

1                   **A European map of groundwater pH and calcium**

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56 **Abstract.** Water resources and associated ecosystems are becoming highly endangered due to

57 ongoing global environmental changes. Spatial ecological modelling is a [widely used](#)

58 [tool promising toolbox](#) for understanding the past, present and future distribution and diversity

59 patterns in groundwater-dependent ecosystems, such as fens, springs, streams, reed beds or wet

60 grasslands. Still, the lack of detailed water chemistry maps prevents [their the use of](#) reasonable

61 [use models to be applied](#) on continental and global scales. Being major determinants of

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62 biological composition and diversity of groundwater-dependent ecosystems, groundwater pH and  
63 calcium are of utmost importance. Here we developed the up-to-date European map of  
64 groundwater pH and Ca, based on 7,577 measurements of near-surface groundwater pH and  
65 calcium distributed across Europe. In comparison to the existing European groundwater maps,  
66 we included a several times larger number of sites, especially in the regions rich in spring and  
67 fen habitats, and filled the apparent gaps in Eastern and Southeastern Europe. We used Random  
68 Forest models and regression kriging to create continuous maps of water pH and calcium at the  
69 continental scale, which is freely available also as a raster map (Hájek et al. 2020;  
70 [10.5281/zenodo.4139912](https://doi.org/10.5281/zenodo.4139912)). Lithology had higher importance than climate for both pH and  
71 calcium. The previously recognised latitudinal and altitudinal gradients were rediscovered with  
72 much refined regional patterns, as associated with bedrock variation. For ecological models of  
73 distribution and diversity of groundwater dependent, but also other many terrestrial  
74 ecosystems, the our new map based on field ground water measurements is more suitable than  
75 previously used maps of soil pH, unlike, which it mirrors mirror not only bedrock chemistry more  
76 than, but also vegetation-dependent soil processes.  
77  
78 **Copyright statement**  
79 No copyright statement is needed.  
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81 **1. Introduction**  
82 The Earth system is currently undergoing unprecedented changes in climate, global  
83 biogeochemical cycles, and land use, resulting in biodiversity loss (Ceballos et al. 2017, Song et

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84 al. 2018, Blowes et al. 2019, Brondizio et al. 2019). Freshwater systems belong to the most  
85 endangered habitats (Cantonati et al. 2020a, Tickner et al. 2020) and, among them, groundwater-  
86 dependent ecosystems, such as fens and springs, hold primacy (Janssen et al. 2016, Chytrý et al.  
87 2019, Hájek et al. [2020](#), Stevens et al. [2020](#)). Species composition and richness of spring systems  
88 are generally governed by water pH and calcium concentration ( $\text{Ca}^{2+}$ ), which are highly variable  
89 at different spatial scales (Malmer 1986; Rydin and Jeglum 2013; Peterka et al. 2017; Horskáková  
90 et al. 2018; Cantonati et al. 2020a,b). Therefore, understanding the spatial patterns in  
91 groundwater pH and  $\text{Ca}^{2+}$  is important not only for general geochemical knowledge and for water  
92 resource management, but to the same extent for the conservation of freshwater systems and  
93 associated biodiversity.

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94 In Earth and biodiversity sciences, ecological modelling is a widely used tool for  
95 understanding the distribution and diversity patterns of ecosystems and habitats, and for  
96 predicting their future development under global change. Ecological models usually incorporate  
97 environmental or historical predictors extracted from thematic maps (Jiménez-Alfaro et al.  
98 2018a, Večeřa et al. 2019, Divíšek et al. 2020), including soil properties for terrestrial  
99 ecosystems (Hengl et al. 2017). However, soil parameters as soil pH contribute negligibly to the  
100 models for groundwater-dependent habitats, even for those strongly controlled by pH and  $\text{Ca}^{2+}$ ,  
101 such as base-rich fens (Jiménez-Alfaro et al. 2018b). This is due to a poor correlation between  
102 groundwater chemistry and pH or  $\text{Ca}^{2+}$  in soil, disrupted mainly by mineral leaching or  
103 accumulation of organic matter in soil. For this reason, there is a strong need to produce maps for  
104 groundwater pH and  $\text{Ca}^{2+}$  concentration at the European scale that would allow producing the  
105 continental-scale ecological models useful for enforcing conservation strategies in groundwater-  
106 dependent habitats. [Ideally, such models should include lithology as a dominant factor](#)

107 determining groundwater pH and Ca<sup>2+</sup> concentration (Hem 1985, Chapelle 2003, Tahvanainen  
108 Stevens et al. 2020).

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109 In spite of important mapping efforts of groundwater (Duscher et al. 2015) and karst  
110 aquifers (Chen et al. 2017) at the European and global level, the only available European-scale  
111 maps of groundwater pH and Ca<sup>2+</sup> concentration are those included in the FOREGS Geochemical  
112 Atlas of Europe (Salminen et al. 2006). These maps are based on 808 stream-water  
113 measurements distributed relatively equally across Europe. However, they show a large gap in  
114 Eastern and Southeastern Europe (Romania, Bulgaria, Belarus, Russian Federation, Ukraine,  
115 Moldova, Serbia, Kosovo, Montenegro, Bosnia and Herzegovina, Northern Macedonia). In  
116 addition, ~~the~~those maps are ~~constructed~~based on insufficient data density in some areas rich in ~~the~~different  
117 ~~different~~groundwater-dependent ~~ecosystem~~ecosystem types, but heterogeneous in terms of  
118 lithology (the Alps, the Carpathians, Bohemian Massif, the Cantabrian Mountains and the  
119 Pyrenees, and some regions of Fennoscandia). We therefore aimed at substantial improvement of  
120 the existing data by creating a database with field data measurements across the entire European  
121 continent, and at creating a model-based map representing major patterns of groundwater pH and  
122 Ca<sup>2+</sup> concentration at local and continental scales. Our data will allow better understanding of the  
123 patterns and causes of groundwater conditions in freshwater systems, ~~and~~strongly  
124 ~~improve~~improving the ~~databases~~spatial information suitable for European-scale modelling of  
125 ~~the~~biodiversity in groundwater-dependent and related ecosystems.

