Synoptic Analysis of a Decade of Daily Measurements of SO₂ Emission in the Troposphere from Volcanoes of the Global Ground-Based Network for Observation of Volcanic and Atmospheric Change

Santiago Arellano [1], Bo Galle [1], Fredy Apaza [2], Geoffroy Avard [3], Charlotte Barrington [4], Nicole Bobrowski [5], Claudia Bucarey [6], Viviana Burbano [7(+)], Mike Burton [8,a], Zoraida Chacón [7], Gustavo Chigna [9], Christian Joseph Clarito [10], Vladimir Conde [1], Fidel Costa [4], Maarten De Moor [3], Hugo Delgado-Granados [11], Andrea Di Muro [12], Deborah Fernandez [10], Gustavo Garzón [7], Hendra Gunawan [13], Nia Haerani [13], Thor H. Hansteen [14], Silvana Hidalgo [15], Salvatore Inguaggiato [8], Mattias Johansson [1], Christoph Kern [16], Manne Kihlman [1], Philippe Kowalski [12], Pablo Masias [2], Francisco Montalvo [17], Joakim Möller [18], Ulrich Platt [5], Claudia Rivera [1,b], Armando Saballos [19], Giuseppe Salerno [8], Benoit Taisne [4], Freddy Vásconez [15], Gabriela Velásquez [6], Fabio Vita [8], Mathieu Yalire [20]

[1] Chalmers University of Technology (Chalmers), Department of Space, Earth and Environment, Sweden
[2] Instituto Geológico, Minero y Metalúrgico (INGEMMET), Peru
[3] Observatorio Vulcanológico y Sismológico de Costa Rica (OVSICORI), Costa Rica
[4] Nanyang Technological University, Earth Observatory of Singapore (EOS), Singapore
[5] Heidelberg University, Institute of Environmental Physics, Germany
[6] Servicio Nacional de Geología y Minería (SERNAGEOMIN), Chile
[7] Servicio Geológico Colombiano (SGC), Colombia
[8] Istituto Nazionale di Geofisica e Vulcanologia (INGV), Italy
[9] Instituto Nacional de Sismología, Vulcanología, Meteorología e Hidrología (INSIVUMEH), Guatemala
[10] Philippine Institute of Volcanology and Seismology (PHIVOLCS), Philippines
[11] Universidad Nacional Autónoma de México (UNAM), Instituto de Geofísica, Mexico
[12] Institut de Physique du Globe de Paris (IPGP), Observatoire Volcanologique du Piton de la Fournaise, France
[13] Center for Volcanology and Geological Hazard Mitigation (CVGHM), Indonesia
[14] GEOMAR Helmholtz Centre for Ocean Research Kiel, Germany
[15] Escuela Politécnica Nacional, Instituto Geofísico (IGEPN), Ecuador
[16] United States Geological Survey, Volcano Disaster Assistance Program (USGS/VDAP), United States
[17] Servicio Nacional de Estudios Territoriales (SNET), El Salvador
[18] Möller Data Workflow Systems AB (MolFlow), Sweden
[19] Instituto Nicaragüense de Estudios Territoriales (INETER), Nicaragua
[20] Observatoire Volcanologique de Goma (OVG), DR Congo
[+] Deceased
[a] Now at University of Manchester, School of Earth, Atmospheric and Environmental Sciences, United Kingdom
[b] Now at Universidad Nacional Autónoma de México, Centro de Ciencias de la Atmósfera, Mexico

Correspondence to: Santiago Arellano (santiago.arellano@chalmers.se)

Abstract. Volcanic plumes are common and far-reaching manifestations of volcanic activity during and between eruptions. Observations of the rate of emission and composition of volcanic plumes are essential to recognize, and in some cases predict, the state of volcanic activity. Measurements of the size and location of the plumes are important to assess the impact of the emission from sporadic or localized events to persistent or widespread processes of climatic and environmental importance.

These observations provide information on volatile budgets on Earth, chemical evolution of magmas and atmospheric circulation and dynamics. Space-based observations during the last decades have given us a global view of Earth's volcanic emission, particularly of sulphur dioxide (SO_2). Although none of the satellite missions were intended to be used for measurement of volcanic gas emission, specially adapted algorithms have produced time-averaged global emission budgets. These have confirmed that tropospheric plumes, produced from persistent degassing of weak sources, dominate the total emission of volcanic SO_2. Although space-based observations have provided this global insight into some aspects of Earth's volcanism, it still has important limitations. The magnitude and short-term variability of lower-atmosphere emissions, historically less accessible from space, remain largely uncertain. Operational monitoring of volcanic plumes, at scales relevant for adequate surveillance, has been facilitated through the use of ground-based scanning-differential optical absorption spectrometers (ScanDOAS) since the beginning of this century, largely due to the coordinated effort of the Network for Observation of Volcanic and Atmospheric Change (NOVAC). In this study, we present a compilation of results of homogenized post-analysis of measurements of SO_2 flux and plume parameters obtained during the period March 2005 to January 2017 on 32 volcanoes in NOVAC. This inventory opens a window into the short-term emission patterns of a diverse set of volcanoes in terms of magma composition, geographical location, magnitude of emission, and style of eruptive activity. We find that passive volcanic degassing is by no means a stationary process in time and that large sub-daily variability is observed in the flux of volcanic gases, which has implications for emission budgets produced using short-term, sporadic observations. The use of a standard evaluation method allows intercomparison between different volcanoes and between ground- and space-based measurements of the same volcanoes. The emission of several weakly degassing volcanoes, undetected by satellites, is presented for the first time. We also compare our results with those reported in the literature, providing ranges of variability in emission, not accessible in the past. The open-access data repository, introduced in this article, will enable further exploitation of this unique dataset, with a focus on volcanological research, risk assessment, satellite-sensors validation, and improved quantification of the prevalent tropospheric component of global volcanic emission.

Data sets for each volcano are made available at https://novac.chalmers.se, under license CC-BY 4, or through the DOI-links provided in Table 1.

1 Introduction

Volcanic eruptions are to a large extent triggered or modulated by the intricate dynamics of segregation and escape of volatiles from magmas, making the observation of the rate of gas emission an important component of monitoring efforts to identify and predict the state of a volcanic system (Sparks, 2003; Sparks et al., 2012). The resulting atmospheric plumes are the farthest-reaching products of volcanic activity and constitute rich environments for a number of important processes affecting the physics and chemistry of the atmosphere, the radiative balance of the climate system, or the biogeochemical impact on soils and the ocean (e.g., Robock, 2000; Langmann, 2014; Schmidt et al., 2018).

Volcanoes are sources of many trace atmospheric compounds, such as water vapor (H_2O), carbon dioxide (CO_2), sulphur dioxide (SO_2), carbonyl sulphide (OCS), hydrogen chloride (HCl), hydrogen fluoride (HF), hydrogen sulphide (H_2S), molecular hydrogen (H_2), as well as of solid particles and metals. From these species, SO_2 is the most widely observed by passive optical remote sensing methods (Oppenheimer, 2010). This is a consequence of its low atmospheric background and accessible radiation absorption bands, particularly in the near-ultraviolet (NUV) and mid-infrared (MIR) spectral regions. This is advantageous for several reasons, for example: for (1) the volcanologist, because SO_2 is a reliable tracer of magmatic activity due to its strongly pressure-dependent solubility in magmas. Since H_2O is usually the most abundant volatile species and thus the most important driver of volcanic activity, and has a pressure-dependent solubility, both H_2O and SO_2 fluxes are positively correlated with eruptive intensity. For (2) the climatologist, because SO_2 may be transformed by a series of reactions into sulfuric acid (H_2SO_4)-containing aerosols, which exert a strong radiative forcing, especially when reaching the stratosphere. Or, for (3) the meteorologist, because SO_2 has a long enough residence time in the atmosphere to serve as a tracer of volcanic plume transport at regional or even global scales.

Measurements of the mass emission rate or flux of SO_2 from volcanoes started in the 1970s with the development and application of the Correlation Spectrometer (COSPEC) (Moffat and Millán, 1971; Stoiber et al., 1973). This instrument disperses ultraviolet sky radiation using a grating and employs a mechanical mask to correlate the intensity of diffused solar radiation in the near ultraviolet region at selected narrow bands, matching absorption features of SO_2. With proper calibration using cells containing SO_2 at known concentrations, the COSPEC measures the column density of SO_2 relative to background by the methods of differential absorption. Flux is quantified assuming mass conservation: the volcanic source emission strength is equal to the integrated flux across a surface surrounding the volcano, when no other sources or sinks are enclosed. The integrated flux is measured by scanning through a surface perpendicular to plume transport, integrating the column densities in the plume cross-section, and multiplying this integral by the corresponding transport speed. COSPEC was typically used for sporadic or periodical field surveys, during both volcanic crises and periods of passive degassing. The first global emission budgets for volcanic SO_2 were based on extrapolation of these sporadic measurements on a fraction of globally degassing volcanoes, through a series of non-verified assumptions regarding the statistics of emission for measured and non-measured sources. Halmer and Schmincke (2002) recognized this problem and highlighted the need for increasing: (i) the number of monitored volcanoes, (ii) the periods of observation, (iii) the sampling frequency of the measurements, and (iv) the homogeneity of protocols of measurement by different observers.

In the late 1970s, the first satellite-based sensors, intended primarily for monitoring the stratospheric ozone (O_3) layer, opened up the possibility to map and quantify volcanogenic SO_2 from space (Krueger, 1983; Krueger et al., 1995). The successful Total Ozone Mapping Spectrometer (TOMS) instrument programme was succeeded by a series of optical instruments such as the Global Ozone Monitoring Experiment (GOME/GOME-2), the Scanning Imaging Absorption Spectrometer for Atmospheric Cartography (SCHIAMACHY), the Ozone Monitoring Instrument (OMI), and the Ozone Mapping & Profiler Suite (OMPS). Infrared (IR) sensors, such as Infrared Atmospheric Sounding Interferometer (IASI) or Moderate Resolution Imaging Spectroradiometer (MODIS), have been also used for routine global observation of volcanic emissions (Khokhar,

105 2005; Carn et al., 2013; Theys et al., 2013). More recently, the Tropospheric Monitoring Instrument (TROPOMI), on-board ESA's Sentinel-5 Precursor satellite since 2017, achieves a factor of 3 to 4 better sensitivity than OMI, due to better spatial resolution and sensor performance. This makes detection of weak emissions of SO_2 in the lower atmosphere feasible, every day and with global coverage. Under ideal measurement conditions and knowledge of plume velocity, time series of volcanic SO_2 flux as low as ~1 kg/s (for 1 m/s wind speed) with sub-daily frequency can be derived from TROPOMI (Queiβer et al.,

110 2019; Theys et al., 2019).

During the 1990s and early 2000s smaller, cheaper and more accurate and versatile alternatives to the COSPEC were developed, in particular the Miniaturized Differential Optical Absorption Spectrometer (MiniDOAS) (Galle et al., 2003). This instrument incorporates a grating spectrometer to obtain the spectrum of diffused solar radiation in the UV and retrieves the relative column density of SO_2 by the DOAS method (Platt and Stutz, 2008). This line of research led to the implementation

115 of fully automated Scanning-DOAS (or ScanDOAS) systems (Edmonds et al., 2003), which has enabled volcanological observatories to conduct nearly continuous monitoring of volcanic plumes. A version of this instrument known as Dual-Beam Scanning-DOAS can measure the plume velocity, height and the integrated SO_2 flux in near to real-time, with a time resolution of 1-15 min during daylight hours (Johansson et al., 2009a). Similar spectroscopic instruments have been developed or replicated by different groups (Horton et al., 2006; Mori et al., 2007; Arellano et al., 2008; Burton et al., 2008).

120 Among other methods for ground-based optical remote sensing of integrated volcanic flux we highlight different types of imaging systems such as Imaging-DOAS (I-DOAS) (Bobrowski et al., 2006; Louban et al., 2009), thermal imaging Fourier Transform Infra-Red (FTIR) (Stremme et al., 2012), as well as UV and IR SO_2 cameras based on broad-band filters or interferometry (Mori and Burton, 2006; Bluth et al., 2007; Kern et al., 2010b; Kuhn et al., 2014; Prata and Bernardo, 2014; Platt et al., 2015; McGonigle et al., 2017; Smekens et al., 2018). A crucial advantage of these systems, compared with the

125 ScanDOAS, is their higher temporal resolution and accurate quantification of plume speed by image-correlation techniques. Among the disadvantages we mention that they usually require more restricted measurement conditions with respect to measurement geometry and weather, have higher susceptibility to interference (e.g. aerosols), are usually designed for measurement of a single species and require calibration by another instrument, usually a MiniDOAS.

An important step towards extending the newly available tools for permanent volcanic gas monitoring has been the creation

130 of the Network for Observation of Volcanic and Atmospheric Change (NOVAC) in 2005. The network was established with funding from the European Union during 2005-2010 and it has continued and expanded with resources from volcanological observatories and cooperating research groups, the Deep Carbon Observatory programme (https://deepcarbon.net/), the Volcano Disaster Assistance Programme (VDAP) of the United States Geological Survey (USGS) and the United States Agency for International Development (USAID), and Chalmers University of Technology. The main purpose of the NOVAC

135 project was to set up local monitoring networks of Dual-Beam ScanDOAS instruments. It started with 15 volcanoes monitored by observatories in Latin America, D.R. Congo, La Reunion Island, and Italy; involving 18 different groups with expertise in volcanology, atmospheric remote sensing and meteorology. At the time of writing, NOVAC has included about 160 stations at 47 volcanoes in different regions around the world, now including Iceland, The Philippines, Indonesia, Papua New Guinea,

and Montserrat. The advantages of these instruments with respect to spaceborne sensors include continuous calibration, better temporal and spatial resolution, more direct measurement of flux, and better sensitivity to tropospheric plumes. A key disadvantage is the limited spatial coverage inherent to ground networks. Details of the instrument and operation routines are given in Galle et al. (2010). Figure 1 shows a map with locations of the volcanoes that have been part of NOVAC.

