn-winter along with other characteristics such as the warm,
148 | salty waters from the Iberian Poleward Current (IPC) of subtropical origin (Fraga et al.,
149 | 1982; Alvarez-Salgado et al., 2006) that flows constrained to the Iberian shelf break
150 | (Frouin et al., 1990). The run-off from local rivers also contributes to the presence of
151 | river plumes over the shelf (Otero et al., 2008). These hydrodynamic conditions, the
152 | meteorological forcings and the alternation of periods of upwelling and downwelling
153 | (Álvarez, 1999; Gago et al., 2003c; Cobo-Viveros et al., 2013) stimulate the
154 | development of intense primary production and high rates of recycling and downward
155 | carbon export (Alonso-Pérez and Castro, 2014). The result of this biogeochemical
156 | variability in terms of air-sea CO₂ exchange is that the surface waters act as a net CO₂
157 | sink that is especially intense and variable over the shelf compared to offshore or in the
158 | inner *Rías Baixas* (Padin et al., 2010).

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160 | Besides, the short-term and seasonal variability, significant changes in the long-term
161 | scale have been reported in this region. In addition to changes such as the weakening
162 | and shortening of the upwelling events (Lemos and Sansó, 2006; Pérez et al., 2010;
163 | Alvarez-Salgado et al., 2009), the warming (González-Pola et al., 2005; Pérez et al.,
164 | 2010) and changes in the composition of phytoplankton (Bode et al., 2009; Pérez et al.,
165 | 2010), the acidification in the first 700 metres for the geographical area from the Iberian
166 | Peninsula to the 20° W meridian and from 36°N to 43°N has also been observed at a rate
167 | of -0.0164 pH units per decade (Ríos et al., 2001; Castro et al. 2009).

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169 | 3.2. Distribution of sampling

170 | According to the type of region under study, different areas were identified in order to
171 | classify the measurements gathered in the oceanographic cruises (Fig. 1). The latitude
172 | of 43°N where Cape Finisterre is located was used as the dividing line between northern
173 | and southern waters. Subsequently, a criterion of depth also split the waters to the north
174 | of 43°N into north oceanic (below 250 m), north shelf (between 205 m and 75 m) and
175 | north coast (75 m to the surface). The southern shelf waters were divided by latitude
176 | 42°N into Portuguese and the *Rías Baixas* (RB) shelves, whereas the shallower waters
177 | were identified by the main rias, where three different zones were defined using
178 | longitude boundaries (outer, middle and inner) according to Gago et al. (2003c) in the
179 | *Ría de Vigo*, and just two zones in the other rias (*Ría de Pontevedra*, *Ría de Arousa*, *Ría*
180 | *de Muros*). Southern waters between the isobath at 75 metres and the mouth of the

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188 estuaries were identified as the Portuguese and RB coast.

189

190 The discrete measurements gathered in the ARIOS dataset were mainly found in
191 different regions' waters around 42°N latitude (Fig. 1; Fig. 2a), especially in the outer
192 and middle areas of the Ría de Vigo, which accounted for 15% and 21% of the total
193 measurements respectively due to the proximity to the Instituto de Investigaciones
194 Mariñas (IIM-CSIC). Most of the measurements (85%) carried out by many of these
195 cruises to study the coastal ecosystems concentrated on shallow waters between the
196 seawater surface and 75 metres in depth (Fig. 2b). Although waters below 4,900 metres
197 deep were also sampled, observations below 900 metres only account for 1% of the
198 ARIOS database.

199

200 The observations made over more than 40 years in every region of the ARIOS database
201 were irregular on both an interannual and seasonal scale (Fig. 2a). The period of most
202 sampling activity was the 80s and 90s, whereas samples were especially scarce in the
203 early 2010s. On a seasonal scale, summer and autumn were the preferred seasons to
204 address the different research purposes, with 37% and 36% of the total samples
205 respectively. The observations taken during less favourable winter conditions,
206 especially aboard the coastal vessels usually available, only accounted for the 10% of
207 the ARIOS database.

208

209 **3.3. Data sources**

210 The ARIOS database is a compilation of biogeochemical properties with discrete
211 measurements of temperature, salinity, oxygen, nutrients, alkalinity, pH and chlorophyll
212 that were sampled in waters off the northwest of the Iberian Peninsula from 1976 to
213 2018 and measured by IIM-CSIC (Table 1). This data collection is part of the research
214 by 24 projects and oceanographic cruises conducted in response to different aims. The
215 different sampling strategies built up an irregular biogeochemical database whose
216 particular frequency and spatial coverage is shown in Figure 2.

217

218 The contribution to the ARIOS database from the oceanographic cruises and projects
219 over the different decades is described below.

220

221 **Cruises in the 70s and 80s:**

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225 The first three cruises were carried out over three periods (1976, 1981-1983 and 1983-
226 1984), sampling the *Ría de Vigo*. These cruises were designed to provide environmental
227 information (upwelling events, estuarine circulation, continental inputs, etc.) for
228 research into the biology of some fish species. They measured identical parameters in
229 the Vigo estuary but at different stations and frequency.

230

231 In the summer of 1984, the *Galicia VIII* cruise studied the summer upwelling events
232 occurring on the contact front between the two ENACW water masses off Cape
233 Finisterre from short sections perpendicular to the Galician coast with 85 stations
234 offshore and 35 stations over the shelf. This cruise marked a milestone in the
235 oceanographic research of IIM-CSIC because it was the first time that the parameters of
236 the carbon system were measured on-board in offshore waters. Moreover,
237 measurements of a particular station on the shelf break with a bottom depth of 600
238 metres were taken every two days for a month, including two-day continuous
239 samplings.

240

241 Two years later, the Ria de Vigo 1986 sampled along the main axis of the Ria de Vigo
242 in 7 monthly repetitions during the first half of the year in which the primary production
243 and the organic matter exchange between the estuary and the shelf was studied in
244 relation to the hydrographic regime. Shortly afterwards, the same topic was also
245 researched by the Galicia IX project in September and October 1986 from 145 stations,
246 50 of which were coastal and 80 located in ocean waters (Prego et al., 1990).

247

248 The following year, the 1987 Provigo project (Nogueira et al., 1997) initiated a periodic
249 study from a fixed site (42°14.5'N, 8°45.8'W) located in the main channel in the middle
250 zone of the *Ría de Vigo*. This oceanographic station was selected as suitable for
251 evaluating the main processes that occur in the inner ria associated with external forcing
252 changes (Rios, 1992; Figueiras et al., 1994). Although the Provigo project finished in
253 1996, the fixed station was repeatedly included in subsequent cruises, extending the
254 time series at this location until today, when it is currently sampled every week by
255 INTECMAR (www.intecmar.gal). An example of the subsequent sampling repetition of
256 this station occurred the following year when one of the three stations in the Vigo
257 estuary in the Luna 1988 project (Fraga et al., 1992) took a sample every two weeks to

258 study the environmental control over the phytoplankton populations throughout an
259 annual cycle (February 1988 - February 1989).

260

261 At the end of 80s, the carbon system monitoring by the IIM-CSIC was extended to the
262 *Ría de Arousa* throughout 1989 (Álvarez-Salgado et al., 1993; Perez et al., 2000) in
263 order to learn the effect of upwelling on the water circulation pattern, community
264 production and the fluxes and net budgets of biogenic constituents in this ria with the
265 highest mussel production in Europe. For 5 months, 11 stations' samples were repeated
266 twice a week in the ria that is the most productive, housing intense cultivation of
267 mussels on rafts (Blanton et al., 1984).

268

269 **Cruises of the 90s:**

270 In the first half of this decade, studying the phytoplankton communities was the
271 oceanographic cruises' most relevant aim, concentrating particularly on harmful algae
272 blooms. The hydrodynamic and biogeochemical conditions controlling the growth,
273 development and migration of the phytoplankton were analysed both in the interior of
274 the estuary and on the continental shelf.