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## 127 2. Methods

### 128 Data collection

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129 We assembled the data set of pH and Ca<sup>2+</sup> (or electrical conductivity in µS.cm<sup>-1</sup> at 20 °C;  
130 hereinafter abbreviated as EC) measurements in groundwater, covering the whole of Europe,  
131 with a greater density in the regions rich in endangered groundwater-dependent ecosystems such  
132 as springs and fens. We excluded most of Ukraine and European part of Russian Federation,  
133 because of large data gaps in these areas. We considered all types of shallow groundwater  
134 systems, especially spring, spring-fen, and stream water. The core of our data set is formed by  
135 unpublished pH and Ca<sup>2+</sup> or EC data sets of co-authors (3,618 sites); some of them processed in  
136 ecological papers without presenting original pH and Ca<sup>2+</sup> data (Hájková et al. 2006, 2008; Hájek  
137 et al. 2008; Sekulová et al. 2013; Plesková et al. 2016; Horská et al. 2018; Šimová et al.  
138 2019). The second most important source were vegetation databases registered in GIVD  
139 (Dengler et al. 2011; Table 1) and EVA (Chytrý et al. 2016), from where 1,160 measurements  
140 from freshwater habitats were obtained. Both unpublished data and data from vegetation  
141 databases were filtered using original information or metadata of the sources in a way that only  
142 data from spring-fed fens and springs were considered. The data from ombrotrophic bogs and  
143 clearly topogenic fens (mainly terrestrialised lakes) were omitted because their water chemistry  
144 is governed by the decomposition of organic matter, atmospheric humidity and deposition, algal  
145 photosynthesis (Kann and Smith 1999), and biotic processes such as cation exchange capacity of  
146 mosses (Clymo 1963, Soudzilovská et al. 2010, Vicherová et al. 2015), rather than by bedrock  
147 chemistry. We also obtained data from public data sets ~~from stored in~~ national environmental and  
148 nature conservation agencies of Germany, Slovenia and Bulgaria (1081 sites; see Table 1), data  
149 from FOREGS Geochemical Atlas of Europe (Salminen et al. 2006; 808 sites) and literature data  
150 based on our gap-oriented excerpt (883 sites; Table 1); most data came from Hinterlang

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151 (1992), Tanneberger et al. (2011), Eades et al. (2018), Kadūnas et al. (2017) and Savić et al.  
152 (2017).

153  
154 **Table 1.** Data sources. The name abbreviations are explained in the team list or  
155 acknowledgements: **Table did no change.**

156  
157 In total, we collected 7,577 samples (Table 1). A part Some of the these samples,  
158 however, represented are repeated measurements conducted in the same site, especially in  
159 public data sets. Some, while other samples (from other data sets (vegetation databases, or  
160 literature data) shared share the same coordinates and site name or code, suggesting repeated  
161 measurements as well. We therefore averaged repeated measurements from the same sampling  
162 spots sites. We further deleted samples whose coordinates were obviously erroneous, such as  
163 those in oceans. These steps reduced the number of samples to 6,561, out of which 6,459  
164 samples contained information on water pH value and 5,927 samples contained information  
165 about EC of water or Ca<sup>2+</sup> concentration. Out of these 5,927 samples, 2,988 samples had directly  
166 measured both Ca<sup>2+</sup> and EC ( $\mu\text{S} \cdot \text{cm}^{-1}$  at 20 °C), and for the remaining 2,939 samples we  
167 estimated Ca<sup>2+</sup> concentration by EC of water.

168  
169 *Imputation of missing Ca<sup>2+</sup> values by EC of water*  
170 For imputation of Ca<sup>2+</sup> values based on EC, we first aimed at constructing a simple imputation  
171 equation based on the well-known correlation between EC and Ca<sup>2+</sup> concentration in springs and

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172 fens ([Hem 1985](#), Sjörs & Gunnarsson 2002, Plesková et al. 2016). In our data set of 2,988  
173 samples, as well as in its regional subsets, this relationship was strongly governed by EC values  
174 above ca 1,000  $\mu\text{S.cm}^{-1}$ , although they formed only a small part of the data set (4.7% of the data  
175 set; 139 samples). In the EC range 1,000-10,000  $\mu\text{S.cm}^{-1}$  (an outlier of 17,000  $\mu\text{S.cm}^{-1}$  was  
176 omitted), the correlation between water EC and  $\text{Ca}^{2+}$  concentration was not statistically  
177 significant ( $r = 0.15$ ,  $P = 0.07$ ). The problem of high EC values governing the regression model  
178 was the most apparent in the public data sets. In the data set of Bulgarian Ministry of  
179 Environment, weak correlation between EC and  $\text{Ca}^{2+}$  persisted even when EC values above 1000  
180 were omitted (Supplementary Figure 1). This database further contains many samples which are  
181 not near-surface samples that were measured in other datasets. We therefore finally decided (1)  
182 not to include the database of Bulgarian Ministry of Environment into the imputation model, and  
183 (2) limit the gradient of EC to 1,000  $\mu\text{S.cm}^{-1}$ . We further omitted a few samples from ophiolite  
184 ([Kamberović et al. 2019](#)) where high EC occurred despite low  $\text{Ca}^{2+}$ . The resulting data set of  
185 2,319 samples nevertheless still showed some samples with suspiciously high or low  $\text{Ca}^{2+}$   
186 concentration relative to EC (Supplementary Figure 2), suggesting either the effect of other ions  
187 than  $\text{Ca}^{2+}$  or inconsistent analytical methodology. Because our aim was to create the most  
188 accurate imputation model rather than testing the relationship, we removed these outliers.  
189 Therefore, we calculated the EC:Ca and Ca:EC ratios and removed outliers, i.e., all points  
190 outside the 1.5 x interquartile range. The final imputation model was hence based on 2,062 sites.  
191 We performed a null-intercept linear regression (Figure 1) with  $\text{Ca}^{2+}$  as dependent variable (y)  
192 and EC as predictor (x); the resulting equation  $y = 0.153x$  was obtained ( $R^2 = 0.84$ ). [Such](#)  
193 [relationship between  \$\text{Ca}^{2+}\$  and EC is similar as that found in the abovementioned studies \(Hem](#)  
194 [1985, Sjörs & Gunnarsson 2002, Plesková et al. 2016\)](#). Based on this equation, we imputed  $\text{Ca}^{2+}$