(Fig. 1)

The purpose of this paper is to present an inventory of daily flux measurements of SO_2 obtained in NOVAC from 1 March 2005 until 31 January 2017. These results were obtained by standardized re-evaluation of the collected spectra, incorporating information of wind velocity from a re-analysis dataset provided by the European Centre for Medium-Range Weather Forecasts (ECMWF). We present daily statistics of emission and corresponding information of plume parameters. A database for access to the results is described in detail, providing a substantial basis for further investigations of volcanic degassing patterns over time. We compare the emission inventory of NOVAC with past compilations of degassing intensity on these volcanoes. These topics determine the structure of the paper.

2 Methods

2.1 The dual-beam ScanDOAS instrument and real-time operation

NOVAC is a network of dual-beam ScanDOAS instruments. This is a well-established technique that has been described in detail elsewhere (Johansson et al., 2009a; Galle et al., 2010; Galle et al., 2011). There are two types of NOVAC instruments: "Mark I" are more robust and simpler, designed for routine long-term monitoring; and "Mark-II", with more sophisticated optics and spectrometer, developed for more specific scientific observations (Kern, 2009). The results of this study correspond to measurements with Mark-I systems, which comprise more than 95% of installations (and >99% of collected data).

A typical volcano in NOVAC is monitored by two or three ScanDOAS instruments, located within 10 km distance from the main volcanic vent. The objective is to guarantee as complete azimuthal coverage of the volcanic plumes as determined by wind patterns and permitted by logistical constraints. The selection of the sites for installation should also consider aspects of (i) altitude: neither too high to obtain clear atmospheric spectra outside of the plume, nor too low to avoid obstacles in the viewing directions of the instrument; (ii) distance from the vent: neither too close where turbulence and the optical thickness of the plume may affect the quality of the measurements, nor too far where atmospheric dispersion and depletion processes take a dominant role making difficult quantification of the source emission; and (iii) orientation of the scanning path: flat or conical, to maximize the probability of intercepting the plume with overlapping scanning paths of several stations, which is used for calculation of plume location by triangulation.

The fore-optics of the dual-beam ScanDOAS consists of a scanning telescopic system with left-hand orientation defining roll, that is, scanning angle between −180 and +180 deg in steps of 3.6 deg; pitch, i.e. the conical (60 deg) or flat (90 deg) angles of the scanner; and, yaw, or azimuth angle, usually oriented towards the volcano. The telescope consists of a single plane-convex quartz lens with a diameter of 25.4 mm and a focal length of 7.5 cm, as well as a Hoya (U330) UV filter that reduces

intensity of light with wavelengths longer than 360 nm. The telescope is coupled to one (single-beam) or two (double-beam) quartz optical fibre(s) with a diameter of 600 μm. This combination gives an effective field of view of 8 mrad. The optical fibre is coupled to the entrance slit of the spectrometer, which has a width of 50 μm and height of 1 mm. The spectrometer (SD2000 from OceanOptics) has a crossed Czerny-Turner configuration with a grating of 2400 lines/mm operating in reflection and a UV-enhancement-coated, uncooled, linear Charge-Couple-Device (CCD) detector (ILX511b from Sony) of 2048 14×200 μm effective pixels, as well as an Analog-to-Digital Converter (ADC) of 12 bits. The effective spectral range of the spectrometer is ~275-480 nm, the spectral resolution (FWHM) ~0.5 nm and the pixel resolution for the combination of grating and slit is ~6.5 pixels. Signal-to-Noise (S/N) at 50% of saturation is ~500:1 for an average of 15 spectra taken at typical (~500 ms) exposure time.

Data are transferred via a serial port to the instrument computer. Three versions of the control unit have been developed over the years, all of them industrial-grade, running on a Linux-OS and including serial, USB2 and ETH communication ports. Serial ports are used for communication with the spectrometer and control of the scanner's stepper motor. The USB port can be used for powering the spectrometer, while the ETH port is usually used for data transfer to radio modems. Other peripherals include a digital thermometer (for record of internal temperature), a voltmeter (for control of battery voltage) and a GPS antenna (for recording location and time).

The NOVAC instruments are usually powered by an array of 12 V batteries and solar panels, the power consumption at full operation is 6.9-10 W (depending on the model of computer) and communicated with the observatories by radio modem telemetric networks (usually at 900-930 MHz). A timer is added to interrupt operation of the instrument during night and trigger a reset of the instrument in the morning. Data are collected in-situ (a 4 GB compact flash card can keep compressed-format data for up to ~6 weeks before older data is overwritten) or transmitted for real-time evaluation and display of results with the software called NovacProgram (Johansson, 2009b). Raw and analysed data are archived in a server hosted in Gothenburg and mirrored in Brussels and Heidelberg. This server is accessible to members of the network.

The standard protocol for a measurement of the flux of SO_2 begins with a determination of the exposure time required for an adequate (typically about 65%) saturation of the spectrometer detector. For this, the scanner is moved to 0 degrees scan angle (closest to zenith) and the exposure time is adjusted to a value between 50 and 1000 ms. Next, a preliminary Fraunhofer reference spectrum is measured at 0 deg scan angle, then a dark spectrum is recorded at 180 deg scan angle (closest to obstructed-view nadir) followed by a total of 51 measurements of skylight from scanning angles -90 deg to 90 deg at steps of 3.6 deg. Each measured spectrum consists of 15 co-added spectra to increase S/N. A full scan is collected every 1-15 min, depending on illumination conditions. Data are spectrally analysed using the DOAS method (Platt and Stutz, 2008), evaluated in the spectral range 310.6-324.6 nm. Before the analysis, generic corrections for dark current, electronic offset and wavelength shift (based on absorption features of SO_2) are applied. The spectral analysis model includes absorption spectra of SO_2 at 293K, 1000 mbar (Bogumil et al., 2003) and O_3 at 223K and 100 mbar (Voigt et al., 2001), as well as a Ring-effect pseudo-absorber synthesized from the Fraunhofer spectrum using the software DOASIS (Kraus, 2006) and a 5th degree polynomial to account

for broad-band extinction. Molecular absorption cross-section spectra are retrieved from the MPI-Mainz UV/VIS spectral atlas

205 (Keller-Rudek et al., 2013). For the DOAS analysis, a convolution is applied to the high-resolution spectra with the instrumental function of each instrument, which is approximated from the 302.15 nm emission line of a low-pressure Hg lamp measured at room temperature before installation. All calibration data are archived in the data server.

Along with instrument endurance, data acquisition and analysis has also been designed to guarantee compensation of measurement errors and traceability of measurement conditions. The compensation is achieved by acquiring reference and

210 dark spectra on each scan, which permits an efficient cancellation of instrumental imperfections, which may change over time, because the measurements are taken within minutes of each other. The traceability is ensured by logging all measurement parameters, which allows rigorous scrutiny of the quality of the measurements and offers the possibility to apply more advanced algorithms in the future that make use of this auxiliary information (e.g., instrument line shape models that account for the effect of temperature).

215 The NOVAC instruments have proven to be remarkably robust, particularly given the harsh conditions they are regularly exposed to. Instruments are often installed at high elevation, exposed to large temperature and humidity variations, and experience ash-fall or even exposure to highly acidic volcanic gases. The simple design of the instruments and the separation of the optical scanner from the rest of the instrumentation are the keys to their robustness. This, combined with the strong sense of community within the NOVAC consortium, has led to the growing number of scanners installed at active volcanoes

220 around the world.

2.2 Batch-processing with the NOVAC Post-Processing-Program

As mentioned above, data are transmitted to the observatories, where it is analysed and archived in real-time using the NovacProgram. This evaluation uses meteorological information (wind speed and direction) provided by each operator, it may thus vary enormously in quality among different observatories. Additionally, the combination of nearly simultaneous

225 measurements from intercepting scanning paths of two instruments is used to calculate plume height and direction in real-time. Also plume speed can be derived from measurements of a single instrument by the dual-beam cross-correlation method when certain conditions are fulfilled regarding direction, strength and stability of the plume (Johansson et al., 2009b; Galle et al., 2010).

In order to adopt a standardized methodology for the evaluation of data collected by each station, a program called NOVAC

230 Post-Processing-Program (NovacPPP) was developed by Johansson (2009b). This program retrieves all scan measurements collected at a specified volcano within a given period and proceeds to evaluate them selecting the best information available for each variable. For instance, measurements of plume speed, direction and height are prioritized over information obtained from a meteorological model or from common assumptions (e.g. plume height equal to difference in altitude between volcano summit and scanner).

235 In this work we used wind speed from the ERA-interim re-analysis database of ECMWF, which is based on the IFS Cycle 31r2 4D-Var assimilation system, using a TESSEL land-surface model. This database, with a coverage period since 1979 until

2019, has an assimilation period of 12 h, while the spatial resolution is 79 km (TL255) in the horizontal and 60 vertical levels from sea level pressure up to 10 hPa, with a typical difference equivalent to ~200 m (Dee et al., 2011). For each volcano, horizontal wind vectors, relative humidity and cloud cover are retrieved every 6 h on a horizontal grid of 0.125×0.125 deg (13.9×13.9 km for mean Earth radius) surrounding the location of the main volcanic vent. It is then further interpolated to the vent location and the time of each scan.

The program also applies a correction for spectral shift (i.e., the possible change in the pixel-to-wavelength mapping of the spectrometer during operation), based on correlation of the position of the Fraunhofer lines in the measured spectrum with those of a high-resolution solar spectrum (Chance and Kurucz, 2010) adjusted to the resolution of the spectrometer. Other than these changes, the evaluation follows the same routines as the standard NOVAC program, specified above.

2.3 Uncertainty of SO₂ mass flow rate measurements

It is difficult to assign a "typical" uncertainty to measurements of flux with NOVAC instruments because the flux calculation depends on different variables and assumptions, which are subject to a wide range of conditions (meteorology, distance to plume, content of aerosols, amount of absorber, etc.). Detailed analysis performed by Arellano (2014) shows that the range of uncertainty can be as low as 20-30% or as large as >100%. Categories of uncertainty include model, measurement and parameter uncertainties. Model uncertainty refers to the plausibility that a certain measurement scenario is realized in practice. For example, the assumption that transmittance can be calculated from simple application of Bouguer-Beer-Lambert law may not hold due to radiative transfer effects (Millán, 1980; Mori et al., 2006; Kern et al., 2010a), or the model adopted for the geometrical shape of the plume may not be adequate. Measurement uncertainty could be induced, for example, by inaccurate determination of the viewing direction of the scanner, or variations in the spectrometer response caused by changing environmental conditions. Parameter uncertainty could e.g. be caused by inaccuracy of a laboratory absorption cross-section or the uncertainty in plume speed data derived from a mesoscale meteorological model.

If we split the analysis into the variables involved in the calculation of a single flux measurement, the sources of uncertainty include the uncertainty in the derivation of the column densities, plume speed, plume height, plume direction, orientation or scanning angles, and radiative transfer. If the intention is to quantify the source emission strength from the measurements of the plume mass flow rate, the possible depletion/production of SO₂ downwind the vent, understood as the sum of all processes that reduce/increase the measured amount of SO₂, should be further considered. Measurement and parameter uncertainties can to a large extent be derived from the actual observations and the literature. Analysis presented in Arellano (2014) indicate that ScanDOAS measurements have asymmetric distribution of uncertainty, showing typically high left-skewness, i.e., the mean value of the distribution is most likely an underestimation of the true flux. In the results presented here, we compute statistics of daily flux only for measurements considered to have "good quality", based on several criteria, specified below. By adopting these criteria, we consider that a reasonable minimum estimate of fractional uncertainty lies between -30 to 10%, i.e., the reported values for individual flux measurements correspond to the average value, while the span of the uncertainty has 1σ

limits of confidence between 70 and 110% of the average value. The reduction in the number of valid results is usually large (40-60% of the total number of measurements, depending on the site). As the intention of this paper is to improve the statistics of measurements of SO$_2$ emission, we think that negotiating more quality for less quantity is justified. For details of the available raw data, see fig. S1 in the Supplements.

2.4 Criteria for data selection

Each measurement considered for the statistics and analyses presented in this paper has been validated according to a few quality criteria. As a prerequisite, valid spectra have position, time, total duration ($\leq$ 15 min), and observation geometry all within normal ranges. We tracked the history of installation of each station, determining the locations, orientations and scanning geometries of each spectrometer over the years, and checked them for accuracy. Then, each spectrum should have adequate intensity, neither saturated nor over-attenuated ($\leq$ 10% of saturation) in the region of evaluation (~310-325 nm). The DOAS-fit threshold for retrieval of SO$_2$ corresponds to a chi-square (χ^2) value of 9×10^{-3}. For a plume scan to be included in the analysis, we required it to have a 'plume completeness', calculated according to the algorithm described in Johansson (2009b), of at least 0.8. The absolute value of the scan angle with respect to zenith of the column density-weighted, centre of mass of the plume should not be larger than 75 deg and the calculated plume geometry should be reasonable (e.g., measurements which retrieved distance to the plume are larger than 10 km are not considered for further analysis). From the set of valid flux measurements in a day, time-averaged statistics are computed, if at least five measurements passed the quality checks. Figure 2 shows a flow chart of the steps followed in the evaluation of data.