275

276 For five days in September, the 1990 *Ría de Vigo* cruise (Figueiras et al., 1994) sampled
277 five stations distributed along the longitudinal axis of the ria and one at the northern
278 mouth. The next year, the cruise Galicia XI was carried out in May, sampling at 39
279 stations along eight transects perpendicular to the coastline; and Galicia XII (Alvarez-
280 Salgado et al., 1998, 2002, 2003; Castro et al., 1994) in September, sampling at 37
281 oceanic stations and 7 coastal stations.

282

283 The *Ría de Vigo* cruise in 1993-4 (Miguez et al., 2001), with four stations using 24
284 repetitions with a CTD-SBE25, investigated the hydrodynamic and biogeochemical
285 effect on the evolution of phytoplankton communities in the *Ría de Vigo*. Six samples
286 were taken in approximately two weeks corresponding to two different periods
287 (27 September to 8 October 1993, and 6 March to 24 March 1994).

288

289 *Ría de Vigo* 1994-95 (Alvarez et al., 1999; Doval et al., 1998, 1997a, 1997b) and *Ría de*
290 *Vigo* 1997 (Gago et al., 2003a, b, c) were two cruises that took place in the second half
291 of the decade. These campaigns' objective was no longer the ecology of the plankton,

292 but the factors behind the variation of the carbon pools during the upwelling and
293 downwelling events along the central axis of the *Ría of Vigo*. During the 1997 cruises
294 on board the *B/O Mytilus*, a systematic observation of the pCO₂ was carried out for the
295 first time in Spanish coastal waters, using an autonomous continuous system with
296 additional measurements of temperature, salinity and chlorophyll.

297

298 **Cruises in the 2000s and recent years:**

299 After a period of poor sampling at the end of 90s, the first decade of the 21st century
300 gave new impetus to biogeochemical monitoring of Galician waters. As shown below,
301 several projects dealt with various objectives, focussing on particular issues in the
302 dynamics of these waters:

303

304 The DYBAGA project (Galician Platform's Annual Dynamics and Biochemistry: short-
305 scale variation) (Álvarez-Salgado et al., 2006; Castro et al., 2006; Nieto-Cid et al.,
306 2004) analysed the phenomena of upwelling and downwelling in the Galician shelf
307 opposite the *Ría de Vigo* weekly and their impact on the different biogeochemical and
308 carbon system variables including organic dissolved matter. Three stations were
309 sampled weekly from May 2001 to April 2002 between the shelf break (1,200 m deep)
310 to the middle of the *Ría de Vigo* (45 m deep).

311

312 The REMODA (Reactivity of dissolved organic matter in a coastal upwelling system)
313 (Álvarez-Salgado et al., 2005; Piedracoba et al., 2005; Nieto-Cid et al., 2006) project
314 concentrated on learning the origin and destination of dissolved organic matter in the
315 *Ría de Vigo* as well. Three stations along the main axis of the *Ría de Vigo*, including the
316 fixed station as the central one, took samples with short (3-4 days) and seasonal time
317 scales.

318

319 The FLUVBE project (Coupling of benthic and pelagic fluxes in the *Ría de Vigo*) added
320 to knowledge about the productivity and the benthic fluxes of oxygen and inorganic
321 nutrients in the *Ría de Vigo* from 16 oceanographic surveys with four stations between
322 April 2004 and January 2005.

323

324 The ZOTRACOS project studied the biogeochemical and hydrodynamic
325 characterization of the coastal transition zone in NW Spain during the downwelling
326 period (Teira et al., 2009).

327

328 The CRIA (Circulation in a RIA) (Barton et al., 2019) project examined the layout of
329 the two-layer circulation and propagation of upwelled and downwelled waters in order
330 to estimate the flushing and vertical velocities in the *Ría de Vigo* in repeated
331 hydrographic surveys between September 2006 and June 2007 (Barton et al., 2015,
332 2016; Alonso-Perez and Castro, 2014; Alonso-Perez et al., 2010; Alonso-Perez et al.,
333 2015).

334

335 The RAFTING project (Impact of mussel raft cultivation on the benthic-pelagic
336 coupling in a Galician ria) (Frojan et al., 2018; Frojan et al., 2016; Froján et al., 2014)
337 assessed for the first time how mussel cultivation influences the quality of particular
338 organic carbon fluxes in the *Ría de Vigo*. Over the four seasons, two stations were
339 visited every two to three days during each period, meaning 24 oceanographic cruises in
340 2007 and 2008.

341

342 The CAIBEX (Continental shelf-ocean exchanges in the marine ecosystem of the
343 Canary Islands-Iberian Peninsula) (Villaceros-Robineau et al., 2019) project compared
344 the dynamics and biogeochemical activity between the coastal zone and the adjacent
345 ocean in the study zone during the summer upwelling events. As part of the CAIBEX
346 project, a mooring at the LOCO (Laboratory of Ocean and Coastal Observation)
347 (Zuñiga et al., 2016, 2017) site located on the continental shelf was deployed and visited
348 monthly for one year to monitor the vertical profiles of biogeochemical variables.

349

350 After these projects were completed in 2009, new measurements were not provided
351 until 2018. The aim of the ARIOS project (Acidification in the rias and on the Iberian
352 continental shelf) was to evaluate the impact of ocean acidification and learn about
353 potential impacts on the mussels and their adaptation (Lassoued et al., 2019) to the new
354 climate change.

355

356 | **3.4. Methods**

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358 To assess of the level of acidification in the ocean adjacent to the Galician coast,
359 variables of the carbon system (pH and alkalinity), nutrient concentration, dissolved
360 oxygen, chlorophyll-a, salinity and temperature were measured in each cruise. The
361 variables measured in each oceanographic cruise gathered in the ARIOS dataset are
362 shown in Table 1. The main changes in the material and methods throughout these years
363 are detailed below.

364

365 **T-S measurements**

366 Temperatures from 1976 to 1984 were measured using a Wallace and Tiernan
367 bathythermograph. Reversing thermometers that had a precision of 0.02°C were used,
368 attached to the water samplers between 1984 and 1990, correcting the temperature
369 between the protected and unprotected thermometers according to Anderson (1974).
370 During those years, the depth was calculated from the thermometric readings, rounding
371 the result off to the nearest ten. After 1990, different models of CTD instruments that
372 measured the seawater temperature with a precision of 0.002°C were used to obtain the
373 thermohaline profile.

374

375 The first measurements of salinity were determined with a Plessey Environmental
376 Systems 6230N inductive salinometer calibrated with normal IAPSO water and
377 calculated from the equations given in the NIO and UNESCO International
378 Oceanographic Tables (1981). The precision of these salinity measurements was 0.005
379 psu. After using this equipment, the salinity was determined with an AUTOSAL 8400A
380 inductive salinometer calibrated with normal IAPSO water whose estimated analytical
381 error was 0.003 psu, using the equations given by UNESCO (1981). CTDs began to be
382 used in 1990 to record the vertical salinity profiles, calibrated using the salinity samples,
383 whose possible deviations in the measurements were estimated from the discrete
384 measurements from the AUTOSAL salinometer.

385

386 **pH measurements**

387 The pH measurements were originally taken with a Metrohm E-510 pH meter with a
388 glass electrode and a Ag/ClAg reference one calibrated with 7.413 NBS buffer. All pH
389 values were converted to values at 15 °C using the temperature correction from the
390 Buch and Nynas tables published by Barnes (1959). In 1984, the method was modified
391 and the temperature normalization was carried out following Pérez and Fraga (1987b).

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394 Two years later, the measurement equipment was the Metrohm E-654 pH meter with an
395 Orion 81-04 Ross combined glass electrode, with the pH converted to the SWS scale
396 using the hydrogen activity coefficient given by Mehrbach et al. (1973) at 25°C with the
397 parameterization given by Pérez and Fraga (1987b). The error in this potentiometric
398 method was 0.010. In 2001, the seawater pH measurements were determined with a
399 spectrophotometric method following Clayton and Byrne (1993), subsequently adding
400 | 0.0047 to the pH value according to DelValls and Dickson (1998). The precision of the
401 spectrophotometric measurements was 0.003 pH units.