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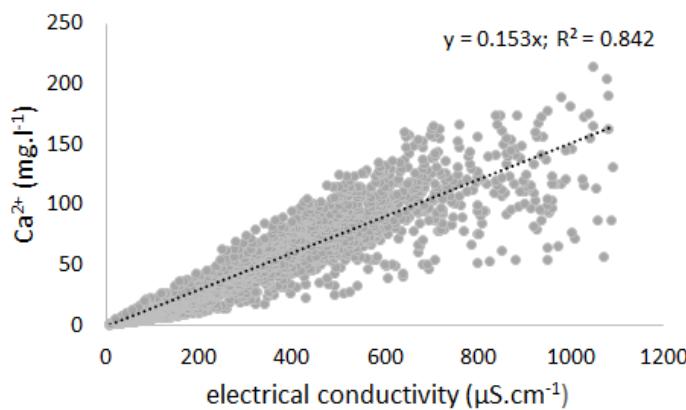
195 concentrations to all samples where only EC was measured. The imputed  $\text{Ca}^{2+}$  values show a  
196 somewhat narrower range (Supplementary Figure 3) than originally measured values. Both  
197 subsets show minimum  $\text{Ca}^{2+}$  value below 1  $\text{mg.l}^{-1}$ , but imputed data show lower non-outlier  
198 maximum (125.5  $\text{mg.l}^{-1}$ ) than measured data (197.2  $\text{mg.l}^{-1}$ ). Absolute maximum value was also  
199 lower for the subset with imputed values. Imputation of  $\text{Ca}^{2+}$  values to all samples, including  
200 high-EC ones ( $> 1,000 \mu\text{S.cm}^{-1}$ ), hence did not skew the imputed data to higher values.

201

202 **Figure 1:** The final regression model to impute  $\text{Ca}^{2+}$  values based on electrical conductivity (EC;  
203  $\mu\text{S.cm}^{-1}$ ;  $n = 2,062$

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206

207 Geographical modelling and selection of the predictors

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208 We used our dataset [with 6,561 sites](#), with measured groundwater pH and either measured or  
209 imputed Ca<sup>2+</sup> concentrations to model expected values across non-sampled areas.  
210 [Groundwater](#) Our aim was to produce continuous maps for groundwater-dependent pH (GW-pH)  
211 [was modeled with using](#) 6,459 samples (pH min = 2.20; max = 11.32; mean = 6.69); and  
212 groundwater-dependent Ca<sup>2+</sup> (GW-Ca) [with using](#) 5,927 samples (min = 0.15; max = 3567.41;  
213 mean = 48.73 mg.l<sup>-1</sup>). Ca<sup>2+</sup> values were ln-transformed. All field samples had geographic  
214 coordinates assigned from GPS or georeferenced with an accuracy between ca 10 [m](#) (precise  
215 field measurements) and 500 m ([some database data from georeferenced sites in databases](#)). We  
216 kept the pH outliers; ten values below 3.5 and nine values above 8.8. Even if these values may be  
217 suspicious, they largely come from published sources (FOREGS Geochemical Atlas of Europe,  
218 British vegetation database). Apart from measurement error, they may be explained by the  
219 influence of mineral waters from deep hydrological circulations (e.g., in a spring in the  
220 Apennines, very high pH value 11.2 was due to enrichment with sodium and chloride associated  
221 with low temperature reaction between meteoric water and ultramafic rocks; Boschetti and  
222 Toscani 2008, Boschetti et al. 2013, Segadelli et al. 2017, Cantonati et al. 2020c). These values  
223 form only a minor part of the data set and have negligible effect on the results.

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224 [For each site, we obtained environmental predictors from thematic GIS maps \(see below\).](#)  
225 [We focused on the predictors that may causally affect the groundwater pH and calcium richness.](#)  
226 [Aquifer chemistry is of prime importance \(Hem 1985, Fairchild et al. 1994, Frei et al. 2000,](#)  
227 [Chapelle et al. 2003, Tahvanainen 2004, Stevens et al. 2020\), but no such thematic map exists, at](#)  
228 [least at the scale needed for computing our spatial predictions. We therefore included the](#)  
229 [lithological groups from the Hydrogeological Map of Europe \(Duscher et al. 2015\), together with](#)  
230 [soil pH maps \(see below\), for which we anticipated a certain correlation with bedrock chemistry.](#)

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231 Apart from aquifer chemistry, residence time may also affect groundwater chemistry by  
232 impacting dissolution rates. Precipitation amount and frequency affect not only flow paths  
233 activity and redistribution of groundwater, but also its residence time in the aquifer, impacting  
234 carbonate dissolution and precipitation rates (Hem 1985, Crossman et al. 2011, Lewandowski et  
235 al. 2015, Vystavna et al. 2020). Groundwater with a short transit time (1–3 years) or ‘young  
236 water’ (Soulsby et al. 2015) can be particularly sensitive to changes in precipitation amount and  
237 frequency. For each site, we obtained environmental predictors from the thematic GIS maps  
238 (see below). For some sites, important predictors were missing in the maps (e.g., sites at far  
239 north in the arctic zone, or close to sea or water bodies) and these sites were therefore not  
240 included in the final models.