(Fig. 2)

3 Results

3.1 SO$_2$ emission rate

In this study we report daily statistics of SO$_2$ emission. These are derived from minute-scale scan measurements, but we regard the daily emission as more representative of volcanic degassing, because the sub-daily values are subject to large variability introduced by meteorological effects, tidal influences and other reasons (e.g., Bredemeyer and Hansteen, 2014; Dinger et al., 2018). We report the daily average SO$_2$ emission rate, standard deviation, different quantiles, and number of measurements in each day; as well as similar statistics for plume location, velocity, and cloud cover. Figure 3a-h shows the time-series of daily SO$_2$ flux between 1 March 2005 and 31 January 2017 for 32 volcanoes in NOVAC which produced a reasonable amount of valid data. Figure 4 shows the mean emission and 25-75% quantiles calculated from measured fluxes for all volcanoes during the same period. Results in numerical format are presented in the Supplement S2. An important exception in this compilation is Bárðarbunga volcano in Iceland; its Holuhraun eruption in 2014-2015 was monitored in detail, but the analysis of its data

required special handling, not apt for the procedure described here due to extreme measurement conditions and enormous amount of gas (Pfeffer et al., 2018).

300 **(Fig. 3a-h, 4)**

(Table 1)

3.2 Long term emission budgets and comparison with satellite-based data

The analysis of long-term data from automatic instrumental networks of this type presents a challenge for the extrapolation of (often irregular) sets of measurements to produce an estimation of time-averaged emissions. This challenge has to do with
305 distinguishing periods of null observations, in the sense described above (i.e. when less than five measurements of good quality were obtained within a day), or that are caused by instrumental (e.g., when no measurements were acquired) or observational (e.g., winds drifting the plume beyond zone of observation) causes, from periods of legitimate low emission (i.e., absence of a plume). To account for these periods, we need additional information about the level of activity, visual observations, photographic records, etc., which is not always available.

310 To deal with this problem we have adopted the following strategy: if there are no statistics of emission for a given day, and there were either no scanning measurements conducted or the mean plume direction (obtained from the meteorological model) lies outside the 5-95% range of historical plume directions observed by the instruments, then no inference can be made about the actual emission on that day and the value is simply interpolated linearly between the nearest data points with valid observations. On the other hand, if measurements were done and the modelled wind data indicates that the plume should have
315 been observed by the instrumental network, we attribute the lack of data for that day to low volcanic emission. The actual value of this low emission is chosen as the 5% quantile of valid historical observations. This value is chosen arbitrarily to represent an effective detection limit, noticing that the flux depends not only on the actual gas column density detection limit but also on the size and speed of the plume.

By filling-in data in this way we can obtain a more regular and accurate representation of the actual emission for prolonged
320 periods of time and calculate the corresponding statistics. The results of this procedure are shown in fig. 5a-c, where we present only the time-series for volcanoes which have corresponding observations by OMI in the same period (12 out of 32), as reported by Carn et al. (2017). Notice also the time series of 'only observed' data along with the corresponding time series of emission from the OMI sensor. Results in numerical form are presented in the Supplement S1.

3.3 NOVAC emission data repository

325 The results presented here are made public through a data repository hosted on the website https://novac.chalmers.se/. The site shows a map with the location of the volcanoes for which valid data has been produced. The dataset produced according to the methodology described here is labelled Version 1, and updates (temporal increments) and upgrades (different versions of data produced with improved methodology) are planned in the future. The dataset shows a summary of available raw data (i.e., scans) collected by the instruments, along with a summary of valid fluxes derived from those measurements.

330   After selection of a volcano, a dedicated window presents a map with the setup of monitoring instruments, including coordinates and measurement parameters, a link to generic information about the volcano hosted on the Smithsonian Institution's Global Volcanism Program website (https://volcano.si.edu/), information on the responsible observatory and contact details, and the time series of daily mean SO_2 emissions with associated statistics. The plots are easy to explore through different scaling and textual information. From each volcano-page, data can be downloaded, after registering basic contact

335   information and accepting the Data Use Agreement, which states (e.g. for the case of Popocatépetl volcano) the following:

> *"Large efforts have been made by the volcano observatories and institutions*
> *responsible for data collection and evaluation. Thus, data presented here can*
> *be used on the condition that these organizations and people are given proper*
> *credit for their work, following normal practice in scientific communication:*

340
> *(1) If data from this repository contributes an important part of the work,*
> *co-authorship should be offered to the listed contributors and the data-set*
> *should be cited.*
> *(2) If data from this repository contributes only a small, but still important*
> *part of the work, the data-set should be cited.*

345
> *To cite this data-set include this information:*
> **
> *Delgado, H., Arellano, S., Rivera, C., Fickel, M., Álvarez, J., Galle, B.,*
> *SO2 flux of -POPOCATEPETL- volcano, from the NOVAC data-base;*
> *2020; [Data set]; v.001; doi:10.17196/novac.popocatepetl.001*

350
> **
> *Additional data, data with higher time resolution and raw data may be*
> *made available upon request to the respective contacts, listed below.*
> *This data-set has license: CC-BY 4.0"*

    Notice that each dataset is assigned a registered and permanent Digital Object Identifier (DOI).

355   The data files were prepared following the guidelines of the Generic Earth Observation Metadata Standard (GEOMS) (Retscher et al., 2011), which are generic metadata guidelines on atmospheric and oceanographic datasets adopted for global initiatives, such as the Network for Detection of Atmospheric Composition Change (NDACC). The GEOMS standard requires a file format such as netCDF. This data format can be explored using openly available tools such as Panoply (https://www.giss.nasa.gov/tools/panoply/). For users not familiarized with the netCDF format, a text format file, easily

360   accessible through standard workbook or text editor applications and containing the same information that the netCDF file, is also available. An example of such a file is presented in the Supplement S3 and a list of all files is listed in the Supplement S1. The GEOMS standard requires data and metadata to be included in the same file. In the case of NOVAC, the metadata, required for the description and interpretation of the data, includes the following:

- General information about the dataset
365 - Data Use Agreement
- Data Set Description (site, measurement quantities, processing period, processing level, data version, DOI, accompanying file, date of file production)
- Contact Information
- Reference articles
370 - Instrument(s) Description (instrument type, spectrometer specification, fore-optics specifications, control unit specifications, instrument(s) ID(s), site(s) name(s), site(s) coordinates, site(s) measurement parameters, instrument(s) serial numbers)
- Measurement description
- Algorithm description (slant column densities, vertical column densities, SO_2 flux and plume parameters, statistics)
375 - Expected uncertainty of measurement
- Description of appended results

The data include the following:

- Date UT
- Daily mean, standard deviation, quartiles, and number of valid SO_2 flux measurements
380 - Daily mean and standard deviation of plume speed
- Daily mean and standard deviation of plume direction
- Daily mean and standard deviation of plume height
- Daily mean and standard deviation of plume distance to instruments and width
- Daily mean and standard deviation of cloud cover (from re-analysis meteorological model)

385 Additional pages in the data repository provide details about the database, the Data Use Agreement, technical details of the instrument, description of algorithms used for available data versions, contact information and acknowledgements.

The NOVAC data repository will be linked to other thematic databases such as the Database of Volcanic Unrest of the World Organization of Volcano Observatories (WOVOdat), the database of the Global Volcanism Programme of the Smithsonian Institution, the EarthChem data repository, the Global Emission InitiAtive (GEIA), the database of the Emissions of 390 atmospheric Compounds and Compilation of Ancillary Data (ECCAD), and the database of the EU Copernicus Atmospheric Monitoring Service (CAMS).

4 Discussion

4.1 Comparison of emission from different volcanoes

Figures 3a-h and 4 summarize the statistical information of the time series of emission for 32 volcanoes in NOVAC during 395 2005-2017. The plots show the daily and annual means and 25-75% quantiles of daily SO_2 emission, to represent variability. We highlight three main characteristics from these results: (i) the relatively large range of variation of emissions, spanning typically up to three orders of magnitude in variability, for the same volcano at different times (fig. 3a-h); (ii) the skewed nature of the distributions, with a dominance of low emission values (i.e., more frequent low emission rate values and a few

large emission values that account for a considerable fraction of the total emission); and (iii) the large difference between the characteristic emission of different volcanoes (fig. 4).

With respect to the intra-variability (for a particular volcano), we consider this to be one of the most important findings of long-term monitoring. As mentioned in the introduction, the production of high-sampling rate, long-term measurements is relatively recent. Most compilations of measurements in the past include campaign-based estimates of gas emissions, typically during periods of enhanced activity, when a plume was visible and during short periods of time. The skewed, large range distributions of emission seem to be a general feature of degassing volcanoes and merit more attention. Our analysis, using only measured, i.e. not 'in-filled' data, indicates that the ratio of the first quartile to the mean of daily SO_2 fluxes reaches 43 ± 14% (±1σ), which means that the distribution of daily emission is dominated by low values,. An important implication of this finding is that the low-emission spectrum of the distribution, which has usually not been measured in the past, contributes a significant amount of the total emission and should therefore be better characterized. Another is that short-term measurements may be skewed and could therefore not be representative of the long-term emission of a volcano.

Regarding the inter-variability (among different volcanoes), the observation of a large variance between sources is not new. Indeed, it has been speculated and partially shown by several authors (e.g., Brantley and Koepenick, 1995; Andres and Kasgnoc, 1998; Mori et al., 2013; Carn et al., 2017) that the partition between sources of volcanic degassing, particularly quiescent degassing, seem to follow either a log-normal or a power-law distribution. These distributions may seem similar but choosing one over the other results in significant differences in estimating the global volcanic flux. The relative importance of low vs. high emitters is also different for log-normal or power-law distributions. Evidently, with 32 volcanoes, out of perhaps 90-150 degassing volcanoes, it is not possible to verify these speculations with certainty. In any case, our measurements provide bounds for the contribution of weak emission sources, which have escaped observation by satellites during the same period.

4.2 Ground-based vs. space-based observations

The recent compilation of global volcanic degassing from satellite-based measurements of OMI (Carn et al., 2017) offers an excellent opportunity for comparison with the measurements obtained from the ground with NOVAC instruments. First, both methods have operated for about the same period (since 2005); second, both sets of measurements are analysed independently in a consistent manner; and, third, the two datasets are focused on passive degassing. The quantification of SO_2 flux from OMI observationeswas achieved by stacking of wind-rotated images over the course of a year, discarding both pixels contaminated by clouds and pixels with elevated column densities resulting from explosive eruptions. The co-added annual image is then fitted to a Gaussian distribution, following Fioletov et al. (2016), and the goodness of this fit is expressed as an "uncertainty", but the actual uncertainty in the reported emission is not quantified but assumed in the order of 50% (Fioletov et al., 2016; Carn et al., 2017).

430 As mentioned above, it is necessary to fill-in the measured SO_2 emissions at each volcano during times when degassing was not detected but the instruments would have picked it up had it been occurring. The original, "un-filled" time series and the time-series "filled" with low-emission values are presented in fig. 5a-c, along with the corresponding time-series for OMI.

(Fig. 5a-c)

This comparison shows a general agreement in the temporal trends of annual emission for ground- and space-based methods,
435 but with differences in magnitude, which in some cases are considerable. Only 12 out of 32 volcanoes from NOVAC have corresponding detections from OMI. This is not surprising, as all volcanoes not observed by OMI are weak sources of emission, confined to the lower atmosphere, and in some cases located in areas of persistent cloud cover. Consequently, our dataset provides new data for several volcanoes, such as Sangay, Cotopaxi, Planchón Peteroa and Sinabung, and the largest dataset for all other volcanoes ever published.

440 Figure 5a-c, as well as the Supplement S1, also show that the difference between ground- and space-based observations is reduced by the method of filling-in low-emission values in the patchy time-series in NOVAC. Notwithstanding this better convergence, the differences are, in general, biased towards higher emission observed from satellites. There are many possible reasons for this, for example, that the selection of data for OMI may tend to pick images from higher plumes that reach altitudes above low-level clouds, which may be the result of more explosive activity. Another reason could be that the data selection
445 for NOVAC favours plumes with clear boundaries, and in some instances, high gas-content plumes may be too wide and thus completely overcast the instruments, and, as a consequence, are filtered out by the strict quality control filters applied to the dataset during our analysis.

However, the larger differences are caused by obvious reasons: for example, in the case of two nearby volcanoes, such as Nyiragongo-Nyiamulagira (with a footprint of 13×24 km^2), OMI cannot separate completely the contributions of each source,
450 so they are reported as a complex. In this respect, NOVAC can aid to discriminate between these sources, since the stations are deployed with a focus on Nyiragongo and the finer time resolution allows disentangling contributions, especially during periods of heightened activity at any of them. Other reasons for discrepancy are to be found in the different periods covered by the instruments, i.e., only daytime measurements for NOVAC, whereas OMI could in principle detect the emission occurring while overpassing at 13:30 LT in addition to remaining gas that was emitted emissions during the previous hours,
455 potentially even during night. Other factors are the relatively large measurement uncertainties of both methods and different radiative transfer effects depending on altitude of surrounding plumes. A more in-depth study of these discrepancies is highly needed.

Finally, the method proposed here to account for days with null observations improves considerably the comparison with OMI in general. This is more obvious for volcanoes with constant emissions and good instrumental coverage, resulting in more
460 valid measurements (represented by the size of the circles in fig. 5a-c), which give us confidence in the validity of this approach.

4.3 The NOVAC inventory and past compilations of emission

It is interesting to compare the emission statistics obtained from the NOVAC data with past compilations of emissions presented in other studies. We refer in particular to Andres and Kasgnoc (1998), who report the volcanic input during 1970-1997 to the Global Emission Inventory Activity (GEIA) database.

