

## Summary of Changes:

Revisions made to address the referee comments improved the clarity and accuracy of the paper. A response is given below to each comment where a response was required. Minimal changes were made to the paper beyond what was done to address the comments. A revised version of the dataset is now available at Ag Data Commons. The revision to the dataset corrected two metadata issues: first, as the referees point out, the record length (10, 20 or 30 years) was not correctly shown in the headings of the 10 and 20-year .par files, and this is now corrected; second, improvements were made to the formatting of locality names in the headings of the .par files. There were also changes made to correct inconsistencies in MX.5P and transition probability parameter values for very dry climates or months. Some values were predicted to be smaller than  $10^{-3}$  (or two decimal places), which is the minimum value allowed by CLIGEN, and were therefore set to zero. This sometimes corresponded with non-zero MEAN P values. If this was the case, the MX.5P or transition probability was set to the minimum allowed value of  $10^{-3}$ . This had a small effect on the validation metrics in Table 4.

## RC1 Key Points:

**RC1 #1) While data sources for individual weather variables were described in reasonable details. What is missing is the table summarizing the spatial resolution for each.**

Since the precipitation parameters were point-scale (or site-specific), spatial resolution is relevant only for the temperature and solar radiation parameters derived from gridded products. The climate models that were used, ERA and GLDAS, both have  $0.25^\circ \times 0.25^\circ$  resolution, and this is now stated in the text instead of in a table.

**RC1 #2) How were temperature parameter values prepared for individual sites? Were gridded temp data mapped to individual precipitation sites through spatial interpretation techniques? The same applies to solar radiation data.**

This relates to the previous comment. In the same revision to include the spatial resolution of the climate models, it is now stated that no weighting of values based on proximity of a station to neighbouring cells, or other forms of interpolation, is done. The  $0.25^\circ$  resolution translates to ~28 km resolution at the equator, which is reasonable resolution for this application. Using the original values from the models also enables easier interpretation of existing uncertainty information for respective models.

**RC1 #3) Recognition of the issue with record length is useful, but not critical. In CLIGEN parameter file, the number of years of data is recorded. I would leave at that, Caveat Emptor!**

We feel that the record length issue is worth explaining in the text because it may complicate the interpretation of the data and any future validation that is done through comparison to other climate records, particularly for non-stationary climates or climates with long-term cycles.

**RC1 #4) I have had a close look at the parameter files generated. For 20-year data set. The years used are still 30.**

This metadata issue is now corrected in the revised dataset.

**RC1 #5) Fig.1 Why are there more sites with 30-year data than those with 10? Any station with 30 year would also have 10 years of data?**

Correct, any station with 30 years of data would be viable for the 10-year dataset. The longest possible record length (of 10, 20, or 30 years) was used for a given site, such that if a 30-year dataset was possible, a 10 and 20-year dataset were not made in addition. So, no site had multiple datasets created for it. This is now stated in the text, and this partly explains why the 30-year dataset has the most locations. It is also the case that many NOAA-GHCN sites included a long backlog of data at the time of being added to the network, for which 30-year datasets were possible, while the 10-year datasets tended to come from newer installations without a long backlog.

**RC1 #6) (Equation) (1) 'n days' 'MEAN P' should not be used as variables in the equation. Equations need to be readable, clear, precise.**

In the context of the equation, variable names for n days and MEAN P were changed to shorter names with no spaces, making the equation easier to read. The identities of the new variable names are explained in the text under the equation.

**RC1 #7) (Equation) (2) The equation is wrong, once the summation sign is used. there is no need for all other terms. That is what the summation is for.**

Correct, the "+...+" inside of the summation operator should not be shown. This was meant to be "...," which clarifies what the set of terms is being summed.

**RC1 #8) (Equation) (3) Again 'Time Pk(i)', any variables with a space ' ' in them can lead to confusion One bracket is missing from the equation in the third row in Table A2. If there is space in the variable name, use ' ' for the variable.**

The space was removed from the Time Pk variable name in the equation and where it is used in sections of the text. The space was previously used to be consistent with what is shown in CLIGEN .par files, but the connection between the two labels is evident and shouldn't lead to confusion. Several of the CLIGEN

parameter names have spaces in them but are not used frequently in the text like Time Pk is. So, spaces in the other names were kept. Also, the percent bias equation was missing a bracket around  $(O - P)$ , which is now fixed.