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241 We therefore also considered climatic parameters associated with precipitation in the  
242 models (see below). Although there are some other potential predictors of minor importance that  
243 may affect groundwater chemistry (Hem 1985, Stevens et al. 2020), no corresponding thematic  
244 map is available to be included into our models. For some sites, the selected predictors were  
245 missing in the maps (e.g., sites at far north in the arctic zone, or close to sea or water bodies) and  
246 these sites were therefore not included in the final models. We finally collected topographic data  
247 to test the potential effect of elevation and slope as indirect factors potentially influencing  
248 groundwater chemistry.

249  
250 Numerical analyses  
251 Numerical analyses were done in R version 3.6.3 (R Core Team, 2020), with the support of  
252 ArcGIS 10.2 (ESRI, Redlands, CA) for geoprocessing and map production. We first conducted

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253 exploratory analyses to test the [Ca and pH](#) prediction ability of [different](#) GIS layers related to soil  
254 bedrock, climate, and topography-[on the variation of both Ca and pH](#). We focused on layers  
255 with a complete coverage of Europe, with an eastern border from the Black Sea in Turkey to the  
256 White Sea in Russian Federation, thus including the regions with a relatively good cover of field  
257 measurements (Figure 2). We performed Linear Models for individual variables to select those  
258 providing significant relationships and > 1% of explained variance. A variable for soil pH  
259 (measured in water solution) at 15 cm depth for a 250 m grid resolution provided by the soilgrids  
260 project ([www.soilgrids.org](http://www.soilgrids.org)) had the highest explanatory power for GW-pH ( $R^2 = 0.22$ ) and GW-  
261 Ca ( $R^2 = 0.16$ ). The same results were obtained when using the same variable for 5 or 10 cm  
262 depth. We also tested soil estimates from Ballabio et al. (2019), but they provided weaker  
263 relationships for both GW-pH ( $R^2 = 0.14$  using soil pH as a predictor) and GW-Ca ( $R^2 = 0.01$   
264 using soil pH,  $R^2 = 0.01$  using soil  $\text{CaCO}_3$ ). To account for lithology, we used the lithological  
265 groups (litho3 level) included in the polygon layer of the Hydrogeological map of Europe  
266 (Duscher et al. 2015) as a categorical variable. We also selected annual precipitation (Bio12) as  
267 provided in CHELSA (Karger et al. 2017) to account for precipitation gradients which are  
268 expected to influence groundwater regimes. Other [CHELSA](#) variables related to precipitation  
269 were highly correlated with annual precipitation (Pearson  $r > 0.75$ ) and omitted. [Slope and](#)  
270 [elevation showed negligible effects on both GW-pH and GW-Ca \(linear regression, adjR2 <](#)  
271 [0.05, P < 0.001\)](#). Since preliminary models showed no differences with these variables were  
272 included, they were discharged.

273 The variables of lithology and soil pH were aggregated to the same grid extent of CHELSA at 1  
274 km resolution, [using the dominant unit and a bilinear interpolation, respectively as the most](#)  
275 [appropriate scale to balance the original scales of both layers. This grid extent is also the most](#)

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276 suitable spatial scale to be used in the context of further ecological modeling, which is in many  
277 cases combined with climatic data from e.g. CHELSA or WorldClim ([www.worldclim.org](http://www.worldclim.org)) for  
278 making temporal climatic projections. The lithological map (originally at 1:1,500,000 scale,  
279 which corresponds to a raster resolution of c. 1 km) was converted to a grid resolution using the  
280 dominant unit. Soil pH was converted from the original 250 m to 1 km grid resolution using a  
281 bilinear interpolation to create a smooth surface based on the weighted average of the four  
282 nearest cells.

283

284 **Figure 2.** Spatial distribution of the three groups of calibration data collected for modelling  
285 groundwater pH and Ca<sup>2+</sup> in European fens (original and literature data from springs and fens;  
286 data from streams from FOREGS Geochemical Atlas of Europe; other data). Other data include  
287 public data from national groundwater monitoring of Bulgaria and Slovenia. For separate maps  
288 of pH and Ca<sup>2+</sup> see Supplementary Figure 5.

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292 Spatial predictions were

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295 Because even the combination of the selected variables might not lead to precise fine-scale  
 296 indication of aquifer chemistry, we further employed a kriging approach to data analysis,  
 297 assuming an effect of spatial correlation to estimate the values close to the original samples. The  
 298 final spatial predictions were therefore, based on regression kriging (RK), a technique that  
 299 combines a regression model based on explanatory variables with the interpolation of model  
 300 residuals with ordinary kriging (Hengl et al. 2007; Meng et al. 2013). RK is especially  
 301 appropriate for modeling soil attributes at medium and large scales, combining the spatial  
 302 autocorrelation of soil variables with the explanatory power of auxiliary variables (Keskin and