465 The results of a one to one comparison between the emissions reported for quiescent degassing volcanoes in GEIA and NOVAC is presented in fig. 6. There are a few volcanoes reported in GEIA which are not part of NOVAC yet; conversely, some volcanoes monitored in NOVAC that were not active and thus not considered during the period reported in GEIA. A comparison can only be done for the 16 volcanoes present in both datasets. Undoubtedly, the reported values are not expected to coincide, considering that the measurements were not obtained during the same periods and, as revealed by the NOVAC

470 results, volcanic gas emission is by no means a stationary process over time. However, it is important to highlight that the recent measurements from NOVAC provide a characteristic range of variation for the volcanic sources, which in most cases, but not all, accommodate the results of past, punctuated observations. But we notice also that, except for Momotombo, San Cristóbal and Telica, the mean emissions reported in GEIA lie in the upper end or higher than those reported here. We speculate that such systematic difference may be due to biased sampling during periods of high emission that was the basis for most of

475 the Andres and Kasgnoc (1998) compilation. Long-term observation also captures periods of quiescence that may be the reason for lower values.

(Fig. 6)

5 Conclusions and outlook

In this study, we report the results of post-processing of SO_2 mass emission rate measurements at 32 volcanoes of the NOVAC

480 network during 2005-2017. This is, to our knowledge, the densest (~10-50 measurements per volcano per day during up to 12 years, 32 volcanoes) database of volcanic degassing obtained by a standardized method. Since the ScanDOAS method is subject to multiple and potentially large sources of uncertainty, considerable attention has been given to the selection of high-quality measurements on which to base the reported statistics.

Independent studies (e.g., Stoiber and Jepsen, 1983; Krueger et al., 1995; Halmer et al., 2002; Andres and Kasgnoc, 1998;

485 Carn et al., 2017) have demonstrated over the years that passive degassing dominates, in time and magnitude, the time-averaged global volcanic emission. At the same time, this component of volcanic emission can produce a persistent impact on local, regional or global scales, and potentially affect the climate system. Observational limitations have hindered quantifying the magnitude and variability of global volcanic degassing in the past at the level of detail obtained by a global ground-based network like NOVAC. This database will therefore represent an important contribution to global emission inventories, which

490 are typically based on sporadic and short-term investigations of emission during periods of heightened activity. Moreover, the results from measurements in NOVAC complement the observations from satellite platforms, which on an operational basis during the past decade were more suited for quantification of explosive degassing.

The measurements performed in NOVAC provide more information than the gas emission rate of SO_2. First, spectroscopic analysis of the data can be used for retrieving the abundances of other species, as it has been proven most systematically for the case of BrO (Lübcke et al., 2014; Dinger et al., 2018, Warnach et al., 2019). Second, in principle all the variables involved in the calculation of the mass flow rate can be obtained from the measurements (plume location, dimensions and even velocity), which are valuable for modelling studies of volcanic activity and risk, environmental impact, or atmospheric transport.

An important finding of long-term measurements, as in the case presented here, is the empirical distribution of degassing for individual volcanoes. It has often been assumed that the emission of individual volcanoes exhibits a typical value and a symmetric distribution, and global estimates are computed by assuming a skewed distribution for the global volcanic emission. We have found that individual volcanoes, notwithstanding their completely different volcanological characteristics or states of activity, conform to a distribution that looks nearly symmetrical in logarithmic units. This is to say, the logarithm of the emission rate tends to converge to a central value, and the bulk emission is composed of a vast majority of low-emission events and a few, but significant, high-emission events. It is therefore adequate to typify volcanoes' gas emission in these units, akin to the Volcanic Sulphur Index (VSI) proposed by Schnetzler et al. (1997), which is based in the total emitted sulphur for eruptions rather than the emission rate in general.

We would also like to point out that our standardized processing of data from all volcanoes in the network may not necessarily be the best strategy with regards to obtaining more continuous time-series of gas emission. The procedure presented here applies very strict criteria to validate individual SO_2 flux measurements. These criteria make sure that all assumptions behind the method of scanning the plume to derive the flux are fulfilled. However, for volcanoes like Piton de la Fournaise or Vulcano, the result of applying these criteria is a drastic reduction in the number of valid measurements. The reason is that most recent eruptions of Piton de la Fournaise have taken place in flank vents located inside the Enclos Fouqué caldera at altitudes below the location of the NOVAC stations. If the plume does not reach enough altitude above the stations, the coverage will not be complete. Still, measurements will be able to pinpoint the location of the plume, and under some circumstances also deliver quantitative information of the flux (e.g. adopting a different method of integration of gas column densities). Gas emission is minimal during inter-eruptive periods (Ref). For Vulcano, the dynamics of degassing is different and characterized by weak passive emission. Even if the emission is detectable by the instruments, winds should be strong and stable enough to produce a plume that fulfills the requirements for scanning. Certain criteria, such as the number of spectra with valid gas detection in the same scan, must be relaxed to obtain a gas flux datum, with a detection of SO_2 down to <10 t/d (Vita et al., 2012; Granieri et al., 2017). On the other extreme we find episodes when the plume is too wide to completely overcast the scanning path of the instruments. This could happen for volcanoes with strong emission and relatively low altitude difference between the plume and the station. For example, for Nevado del Ruiz, some stations are placed high in the flanks and due to strong activity of the volcano, the stations will not produce data validated by our method. Using data from more distant stations and novel ways to derive the background SO_2 can help to solve this issue (Lübcke et al., 2016). Another example is the Holuhraun eruption, where very extreme conditions in terms of gas column density and low altitude required a different approach to flux estimation (Pfeffer et al., 2018).

There is still a need to improve the characterization of plume transport and radiative transfer effects in remote sensing of volcanic plumes.. One of the main advantages of NOVAC is that each improvement in the software or hardware can be easily implemented in the several sites of the network and historical data can be re-analysed retrospectively using more advanced algorithms.

Data presented here will be of interest for different applications, but certain considerations are important to mention. First, the reported statistics are considered representative of gas emissions for volcanoes in a state of passive degassing or moderate explosive activity. Emissions resulting from large explosive events may not be properly captured by the near-field, ground-based methods used in NOVAC, since such emissions could either reach several km in the atmosphere, beyond the effective range of the instruments, or they can be too optically thick (due to extreme gas concentrations or aerosols) and thus render the measurements inaccurate. Second, our method to derive statistics will be more representative of periods of continuous or frequent degassing. Null plume detections could be the result of either low or null volcanic emission or due to lack of observations, caused, for example, by meteorological conditions or not-covered plume directions. Therefore, a simple average of the daily means reported here over a longer period may overestimate the emission, unless null emission is distinguished from null observation, as we propose. On the other hand, the analysis of uncertainty indicates that the effective detections are most likely showing minimum emission values, because of the mostly reducing effect of radiation scattering in the retrieved column density values. We think that complementary information to SO_2 flux reported here, namely the statistics of plume location, velocity, dimensions, and general weather conditions will be valuable not only to interpret the emission patterns, but also to assess their impact downwind the volcano.

NOVAC has expanded since its inception in 2005 and to date includes about a third of all volcanoes which have exhibited degassing detectable from space in the last decade. The network has grown from the initiative of volcanological observatories themselves, but also from the important support of the USGS/USAID VDAP initiative, especially in Latin America, South East Asia, and Oceania. Recent initiatives, such as the Deep Carbon Observatory, have built upon the measurements in NOVAC to quantify the global volcanic emission of CO_2, by combining continuous emission rate measurements with punctuated molar ratio measurements (Fischer et al., 2019). Improving the estimates of SO_2 emission will, in this way, result in a better estimate of the emission budgets of other volcanic species. Since volcanic aerosols, to a large extent seeded by primary emission of SO_2, are one of the most important but poorly quantified sources of natural radiative forcing in the climate system (Myhre et al., 2013), there is a need to better quantify their magnitude and location. A global network for the observation of volcanic plumes is of great importance to quantify the magnitude and temporal and spatial variability of volcanic emissions.

Data availability

More information about NOVAC can be found in the website https://novac-community.org/. Raw data from NOVAC are accessible by request to the local observatory responsible for the measurements. The dataset obtained for this study can be accessed, free of charge, through a dedicated website https://novac.chalmers.se/. The datasets of individual volcanoes can be

accessed through the DOI-links provided in Table 1. Updates of the time series, addition of new volcanoes and release of data versions resulting from improved analysis is planned.

Code availability

The NOVAC Post Processing Program and other software used in NOVAC are open source projects available in GitHub (https://github.com/NOVACProject/).

Author contribution

SA conceived the study, analysed data and wrote the manuscript with inputs from all authors; BG initiated and coordinated NOVAC and supervised research and database implementation; BG, CK, UP and TH supervised research for this work at Chalmers U., USGS/VDAP, U. Heidelberg and GEOMAR, respectively; MJ wrote the spectral analysis software; MK designed the electronics of the instrument; JM supervised development of the database; FA, GA, CBa, NB, CBu, VB, MBu, ZC, GC, CJC, VC, FC, MdM, HDG, AdM, DF, GG, HG, NH, SH, SI, PK, PM, FM, CR, AS, GS, BT, FV, GV, FV and MY supervised research, installed, maintained and operated instruments, and analysed real-time data at their respective observatories.

Competing interests

The authors declare that they have no conflict of interest.

Disclaimer

Any used of trade, firm, or product names is for descriptive purposes only and does not imply endorsement by the authors or by the governments of their countries, including the United States.

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Tables

Table 1. Statistics of measured SO2 flux during 2005-2016 for 32 volcanoes in NOVAC. The reported values are the arithmetic mean (average) of the different daily statistics (mean, standard deviation, quartiles) included in the NOVAC data files.

Volcano	Statistics of measured SO2 flux during 2005-2016 [kg/s]					Number of days with valid flux statistics respect to days with measurements	Reference	DOI link
	Mean	Std. dev.	First quartile	Median	Third quartile			
Arenal	1.59	0.80	1.04	1.43	1.98	38/350	Avard et al., 2020	https://doi.org/10.17196/novac.arenal.001
Concepcion	6.10	2.01	4.66	5.92	7.30	186/795	Saballos et al., 2020	https://doi.org/10.17196/novac.concepcion.001
Copahue	9.02	3.61	6.51	8.37	11.03	288/476	Velasquez et al., 2020	https://doi.org/10.17196/novac.copahue.001
Cotopaxi	8.94	5.31	5.24	8.38	11.76	447/2829	Hidalgo et al., 2020	https://doi.org/10.17196/novac.cotopaxi.001
Etna	40.72	12.18	31.98	38.88	47.94	192/875	Salerno et al., 2020	https://doi.org/10.17196/novac.etna.001
Fuego de Colima	2.72	1.66	1.54	2.41	3.60	185/2750	Delgado et al., 2020	https://doi.org/10.17196/novac.fuegodecolima.001
Fuego	3.53	1.30	2.61	3.41	4.24	30/407	Chigna et al., 2020	https://doi.org/10.17196/novac.fuegoguatemala.001
Galeras	8.99	3.80	6.35	8.46	11.11	704/3340	Chacon et al., 2020	https://doi.org/10.17196/novac.galeras.001
Isluga	8.11	3.65	5.46	7.35	10.18	230/497	Bucarey et al., 2020	https://doi.org/10.17196/novac.isluga.001
Lascar	2.62	1.43	1.61	2.32	3.37	75/919	Bucarey et al., 2020	https://doi.org/10.17196/novac.nyiragongo.001

Llaima	10.08	3.42	7.66	10.08	12.49	4/1308	Bucarey et al., 2020	https://doi.org/10.17196/novac.llaima.001
Masaya	3.87	1.38	2.91	3.74	4.70	500/772	Saballos et al., 2020	https://doi.org/10.17196/novac.masaya.001
Mayon	7.01	2.50	5.30	6.59	8.29	102/1173	Bornas et al., 2020	https://doi.org/10.17196/novac.mayon.001
Momotombo	1.85	0.60	1.45	1.75	2.18	35/158	Saballos et al., 2020	https://doi.org/10.17196/novac.momotombo.001
Nevado del Ruiz	7.76	4.96	4.23	6.91	10.43	1281/2431	Chacon et al., 2020	https://doi.org/10.17196/novac.nevadodelruiz.001
Nyiragongo	19.14	8.89	12.71	17.71	24.52	432/1758	Yalire et al., 2020	https://doi.org/10.17196/novac.nyiragongo.001
Piton de la Fournaise	9.58	6.77	4.65	8.62	13.20	22/3402	Di Muro et al., 2020	https://doi.org/10.17196/novac.pitondelafournaise.001
Planchón Peteroa	0.05	0.02	0.03	0.04	0.06	4/22	Velasquez et al., 2020	https://doi.org/10.17196/novac.planchonpeteroa.001
Popocatepetl	24.48	12.28	15.81	23.27	31.96	1207/3306	Delgado et al., 2020	https://doi.org/10.17196/novac.popocatepetl.001
Sabancaya	11.86	5.85	7.56	10.76	15.10	126/162	Masias et al., 2020	https://doi.org/10.17196/novac.sabancaya.001
San Cristobal	8.87	3.41	6.44	8.34	10.86	1028/1557	Saballos et al., 2020	https://doi.org/10.17196/novac.sancristobal.001
San Miguel	22.51	6.17	19.42	22.25	25.25	4/160	Montalvo et al., 2020	https://doi.org/10.17196/novac.sanmiguel.001
Sangay	7.49	3.34	4.87	7.15	9.73	7/536	Hidalgo et al., 2020	https://doi.org/10.17196/novac.sangay.001
Santa Ana	1.96	0.56	1.55	1.86	2.36	22/868	Montalvo et al., 2020	https://doi.org/10.17196/novac.santaana.001
Santiaguito	3.22	1.80	1.96	2.83	4.08	170/570	Chigna et al., 2020	https://doi.org/10.17196/novac.santiaguito.001
Sinabung	4.42	2.07	3.05	3.95	5.14	108/173	Kasbani et al., 2020	https://doi.org/10.17196/novac.sinabung.001
Telica	0.83	0.31	0.60	0.78	1.00	205/460	Saballos et al., 2020	https://doi.org/10.17196/novac.telica.001
Tungurahua	17.12	7.93	11.54	16.17	21.81	1100/3463	Hidalgo et al., 2020	https://doi.org/10.17196/novac.tungurahua.001
Turrialba	11.50	4.98	7.98	10.96	14.48	614/1878	Avard et al., 2020	https://doi.org/10.17196/novac.turrialba.001