## **RC2 Key Points:**

**RC2 #1) In lines 83-84 and 115-118 slightly mismatching statement are proposed. Using complete months in non continuous series could drive to incompatibilities in temporal comparison of the proposed parameters?**

The statement on 83-84 does imply continuous records should be used. The handling of the pervasive data gaps in NOAA-GHCN records becomes a source of error, and complicates validation using other datasets with the same temporal ranges. A stronger statement is made about the uncertainty from using non-continuous records with mention of the uncertainty associated with non-stationary climates, long-term climate cycles, and the complication that arises when comparing to other climate data.

**RC2 #2) Overplotting occurs in figure 1, maybe a thematic raster "distance from nearest location" can enhance the information provided? Furthermore details on spatial coverage of the proposed parameters could be provided.**

A new figure (Fig. 2) was added that shows a raster of station density. This allows the reader to see the relative density of stations in places where overcrowding makes this impossible.

**RC2 #3) The collection and harmonization of international climate data encounters notorious obstacles, their once for all overcome falls in the goals of this work. The methods used are soundly reported, but -I miss an explicit criterion for gridded model temperature values: radiation is in ERA5 dataset. Why use gridded temperature and not radiation? -About SD RAD, the 2.6 closes on GLDAS-Ameriflux comparison, the proposed global parameters rely on the continue model of GLDAS 3h values?**

The temperature and solar radiation parameters come from ERA and GLDAS gridded climate models, respectively, and this is now made more explicit in the text. A lack of clarity about this may have been leading to a misunderstanding originating from the discussion of the use of point-scale Ameriflux ground observations for validation that solar radiation parameters did not come from a gridded product. So, answering the question in the comment, it is correct that the parameters rely on the GLDAS 3h values.

**RC2 #4) I would expect 30y points to be fulfilling the requirements of 10y ones, it doesn't look so in figure 1.**

The same question was addressed in response to RC1 #5.

**RC2 #5) May the definition of time-series at line 11 be improved using "maximum" or "available" rather than "minimum"?**

“Minimum” was changed to “available”. Use of “available” in the context of record length may suggest that some screening or filtering was done, which is the case.

**RC2 #6) Somewhere in the text CLIGEN parameters are referred with no introduction (ie. lines 92, 236), their presentation would get the text easier to read.**

Table 2 gives all parameter definitions, so rather than restating their definitions in the text, reference to Table 2 is made at the beginning of each of the major methods sections where the respective parameters being calculated are generally listed.

# 1 **Climate benchmarks and input parameters representing locations in** 2 **68 countries for a stochastic weather generator, CLIGEN**

3 Andrew T. Fullhart<sup>1</sup>, Mark A. Nearing<sup>1</sup>, Gerardo Armendariz<sup>1</sup>, Mark A. Weltz<sup>2</sup>

4 <sup>1</sup>Southwest Watershed Research Center, USDA-ARS, 2000 E. Allen Rd., Tucson, AZ, 85719, U.S.A.

5 <sup>2</sup>Great Basin Rangelands Research Unit, USDA-ARS, 920 Valley Rd., Reno, NV, 89512, U.S.A.

6 *Correspondence to:* Andrew T. Fullhart (andrew.fullhart@usda.gov)

