ADHI: The African Database of Hydrometric Indices (1950-2018)

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Abstract

The African continent is probably the one with the lowest density of hydrometric stations currently measuring river discharge, despite the fact that the number of operating stations was quite important until the 70s. This new African Database of Hydrometric Indices (ADHI) provides a wide range of hydrometric indices and hydrological signatures computed from ~~is compiling data from~~ different sources of data ~~carefully checked for~~ after a quality control. It includes ~~about 1500~~1466 stations with at least 10 years of daily discharge data over the period 1950-2018. The average record length is ~~19~~ 33 years and for over 100 stations complete records are available over 50 years. With this new dataset spanning most regions of the African continent, several hydrometric indices have been computed, representing mean flow characteristics and extremes (low flows and floods), and are made accessible to the scientific community. The database will be updated on a regular basis to include more hydrometric stations and longer time series of river discharge. The ADHI database is available for download at: https://doi.org/10.23708/LXGXQ9 (Tramblay and Rouché, 2020).

1. Introduction

There is a growing need for large-scale streamflow archives (Addor et al., 2020; Hannah et al., 2011), that are extremely useful to evaluate continental land-surface simulations (Archfield et al., 2015; Newman et al., 2015; Ghiggi et al., 2019; Do et al., 2020), remote sensing data products (Beck et al., 2017; Brocca et al., 2019; Forootan et al., 2019; Satgé et al., 2020), develop operational flood or drought monitoring systems (Alfieri et al., 2020; Harrigan et al., 2020; Lavers et al., 2019; Thiemig et al., 2011), or evaluate aquifer~~s~~ outflows and characteristics (Dewandel et al., 2003, 2004). In Africa, the density of active monitoring networks is lower compared to other continents and there are challenges in the exchange of hydrometric data across countries (Mahé and Olivry, 1999; Viglione et al., 2010; Mahe et al., 2013; Stewart, 2015; Dixon et al., 2020).

African countries are largely under-represented in large-scale databases such as the Global Runoff Data Center (GRDC) or the recent GSIM initiative (Do et al., 2018; Gudmundsson et al., 2018), and/or the time series are mostly not updated. At the African scale, there is still a lack of coordination for hydrological data collection and dissemination, despite the launch in 1975 of the UNESCO Intergovernmental Hydrological Program (IHP) dedicated to water research, water resources management, as well as education and capacity building. This initiative~~programme~~ enhanced the set up and management of international rainfall and runoff databases at the regional scale of the FRIEND programs (Van Lanen et al., 2014), but these are still largely not updated. There is still not enough partnership between the national hydrological services and in many countries licensing issues prevent the distribution of the data collected.

The density of monitoring networks in Africa has been declining over time; a serious concern for hydrologists since data acquisition and experimental data analysis remain central to understand hydrological processes and their spatio-temporal variability (Hannah et al., 2011; Roudier et al., 2014; Blume et al., 2016; Beven et al., 2020). There are ~~mainly two~~several reasons for this decline: the budgetary austerity measures imposed by the international financial institutions, the lack of permanent funding of national hydrological services, and the typically low number of well-trained technical staff in these departments (Bodian et al., 2016, 2020; Hannah et al., 2011). As a result, hydrological monitoring is now often dependent on research projects that cannot support long term observations. Studies focusing on regional river discharge variability are rare at the scale of Africa due to the lack of data. For instance, Conway et al. (2009) could only present a study ~~of~~ on a reduced number of representative regional basins ~~in their study of African rivers variability during the XXth century~~, and Roudier et al. (2014) compared only published anomaly results ~~rather than direct measurements~~ in their review of climate change impacts on the hydrology of West Africa.

Since in many cases, there are strict conditions related to the redistribution of un-processed data (Do et al., 2018), it is very often not possible to provide the complete time series of discharge data. Nevertheless, hydrological indices or hydrological signatures are useful to characterize the behavior of different components of river discharge, from low flows, annual runoff to floods (Addor et al., 2018; McMillan et al., 2017), and to assess the potential impact of climate change and human activities on river regimes (Mahe et al., 2013~~; Gnann et al., 2020~~). They can be used for various purposes, including basin classifications, aquifer properties characterization, hydrological predictions in ungauged catchments (Westerberg et al., 2016, Gnann et al., 2020) and to investigate long term trends for different hydrological processes (Do et al., 2017; Nka et al., 2015). We introduce here the African Dataset of Hydrometric Indices (ADHI) that aims at giving access to an ensemble of hydrometric indices computed from an unprecedented large ensemble of stations with daily discharge data (Tramblay et al., 2020, Tramblay and Rouché, 2020). Thus, a minimum of useful information regarding the African ~~rivers~~rivers' variability over the last 68 years can be shared with the international community, while respecting the confidentiality of the original records when these are not allowed to be publicly shared by the national authorities.

2. Data sources and processing

2.1 Data collection

The database used in the present work is based on the collection of stations from the Global Runoff Data Center (GRDC) and the SIEREM database (Boyer et al., 2006; Dieulin et al., 2019). The hydro-climatological data contained in SIEREM is the legacy from the former *Laboratoire d'Hydrologie* of the *Office de la recherche scientifique et technique outre-mer* (ORSTOM, now *Institut de Recherche pour le Développement*, IRD, France). It must be noted that in addition to the daily data, the SIEREM database also contains instantaneous rainfall and discharge for hundreds of experimental small catchments mostly established in the 1950s and 1960s. The criterion to include a station in the ADHI database is to have a minimum of 10 full years of daily discharge data between 1950 and 2018. Most of the hydrological stations in French-speaking countries have been set up and managed for decades by the ORSTOM Institute (Mahe and Olivry, 1999). At the time the data were processed, the SIEREM database included a total of 1046 series, with several of them being duplicates of the same monitoring station but for different time periods. There are a total of 101 stations with 2 times series, 42 stations with 3 time series, 24 stations with 4 time series and 7 stations with 5 time series. In most cases, one time series includes the longest record and that one was kept for the analysis in the present paper. For some stations, the different time series were differing substantially during the same period, due to different rating curves. A ~~careful~~ visual inspection of these series led to the elimination of erroneous or doubtful data. Only for 17 stations the time series were concatenated, after making sure the rating curve(s) applied on the different time periods to compute river discharge were adequate, by comparing daily runoff on a common period. Additionally, to these 1046 series, 933 stations have been retrieved from the GRDC database. For 106 of these stations, there was a duplicate station in the SIEREM database with longer time series and the latter were selected. After this data quality processing step, ~~673~~ 672 stations were kept for SIEREM and ~~800~~ 794 for the GRDC database for a total of ~~1473~~ 1466 stations (Figure 1). The stations from SIEREM mostly cover the Western, Central and Northern African regions and stations from the GRDC the Eastern and Southern parts of Africa. Figure 2 depicts the number of stations available per year, showing a sharp decline at the end of the 1980s, and shows the number of stations having from 10 to 69 years of record. It can be seen that, for about 100 stations, complete records are available over 50 years.