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402
403 The pH values were reported on total pH scale at 0 dbar of pressure and both at 25°C
404 | and in situ temperature (pH_T) following the same procedure of GLODAP v2 (Olsen et
405 al., 2019). A total of 12,220 measurements of pH on NBS scale were converted to the
406 total scale using CO2SYS (Lewis and Wallace, 1998) for MATLAB (van Heuven et al.,
407 2011) with pH and total alkalinity as inputs. The conversion was conducted with the
408 carbonate dissociation constants of Lueker et al. (2000) and the borate-to-salinity ratio
409 of Uppström (1974). Whenever total alkalinity data were missing, these values were
410 approximated as 66 times salinity that is the mean ratio between the total alkalinity and
411 the salinity of every in situ measurements compiled in the ARIOS database. Data for
412 phosphate and silicate are also needed and were, whenever missing, a constant values of
413 $10 \mu\text{mol kg}^{-1}$ for silicate and $1 \mu\text{mol kg}^{-1}$ for phosphate were used. These
414 approximations were tested on 8,296 samples with complete biogeochemical
415 information showing a bias of less than 0.0004 pH units for 99.95% of the samples.

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416 417 **Alkalinity measurements**

418 The seawater alkalinity was measured for the first time in 1981 by potentiometric
419 | titration with HCl 0.1M, at final pH 4.44 following Pérez and Fraga (1987a) with an
420 analytical error of $2 \mu\text{mol kg}^{-1}$ and a precision of 0.1%. Sodium tetraborate decahydrate
421 (Borax, $\text{Na}_2\text{B}_4\text{O}_7 \cdot 10\text{H}_2\text{O}$, Merck p.a.) was used for standardizing the HCl (0.13 M). The
422 pH measurements were carried out with a combined glass electrode (Metrohm E-121)
423 with Ag/AgCl (KCl 3M) as the reference. The pH was calibrated using the NBS buffers
424 assuming the theoretical slope. As of 2001, the accuracy of alkalinity measurements
425 was determined using samples of certified reference material (CRM) provided by Dr. A.
426 Dickson, University of California, improving the precision to $\pm 1.4 \text{ mol kg}^{-1}$ and an

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430 accuracy of <0.1% recently established by (Ríos and Pérez, 1999) from cross-
431 calculation with measured Certified Reference Materials (Dickson et al., 2007).

432

433 Nutrient measurements

434 Except for the Galicia cruises (Table 1), in which nutrient samples were analysed on
435 board, samples were kept in the dark and cold (4°C) after collection for further analyses
436 in the shore based laboratory. Nutrient concentration was determined by a flow-
437 segmented autoanalyzer (Technicon AAII and Alpkem after 1995) as described in
438 Strickland and Parsons (1968) with the particularity that the reduction of nitrate to
439 nitrite with Cd column was done using a citrate buffer according to Mouriño and
440 Fraga's modification (1985). Phosphates and silicates were measured following
441 Grasshoff (1983), and ammonium as described by Grasshoff and Johannsen (1972).
442 This method was maintained in the subsequent cruises, achieving a precision of 0.02
443 $\mu\text{mol/kg}$ for nitrite, 0.1 $\mu\text{mol/kg}$ for nitrate, 0.05 $\mu\text{mol kg}^{-1}$ for ammonium and silicate,
444 and 0.01 $\mu\text{mol/kg}$ for phosphate.

445

446 Oxygen measurements

447 The dissolved oxygen was determined via the Winkler titration method for the first time
448 in 1981 following the procedure published later by Culberson et al. (1991). The oxygen
449 concentration in the samples in this method was fixed with MnCl₂ and NaOH/NaI,
450 which were kept in the dark until analysis in the laboratory 12-24 hours later. The
451 measurements were made by titration of iodine with thiosulfate using an automatic
452 titrator. During the 80s and early 90s, the titration was carried out with Metrohm
453 instruments (E-425 or E-473), which had an analytical error of 1 $\mu\text{mol kg}^{-1}$. The oxygen
454 concentration after 1997 was estimated using a Titrino 720 (Metrohm) analyser with an
455 accuracy of 0.5 $\mu\text{mol kg}^{-1}$.

456

457 Chlorophyll measurements

458 The chlorophyll-a values were measured following SCOR-UNESCO (1966) using a
459 6 cm diameter Schleicher and Scholl 602eh filter covered with magnesium carbonate.
460 The absorption was measured in 1 cm optical path cuvettes using a Beckman DU
461 spectrophotometer. In 1984, discrete water samples of the chlorophyll-a samples were
462 filtered through Whatman GF/F₂ filters of 2.5 cm, which were preferred from then on,

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473 and measured fluorometrically following Strickland and Parsons (1972) without
474 correction for concentration by pheophytes. The fluorescence readings were carried out
475 with a Turner Designs 10,000 R fluorometer (Yentsh and Menzel, 1963) obtaining a
476 precision of 0.05 g L⁻¹.

477

478 **Quality control**

479 Every cruise gathered in Table 1 passed 1st quality control (QC1) to ensure truly
480 confident results. The GO-SHIP software for quality control of hydrographic data (Velo
481 et al., 2019) that compile several QC1 procedures was applied to ARIOS dataset. That
482 procedures consist on reviewing the property profiles and property-property plots
483 generated by that application, adequate for each variable. A quality flag was assigned to
484 each measurement available from the repository sites (Table 1). This method was
485 preferred over applying a very stringent flagging process because it is difficult to rule
486 out some extreme values associated with low salinities or that could be supported by the
487 high variability of an ecosystem characterized by an intense biological activity.
488 Nutrients and chlorophyll with values inferior to the precision were flag=6. Some very
489 low pH associated to very low salinity waters were flagged as doubtful.

490

491 The ARIOS database includes the cruise corrections for pH data of the -0.017 for the
492 Galicia VIII cruise (29GD19840711) and +0.032 for Galicia IX cruise
493 (29GD19860904) detected during the second level quality control of CARINA project
494 (Velo et al., 2010).

495

496 **4. Results**

497 Some of the most obvious results provided by the ARIOS database are shown below.
498 The purpose is to describe the environmental context and the main oceanographic
499 processes that affect the variability of these discrete measurements and offer
500 preliminary information for future detailed biogeochemical research.

501

502 **4.1. Vertical distribution**

503 The vertical profile of the temperature, salinity, pH on total scale at in situ temperature
504 (pH_T), NO₃⁻ and oxygen concentration in the ocean region between 41°N and 43°N was
505 estimated for each oceanographic station as the mean value of the depth ranges
506 described in Figure 2b. These measurements were gathered attending to the collection

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517 periods (December-February, March-May, June-August and September-November) and
518 averaged to describe winter, spring, summer and autumn respectively (Fig. 3).

519

520 The vertical distribution of the temperature (Fig. 3a) showed the presence of warmer
521 saline waters throughout the water column in winter with the exception of the surface
522 waters during summer, which showed intense heating due to the radiant solar energy.
523 Below the maximum temperature observed during the summer, cold central waters of
524 subpolar origin occupied the water columns with lower salinity (Fig. 3b). The vertical
525 variation of temperature is typical for a temperate region with relatively homogenous
526 deep water below the seasonal thermocline, reaching maximum SST values in summer
527 and autumn, and minimums in spring and winter. The winter temperature profile is
528 relatively warmer than in spring because of the presence of the IPC (Alvarez-Salgado et
529 al., 2006), which reaches a depth of 300 metres. The maximum salinity is also found in
530 winter due to the presence of the IPC, whereas the minimum values are found in autumn
531 (Fig. 3b). Below 500 metres in depth, the increase in salinity points to the presence of
532 Mediterranean Water. These differences reach a minimum at 500 metres deep, where
533 the salinity values coincided. From this depth down to 1,100 metres, the differences in
534 temperature and salinity throughout the four seasons were minimal, with the mean
535 values converging to $11.03 \pm 0.07^\circ\text{C}$ and 36.117 ± 0.009 psu, respectively (Fig. 3ab).