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303 Grunwald 2018). We implemented RK with the GSIF R package (Hengl 2020). As the regression  
304 component, we computed Random Forests since a preliminary analysis with our data showed  
305 better performance than linear models, generalised linear models, or generalised additive models.  
306 Random Forests are ensemble learning methods based on decision trees and an internal  
307 correction of overfitting, which provide high interpretability and good performance when  
308 compared with other algorithms used in soil spatial modeling (Wiesmeier et al. 2011). Another  
309 advantage of Random Forests is that they have no requirements for considering the probability  
310 distribution of soil variables, fitting complex non-linear relationships for spatial extrapolation  
311 (Hengl et al. 2015).[2015\) that ultimately improve the spatial predictions.](#) We fitted the Random  
312 Forests model and the residual variogram for groundwater pH and Ca<sup>2+</sup> separately using the  
313 function fit.gstatModel() in GSIF package. [Effect plots for the predictors were created for the](#)  
314 [same models using partial\(\) function in pdp package \(Greenwell 2017\).](#) Spatial predictions were  
315 then computed with the predict() function using the model object generated previously and a 5-  
316 fold cross-validation. Model evaluation was based on the calculation of the Mean Error (ME)  
317 and the Root Mean Squared Error (RMSE) as the differences between predicted and observed  
318 values (Keskin and Grunwald 2018; Pham et al. 2019). We compared the relationships between  
319 the models [produced for both](#) groundwater pH and Ca<sup>2+</sup> by using a random sampling of 5,000  
320 points to extract cell values and computing a Pearson correlation. To assess regional differences,  
321 we correlated values grouped in 25 neighboring cells of each single cell using the  
322 rasterCorrelation() function in the [package spatialEco- R package \(Evans 2020\).](#)

323

324 **3. Results**

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325 In measured data, ranges and medians of pH and  $\text{Ca}^{2+}$  concentration ~~was~~were similar  
326 across Europe (Supplementary Figure 4), with ~~the~~the lowest pH values found in the Atlantic and  
327 Iberian regions~~and~~and ~~the~~the highest pH values found in southern Europe except Iberian Peninsula.  
328 ~~lowest~~The lowest  $\text{Ca}^{2+}$  values were found in boreal Europe, while the highest in Central and  
329 Southern Europe. The Random Forest models computed with the lithology, soil pH, and  
330 precipitation explained 40% and 55% of the variance for GW-pH and GW-Ca, respectively.  
331 Lithology was the variable with the highest importance in both models (Figure 3), although its  
332 effect was higher in the model computed for  $\text{Ca}^{2+}$  than for pH. ~~Conversely, soil~~These effects were  
333 mainly associated with the lithological units reflecting calcareous bedrocks, followed by  
334 categories with coarse and fine sediments such as flysch (Figure 4). SoilpH had higher relative  
335 importance in GW-pH than GW-Ca, ~~while~~although in both cases the variable had a similar  
336 positive effect. Finally, annualprecipitation had the lowest contributions in ~~both~~the two models  
337 (Figure 3), with both  $\text{Ca}^{2+}$  and pH dropping suddenly after the threshold of annual precipitation  
338 of ca 1800 mm (Figure 4), although the highest pH values occur under the lowest precipitation  
339 and tend to decrease towards high-precipitation areas.  
340  
341 **Figure 3.** Variable importance of Random Forest models computed for groundwater pH and  
342  $\text{Ca}^{2+}$ . MSE = Mean Standard Error.  
343  
344  
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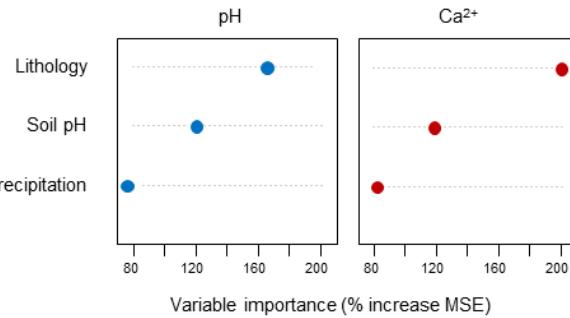
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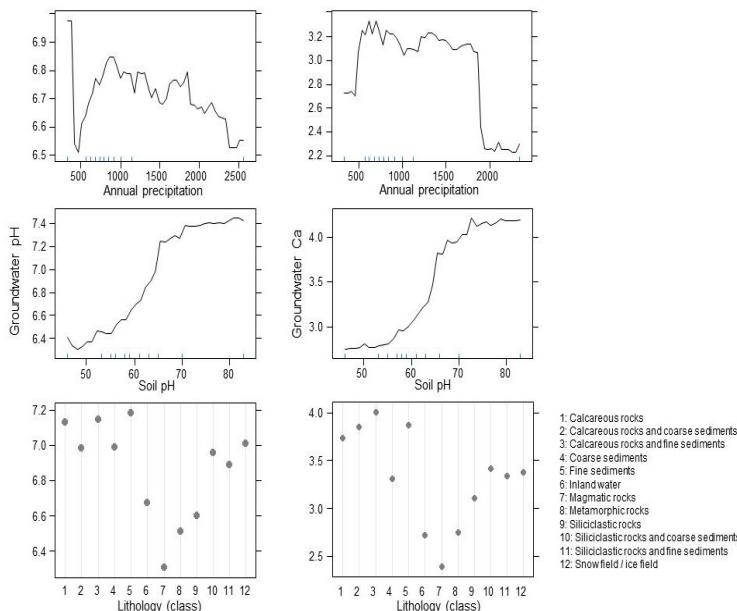
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347 **Figure 4.** Partial dependence plots showing the effects of the variables used in the Random348 Forest models computed for model groundwater pH and Ca<sup>2+</sup> in Europe.

349

350 When adding the kriging component, model predictions reached 65% and 74% of explained  
351 variance for GW-pH and GW-Ca, respectively. The mean values of standard errors (SE; 0.00058  
352 for pH; -0.0009 for Ca) and Root Mean Squared Errors (RMSE; 0.588 for pH; 0.690 for Ca)  
353 were higher in the models for pH, but in both cases showed low values and that suggest<sup>more</sup> accurate  
354 predictions, in agreement with their total explained variance.