2.2 Data quality

Since the data collected are sometimes from manual records, they are subject to possible errors in the reporting of discharge values. ~~About~~ For outlier detection, no single method can outperform visual inspection and local expert knowledge (Crochemore et al., 2020). Indeed, in rivers with a strong variability in the annual regime and extremes, the most important flood peaks may be wrongly reported as outliers. Consequently, we carried out a visual inspection of the data when the maximum value was exceeding 5 times the median discharge. For only a few data points in the discharge time series, some obvious errors were detected~~,~~ with daily discharge exceeding by several orders of magnitude the median flow. In these~~, and for these~~ cases, the data has been reported as missing data in an absence of an objective criterion to correct the record. In addition, through visual inspection it was possible to identify stations where some gap filling methods have been applied (13 stations) or where the data are suspicious (28 stations). A flag has been added in the metadata to identify these stations. It is worth noting that, for the stations of the SIEREM database, most of the data were analyzed and criticized prior to the inclusion in the database by the former ORSTOM hydrology laboratory, with therefore a ~~much~~ reduced level of error in the archived data.

In addition, to detect possible shifts in the data due to non-natural influences, such as an artificial drift in the monitoring devices, changing instrumentation, recalibration of the rating curve, or river regulation by dams or reservoirs, the Pettitt test (Pettitt, 1979) was applied to mean annual runoff series. We reported the cases when the null hypothesis of homogeneity was rejected, at the 5% significance level. 14 stations are reported with homogeneity breaks in the metadata and this result was consistent with a visual inspection. Since the possible causes of these changes in flow regime could be ~~many~~manyfold~~ folds~~ and should be investigated with a more detailed case-by-case analysis, we choose to keep these stations in the database, but to flag them accordingly.

2.3 Climate characteristics

This data collection results in the largest ever built database of daily discharge data in Africa. These stations belong to different climate zones (Figure 1), according to the Köppen-Geiger climate classification (Peel et al., 2007). The main climate zone represented is Savannah (class Aw) for 687 stations corresponding to west and central Africa basins. The second most represented climate zone is Steppe-hot (Bsh) for 207 stations located in the Sahel region and southern Africa (Botswana, Namibia). The temperate with dry winter classes (Cwa and Cwb) include~~s~~ 187 and 125 stations, respectively located in southern Africa (Zambia, Angola, Rwanda, Mozambique, South Africa and Zimbabwe). The 98 stations belonging to the Desert-hot class (Bwh) are mostly located in the northern and southern boundaries of the Sahara Desert. 87 Stations under a temperate climate with dry hot summer, corresponding to Mediterranean climate (Csa) are found in North Africa and the southwestern part of South Africa. Thus, the selected river basins are representative of most of the climate zones in Africa. It must be noted that for large basins, such as the Congo, Niger or even the Orange rivers, the climate type at the outlet may not be representative of the whole catchment, that may span over diverse climate zones.

To document the mean annual precipitation and evapotranspiration at the catchment scale, the CRU4 dataset has been considered (Harris et al., 2020). However, without long-term and homogeneous ground monitoring networks over the African continent, no best precipitation database could be identified for Africa as a whole (Sylla et al., 2013; Beck et al., 2017; Awange et al., 2019; Satgé et al., 2020). For some regions, such as Northern or Equatorial Africa, there are large differences between different remote sensing or gauged-based precipitation products (Gehne et al., 2016; Harrison et al., 2019; Nogueira, 2020), in particular for extreme precipitation events. This is the reason why we choose to provide only mean annual precipitation, evapotranspiration and temperature. This implies that the ADHI dataset does not provide metrics relying on time series of precipitation or evapotranspiration, such as the runoff ratio, streamflow-precipitation elasticity or catchment response time. To calculate these indices requiring climatic time series for a given catchment, the user is advised to check first the best available data for that area.

2.4 Catchment delineation ~~and information about river regulation~~

Station catchments areas have been delineated with the Hydroshed Digital Elevation Model (DEM) at 15 sec resolution using the TopoToolbox2 algorithm (Schwanghart and Scherler, 2014). The map of the catchments is shown in Figure 3. Despite a careful check of the geographic coordinates of the stations, this type of automatic catchment delineation procedure is prone to some errors, in particular in regions with low elevation and flat terrain properties. This is particularly the case of catchments with endoreic areas, such as the Niger, Chari and Logone basins, where the precision of the DEM is crucial to identify these areas. Since the gauge locations are not necessarily located on the streams derived from the DEM, The TopoToolbox2 makes possible to re-locate automatically the gauges on the nearest river stream. However, this procedure did not work for 61 catchments, with a catchment area error exceeding 10% compared to the available metadata. For these basins, a manual procedure with the Arcmap® software has been implemented to delineate the catchment boundaries from flow direction maps. For several hundred of ~~catchments~~catchments, it is possible to compare the results of the automatic delineation procedure with the catchment ~~boundaries~~ areas available in the SIEREM database and the ORSTOM reports (available online at the adress: https://horizon.documentation.ird.fr), which have been most often individually delineated and carefully checked from ground knowledge over the years (Dieulin and Boyer, 2005). For 37 stations in the SIEREM database, the catchment areas where not correct in the metadata, by comparing the delineated catchments.