536

537 The vertical profiles of pH_T , NO_3^- and oxygen concentration (Fig. 3cde) also showed a
538 variation lower than 1% within this depth range with annual means of $15.2 \pm 0.1 \mu\text{mol}$
539 kg^{-1} , 8.025 ± 0.005 and $188 \pm 1 \mu\text{mol kg}^{-1}$ respectively. The pH values from a maximum
540 subsurface located at around 40 metres deep showed a clear inverse correlation with the
541 depth down to a depth of 500 metres throughout the seasonal cycle, where the annual
542 minimum value of 8.018 ± 0.005 was reached. The higher pH values could be attributed
543 to the biological reduction of CO_2 by phytoplankton activity, which brought the pH to a
544 maximum value of 8.13 to 40 meters during the spring bloom. After the intense
545 photosynthetic activity observed in surface waters during spring and summer, pH values
546 reached minimum values in the first 200 metres of depth during autumn due to
547 respiration of organic matter. However, it was at a depth of 500 metres that the
548 minimum pH values were measured in all seasons where is found the subpolar Eastern
549 North Atlantic Central Water proceeding from the northeastern cyclonic gyre (Harvey,
550 1982; Ríos et al., 1992). The influence of phytoplankton growth on biogeochemistry

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553 during spring can be also evidenced by the oxygen concentration pattern (Fig. 3e). In
554 the upper layer above 250 metres depth, spring oxygen levels exceeded those in winter,
555 whereas a decrease in oxygen concentration was found from this depth down to 1000
556 metres, possible due to enhanced respiration from cascading organic matter. The impact
557 on the growth of the phytoplankton community during the spring was also evident,
558 judging by the oxygen concentration. So, in the upper waters the spring oxygen
559 concentration values exceeded those of the winter values, while oxygen consumption
560 was found from a depth of 300 metres to 1,000 metres due to respiration from organic
561 matter arriving from above. The minimum values for oxygen concentration throughout
562 the water column were found during summer and autumn. The nitrate concentration
563 displayed a particularly vertical distribution, growing with depth from minimum values
564 in the upper layer of the ocean region, which was practically zero during the first 50
565 metres. Below 100 metres, the nitrate concentration showed the maximum values in the
566 vertical distribution during summer and autumn coinciding with the presence of waters
567 of subpolar and subtropical origin respectively, whereas the minimum values appeared
568 in winter. Towards the bottom, the seasonal values of NO_3^- concentration were almost
569 coincident at a mean value of $15.2 \pm 0.1 \mu\text{mol kg}^{-1}$.

570

571 4.2. Seasonal cycle

572 The seasonal cycle of the biogeochemical properties (temperature, salinity, pH_T , oxygen
573 concentration, nitrate concentration and chlorophyll) in the surface waters (0 to 5
574 metres) of five geographical boxes was estimated as a monthly average previously
575 filtering values outside of two standard deviations of the mean. Five regions that were
576 located as a longitudinal transect between the inner Ría de Vigo and the ocean zone are
577 shown in Fig. 4.

578

579 In general terms, the seasonal variability of the temperature was very similar in every
580 area, ranging between 12 and 19°C (Fig. 4a). Only particular features observed on a
581 short-term scale as in the examples below differ between each region. The warmer
582 waters were usually found in the oceanic zone, reaching a maximum monthly averaged
583 temperature of 18.6°C in September, while the coldest surface waters of 12.6°C were
584 located in the inner stations closer to the mouth of the *Ría de Vigo* in January. Another
585 secondary minimum averaged temperature was also found in the shelf and the outer area
586 of the *Ría de Vigo*, which was remarkably low in August due to the entry of cold

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Eliminado: The highest values were related to the biological CO_2 drawdown, which brought the pH to a peak value of 8.13 at 40 metres deep during the spring bloom. Underneath this intense photosynthesis activity between the surface and 100 metres, the respiration of organic matter took the pH to lower values than those measured in winter between 200 to 500 metres, a depth at which the spring and winter values were practically equal.

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Eliminado: The seasonal cycle of the surface waters (0 to 5 metres) was estimated as a monthly filtered means, accepting only values within two standard deviations. Five regions that were located as a longitudinal transect between the inner *Ría de Vigo* and the ocean zone are shown in Fig. 4.

605 upwelled waters in the surface layer (Alvarez-Salgado 1993).

606

607 The monthly salinity averages (Fig. 4b) clearly showed significant differences between
608 the offshore and coastal waters. Sharp salinity changes were seen in the estuary during
609 winter, especially in the inner area where values lower than 28 psu were reached with
610 the arrival of continental inputs in December. The weak seasonal cycle of salinity in the
611 shelf and ocean waters showed high values in December due to the influence of warm
612 saline water from the IPC, usually located on the shelf slope even though it may even
613 enter the rias depending on the relative intensity of shelf winds and the intensity of the
614 continental runoff (Alvarez-Salgado et al., 2003). In this sense, the slight salinity
615 minimum observed in the shelf waters in March could be consequence of the offshore
616 spreading of the maximum discharges from the River Miño and Douro (Otero et al.,
617 2010) at the end of downwelling season. After this, the shelf and ocean waters showed
618 minimum values in summer due to the arrival of cooler and fresher subpolar waters
619 (Rios et al., 1992; Alvarez-Salgado et al., 2003, 2006). In August, coinciding with the
620 maximum salinity of the surface waters in the interior of the *Ría de Vigo* due to the
621 minimum river runoff, the surface waters between the inner *Ría de Vigo* and the ocean
622 region were almost homogeneous, with minimum differences in salinity of 0.2 psu.

623

624 Like salinity, there was little seasonal variability in pH in the offshore waters, but large
625 seasonal variability in coastal waters, with maximum and minimum pH values in spring
626 and autumn, respectively, and in all regions (Fig. 4c). The net balance between
627 production and respiration of organic matter and the estuarine circulation caused a
628 maximum pH of 8.19 in the outer region of the *Ría de Vigo* in May and a minimum of
629 7.96 in the inner waters in November.

630

631 The oxygen concentration (Fig. 4d) in the coastal ecosystems is also controlled by the
632 remineralization of the organic matter and photosynthetic activity of the phytoplankton
633 community, with the effect of salinity and temperature on the oxygen saturation level.
634 The variability in the oxygen concentration, like the pH distribution, showed a growing
635 seasonal amplitude towards the coastline, with maximum values in the outer and middle
636 *Ría de Vigo* and lower values in the inner waters, especially during the second half of
637 the seasonal cycle. Hence, the dissolved oxygen concentration mirrored the seasonal
638 cycle of pH, showing growing seasonal amplitude towards the coastline with a range

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643 between 284 $\mu\text{mol kg}^{-1}$ found in the outer region of the *Ría de Vigo* in May and 205
644 $\mu\text{mol kg}^{-1}$ in the inner waters in November. These results seem to reinforce the
645 importance of the oxygen consumption in this shallow area, where the water column is
646 less than 10 metres deep and therefore, it would also be influenced by benthic respiration
647 (Alonso-Pérez and Castro, 2014).

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648
649 The monthly means of nitrate concentration (Fig. 4e) could be summarized as high
650 values during autumn and winter due to the nutrients delivered from the continent and
651 the vertical mixing, and as minimum nitrate values from March to September because of
652 phytoplankton consumption. The nitrate concentration was markedly higher in the inner
653 *Ría de Vigo*, where it exceeded 9 $\mu\text{mol kg}^{-1}$ in February and decreased towards the open
654 ocean, where the highest monthly value was seen to be 2.5 $\mu\text{mol kg}^{-1}$. Some notable
655 aspects can be seen in Fig. 5d, such as water poor in nitrate in the ocean region between
656 the two peaks of 3.5 $\mu\text{mol kg}^{-1}$ in March and 1.3 $\mu\text{mol kg}^{-1}$ in October. This shows the
657 presence of the IPC waters, which are warmer and saltier than the shelf waters. Also
658 noteworthy was the particular fact that while the nitrate concentration in other areas was
659 practically zero in summer, the nitrate amount in the surface waters within the *Ría de*
660 *Vigo*, and especially in the inner *Ría de Vigo*, was not completely consumed. This
661 indicates a constant supply throughout the year, either through upwelling events or the
662 continental inputs. This in turn means that while the chlorophyll values were at a
663 minimum in the offshore waters in summer, the phytoplankton community in the
664 estuary grew in summer during the upwelling relaxation periods (Pérez et al., 2000).