Ubinas	3.52	2.00	2.09	3.20	4.62	375/622	Masias et al., 2020	https://doi.org/10.17196/novac.ubinas.001
Villarrica	6.75	2.58	4.90	6.28	8.23	203/2001	Velasquez et al., 2020	https://doi.org/10.17196/novac.villarrica.001
Vulcano	0.20	0.06	0.15	0.19	0.24	34/1180	Vita et al., 2020	https://doi.org/10.17196/novac.vulcano.001
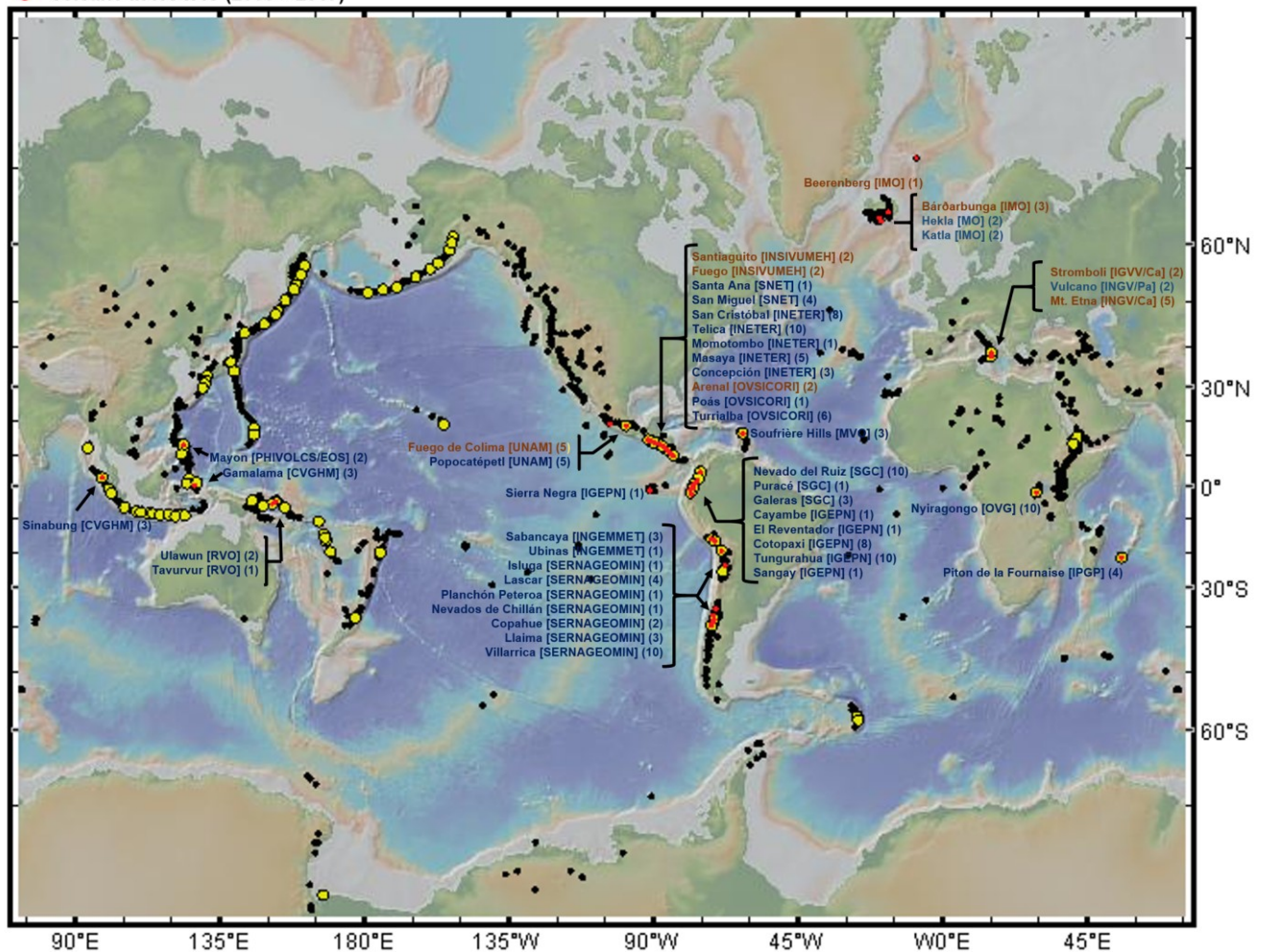
Figure 1. Topographic map showing the locations of volcanoes in NOVAC (red circles). The locations of Holocene volcanoes from the Global Volcanism Program of the Smithsonian Institution (2013) are shown with black circles. The locations of volcanoes detected by OMI during 2005-2015 (from Carn et al., 2017) are shown with yellow circles. Beside the names of the volcanoes are the

600
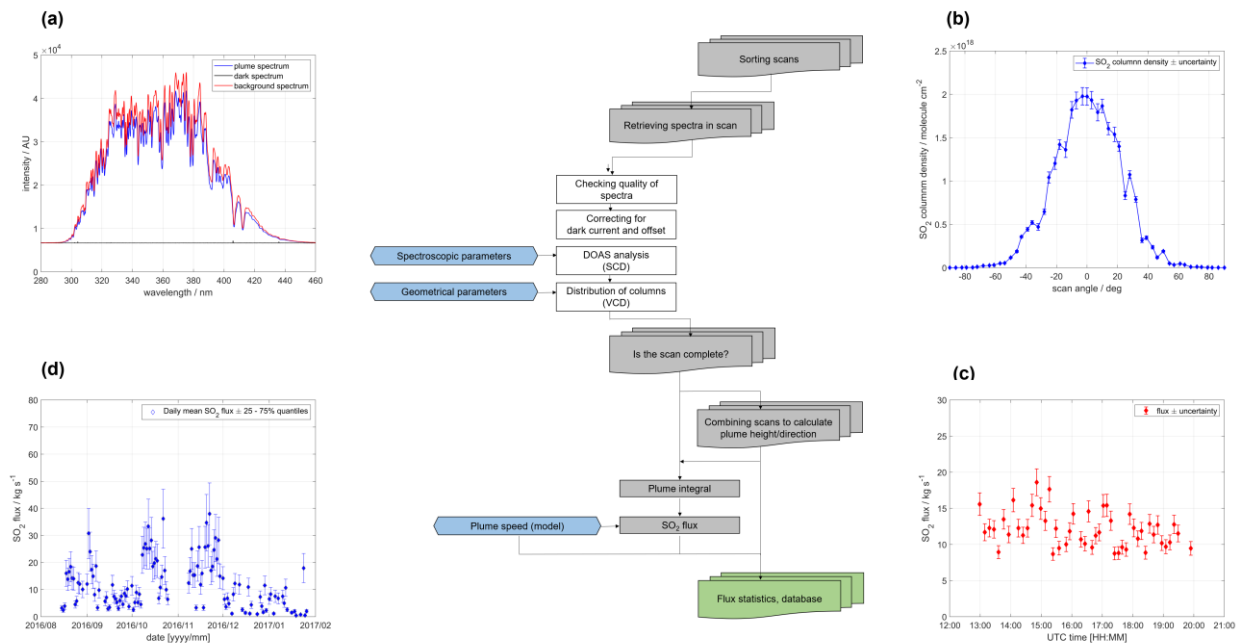
Figure 2. Schematics of the algorithm used to derive time-average emission of volcanoes in the NOVAC database. (a) Scattered
605 sunlight spectra (shown in figure as uncalibrated spectral radiance in arbitrary units -AU-) are checked for quality and combined to correct instrumental effects and to derive differential slant column densities (SCD) of SO_2 through the non-linear DOAS method. (b) A collection of column densities in the scan is used to determine the baseline column density, the angular position of the centre of mass of the plume, to convert slant to vertical column densities (VCD), and to estimate the completeness of the scanned plume. (c) Pairs of scans taken close in time by different instruments are used to derive plume altitude and direction. This is combined with
610 plume speed using a meteorological model to derive the flux, including uncertainty. Individual flux measurements are chosen considering uncertainty, completeness, and other criteria. (d) If at least five valid measurements exist on a given day, statistics of daily emission are computed and reported in the NOVAC database. The background colour of the boxes indicate processing at the spectral level (white), scan level (gray), flux level (green), and external parameters (blue).

615

(a)

(b)

(c)