7 **Abstract.** This dataset contains input parameters for 12,703 locations around the world to parameterize a stochastic weather  
8 generator called CLIGEN. The parameters are essentially monthly statistics relating to daily precipitation, temperature and  
9 solar radiation. The dataset is separated into three sub-datasets differentiated by having monthly statistics determined from 30-  
10 year, 20-year, and 10-year ~~minimum~~ record lengths. Input parameters related to precipitation were calculated primarily from  
11 the NOAA GHCN-Daily network. The remaining input parameters were calculated from various sources including global  
12 meteorological and land-surface models that are informed by remote sensing and other methods. The new CLIGEN dataset  
13 includes inputs for locations in the U.S., which were compared to a selection of stations from an existing U.S. CLIGEN dataset  
14 representing 2,648 locations. This validation showed reasonable agreement between the two datasets, with the majority of  
15 parameters showing less than 20% discrepancy relative to the existing dataset. For the three new datasets, differentiated by the  
16 minimum record lengths used for calculations, the validation showed only a small increase in discrepancy going towards  
17 shorter record lengths, such that the average discrepancy for all parameters was greater by 5% for the 10-year dataset. The new  
18 CLIGEN dataset has the potential to improve the spatial coverage of analysis for a variety of CLIGEN applications, and reduce  
19 the effort needed in preparing climate inputs. The dataset is available at the National Agriculture Library Data Commons  
20 website at [https://data.nal.usda.gov/dataset/international-climate-benchmarks-and-input-parameters-stochastic-weather-](https://data.nal.usda.gov/dataset/international-climate-benchmarks-and-input-parameters-stochastic-weather-generator-cligen)  
21 [generator-cligen](https://data.nal.usda.gov/dataset/international-climate-benchmarks-and-input-parameters-stochastic-weather-generator-cligen) and <https://doi.org/10.15482/USDA.ADC/1518706> (Fullhart et al., 2020c).

22  
23  
24  
25  
26  
27

## 29 **1 Introduction**

30 Essential climate variables defined by the World Meteorological Organization are physical, chemical, or biological  
31 variables, or groups of linked variables that critically contribute to the characterization of Earth's climate (Bojinski et al.,  
32 2014). Aside from their use in climate studies, basic essential climate variables like precipitation and temperature are important  
33 for water resource management, drought monitoring, agricultural engineering, and other applications (Hollmann et al., 2013).  
34 The temporal resolution of climate data varies for these applications. Climate data reduced to monthly statistics may facilitate  
35 analysis of multi-decadal climate trends and serve as benchmarks of climate normals (Menne et al., 2012; Hollmann et al.,  
36 2013). In this paper, it is discussed how a stochastic weather generator may be parameterized with a new dataset of monthly  
37 climate statistics to simulate daily weather outputs for locations around the world.

38 Stochastic weather generators are used for a variety of applications that include model forcing, statistical downscaling  
39 of climate models, and study of climate change scenarios (Vaghefi and Yu, 2017). CLImate GENerator (CLIGEN) is one such  
40 point-scale weather generator that produces daily outputs based on input parameters that are essentially observed monthly  
41 statistics. CLIGEN is regularly used to provide soil erosion models with realistic trends and statistical distributions of weather  
42 parameters (Kinnell 2019). Such models include the Rangeland Hydrology and Erosion Model (RHEM); the Water Erosion  
43 Prediction Project Model (WEPP); and the Revised Universal Soil Loss Equation 2 Model (RUSLE 2). CLIGEN can generate  
44 long-term realizations of stationary climate, subsequently enabling long-term erosion simulations, and ensuring that average  
45 annual erosion rates reach convergence (Baffaut et al., 1996). CLIGEN has been validated in a number of countries, under a  
46 variety of climates, and for different outputs that include daily precipitation, peak intensity, time-to-peak intensity, storm  
47 duration, and storm frequency. For example, Mehan et al. (2017) showed that the mean of all daily precipitation values was  
48 within 0.1 mm of observations, and minimum and maximum daily temperatures within 0.1 °C for locations in the western  
49 Lake Erie basin. A particularly important CLIGEN output is precipitation intensity because of its high model sensitivity in  
50 erosion and runoff modeling (Nearing et al., 2005). Zhang et al. (2008) validated intensity for the loess plateau of China based  
51 on distributions of maximum 30-min intensities ( $I_{30}$ ) that were derived from CLIGEN's peak intensity. They found that  
52 differences with observed distributions were statistically insignificant, suggesting that rainfall erosivity could be accurately  
53 estimated using CLIGEN.