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355 Model predictions for groundwater<sub>GW</sub>-pH reflected the lowest values in Scandinavia, Iceland,  
356 northern UK, and some regions of Central and Eastern Europe (Figure 45). The highest values  
357 were predicted in eastern Iberia and many regions of Central and Eastern Europe, although a big  
358 part of the study area was dominated by neutral pH values (6 to 7). The spatial patterns for GW-  
359 Ca (Figure 45) were rather similar to pH. The overall correlation between the two models was  
360 0.83 (Pearson r, P < 0.001), but they showed differences in some regions. This was supported by  
361 the spatial correlation computed for each cell (Figure 56), reflecting different magnitudes of  
362 correlation across the study area, especially in the eastern Iberian Peninsula and Southeastern  
363 Europe.

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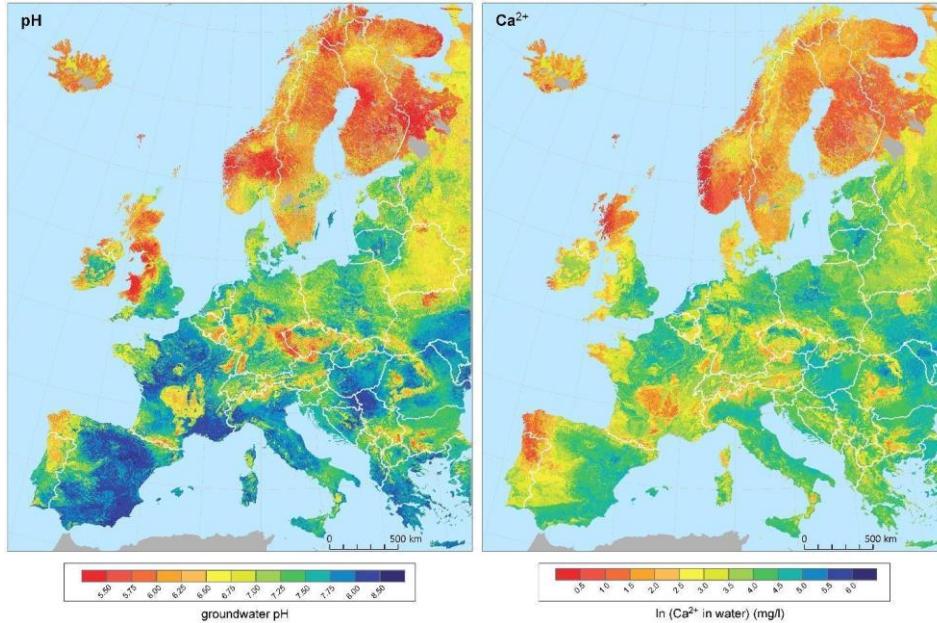
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364  
365 **Figure 45.** Model predictions based on Regression Kriging. Note the Ca<sup>2+</sup> concentration is on ln-  
366 scale.

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371 **Figure 56.** Spatial correlation between the models computed for groundwater pH and Ca<sup>2+</sup>.  
372 Values show Pearson correlation coefficient computed over every single cell by using a sampling  
373 of 25 neighboring cells.

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## 376 **Discussion**

### 377 3.1. Spatial patterns in groundwater pH and Ca<sup>2+</sup> concentration in Europe

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378 As expected, the recent values of water pH and Ca<sup>2+</sup> concentration are largely shaped by  
379 lithology in groundwater-dependent habitats across Europe. Indeed, it has been recognised by  
380 regional studies that the distribution of major spring and fen habitats, of which the species  
381 composition largely depends on pH and Ca<sup>2+</sup>, is well determined by bedrock type (Hájek et al.  
382 2002, Tahvanainen 2004, Hinterlang 2017, Peterka et al. 2017, Cantonati et al. 2020c). Since the  
383 European-scale geological map we used here is not precise enough to capture differences in  
384 bedrock chemistry within the major geologicallithological units that are defined largely by  
385 agegeological stratification, the contribution of soil pH in the model probably also reflected  
386 lithological variation, as soil pH generally correlates with regional bedrock chemistry (Chadwick  
387 and Chorover 2001). On the other hand, soil pH is also affected by climate-dependent  
388 pedogenesis, which incorporates a climate-zonal geographical component in this effect  
389 (Duchaufour 2012, Maxbauer et al. 2017). Precipitation is another important determinant of  
390 groundwater chemistry. Precipitation amount and frequency affect not only flow paths activity  
391 and redistribution of groundwater, but also its residence time in the aquifer, impacting  
392 carbonate dissolution and precipitation rates (Crossman et al.

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393 Precipitation is another determinant of groundwater chemistry in our study. High annual ←  
394 precipitation above ca 1800 mm obviously reduced2011, Lewandowski et al. 2015, Vystavna et  
395 al. 2020). Groundwater with a short transit time (1–3 years) or ‘young water’ (Soulsby et al.  
396 2015) can be particularly sensitive to changes in precipitation amount and frequency. Here,  
397 intensive precipitation may reduce an interaction time of groundwater with Ca<sup>2+</sup> and carbonates  
398 deposited in rocks (Fairchild et al. 1994, Segadelli et al. 2017, Cantonati et al. 2020b), resulting  
399 in lower Ca<sup>2+</sup> concentration of elements in groundwater. In snow-influenced ecosystems, This  
400 effect is more pronounced in snowy regions, where seasonal snowmelt can also

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401 modulates the recharge patterns of groundwater. Particularly, the duration of the  
402 snowmelt period can impact the occurrence and dynamic of preferential flow, and prolong or  
403 reduce the interaction of the seepage with soil and bedrock materials (Mohammed et al. 2019).  
404 We therefore suggest that fast hydrological pathways and short transit time driven by snowmelt  
405 and precipitation can explain the lowest Ca<sup>2+</sup> in hyper-oceanic cold regions of SW Norway or W  
406 Scotland. It may further explain lower pH and Ca<sup>2+</sup> values on windward slopes of high  
407 mountains, even if bedrock is moderately calcium-rich. Nevertheless, understanding the  
408 complementary effects of precipitation and slope will need to account for more accurate models  
409 based on GPS data, better precipitation data, and high-resolution (<250 m) topographic  
410 predictors.