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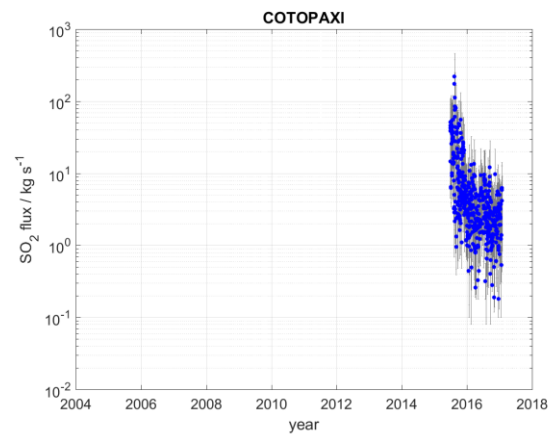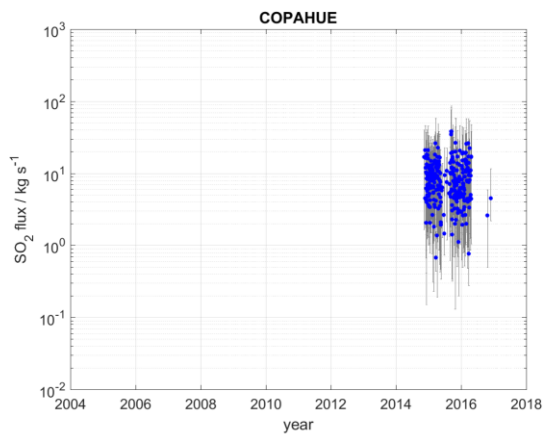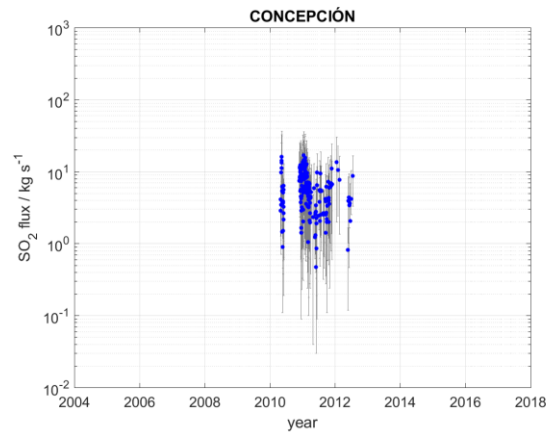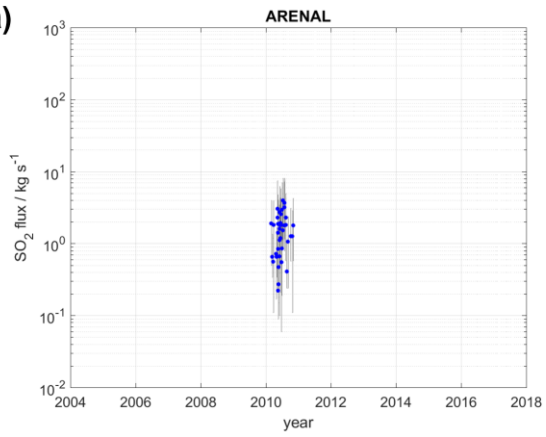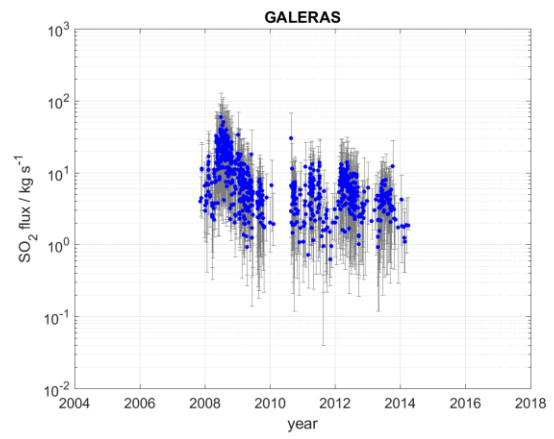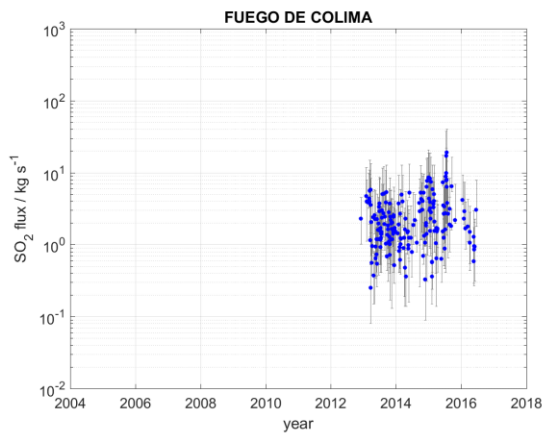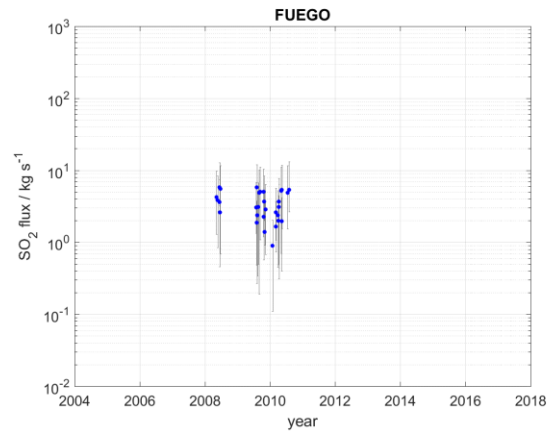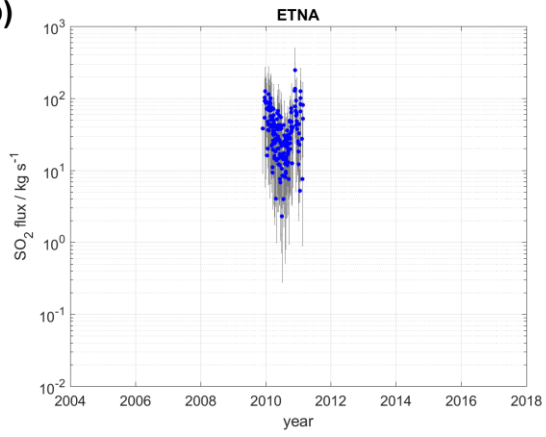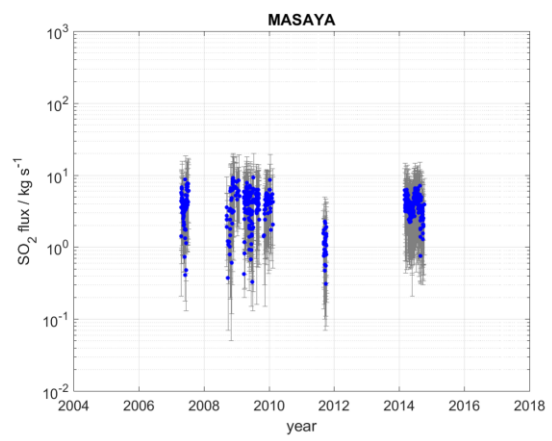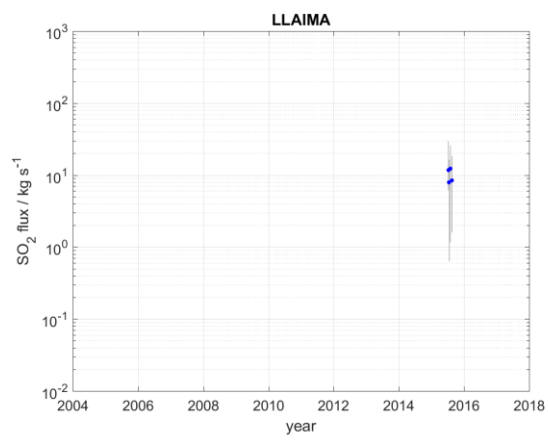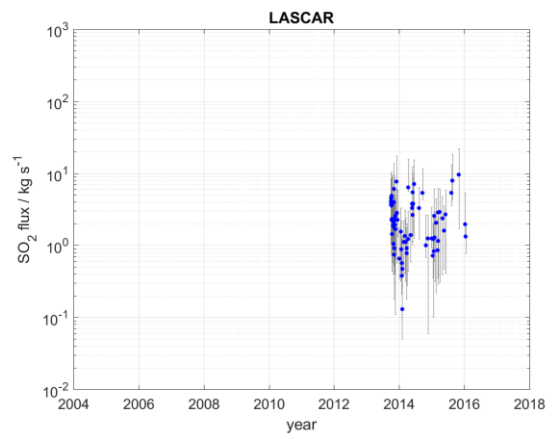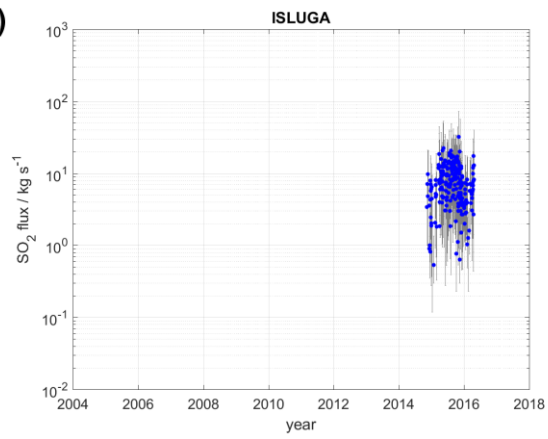
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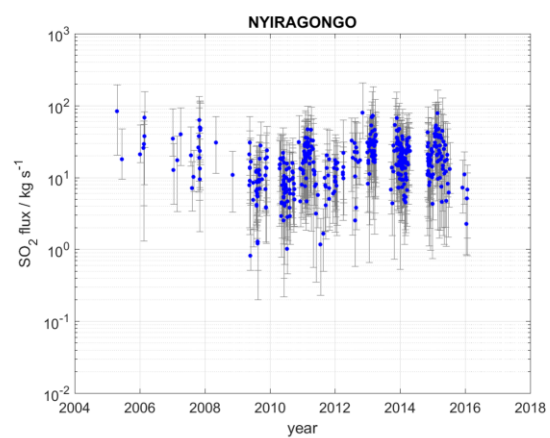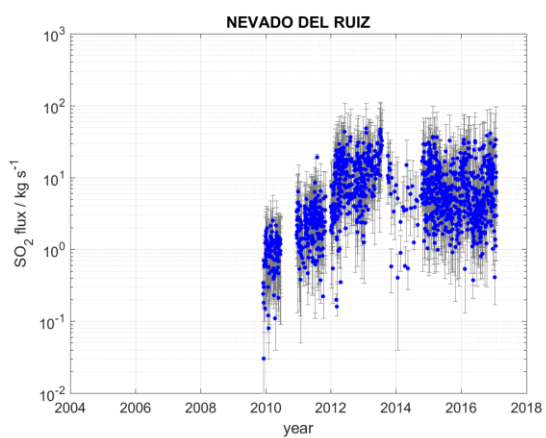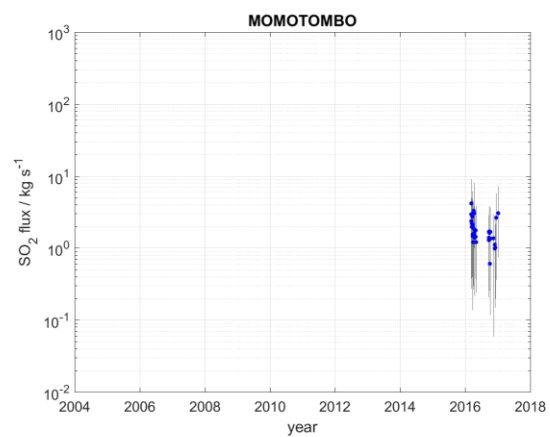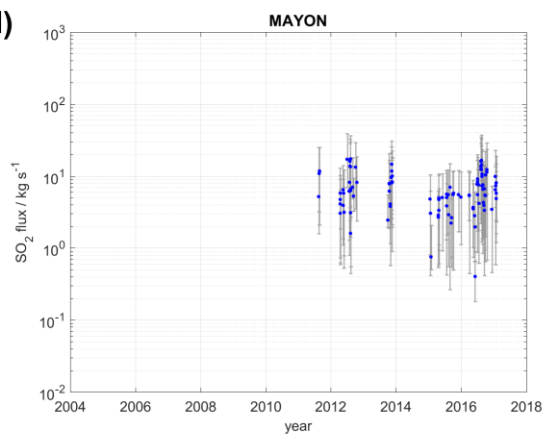
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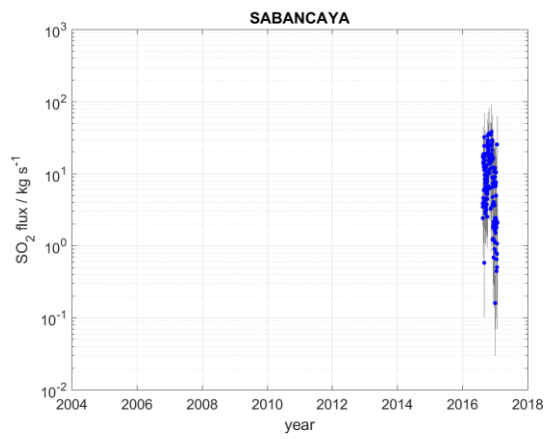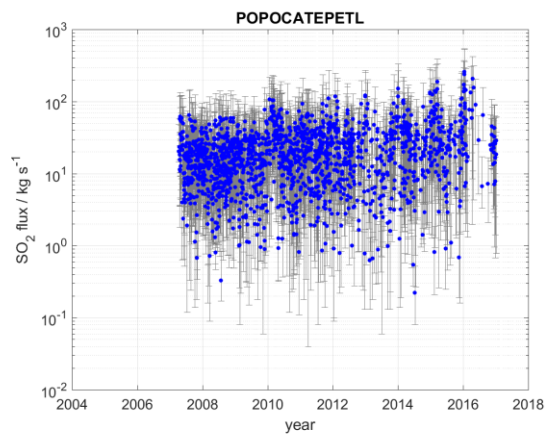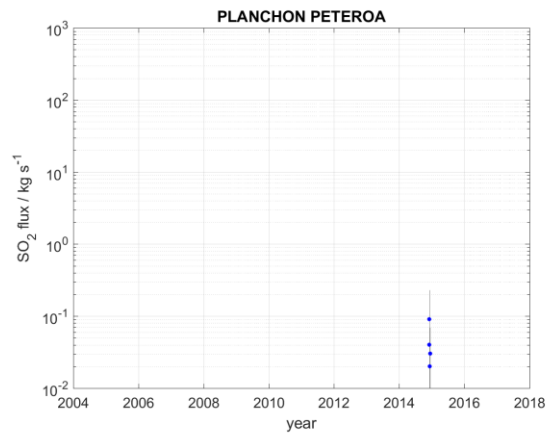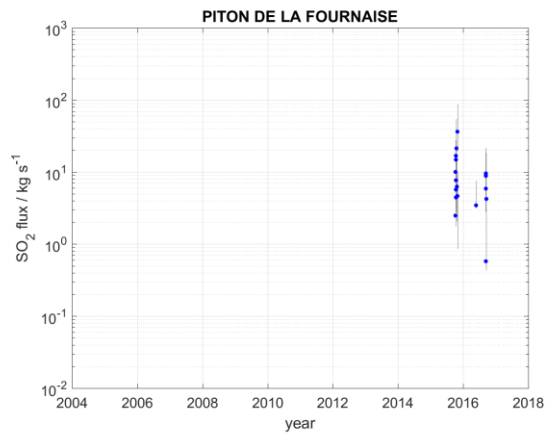
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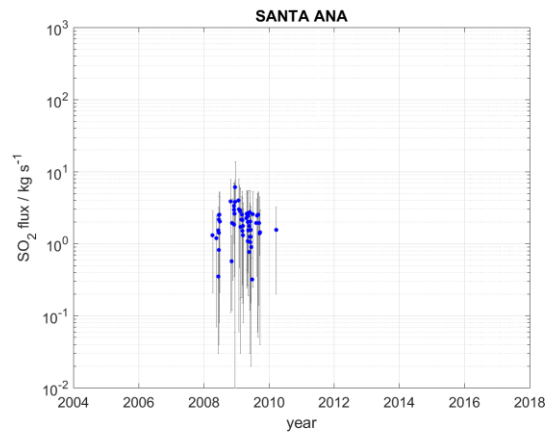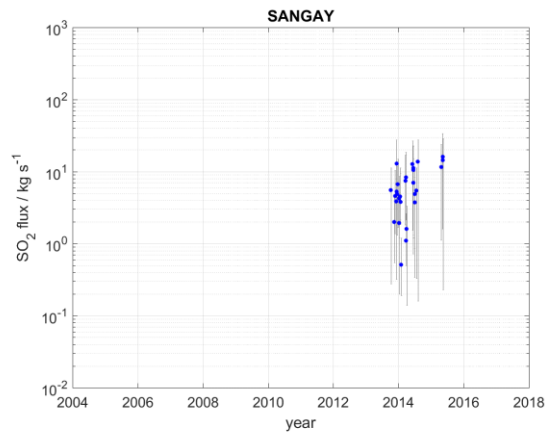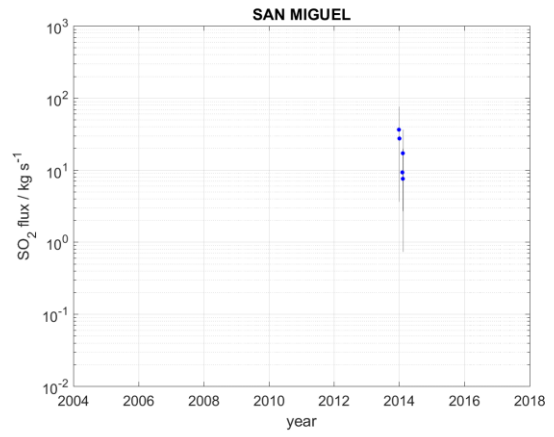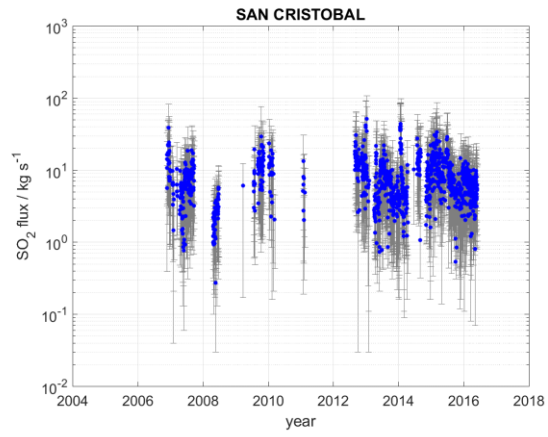