665 The nutrient concentration during spring and summer was only detectable in the newly
666 upwelled waters that can show values up to 6 $\mu\text{mol L}^{-1}$ (Fraga, 1981; Castro et al.,
667 1994). During the cessation of the upwelling season in September and October, the
668 chlorophyll concentration (Fig. 5f) increased again, sustained by nutrients that entered
669 from deeper waters through vertical mixing. It should be noted that there was a
670 coincidence of high chlorophyll in the water column and low oxygen concentration in
671 the inner *Ría de Vigo* from May to November, indicating the potential importance of
672 benthic fluxes and vertical fluxes (reference).

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674 4.3. Long-term trends

682 The long-term trends of the biogeochemical properties in the surface waters along the
683 described longitudinal transect between the inner Ría de Vigo and the ocean zone were
684 estimated to be the interannual linear rate of the deseasonalized time series, previously
685 removing the monthly means in these regions and assuming a null spatial variability.
686 The significant trends in the ARIOS database, meaning long-term variability, should be
687 interpreted as a combination of the natural variability on a decadal scale (Pérez et al.,
688 2010; Padin et al., 2010) and anthropogenic forcings (Wolf-Gladrow et al., 1999;
689 Anderson and Mackenzie 2004; Bakun et al., 2010).

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690
691 No long-term temperature variability was found in the surface waters of any region
692 despite the known warming previously reported on the Northern Iberian coast (Pérez et
693 al., 2010; Gesteira et al., 2011; González-Pola et al., 2005). Unlike the temperature, the
694 other expected consequence of climate change in marine ecosystems, namely ocean
695 acidification (Caldeira and Wickett, 2003) was observed along the longitudinal transect,
696 with a greater decrease in the long-term trend of pH towards the coast (Table 2). The
697 long-term pH variation of $-0.0039 \pm 0.0005 \text{ yr}^{-1}$ in the inner waters was about three fold
698 higher than the change observed in the ocean zone, equivalent to $-0.0012 \pm 0.0002 \text{ yr}^{-1}$ in
699 the ocean zone, explaining the 34% and 22% variation in pH in situ respectively, and
700 representing 1-3% of the seasonal pH variation in all zones. These pH decrease rates
701 found in both coastal and open ocean regions of the Iberian Upwelling System lie
702 within the range of other acidification rates estimated in different sites of the North
703 Atlantic Ocean (Lauvset and Gruber, 2014; Bates et al., 2014), being also coherent with
704 the mean rates calculated for the global ocean and for the Eastern North Atlantic and
705 equal to -0.018 and $-0.0164 \text{ decade}^{-1}$, respectively (Lauvset et al., 2015; Rios et al
706 2001).

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Eliminado: Other acidification rates estimated in different sites of the North Atlantic Ocean (Lauvset and Gruber, 2014; Bates et al., 2014) including $-0.018 \text{ decade}^{-1}$ in the mean global ocean pH (Lauvset et al., 2015) or $-0.0164 \text{ decade}^{-1}$ in the Eastern North Atlantic by Rios et al. (2001) are within the acidification range found in the ocean and coastal zones of these waters.

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707
708 The long-term trend in salinity was also seen to be evidently dependent on the distance
709 to the mouth of the Ría de Vigo. The interannual rate of sea surface salinity in the outer
710 and inner ria previously reported by Rosón et al. (2009) was $0.0426 \pm 0.016 \text{ psu yr}^{-1}$ and
711 $0.0193 \pm 0.0056 \text{ psu yr}^{-1}$ respectively. These changes were observed in parallel to an
712 interannual alkalinity increase that is cancelled out in the normalized alkalinity,
713 estimated as the difference between the alkalinity measured and the alkalinity calculated
714 using the linear regression with salinity in each region. Therefore, the interannual
715 salinity increase was the forcing that explains the increase in the buffer capacity of the

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734 surface waters (Sarmiento and Gruber, 2006).

735

736 Other significant long-term variations were found in other biogeochemical parameters
737 in the ARIOS database. The long-term trend of the concentrations of nutrients in the
738 inner Ría de Vigo that had been previously reported for the period 2001-2011 by Doval
739 et al. (2016) showed a significant increase in nitrate, phosphate and ammonium
740 concentrations of $0.0559\pm 0.0158 \mu\text{mol kg}^{-1} \text{yr}^{-1}$, $0.0076\pm 0.0016 \mu\text{mol kg}^{-1} \text{yr}^{-1}$ and
741 $0.0560\pm 0.0011 \mu\text{mol kg}^{-1} \text{yr}^{-1}$ respectively. This fertilization on a long-term scale in the
742 surface waters of the inner ria estimated from ARIOS database was observed in parallel
743 to the deoxygenation of $-0.7\pm 0.2 \mu\text{mol kg}^{-1} \text{yr}^{-1}$. The apparent oxygen utilisation
744 (AOU), calculated using the concentration of O_2 at saturation calculated according to
745 Benson and Krause (1984), underwent an equivalent significant long-term change of
746 $0.7\pm 0.2 \mu\text{mol kg}^{-1} \text{yr}^{-1}$, indicating that either the biological consumption rates, or a
747 change in the amount of time that the waters are ventilated, or even its interaction or
748 exchange with the sediment, cause the the long-term reduction of oxygen.

749

750 This fertilization on a long-term scale estimated from ARIOS database in the surface
751 waters of the inner ria was observed in parallel to the deoxygenation of $-0.7\pm 0.2 \mu\text{mol}$
752 $\text{kg}^{-1} \text{yr}^{-1}$. The apparent oxygen utilisation (AOU), calculated using the concentration of
753 oxygen at saturation calculated according to Benson and Krause (1984), underwent a
754 long-term change of $0.7\pm 0.2 \mu\text{mol kg}^{-1} \text{yr}^{-1}$ equal to the observed in the measurements
755 of oxygen concentration. This coincidence may indicates that the long-term reduction of
756 oxygen is due to the changes in the biological consumption rates, in the rates of the
757 waters ventilation or even in sediment-water interactions rather than due to the effect of
758 temperature. and salinity on oxygen saturation.

759

760 These findings found in the shallower waters of the Ría de Vigo allow us to hypothesize
761 that the long-term increase in salinity would produce an increasingly weak vertical
762 salinity gradient in the water column that would favour the vertical fluxes between the
763 bottom and surface waters. Therefore the observed changes of oxygen and
764 remineralized nutrient inputs in the surface waters could be due to an increasing
765 footprint of benthic respiration, that has a major importance in the net ecosystem
766 metabolism of this coastal region (Alonso-Pérez et al., 2015). This hypothesis would
767 also explain the intense acidification in the inner waters in spite of growing alkalinity

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Eliminado: The long-term trend of the nutrient concentration in the inner Ría de Vigo showed an increase in the nitrate, phosphate and ammonium concentrations of $0.0158\pm 0.006 \mu\text{mol kg}^{-1} \text{yr}^{-1}$, $0.0076\pm 0.0016 \mu\text{mol kg}^{-1} \text{yr}^{-1}$ and $0.0560\pm 0.0011 \mu\text{mol kg}^{-1} \text{yr}^{-1}$ respectively (Doval et al., 2016).

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777 [buffering](#).

778

779 The mean values at each station of the ARIOS database estimated for each depth range
780 described in Figure 2, resulting in 8,384 values, were used to estimate a general value of
781 the long-term trend in pH. The historical pH values in situ from the ARIOS database
782 showed a general decrease in seawater pH in the Iberian Upwelling between 1976 and
783 2018, with an acidification rate of $-0.012 \pm 0.002 \text{ yr}^{-1}$ that significantly explains 2% of
784 the total pH variation (Fig. 5a). The apparent oxygen utilisation was also shown as
785 function of pH over time, revealing the association of higher AOU values with lower
786 pH. The relationship between pH and AOU (Fig. 5b) showed an inverse linear
787 correlation of $-399 \pm 5 \text{ } \mu\text{mol kg}^{-1}$ [and a coefficient of determination \(r-squared\) of 0.52](#).