411 The resulting pattern at the European scale is governed by the strong latitudinal and  
412 altitudinal gradients, i.e. decreasing pH and Ca<sup>2+</sup> northwards, and regionally also towards  
413 mountain regions. This pattern largely follows bedrock chemistry, with crystalline rocks  
414 prevailing, and most carbonate rocks being eroded by glaciers, in high latitudes and altitudes.  
415 The excess of precipitation over evaporation, and theoretically also slower weathering rates in  
416 colder regions (White & Blum 1995), contribute as well. Although this pattern is well known  
417 (Økland et al. 2001, Hájek et al. 2006, Hinterlang 2017, Peterka et al. 2017) and has been  
418 captured also by the FOREGS Geochemical Atlas of Europe (Salminen et al. 2006), our  
419 improved model provides much finer regional patterns. In Southern Europe, low pH and Ca<sup>2+</sup>  
420 values were modelled in the Pyrenees, the Balkans, SW Corse, and Calabria, i.e. the regions  
421 where boreal or endemic types of fen communities occur as relicts (Chytrý et al. 2020). The  
422 Alps, the Apennines, the Carpathians, and the Baltic region show a fine-scaled mosaic of  
423 alkaline (calcium-rich) and acidic (calcium-poor) groundwater that contributes to the high

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424 diversity and conservation value of groundwater-dependent ecosystems, such as fens (Cantonati  
425 et al. 2009, 2011, Gerdol et al. 2011, Joosten et al. 2017, Horskáková et al. 2018). The most  
426 apparent “acidic island” in Central Europe is located in the SW part of the Bohemian Massif  
427 (Czech Republic, Germany), where acidic types of springs and fens are quite frequent and some  
428 studies further document anthropogenic acidification on siliceous bedrock in 1970–80s, being re-  
429 emerged recently because of extreme climatic events (Kapfer et al. 2012, Schweiger et al. 2015).  
430 It is, however, possible that particularly this acidic island is picked out mainly because of the  
431 high amount of available data.

432 Clearly, most of Fennoscandia is markedly acidic and calcium-poor mainly due to glacial  
433 history. Yet, the model identified small alkaline and calcium-enriched islands in NE and Central  
434 Sweden, and NW Norway, which are associated with rare types of calcareous fen and spring  
435 communities (Dierssen 1982, Vorren et al. 1999, Udd et al. 2015, Miller et al. 2020). More  
436 localised pockets of calcareous habitats are however known from most parts of Fennoscandia  
437 that are not recognised with the grain of our European-wide analysis. With our results, the future  
438 modelling of diversity and distribution of individual habitat types of groundwater-dependent  
439 wetlands will be more reliable, regionally. Regionally rare habitat conditions can/will be better  
440 recognised better, and the disentangling of the climate and pH effects will be more easily,  
441 feasible.

### 442 3.2. Persisting data gaps and further improvements

443 Although being based on the hitherto most comprehensive field data set currently  
444 available, the presented map cannot be considered definitive. Surely there are many pH and EC

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446 or Ca<sup>2+</sup> measurements conducted across Europe that we could not include into the data set  
447 because they are hardly accessible. Except for Russian Federation and Moldova, largest gaps still  
448 occur in the southern parts of the Pannonian plain (southeastern Hungary, northern Serbia, and  
449 western Romania), in SE Belarus, and eastern Ukraine. We dispose ofhave available some data  
450 from the latter region (Vystavna et al. 2015; Supplementary Table 2), but a large gap in the rest  
451 of the data set prevented reliable geospatial modelling. These data might be used in future  
452 updates of the map once the gap in Central Ukraine is filled. The lack of data in the Pannonian  
453 plain has led to poor correlation between predicted pH and Ca<sup>2+</sup> values (Fig. 4). Such a poor  
454 correlation and sometimes low density of data apply also for some other lowland regions, such as  
455 the Danube plain in S Romania, Po valley in Italy and valleys around the Duero, Ebro, and Tagus  
456 rivers in Spain. Apart from eutrophication, this resultpattern may be caused by the imbalanced  
457 distribution of groundwater-dependent habitat types in our data set. Unlike mountain regions, the  
458 data for these lowlands were largely taken from the FOREGS Geochemical Atlas of Europe  
459 (Salminen et al. 2006) and national groundwater databases, i.e. largely from stream water.  
460 Considering the major purpose of our map, the (creating new data sources for ecological  
461 modelling of fens and springs), these regions are generally not as importantless crucial for  
462 biodiversity modeling because of the low number of existing target habitats, as they have  
463 largely been transformed to arable land or they are too dry. On the other hand, caution is needed  
464 when interpreting the maps in an ecological sense. The extremely high pH (> 8) and Ca<sup>2+</sup> ( $\text{In}$   
465  $[\text{Ca}^{2+}] > 4$ ; i.e.,  $\text{Ca}^{2+} > 55 \text{ mg.l}^{-1}$ ) values that occur in lowlands visually govern the map, but for  
466 ecological differentiation of groundwater-dependent habitats in Europe the differences within the  
467 middle part of the gradient, i.e. between pH 5.5 and 7.0, are much more important (Malmer  
468 1986, Wheeler & Proctor 2000, Hájek et al. 2006, Rydin and Jeglum 2013). In any case, our