54 ~~CLIGEN has a dataset of~~ location-specific input parameters for the United States with dense coverage, but  
55 on a global scale, input parameters are sparsely available. This is partly because of the labor-intensive nature of determining  
56 the parameters, and because of numerous data requirements, e.g., high-frequency precipitation measurements. For erosion  
57 modeling, the lack of widely available CLIGEN inputs has hindered progress towards increasing the spatial scale and coverage  
58 of analysis that other aspects of soil erosion research have brought to the global scale, one example being the development of  
59 global maps of annual rainfall erosivity (Panagos et al., 2017). Hence, in the interest of increasing the availability of CLIGEN

60 inputs for soil erosion modeling and other applications, we present a dataset of CLIGEN input parameter files. The dataset  
61 represents 12,703 locations in 68 countries. Besides providing the necessary parameters to run CLIGEN simulations, the  
62 dataset also serves to provide statistics for representing climate normals. The parameters are validated using an existing  
63 CLIGEN input dataset for the United States, and differences are discussed.

## 64 2 Datasets

### 65 2.1 Overview

66 Three sets of CLIGEN v5.3 input files for international locations are presented, differentiated by having monthly  
67 parameters determined from minimums of 30-year, 20-year and 10-year records (note that assumptions were made to handle  
68 data gaps which are discussed in Sect. 2.2) (Fullhart et al., 2020c). The distribution of locations for the three datasets are in  
69 Fig. 1, which shows 7,673 parameter sets based on 30-year records (left panel), 2,336 parameter sets based on 20-year records  
70 (middle panel), and 2,694 parameter sets based on 10-year records (right panel). All locations are unique, with no overlap in  
71 locations between the three datasets. As may be seen in Fig. 1, there is relatively sparse coverage for South America, Africa  
72 and southern Asia, while North America, Europe and Australia have relatively dense coverage. [The spatial density of all](#)  
73 [stations is shown in Fig. 2 so that density may be judged in places where overcrowding of points occurs in Fig. 1, and Table 1](#)  
74 enumerates the number of stations on each continent. Furthermore, a .kmz map layer is available on the Ag Data Commons  
75 website (link given in Sect. 4) that can be imported into Google Earth as an interactive map and allows the CLIGEN station  
76 nearest to an area of interest to be found.

77 As 30 years is traditionally the minimum record length needed to represent climate, the 30-year dataset may be used  
78 to characterize climate normals (Bojinski et al., 2014). The 20-year and 10-year datasets, reflecting the most recent monthly  
79 records available at each location, may be more representative of current climates in some cases considering the non-  
80 stationarity of current and projected climate conditions (IPCC 2013). In soil erosion modeling, a 20-year record has been  
81 suggested as the minimum length needed to represent rainfall erosivity (Wischmeier and Smith, 1978), which may be estimated  
82 using CLIGEN (Lobo et al., 2013). It should be noted that in non-stationary climates, CLIGEN inputs may be adjusted to  
83 represent departures from climate normals (Pruski and Nearing 2002; Zhang 2005; Vaghefi and Yu, 2016). For example,  
84 Zhang et al. (2013) determined how CLIGEN's precipitation intensity and skewness factors scale with monthly precipitation  
85 to correct for future changes in precipitation.

86 A list of parameters and their definitions that were determined for each input file are given in Table 2. These  
87 parameters are used to model statistical distributions that are randomly sampled by CLIGEN to derive daily outputs. Some  
88 parameters such as *TMAX AV* and *TMIN AV* ([refer to Table 2 for definitions](#)) are also typical climate benchmarks. Another  
89 climate benchmark, average monthly precipitation, may be determined by the following calculation from input parameters:

90

$$\text{avg. monthly precip.} = \frac{n_{\text{days}} * \text{MEAN } P * \{n * P_{\text{avg}} * \{P(W|D) / [1 - P(W|W) + P(W|D)]\}}}{n_{\text{days}}} \quad (1)$$

where  $n_{\text{days}}$  is the number of calendar days in the month being considered, and  $P_{\text{avg}}$  is the *MEAN P CLIGEN* parameter.

The various input parameters were derived from an assortment of data sources. In general, there were two main categories of sources: (1) ground-based precipitation networks, and (2) land-surface and meteorological models that assimilate remote sensing data and ground observations, and which reproduce historical time-series of variables of concern. The sources of data had various temporal resolutions. TheIn most cases, the data was used to make direct calculation of parameters, andbut for parameters where the available data was insufficient for direct calculation, parameter estimations were done. Each data source and the resulting parameters are discussed in detail in the following sections.