From the catchment delineated, the mean, maximum and minimum altitude from the Hydroshed DEM have been extracted and included in the metadata. In addition, the European Space Agency Climate Change Initiative Land cover data (ESA-CCI LC) (ESA, 2017) has been extracted for each catchment for the year 2015. This database contains land cover maps at a 300m spatial resolution for 38 classes, compliant with the UN Land Cover Classification System (LCCS). The classes have been grouped into 8 new classes: forest, urban areas, cropland, irrigated croplands, grassland, shrubland, sparse vegetation and bare land. Overall, the basins are characterized by a low proportion of urban areas, a large proportion of forests, especially in the intertropical zone (mean = 41%, median = 37%), and a majority of non-irrigated cultivated area, on average covering 31% of the total area of the basins. Indeed, the irrigated crops represent only 0.43% on average.

2.5 River regulation

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Dams and reservoirs have also been extracted and added in the metadata of the stations. The Global Reservoir and Dam Database (GRanD) v1.3 (Lehner et al., 2011) has also been considered to identify regulated basins. The number of dams included in each river basin has been extracted using the catchment boundaries. As shown in the metadata of GRanD, most of the dams in Africa basins have been constructed around the 1970s (Figure 4). The rivers could be considered regulated if at least one dam exists in the catchment area, otherwise the river is considered natural (Figure ~~1~~5). However, the influence of dams and reservoirs on the flow regime are linked to the location of the regulation structure, the portion of the basin controlled, and the management strategies. For instance, in a large basin with only one dam located on a small headwater catchment, its influence may not be distinguishable at the river outlet. On the other hand, a station located immediately downstream a dam outlet may have its flow regime strongly impacted by dam operations. It should be also noted that other regulation structures like small dams or water diversion channels that may not be included in the GRanD database could be present in the catchments considered natural (Lehner et al., 2011; Pekel et al., 2016). This is particularly the case in semi-arid areas where earthenmade -channels, often informal, draw their water supply from the river itself, by building small diverting structures (Underhill, 1984; Kimmage, 1991). They can represent a large number of structures, but a variable amount of water withdrawal at the basin scale (Barbier et al., 2009; Bouimouass et al., 2020). Similarly, no data is available yet on the importance and impact of groundwater abstraction, if any, on the flow regime measured at the stations.

3. Hydrometric indices

Here is presented the list of indices computed from daily discharge data. Most of the indices are computed with the Toolbox for Streamflow Signatures in Hydrology (TOSSH, available at the address: https://github.com/TOSSHtoolbox/) (Gnann et al., 2021). The indices and signatures selected spans a large variety of runoff characteristics from high to low flows, from previous literature (Poff et al., 1997; Richter et al., 1996; Baker et al., 2004; Yadav et al., 2007; Clark et al., 2009; Estrany et al., 2010; Sawicz et al., 2011; Euser et al., 2013; Safeeq et al., 2013; Addor et al., 2018; McMillan, 2020).

3.1 Methodological considerations

Several signatures charactering baseflow rely on the application of a base flow filter. Since the choice of the baseflow separation method can introduce uncertainties in the calculation of these signatures (Su et al., 2016), two baseflow filtering methods are compared: the Lyne and Hollick recursive digital filter (Ladson et al., 2013), with the default values for the filter parameter (0.925) and the number of passes (3), and alternatively the UKIH smoothed minima method (UKIH, 1980), that does not require any calibration parameter. The base flow index (BFI) is the ratio between the baseflow volume and the total streamflow volume. The baseflow recession (BaseflowR) is the baseflow recession constant assuming an exponential recession behavior (Safeeq et al., 2013). The base baseflow magnitude calculates the difference between the minimum and the maximum of the baseflow regime, defined as the average baseflow on each calendar day. The two base flow separation method compared to compute the baseflow-related indices provide very similar results, with a correlation above 0.9 for all indices obtained with the two approaches.

To compute the mean half flow date and the mean half flow interval, the beginning of the hydrological year has been defined as the month following the month with the minimum average runoff. Indeed, the hydrological year has different starting dates across the African continent, in North Africa the hydrological year usually starts in September, in western Africa around March-April and in January for southern Africa.

The mean half flow date is the day when the cumulative discharge reaches half of the annual discharge. The mean half flow interval is the time span between: i) the date on which the cumulative discharge since the start of water year reaches a quarter of the annual discharge and ii) the date on which the cumulative discharge since the start of water year reaches three quarters of annual discharge.

Some metrics are derived from the calculation of the Flow duration Curve (FDC), such as its slope between the 33rd and 66th flow percentiles (McMillan et al., 2017), the peak distribution, the slope between the 10th and the 50th percentiles of the FDC constructed only with hydrographs peaks (Euser et al., 2013) and the variability index, the standard deviation of the logarithms of discharge from 10th to the 90th percentiles of the FDC (Estrany et al., 2010). It must be noted that 194 rivers have more than 50% of days with zero-flow and for these stations, but also all the others with an intermittent regime, several metrics derived from the Flow Duration Curve (FDC) are not adapted. For these basins, specific methods to estimate the FDC should be applied (Rianna et al., 2013). Similarly, there is no baseflow in these basins. Consequently, the indices relying to base flow or the flow duration curve are not computed for these basins.

In addition, different hydrological signatures describing the hydrologic responses of the basins are also provided. The flashiness index is defined as the sum of absolute differences between consecutive daily flows (Baker et al., 2004), it reflects the frequency and rapidity of short term changes in streamflow, especially during high runoff events. The number of master recession curves (MRC) is computed from the changes in recession slopes, and represent different reservoirs contributing to the runoff response (Clark et al., 2009; Estrany et al., 2010). This signature can help to understand the functional forms of storage–discharge relationships and identify model structures adapted to represent it. The rising limb density is the ratio between the number of rising limbs and the total amount of timesteps in the hydrograph (Sawicz et al., 2011). It is a descriptor of the hydrograph shape and smoothness, without consideration for the flow magnitude. Small values of the rising limb density indicate a smooth hydrograph.