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469 Our data set is expected to be amended in the future, as more studies will be published and more  
470 data will be available, so further versions will be accessible in the open repository. New data  
471 will help to improve predictions for those regions with relatively lower sampling effort, and also  
472 those with a lithologically heterogeneous landscape. Future updates of the model may  
473 also focus at finer spatial resolution (e.g., 100 to 250 m) but this will require to increase the  
474 spatial accuracy of the calibration data and the predictor variables, as in some Central-European  
475 areas (Le et al. 2019, Chuman et al. 2019). Although we tested several variables with potential  
476 predictive effect on groundwater pH and Ca, many of them had lower explanatory power (e.g.,  
477 Ballabio et al. 2019) or they were redundant with the soil pH layer we used (i.e. previous  
478 versions of soil grids using the same data sources). The low predictive value of other predictors  
479 with potential predictive value, like slope, elevation or other precipitation variables, is probably  
480 related to their broad spatial scale, with values averaged at 1 km grid resolution having little  
481 impact to discriminate groundwater variation at the landscape level (Jiménez-Alfaro et al.  
482 2018b). This contrasts with the predictive value of the lithological layer, which is however based  
483 on a similar spatial (but originally vectorial) resolution. However, lithological bedrock is  
484 generally more homogeneous at the landscape scale, with the only exception of certain  
485 geological complex regions. These findings suggest that future improvements of our models will  
486 depend on the quality of new lithological (or related soil chemistry) variables with direct effect  
487 on groundwater pH and Ca. We also note that the lithological map we used here is simplified to  
488 large units, while many regional and national geological maps are being produced at finer details.  
489 The combination of such new predictors with accurate calibration data at the continental level  
490 and at fine scale resolution seems the most likely opportunity to produce significantly better

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491 models, since the influence of spatial-dependent (kriging) effects is rather limited by the  
492 distribution of sample points.

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#### 494 Conclusions

495 Here, we provide the first European map of groundwater pH and Ca<sup>2+</sup> content. We  
496 collected as even as possible distributed field measurements of water pH and Ca<sup>2+</sup> or EC from  
497 European groundwater-dependent habitats. Despite these persisting gaps, our European map of  
498 near-surface groundwater pH and EC provides the best solution for modelling the biodiversity  
499 of groundwater-dependent ecosystems, especially at the continental or supra-regional scale.  
500 We even believe that this map could be more suitable also for ecological modelling of other  
501 than groundwater-dependent habitats. It may mirror the bedrock chemistry better than the  
502 map of soil pH, because soil pH is a resultant of pedogenetic processes, which are tightly  
503 associated with the character of the vegetation cover itself (Miles 1985, Duchaufour 2012).

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504 ▲

505 Conclusions

506 Here, we provide the first European map of groundwater pH and Ca<sup>2+</sup> content. We collected as  
507 even as possible distributed field measurements of water pH and Ca<sup>2+</sup> or EC from European  
508 groundwater-dependent habitats, having high data density with relatively higher sampling effort  
509 in regions rich in endangered groundwater-dependent ecosystems (springs, fens). Despite the  
510 general high accuracy of our models, we note that prediction uncertainties may affect the  
511 reliability of models computed with both Random Forests and kriging (Hengl et al. 2018;  
512 Szatmári & Pásztor 2019). Another source of prediction uncertainty is related to the quality of

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513 the original chemical measurements and used geospatial modelling. The model considered  
514 predominantly the georeferentiation of their geographic position. Moreover, the predictor  
515 variables rely on spatial models (soil pH, precipitation) or broad geographic maps (lithology)  
516 which are based on their own uncertainties and assumptions. Future Improvements of  
517 groundwater pH and Ca estimates should therefore consider a more accurate set of response  
518 variables and fine-scale predictors, preferably including lithology and soil pH (i.e., variables  
519 surrogating bedrock chemistry) and precipitation sum (i.e., residence time of groundwater). Our  
520 results also Despite potential uncertainties and data gaps, this study uses an unprecedented  
521 combination of data to provide a freely accessible map and realistic maps that can be used in any  
522 kind of spatial modelling, showing better resolution and fewer gaps than previously published  
523 maps. The character of our input data, which are also freely accessible, predetermines our map  
524 for being used in ecological modelling to address the distribution and diversity of groundwater-  
525 dependent ecosystems and associated species. The character of our input data, which are also  
526 freely accessible, predetermines our map for being used in ecological modelling to address the  
527 distribution and diversity of groundwater-dependent ecosystems and associated species.  
528 Moreover, we assume the map will represent the best choice also for other types of earth  
529 modeling, because unlike previous maps we included Eastern-European and Balkan countries  
530 and considered lithology in geospatial modelling. We even believe that our maps could be also  
531 suitable for ecological modelling of other than groundwater-dependent habitats. It may mirror  
532 the bedrock chemistry better than the map of soil pH, because soil pH is a resultant of  
533 pedogenetic processes, which are tightly associated with the character of the vegetation cover  
534 itself (Miles 1985, Duchaufour 2012). We conclude that our European maps of near-surface  
535 groundwater pH and EC provides the best solution currently available for modelling the

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536 biodiversity of groundwater-dependent ecosystems, especially at the continental or supra-  
537 regional scale.

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538

539 **Data availability**

540

541 The dataset of georeferenced pH and EC measurements and the resulting maps in GIS-  
542 compatible format (shapefile) are accessible at [www.zenodo.org](http://www.zenodo.org); doi  
543 10.5281/zenodo.4139912 (Hájek et al. 2020).

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544

545 **Code availability**

546 No original R code was used; the used codes are cited.

547

548 **Sample availability**

549 No geoscientific samples registered as International Geo Sample Number (IGSN) have been used  
550 for the manuscript.

551

552 **Team list**

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563

564 **Author contribution**

565 M.H. and B.J-A. contributed equally to the paper. They conceived the research, collected data  
566 and led writing. B.J-A. designed and performed Random Forest and Regression Kriging models.  
567 O.H. prepared the input data and final map outputs. All authors provided unpublished data and  
568 commented on the manuscript.

569

570 **Competing interests**

571 The authors declare that they have no conflict of interest.

572

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587

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