788 The strong biological activity of the upwelling systems is the main driver of pH
789 changes, explaining 52% of the observed variation in the discrete measurements. The
790 distribution of nitrate seen in relation to the distribution of pH and AOU (Fig. 5b)
791 showed the association of higher pH values with negative AOU values and a nitrate
792 decrease, reinforcing the importance of biological processes in these marine carbonate
793 system. [Although the different processes controlling the AOU values were not separated](#)
794 [in this analysis, the oxygen concentration in addition to the remineralization of the](#)
795 [organic matter and the photosynthesis is conditioned by changes in temperature and](#)
796 [salinity, ventilation events, water masses mixing and other processes \(Sarmiento and](#)
797 [Gruber, 2006\). Therefore, the long-term drop in seawater pH measurements estimated](#)
798 [from the ARIOS database presented here confirms that the future evolution of ocean](#)
799 [acidification in this productive region is likely to depend on both the CO₂ increase in](#)
800 [the atmosphere and other long-term changes \(of natural and/or anthropogenic origin\)](#)
801 [affecting the seawater's carbonate system.](#)

802

803 5. Data availability

804 The ARIOS dataset (Pérez et al., 2020) is archived at DIGICAL CSIC under the Digital
805 Object Identifier (DOI): <http://dx.doi.org/10.20350/digitalCSIC/12498>.

806

807 The data are available as WHP-Exchange bottle format (arios_database_hy1.csv). A
808 documentation file (readme_ARIOSDATABASE.txt) provides an description of the
809 material and methods of the measurements and the parameters of the dataset. In both
810 files, a table similar to the Table 1 of this manuscript include the DOI and the

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Eliminado: Attending to the interannual salinity changes in the shallower waters of the *Ria de Vigo*, these findings seem to be related to a weakening of the vertical salinity gradient and a growing exchange between the bottom and surface waters. So, the footprint of oxygen consumption and the remineralized nutrient inputs resulting from benthic respiration seem to reach the upper layer. The metabolic processes in a more homogeneous water column would also explain an intense acidification in the inner waters in spite of growing alkalinity buffering. .

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Eliminado: in this region responding to the intrusion of atmospheric CO₂ may also be due to the impact of other interannual changes affecting the seawater pH, such as biological activity. So, the future evolution of acidification will respond both to the potential increase in CO₂ in the atmosphere and to other long-term changes affecting the seawater's carbonate system.

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838 EXPCODE of the original cruise files gathered in the ARIOS dataset.

839

840 These data are available to the public and the scientific community with the aim of that
841 their wide dissemination will lead to new scientific knowledge about the ocean
842 acidification and the biogeochemistry of the Galicia Upwelling System. The dataset is
843 subject to a Creative Commons License Attribution-ShareAlike 4.0 International
844 (<http://creativecommons.org/licenses/by-sa/4.0/>) and users of the ARIOS dataset should
845 reference this work.

846

847 **6. Conclusions**

848 The ARIOS database is a unique compilation of biogeochemical discrete measurements
849 in the Iberian Upwelling Ecosystem from 1976 to 2018. This data set comprises more
850 than 17,653 discrete samples from 3,357 oceanographic stations (but not always for all
851 parameters) of pH, alkalinity and associated physical and biogeochemical parameters
852 (e.g., temperature, salinity, and chlorophyll and oxygen concentrations). The material
853 and methods varied throughout the sampling period due to logistical and analytical
854 issues such as those described in Table 1, where different sites are mentioned to
855 download these measurements and detailed information.

856

857 Among the results described as preliminary and relevant information to learn the
858 environmental and oceanographic context of the ARIOS database, we can mention the
859 following main points concerning the pH characteristics of the Iberian Upwelling
860 System:

- 861 • A decrease in seawater pH in the Iberian Upwelling between 1976 and 2018,
862 with an acidification rate of $-0.012 \pm 0.002 \text{ yr}^{-1}$ that significantly explains 2% of
863 the total pH variation
- 864 • An interannual pH variation of $-0.0039 \pm 0.0005 \text{ yr}^{-1}$ in the inner waters and -
865 $0.0012 \pm 0.0002 \text{ yr}^{-1}$ in the ocean zone.
- 866 • An inverse linear correlation between pH and AOU of $-399 \pm 5 \text{ } \mu\text{mol kg}^{-1}$ that
867 explained 52% of the observed variation in the discrete measurements.

868

869 This published ARIOS database is a useful and necessary tool to confirm and study
870 biogeochemical changes in the seawater at long term trend. Likewise, we understand
871 that it is a starting point to which to add future observation projects to continue

872 increasing the knowledge about the impact of climate change in the Iberian Upwelling
873 Ecosystem.

874

875 **Acknowledgements.**

876 The compilation of this data set was funded by the ARIOS project (CTM2016-76146-
877 C3-1-R) funded by the Spanish government through the Ministerio de Economía y
878 Competitividad that included European FEDER funds. Part of the processing work was
879 supported by the MarRISK project (European Union FEDER 0262_MarRISK_1_E)
880 funded by the Galicia-Northern Portugal Cross-Border Cooperation Program
881 (POCTEP). This project has also received funding from the European Union's Horizon
882 2020 research and innovation programme under grant agreement No 820989 (project
883 COMFORT, Our common future ocean in the Earth system – quantifying coupled
884 cycles of carbon, oxygen, and nutrients for determining and achieving safe operating
885 spaces with respect to tipping points). This data set encompasses decades of work
886 conducted by an overwhelming number of people. We thank all of the scientists,
887 technicians, personnel, and crew who were responsible for the collection and analysis of
888 the over 22 000 samples included in the final data set. In addition to the PI cited in
889 Table 1 we also thank to Trinidad Rellán, Antón Velo, Miguel Gil Coto, Marta Alvarez,
890 Marylo Doval, Jesus Gago, Daniel Broullón and Marcos Fontela. We also thank Monica
891 Castaño for starting this data compilation more than 10 years ago.

892

893 **References**

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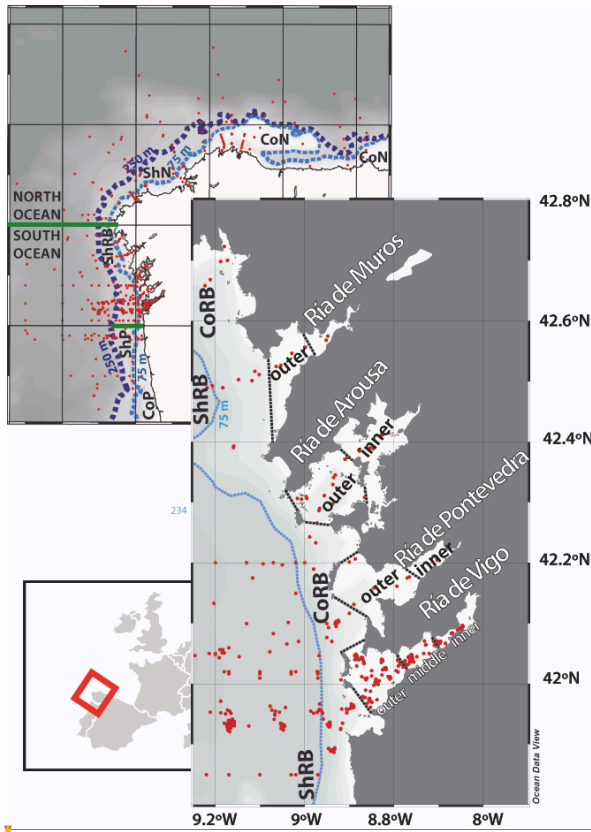
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1381 Figure 1. Map of all stations (red dots) including the geographical areas selected to
 1382 classify the ARIOS database from isobath of 250 m (dark blue line) and 75 metres (light
 1383 blue line), latitudinal criterion (green lines) and geographical lines (black lines).