From the supplied indices, some other useful indicators could be derived. For example, for hydrogeology applications it would be interesting to compute the low stage specific discharge that is the ratio between the low-stage discharge and the area of the watershed. This can be an indicator of aquifers' contribution to river discharge. The main issue is related to the definition of the low-stage discharge. From the indices proposed in the present database, it could be 5th percentiles of daily streamflow or the minimum of 7-days consecutive streamflow, per year. Similarly, the low-flow index could be computed from the ratio of the 90th and 50th percentiles of daily streamflow (Smakhtin, 2001).

3.2 Indices computed on the whole record

These indices have been computed using the whole time series available for each station. Consequently, they are computed on different base periods depending on the stations, with the period of record for each station being made available in the metadata. These indices include:

1. Mean daily streamflow, the arithmetic mean of daily data
2. Standard deviation of daily streamflow
3. Minimum daily streamflow
4. Maximum daily streamflow
5. Mean monthly streamflow (12 values from January to December)
6. 5th, 10th, 25th, 50th, 75th, 90th, 95th and 99th percentiles of daily streamflow
7. BFI_LH = Baseflow index, with the Lyne and Hollick baseflow separation method
8. BFI_UKIH = Baseflow index, with the UK Institute of Hydrology baseflow separation method
9. BaseflowR = Baseflow recession
10. BaseflowM_LH = Baseflow magnitude, with the Lyne and Hollick baseflow separation method
11. BaseflowM_UKIH = Baseflow magnitude, with the UK Institute of Hydrology baseflow separation method
12. CoV = Coefficient of variation of runoff
13. HFD_mean = Mean half flow date
14. HFI_mean = Mean half flow interval
15. AC1 = lag-1 autocorrelation of flow
16. AC7 = lag-7 autocorrelation of flow
17. FDC_slop = Slope of flow duration curve
18. PeakDistri = Peak distribution
19. FlashI = Richards-Baker flashiness index
20. MRC_num = Number of master recession curves
21. Q_skew = Skewness of runoff
22. Q_var = Variance of runoff
23. RLD = Rising limb density
24. VariI = Variability index
25. ~~6.~~ Freq_0 = Frequency of zero-flow days

The basins included in the ADHI database include a wide range of catchment areas, from a few square kilometers to several hundred thousand, in the case of large rivers such as the Congo, Niger, Orange, Zambezi, Senegal, Okavango and Volta. As shown in Figure 6, the average runoff is generally well correlated to the size of the basins with nevertheless a variability linked to local climatic and geological conditions. The mean annual precipitation is one of the explanatory factors of the observed ranges of mean river runoff, but also strongly modulated by local conditions. A large number of basins have an aridity index (ratio between precipitation and potential evapotranspiration) of less than 0.60, indicative of arid to semi-arid conditions (figure 7a). The varying degrees of aridity encountered in the basins are an important explanatory factor for the hydrological response at the African scale. For instance, the coefficient of variation of runoff (figure 7b) or the flashiness index (figure 7c) have greater values under conditions of increasing aridity.

3.3 Indices computed on ~~an~~ monthly or annual basis

These indices have been computed for each calendar year, for consistency with other databases such as GSIM (Do et al., 2018; Gudmundsson et al., 2018). These ~~indexes~~ indices have been computed for the years with less than 5% missing data:

1. Mean annual runoff
2. Minimum of 7-days consecutive streamflow, per year, and corresponding date
3. Annual maximum runoff, and the corresponding date
4. ~~3.~~4. Annual values for the 5th, 10th, 25th, 50th 75th, 90th, 95th and 99th percentiles of daily streamflow

In addition to these annual series, the monthly time series contains for each month the mean, maximum and minimum runoff, the last column being the number of missing days per month. There is one file per station. It is advised to consider the monthly values only for the months with no missing values, or missing values less than 10% or 5%.

These time series make it possible to analyze the long-term evolution of mean and extreme runoff (Tramblay et al., 2020), but can also be useful to validate hydrological modelling results. Focusing on extreme high and low runoff, very different seasonal patterns of occurrence could be observed for different regions of Africa. On figure 8 are plotted the mean dates of annual maximum runoff and the annual minimum of 7-day runoff. This seasonal analysis has been performed with directional statistics (Burn, 1997; Mardia et al., 2015): the dates of occurrence were converted into angular values to compute the mean date of occurrence (θ) together with the concentration index (r), which is a measure of the flood occurrences variability around the mean date. The annual maximum runoff shows three distinct patterns (Figure 8): First, stations with floods occurring during December-February in northern and southern Africa, with a strong variability of their date of occurrence. Second, the stations in western Africa with floods occurring during summer and a low seasonal variability. Third, the stations in central-south Africa, with floods occurring in boreal spring and early summer with various degrees of variability depending on the sub-region considered and the level of aridity. For annual minimum runoff, the patterns are usually reversed, with the low flow period spanning on average during June to October in North Africa, January-March in western Africa, and between September and November in southeast Africa. Yet this global picture hides local behaviors such East-West contrast in southern Africa or the North-South gradient in West Africa (Mahe et al., 2013). Similarly, the observed variability even for some neighboring catchments reflects the local influences of topography, soils and land cover. As noted previously, the seasonal variability of extreme high or low runoff events is also strongly related to the catchment aridity.