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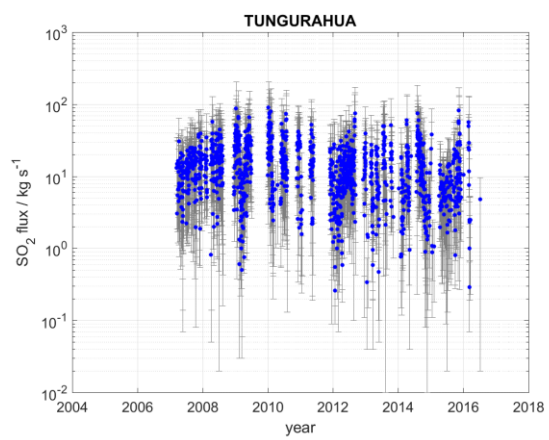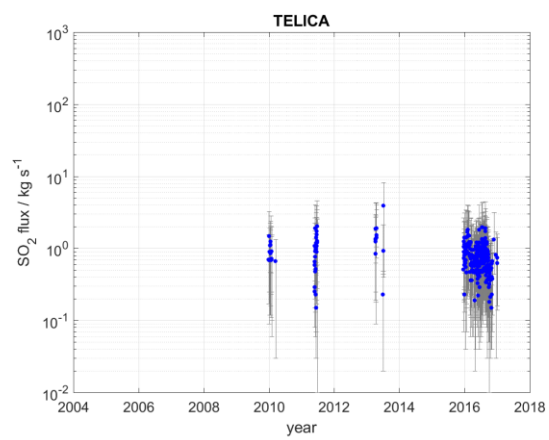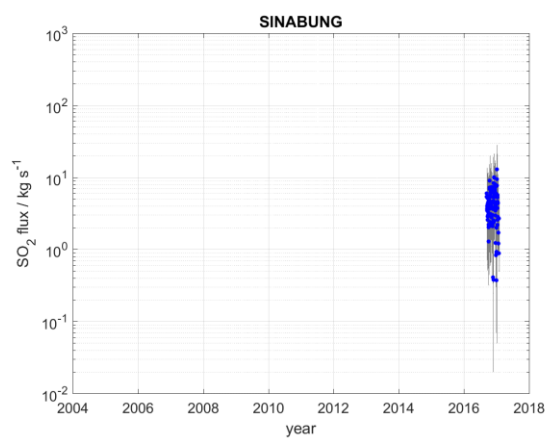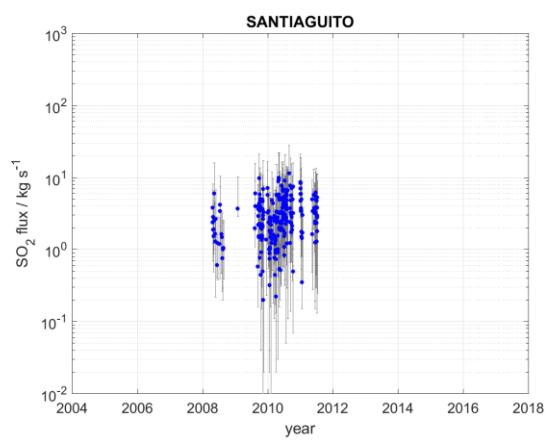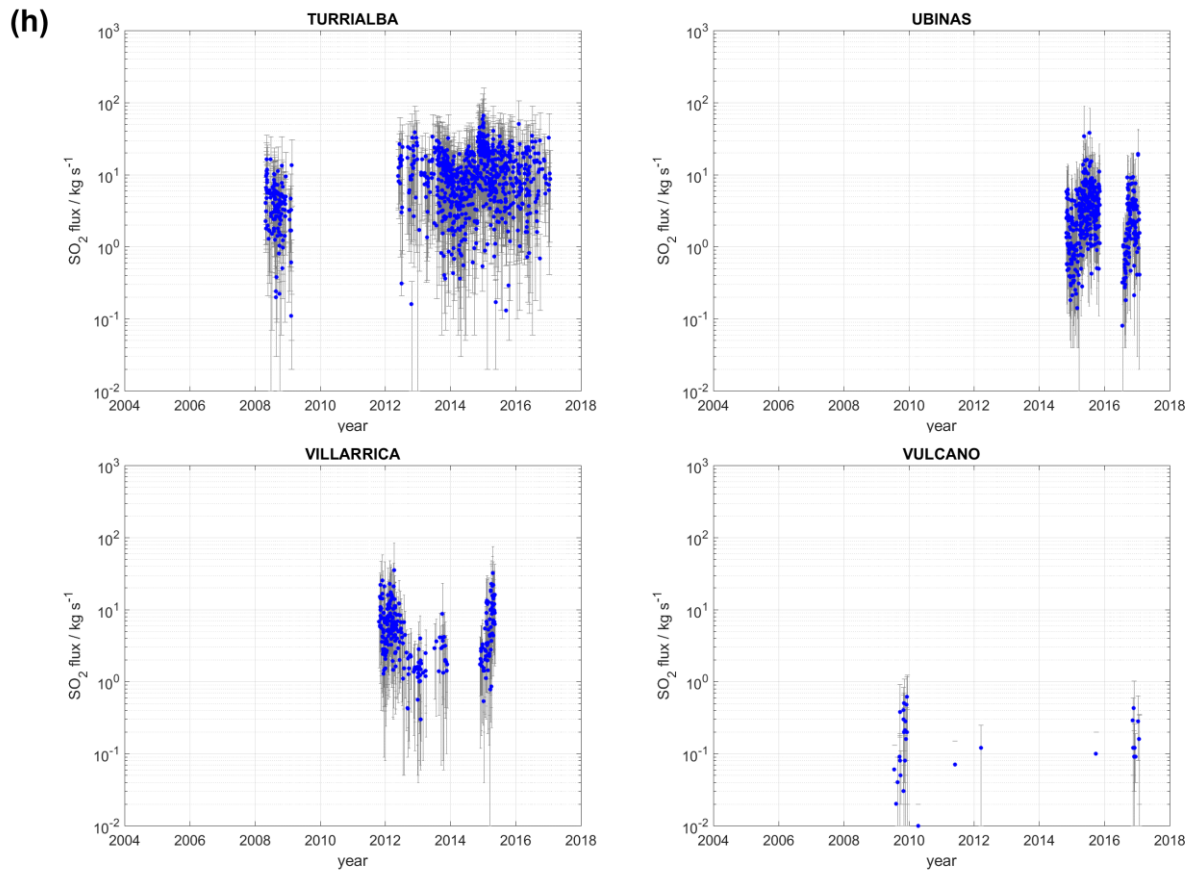
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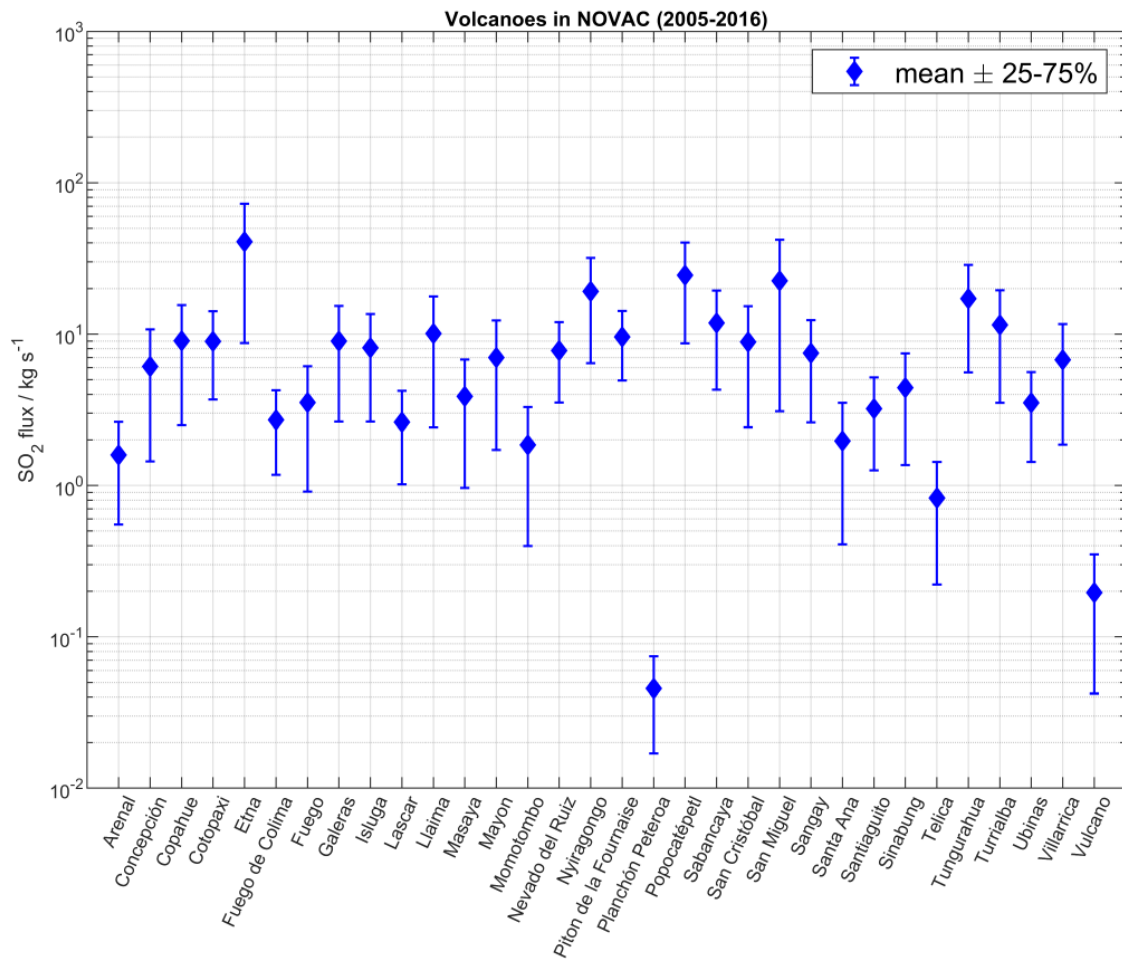
Figure 3. Time series of daily SO₂ emission from 32 volcanoes in NOVAC from 1 March 2005 to 31 January 2017. Blue dots depict the daily mean emission rate and grey lines indicate the 25 and 75% quantiles of daily flux. References for these data-sets are given in Table 1. (a) Data from Arenal, Concepción, Copahue, and Cotopaxi. (b) Data from Etna, Fuego, Fuego de Colima, and Galeras. (c) Data from Isluga, Lascar, Llaima, and Masaya. (d) Data from Mayon, Momotombo, Nevado del Ruiz, and Nyiragongo. (e) Data
630 **from Piton de la Fournaise, Planchón Peteroa, Popocatépetl, and Sabancaya. (f) Data from San Cristóbal, San Miguel, Sangay, and Santa Ana. (g) Data from Santiaguito, Sinabung, Telica, and Tungurahua. (h) Data from Turrialba, Ubinas, Villarrica, and Vulcano. All data is accessible through the links in Table 1.**