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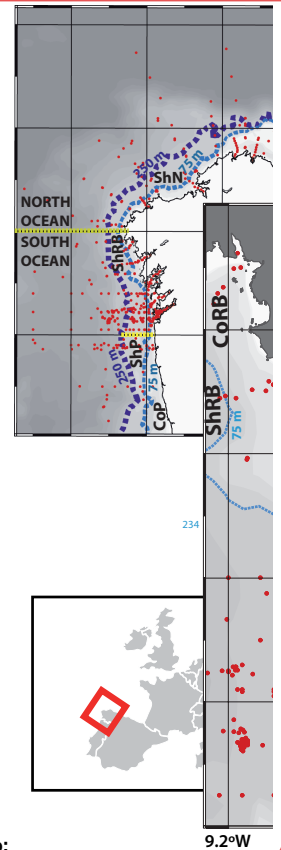
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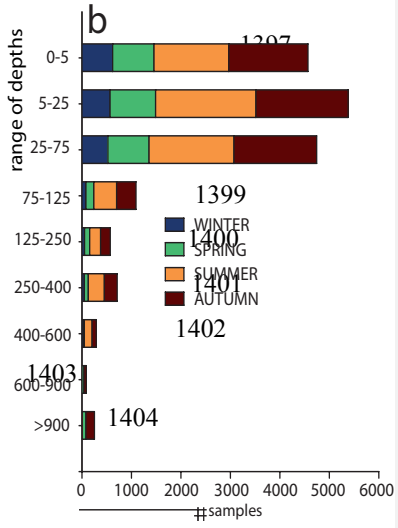
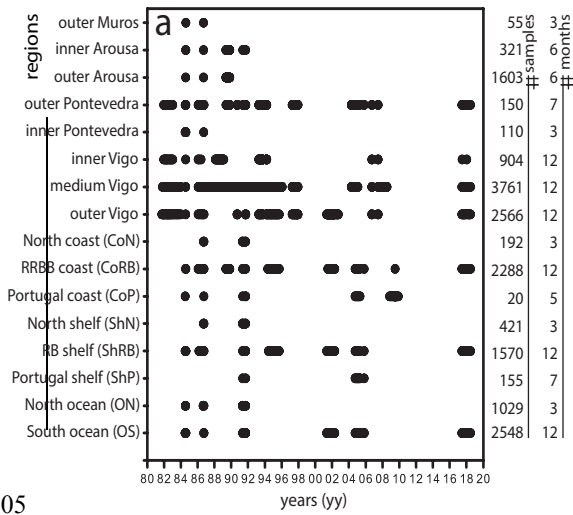
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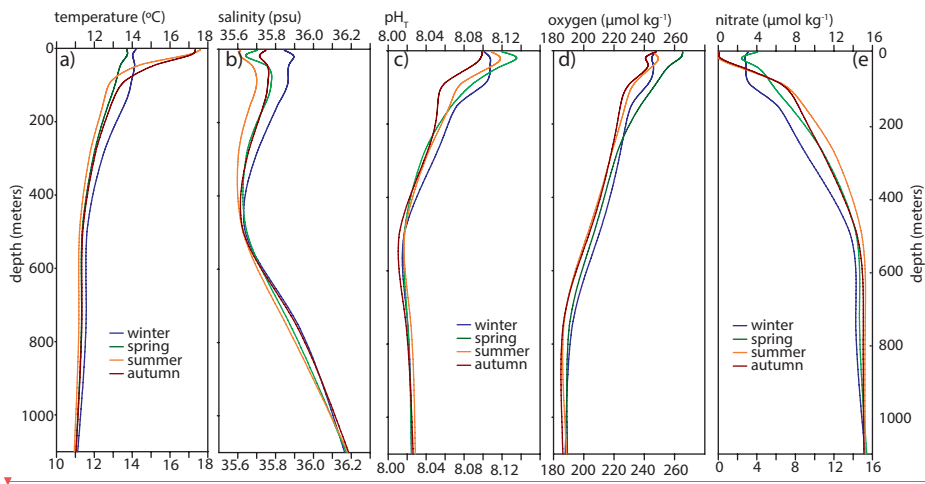


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1406 Figure 2. a) Temporal distribution of the observations in the geographical boxes
 1407 included in the ARIOS dataset. b) Seasonal distribution of the measurements in relation
 1408 to depth.

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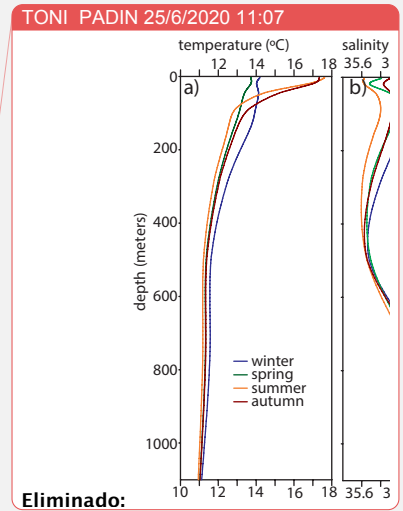
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1431 Figure 3. Profiles of seasonal means of temperature (a), salinity (b), pH_T (c), oxygen (d)
1432 and nitrate concentration (e) in the first 1100 meters of the region South Ocean shown
1433 in Fig. 1.

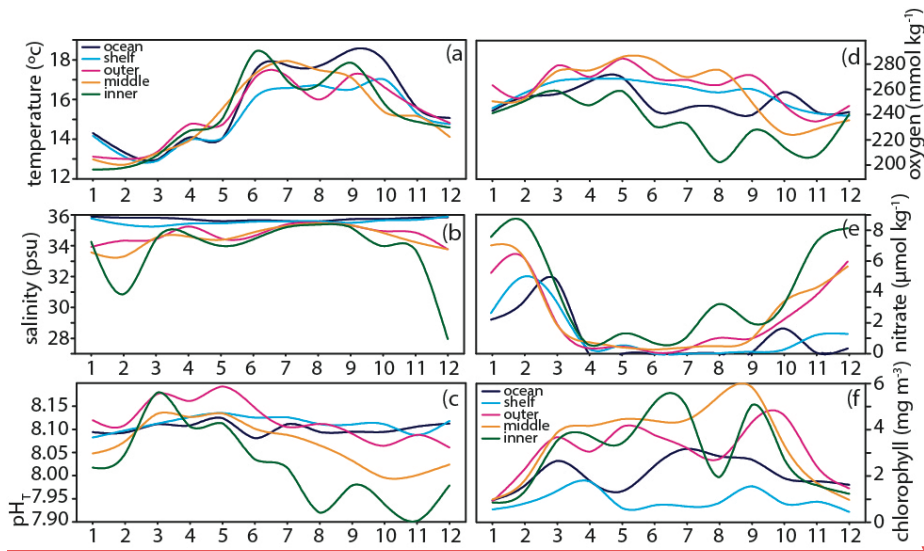
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1457 Figure 4. Sea surface (<5 meters depth) seasonal cycles in 1976 - 2018 of temperature
1458 (a), salinity (b), pH_T (c), oxygen concentration (d), nitrate concentration (e) and
1459 chlorophyll (f) at sea surface for five geographical boxes shown in Fig. 1: South Ocean,
1460 RB shelf and outer, middle and inner Ria de Vigo for the entire period of the ARIOS
1461 database.

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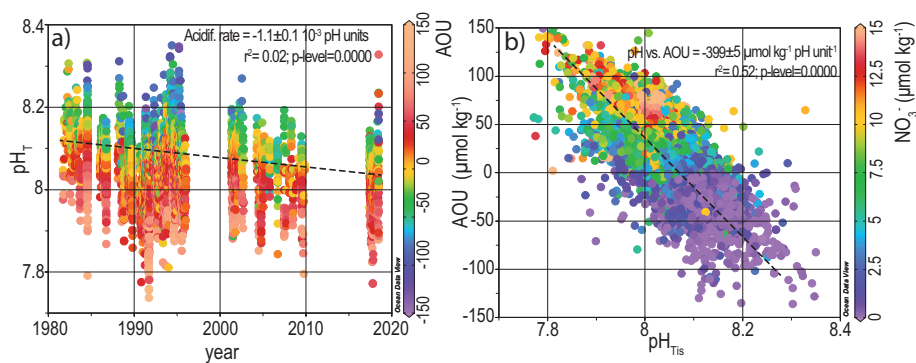
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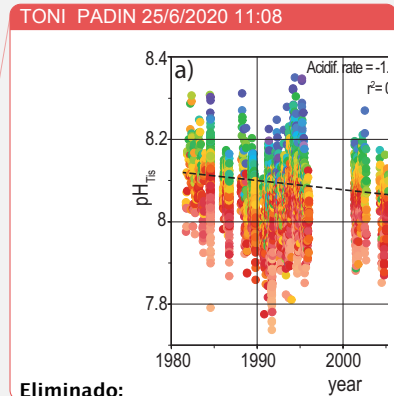


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1480 Figure 5. Time-series of pH ARIOS data. The black line depicts the long-term trend.