4. Data availability

The ADHI database is available for download at: https://doi.org/10.23708/LXGXQ9 (Tramblay and Rouché, 2020). These different files are supplied in the AHDI database:

The ADHI_stations.dat file contains:

-Unique identifier for each station
-Station code (native code from the original datasource)
-Station Name
-Data Source
-Catchment Area (km²)
-Mean Altitude (m)
-Maximum Altitude (m)
-Minimum Altitude (m)
-Mean annual precipitation (mm)
-Mean annual evapotranspiration (mm)
-Mean annual temperature (°C)
-Forest cover (%)
-Urban areas (%)
-Cropland (%)
-Cropland, irrigated (%)
-Grassland (%)

-Shrubland (%)
-Sparse vegetation (%)
-Bare land (%)
-Starting year of the data records
-Ending year of the data records
-Longitude (WGS84)
-Latitude (WGS84)
-Number of dams
-Country
-Flag, 0: no identified data issue, 1: some gap filling detected, 2: suspicious data, 3: Obvious regime break

The ADHI_summary.dat file contains for each station (lines) the following variables (columns):

Mean_q = Mean daily streamflow (m3/s)
Std_q = Standard deviation of daily streamflow
Mini_q = Minimum daily streamflow
Maxi_q = Maximum daily streamflow
Jan_q, Fev_q... Dec_q = Mean monthly streamflow (12 values from January to December)
q5th, q10th, q25th, q50th q75th, q90th, q95th and q99th percentiles of daily streamflow
BFI_LH = Baseflow index, with the Lyne and Hollick baseflow separation method
BFI_UKIH = Baseflow index, with the UK Institute of Hydrology baseflow separation method
BaseflowR = Baseflow recession
BaseflowM_LH = Baseflow magnitude, with the Lyne and Hollick baseflow separation method
BaseflowM_UKIH = Baseflow magnitude, with the UK Institute of Hydrology baseflow separation method
CoV = Coefficient of variation of runoff
HFD_mean = Mean half flow date
HFI_mean = Mean half flow interval
AC1 = lag-1 autocorrelation of flow
AC7 = lag-7 autocorrelation of flow
FDC_slop = Slope of flow duration curve
PeakDistri = Peak distribution
FlashI = Richards-Baker flashiness index
MRC_num = Number of master recession curves
Q_skew = Skewness of runoff

Q_var = Variance of runoff
RLD = Rising limb density
VariI = Variability index
Freq_0 = Frequency of zero-flow days

The compressed folders AnnualMean.zip, AnnualMax.zip, Annual7DayMin.zip, AnnualPercentiles.zip contains time series for mean annual runoff, annual maximum runoff, annual minimum of 7-day discharge and annual values for the 5th, 10th, 25th, 50th 75th, 90th, 95th and 99th percentiles of daily streamflow. There is one file per station.

The compressed folder MonthlySeries.zip contains for each month the mean, maximum and minimum runoff, the last column is the numer of missing days per month. There is one file per station.

The compressed folder Plots.zip contains for each station a plot of the daily discharge data available.

The compressed folder Catchment_boundaries.zip contains the catchment boundaries in the shapefile format (one .shp file per basin).

The compressed folder Catchment_plots.zip contains for each basin a plot of the catchment area in .PNG format.

5. Conclusions and perspectives

This new hydrological database brings together the largest number of African river flow measurement stations, in comparison with other previously published datasets. In this ADHI dataset, we included a total of 1466 stations with at least 10 years of discharge data between 1950 and 2018, for a mean record length of 33.3 years. Half of the stations have more than 30 years of data. By comparison, the recent GSIM database contains 979 stations in Africa, with a record length varying from 1 year to 110 years until 2015, and a mean record length of 33.8 years. This ADHI database ~~It~~ results from a pooling of the GRDC and SIEREM databases, built from contributions of several agencies in African countries in charge of the management of hydrological measurement networks. This database will be regularly updated with data from SIEREM and GRDC. Since most of the pre-processing steps have been automated, it would be possible to increase the number of stations considered or the length of the data series, if more data would become available. The data from the SIEREM database is already regularly updated from contributions of different institutes. In the future, individual contributions from researchers or institutes will be also welcome to increase the spatio-temporal coverage of the data. The FRIEND program~~me~~ (UNESCO/IHP) will also contribute to increase the number of stations through coordinated efforts at the regional level. The dataset provides a series of indices that describes a wide range of mean and extreme runoff properties, allowing the characterization of the hydrological regime and applications linked to the management of water resources and hydrological risks. This database includes ~~a wide range of~~different catchment sizes and rivers with different hydrological regimes that makes possible to analyze the behavior of rivers in very different contexts for a wide range of scales.

More broadly, this ADHI database could contribute to a better knowledge on African hydrology. For instance, the impacts of dams on river discharge remains largely unquantified at the scale of Africa (Biemans et al., 2011). From these indices, various applications can be sought. For example, the percentiles of the daily streamflow could be useful to calibrate hydrological models using the flow duration curve (McMillan et al., 2017) and to constrain model outputs (Tumbo and Hughes, 2015; Ndzabandzaba and Hughes, 2017). Flow duration curves are also useful for catchment classification according to their rainfall-runoff response (Cheng et al., 2012). In the recent years, global runoff simulations have been provided by the Global Flow Awareness System, with land surface or global hydrological model driven by reanalysis data (Alfieri et al., 2020; Harrigan et al., 2020). Yet, due to the small number of stations representing African basins in the currently available databases preventing a robust calibration of the models, the hydrological simulations have a poor performance (Harrigan et al., 2020). More generally, this new ADHI database could open perspectives to apply hydrological models in African basins, in particular combined with recent remote sensing data products (Brocca et al., 2019; Satgé et al., 2020). Beside deterministic hydrological modelling approaches, several statistical methods to estimate the return levels of floods have been proposed, in order to safely design dams, reservoirs, sewers or other water regulation structures. Regional frequency analysis methods have been applied to estimate floods in ungauged basins in several African countries such as Morocco (Zkhiri et al., 2017), Tunisia (Ellouze and Abida, 2008), South Africa (Nathanael et al., 2018; Smakhtin et al., 1997), or the Volta basin (Komi et al., 2016). However studies at a larger regional scale remain very scarce (Farquharson et al., 1992; Padi et al., 2011) while there is a strong need to improve the knowledge on hydrological hazards in African countries (Di Baldassarre et al., 2010). With this recent database becoming available, it could be possible to develop regional frequency analysis techniques for floods or low flows tailored for the African context, taking also into account the impacts of global changes.