635   **Figure 4. Statistics of daily SO₂ emission from 32 volcanoes in NOVAC from 1 March 2005 to 31 January 2017, for periods of time when data were being collected and yielding flux values above the detection limit Blue markers show the average of all measured fluxes for each volcano during this period and the error bars show the corresponding 25 and 75% quantiles.**

(a)

COPAHUE

ETNA

GALERAS

ISLUGA

(b)

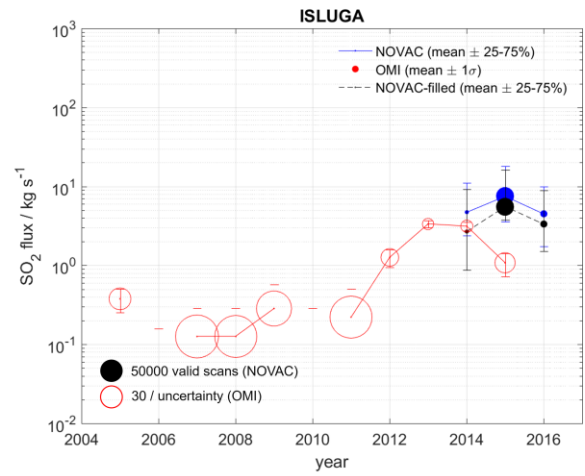
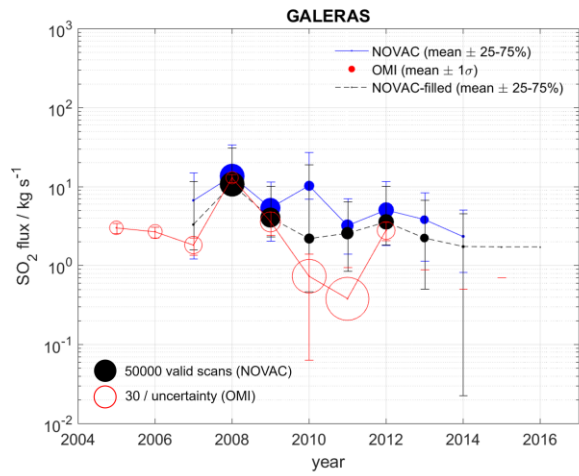
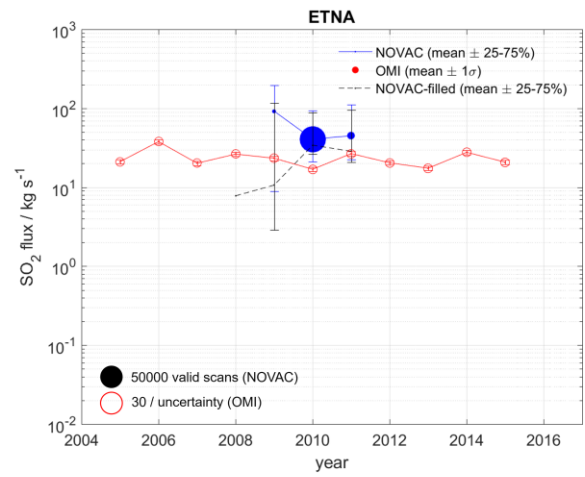
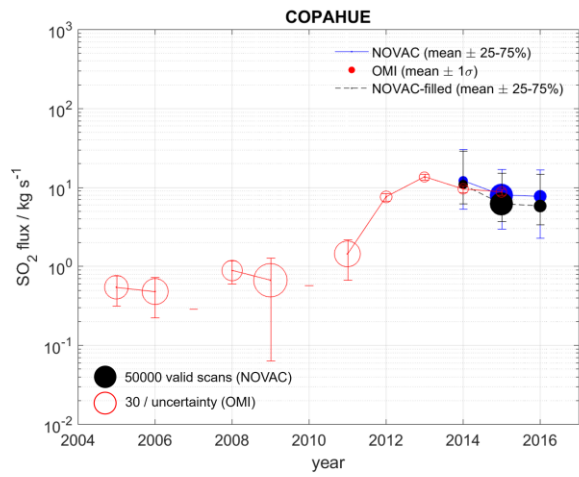
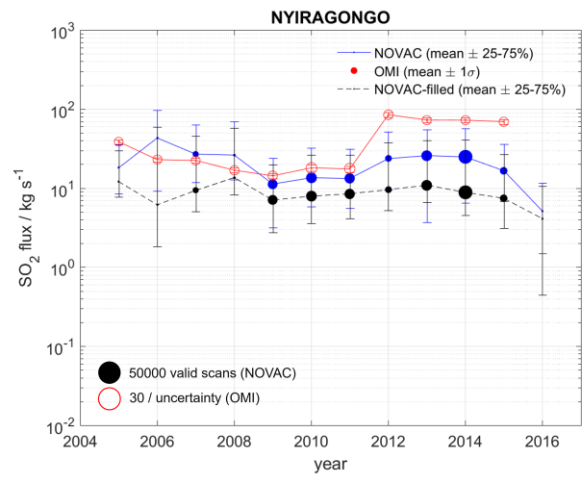
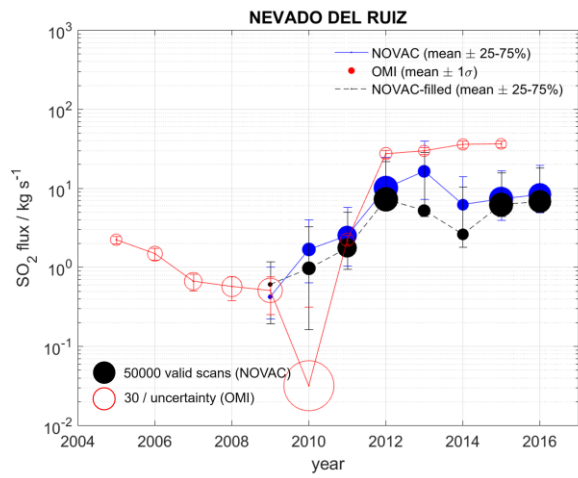
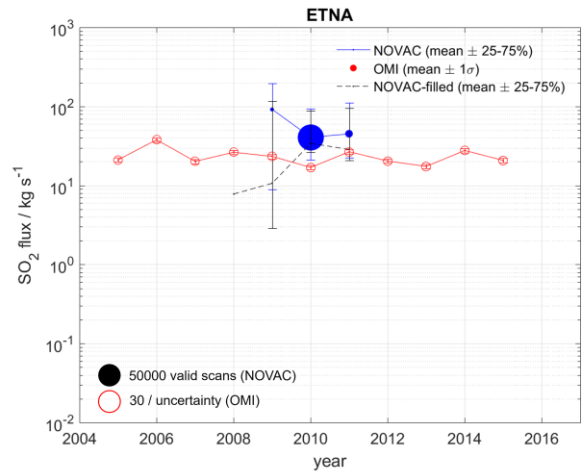
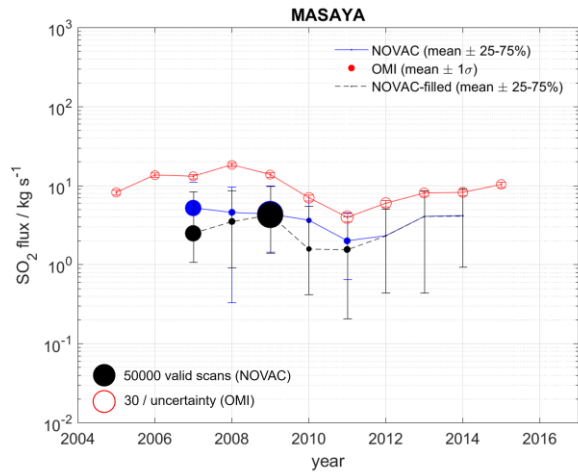
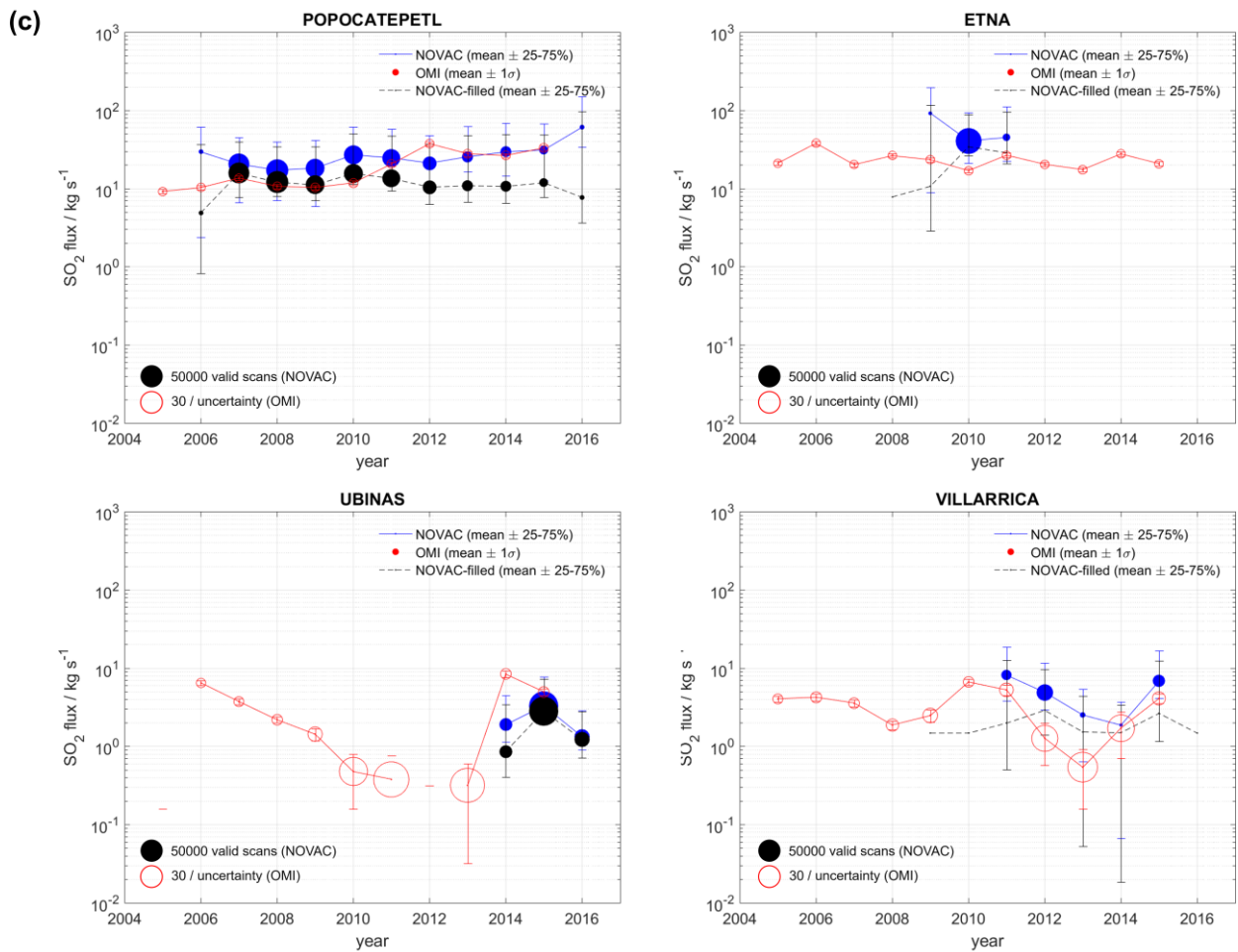



Figure 5. Time series of annual SO₂ emission from 12 volcanoes in NOVAC for which corresponding results from OMI are available for the period 2005-2016. Dots show the annual mean emission rate with size linearly proportional to the number of valid measurements used for the average. The bars indicate 25 and 75% quantiles of the daily means on each year. The series shown in blue correspond to observed plumes, while in black is the emission adjusted for periods of low degassing, when no plumes were observed (see text for details). The series in red is the mean annual emission rate obtained from OMI measurements, with size proportional to the precision (reciprocal of uncertainty) of the estimation and bars showing ±1 standard deviation reported by Carn et al. (2017). (a) Data from Copahue, Etna, Galeras, and Isluga. (b) Data from Masaya, Mayon, Nevado del Ruiz, and Nyiragongo. (c) Data from Popocatépetl, Tungurahua, Ubinas, and Villarrica. For details of the data in numerical format see Supplement S2.

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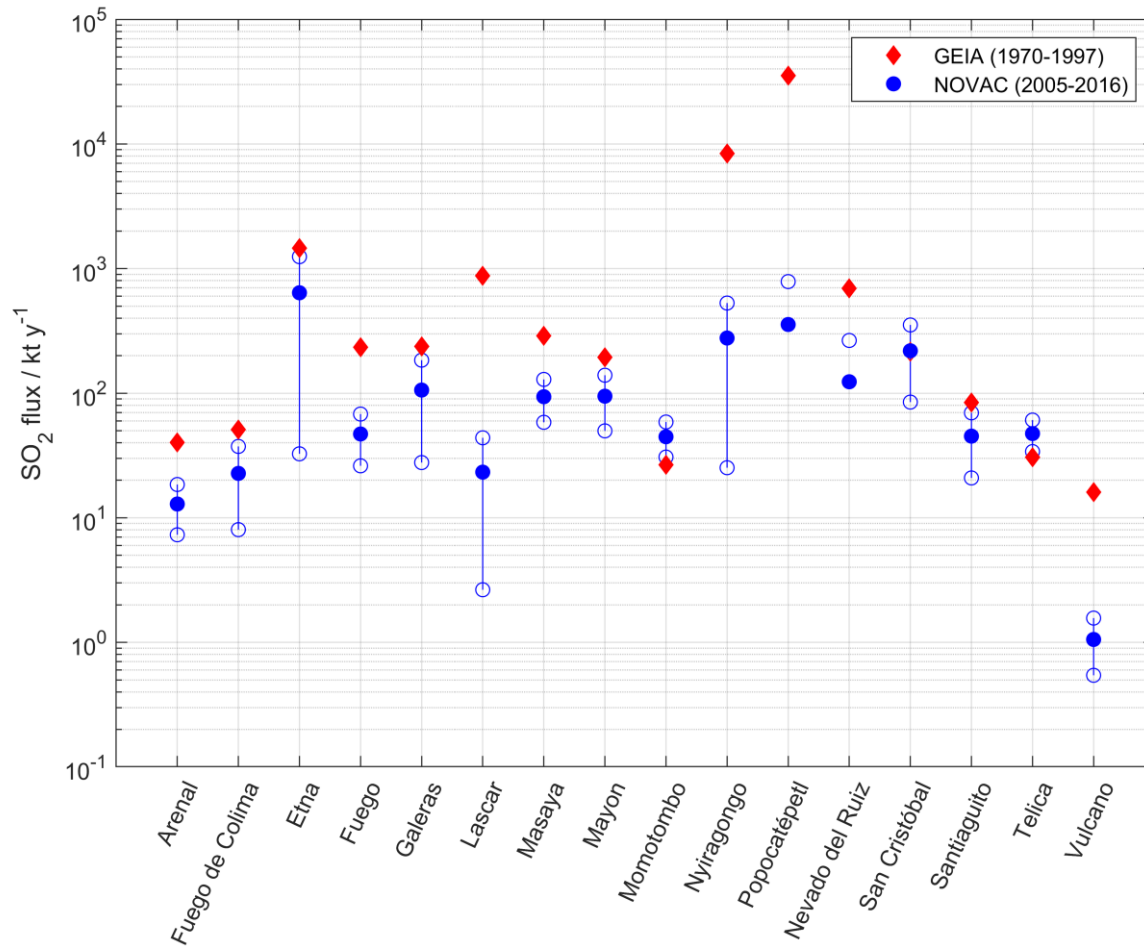
Figure 6. Comparison of emission statistics for 16 volcanoes of the NOVAC and GEIA datasets. For NOVAC the simple averages of annual mean (filled blue dots) and ± 1 standard deviations (un-filled blue dots) for the years 2005-2016 are depicted. The values for GEIA are obtained from Andres and Kasgnoc (1998) for passive degassing and sporadically degassing volcanoes only.

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