1481 Scatter diagram of AOU vs pH_{T} including the nitrate concentration shown as colour of
1482 every dot.

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Table 1. Discrete measurements of projects gathered in the ARIOS database and associated information: including dates, [the number of days between the start and the end of sampling period \(#d\)](#), sample number (#), the principal investigator (PI), measured parameters, link to data repository and the sampled geographical area.

- All projects include measurements of T, S. Others as pH, alkalinity (Alk), nutrient (Nut), oxygen (O₂) concentration, chlorophyll (Chla) are indicated.

- The concentration units of these variables are $\mu\text{mol kg}^{-1}$ or $\mu\text{mol L}^{-1}$ () and the pH measurements in NBS scale (°) or in total scale.*

- Regions are identified as ocean (O), shelf (Sh), coastal (Co), Ría de Vigo (RV), Ría de Pontevedra (RP), Ría de Arousa (RA) and Ría de Muros (RM) while the superscript index means south (S), north (N), Portugal (P), Rías Baixas (RB), outer (O), middle (M) and inner (I).

EXPOCODE	PROJECT	DATE	#d	IP	#	CTD	O ₂	Nut	pH	Alk	Chla	CRM	Data Repository	REGIONS
29LP19761026	Ría Vigo 1977	26/10/76	413	F Fraga	135	N	N	S*	S°	N	N	N	http://dx.doi.org/10.20350/digitalCSIC/9917	Co ^{RB}
29LP19810929	Ría Vigo 1981-83	29/9/81	472	F Fraga	748	N	S*	S*	S°	S	N	N	http://dx.doi.org/10.20350/digitalCSIC/9918	RV ^{O,M,I}
29LP19830215	Ría Vigo 1983-84	15/2/83	322	F Fraga	312	N	S*	S*	S°	S	N	N	http://dx.doi.org/10.20350/digitalCSIC/9919	RV ^{O,M}
29GD19840711	GALICIA-VIII	11/7/84	28	F Fraga	1865	N	S	S	S°	S	S	N	http://dx.doi.org/10.20350/digitalCSIC/9908	O ^{N,S} , Sh ^{RB} , Co ^{P,RB} , RV ^{O,M,I} , RA ^{O,I} , RP ^{O,I} , RM
29GD19860121	Ría Vigo 1986	21/1/86	203	F Fraga	332	N	S	S	S°	S	S	N	http://dx.doi.org/10.20350/digitalCSIC/9910	Sh ^{RB} , Co ^{RB} , RV ^{O,M,I}
29GD19860904	GALICIA-IX	23/9/86	5	F Fraga	1640	N	S	S	S°	S	S	N	http://dx.doi.org/10.20350/digitalCSIC/9911	O ^{N,S} , Sh ^{RB,N} , Co ^{P,RB,N} , RV ^{O,M,I} , RA ^{O,I} , RP ^{O,I} , RM
29LP19870120	PROVIGO	17/9/87	3290	F F Pérez	2317	N	S	S	S°	N	S	N	http://dx.doi.org/10.20350/digitalCSIC/9924	RV ^M
29LP19880212	LUNA 88	12/2/88	367	A F Rios	468	N	S	S	S°	S	S	N	http://dx.doi.org/10.20350/digitalCSIC/9907	RV ^{M,I}
29IN19890512	GALICIA-X	5/5/89	171	F F Pérez	3113	N	S	S	S°	S	S	N	http://dx.doi.org/10.20350/digitalCSIC/9920	Co ^{RB} , RA ^{O,I}
29IN19900914	Ría Vigo 1990	14/9/90	13	FG Figueiras	108	Y	S	S	S°	S	S	N	http://dx.doi.org/10.20350/digitalCSIC/9921	RV ^{O,M,I}
29IN19910510	GALICIA-XI	5/5/91	4	F F Pérez	327	Y	S	S	S°	S	S	N	http://dx.doi.org/10.20350/digitalCSIC/9922	O ^{N,S} , Sh ^{P,RB,N} , Co ^{P,RB,N} , RA ^O
29IN19910910	GALICIA-XII	15/9/91	10	F G Figueiras	663	Y	S	S	S°	S	S	N	http://dx.doi.org/10.20350/digitalCSIC/9923	O ^{N,S} , Sh ^{P,RB,N} , Co ^{P,RB,N} , RV ^{O,M,I} , RA ^O
29LP19930413	Ría Vigo 1993-94	22/3/94	344	F G Figueiras	406	Y	S	S	S°	S	S	N	http://dx.doi.org/10.20350/digitalCSIC/9927	RV ^{O,M,I}
29JN19940505	Ría Vigo 1994-95	5/5/94	504	M Cabanas	669	Y	S	S	S°	S	S	N	http://dx.doi.org/10.20350/digitalCSIC/9926	Sh ^{RB} , Co ^{RB} , RV ^O
29MY19970407	CIRCA-97	7/4/97	248	F F Pérez	547	Y	S	N	S°	S	S	N	http://dx.doi.org/10.20350/digitalCSIC/9928	RV ^{O,M,I}
29MY20010515	DYBAGA	15/5/01	344	F F Pérez	1421	Y	S	S*	S	S	S	Y	http://dx.doi.org/10.20350/digitalCSIC/9929	Sh ^{P,RB} , Co ^{RB} , RV ^O
29MY20010702	REMODA	2/7/01	451	X A Alvarez	203	Y	S	S*	S	S	S	Y	http://dx.doi.org/10.20350/digitalCSIC/9930	RV ^O
29MY20040419	FLUVBE	19/4/04	283	C G Castro	187	Y	S	S*	S	S	S	Y	to be submitted	RV ^{M,I}
29CS20041004	ZOTRACOS	4/10/04	389	M Cabanas	371	Y	S	S	S	S	S	Y	http://dx.doi.org/10.20350/digitalCSIC/9932	Sh ^{P,RB} , Co ^{P,RB} , RP ^O
29MY20060926	CRÍA	26/9/06	275	D Barton	197	Y	S	S*	S	S	S	Y	http://dx.doi.org/10.20350/digitalCSIC/9931	RV ^{O,M,I}
29MY20070917	RAFTING	17/9/07	301	C G Castro	287	Y	S	S*	S	S	S	Y	to be submitted	RV ^M
29MY20081105	LOCO	5/11/08	378	X A Alvarez	72	Y	S	S	S	S	S	Y	http://dx.doi.org/10.20350/digitalCSIC/9936	Co ^{RB}
29AH20090710	CAIBEX-I	16/7/09	11	D Barton	191	Y	S	S	S	S	S	Y	http://dx.doi.org/10.20350/digitalCSIC/9934	Co ^{P,RB}
29MY20170609	ARIOS	9/6/17	382	FF Pérez	1114	Y	S	S*	S	S	S	Y	http://dx.doi.org/10.20350/digitalCSIC/9963	Sh ^{P,RB} , Co ^{RB} , RV ^{O,M,I}

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	SS _{range}	r ² _{ss}	t _{interannual}	r ²	p-value
OCEAN	0.050	0.17	-0.0012±0.0002	0.21	0.0000
SHELF	0.050	0.06	-0.0017±0.0003	0.15	0.0009
OUTER	0.120	0.24	-0.0027±0.0003	0.21	0.0000
MIDDLE	0.130	0.28	-0.0022±0.0005	0.03	0.0000
INNER	0.260	0.47	-0.0039±0.0005	0.34	0.0000

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Table 2: Seasonal amplitude of monthly pH means (SS_{range}) and long-term trends (t_{interannual}) of pH in five regions and significant regression coefficients between the in situ pH measurements and the monthly mean pH values (r²_{ss}) and the regression coefficient of the temporal variability of the deseasonalized pH measurements (r²).