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The database is available from the online repository: https://doi.org/10.23708/LXGXQ9 Additional indices could be computed upon reasonable request to the corresponding author.

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Figures

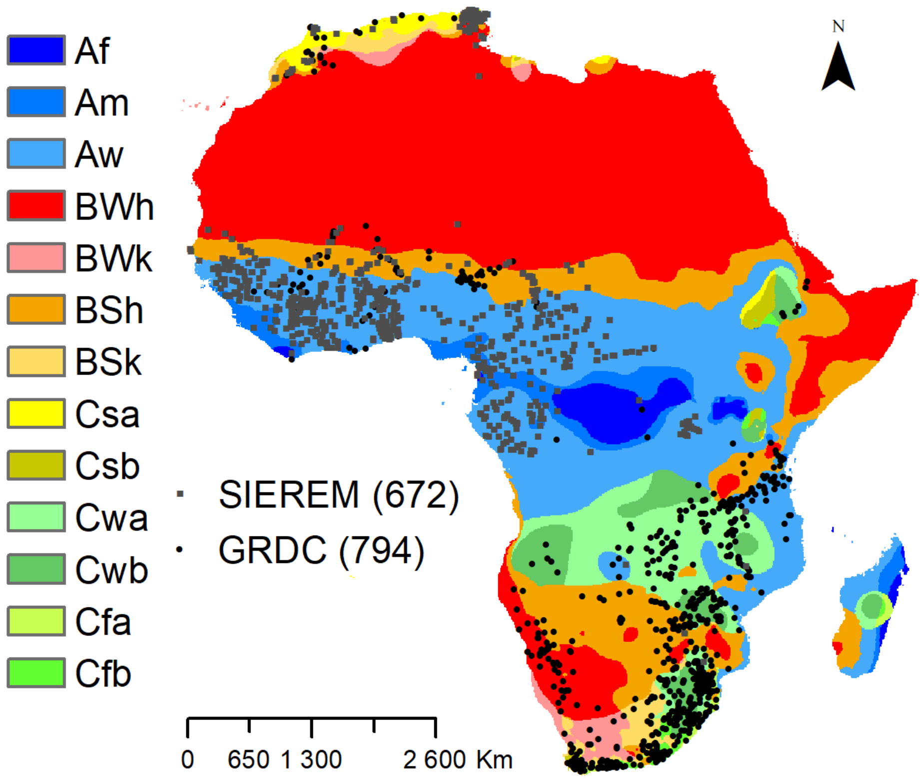

Figure 1: Map of the selected stations for the ADHI database from the SIEREM and GRDC datasets. The different colors represent the main climate zones in Africa from the Köppen-Geiger climate classification (Peel et al., 2007)

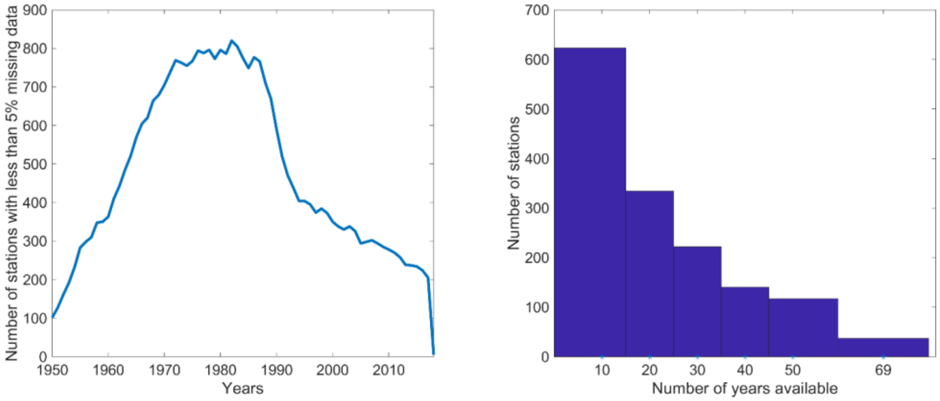

Figure 2: Number of available stations per year with less than 5% missing data (left) and number of stations available for different record lengths (right)

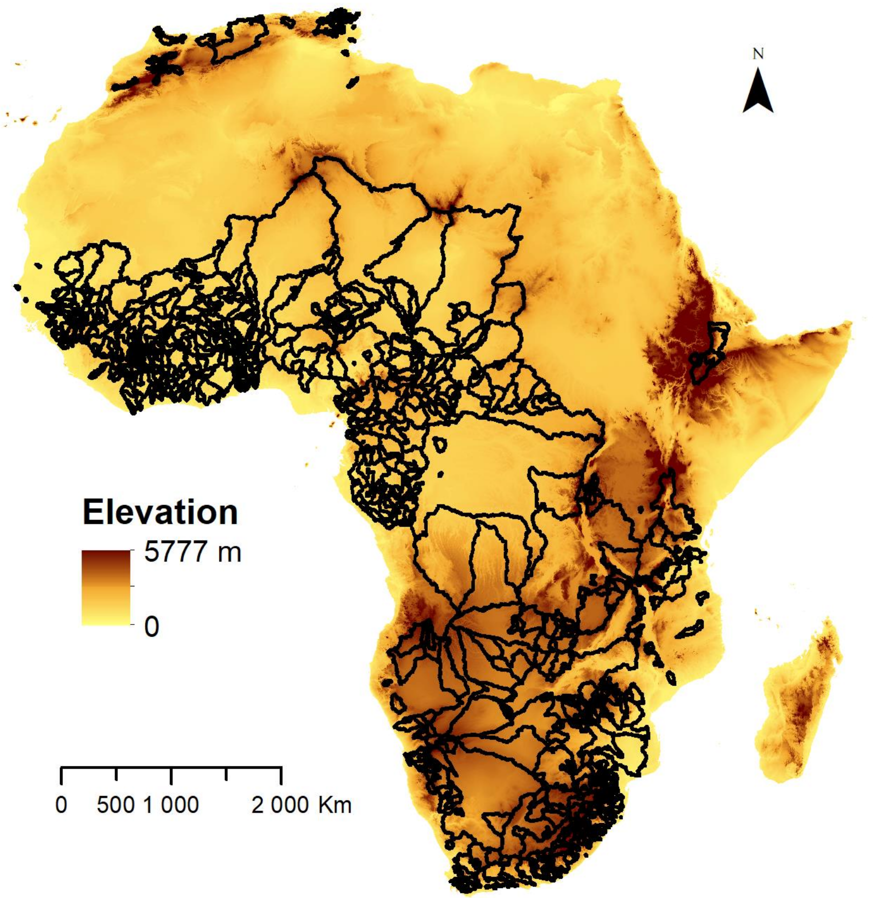

Figure 3: Map of the delineated catchment boundaries in black, with elevation from HydroSheds digital elevation model (https://www.hydrosheds.org/).

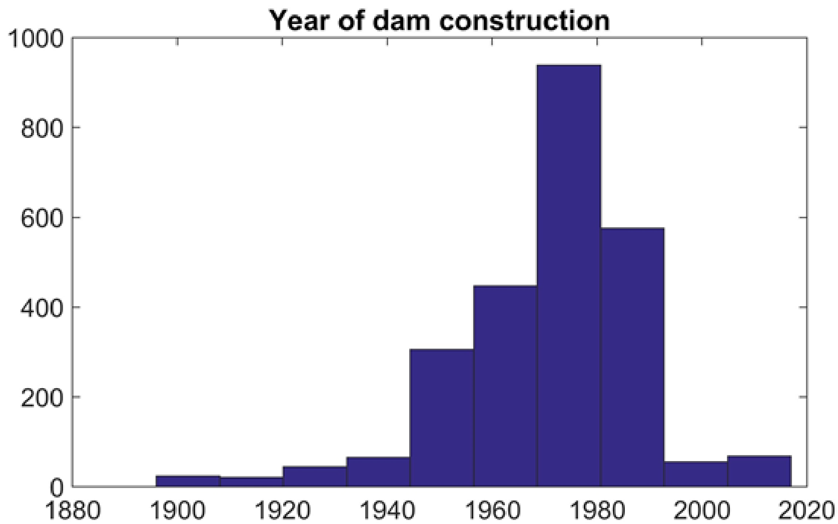

Figure 4: Years of building date for dams located in the catchment database (data from the Global Reservoir and Dam Database v1.3)

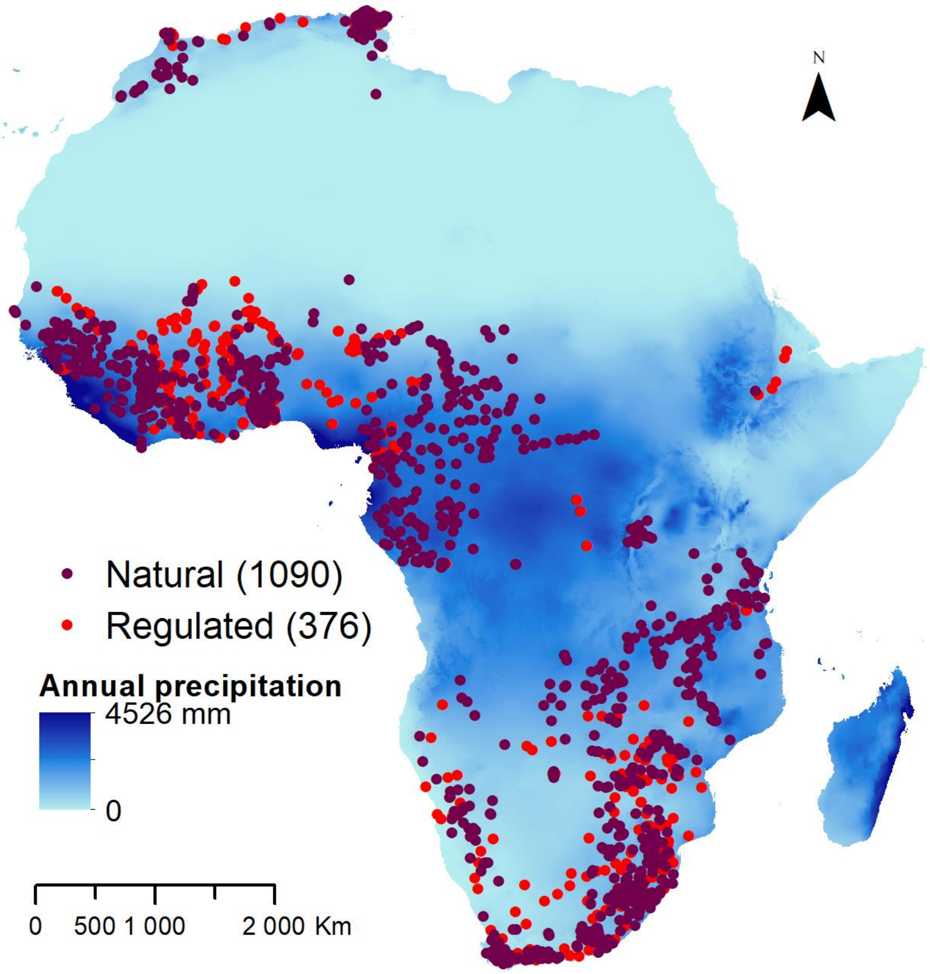

Figure 5: Map of stations with a natural or regulated flow regime. Basins are considered regulated if they contain at least one dam or reservoir from the GRanD database (Lehner et al., 2011). Mean annual precipitation between 1970 and 2000 is provided from the WorldClim database (Fick and Hijmans, 2017).

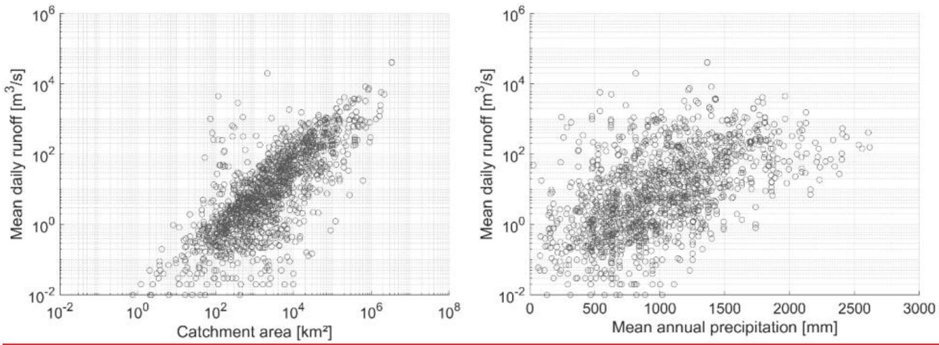

Figure 6: Relationship between mean daily river discharge and catchment area (left) and mean annual precipitation (right)

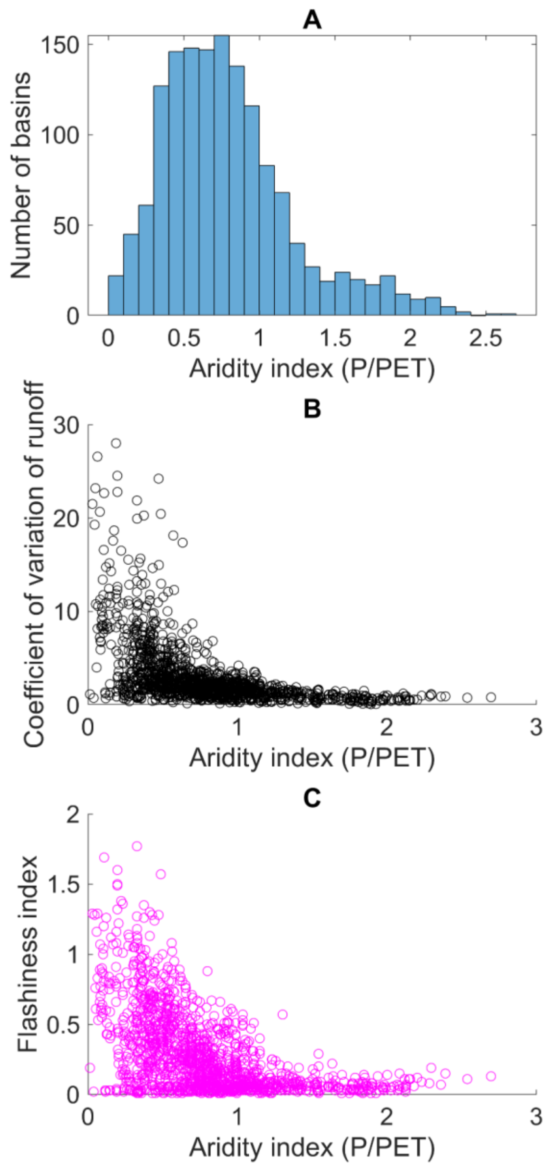

Figure 7: histogram of the aridity index per basin (A), relationship between the aridity index and the coefficient of variation of runoff (B), relationship between the aridity index and the flashiness index (C)

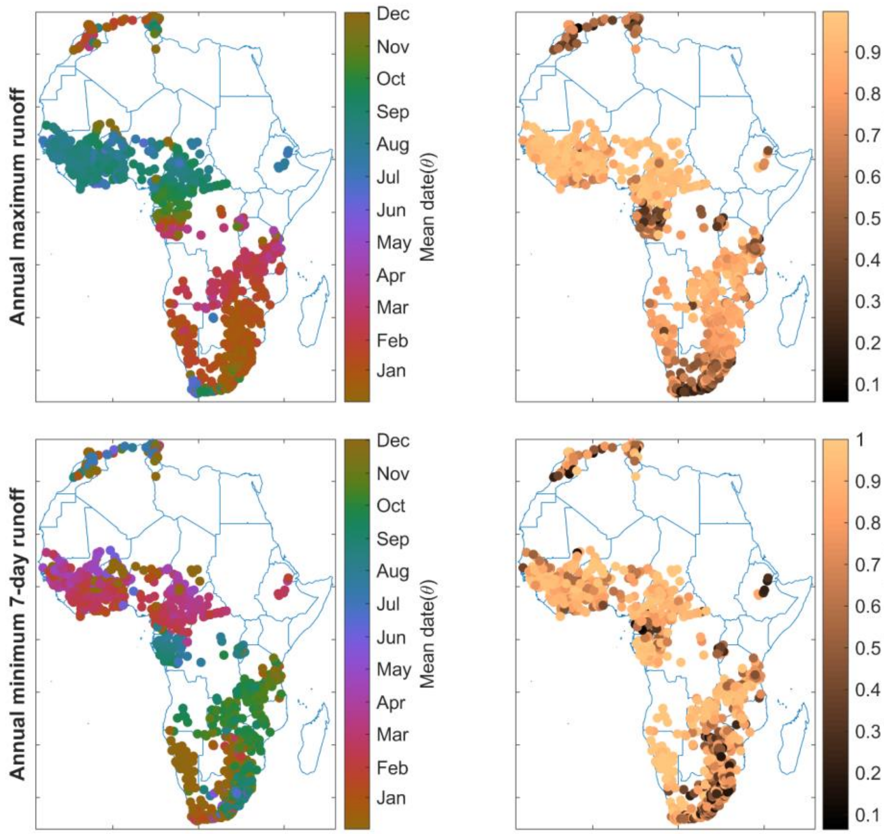

Figure 8: Mean date of occurrence (left) of annual maximum runoff and annual minimum of 7-day runoff, together with the variability around the mean date (right) represented by the concentration index