## 2.2 Precipitation Accumulation

The primary source of precipitation data is the Global Historical Climate Network-Daily (GHCN-Daily) maintained by NOAA (Menne et al., 2012). The locations shown in Fig. 1 correspond to those of selected stations from GHCN-Daily. These ground-based records enabled direct calculation of five parameters related to precipitation accumulation: *MEAN P*, *SKEW P*, *P(W/W)* and *P(W/D)* (see Table 4 for their definitions). The GHCN-Daily dataset undergoes rigorous quality control, both to check for consistency of formatting, and for the integrity of daily values. Values are removed that fail any test in a suite of quality tests which identify a variety of problems. Durre et al. (2010) outlined 19 of the quality tests in detail.

Short record lengths and missing data precluded a wide majority (~90%) of GHCN-Daily stations from being used to create CLIGEN input parameters. A substantial number of data gaps necessitated an assumption for the calculation of the five monthly parameters related to accumulation. To handle gaps, records were queried starting with the most recent year available and going backwards in each time-series until the number of months needed could be produced by replacing gaps with existing records from earlier in the time-series. Therefore, it was assumed that time-series do not need to be temporally continuous. This means that records were accepted which did not necessarily come from sequential months, but which had at least 30, 20 and 10 complete individual months for each calendar month, in order to derive the 30-year, 20-year and 10-year monthly statistics, respectively. As a result, record lengths were queried that were often longer than the number of years needed. Also, since representing recent data was a priority, 96% of stations included at least some data after the year 2000, and 81% included some data after the year 2010. Ranges of years queried for each station are given in an extensive table available on the Ag Data Commons website (link given in Sect. 4). The ranges are defined by the first and last year with at least one monthly record accepted for use. Ranges in excess of the 30, 20 and 10-year minimum record lengths are due to data gaps for respective datasets. The longest viable record length (of 30, 20 and 10 years) was used for each station, such that if a 30-year record was possible, 10 and 20-year records were not created. Therefore, no stations have multiple datasets created for them. This

Formatted: Font: Not Italic

123 treatment of data gaps complicates the validation of the determined climate benchmarks against other datasets with similar  
124 temporal ranges, and the effect of non-stationarity and long-term climate cycles should also be considered.

### 125 2.3 Precipitation Intensity

126 In soil erosion and runoff modeling, precipitation intensity is a critical factor (Pruski and Nearing, 2002; Nearing et  
127 al., 2005). The two parameters related to precipitation intensity,  $MX.5P$  and  $Time-Pk, TimePk$  (refer to Table 2 for definitions),  
128 require data with high frequency measurements such that hyetographs for a single precipitation event may be resolved. Since  
129 GHCN-Daily did not have adequate temporal resolution,  $MX.5P$  was estimated from the daily data using a temporal  
130 downscaling model, and  $Time-Pk, TimePk$  was assumed to follow known-average- $Time-Pk$  representative  $TimePk$  values for  
131 given Köppen-Geiger climate classifications. The development of these procedures is discussed in Fullhart et al. (2020a) and  
132 Fullhart et al. (2020b). High resolution data needed for these procedures came from the Automated Surface Observing System  
133 (ASOS) maintained by NOAA with stations distributed across the United States and its territories (Doesken et al., 2002).

134 In CLIGEN, the  $MX.5P$  input parameter is used to parameterize statistical distributions of normalized peak intensity.  
135 The definition of  $MX.5P$  is as follows:

$$136 \quad 137 \quad 138 \quad 139 \quad 140 \quad 141 \quad 142 \quad 143 \quad 144 \quad 145 \quad 146 \quad 147 \quad 148 \quad 149 \quad 150 \quad 151 \quad 152 \quad 153 \quad 154 \quad 155$$
$$MX.5P = \frac{1}{k} \frac{\sum_{i=1}^{n=k} maxI_{30i} + \dots + maxI_{30n}}{\sum_{i=1}^{n=k} maxI_{30i}, \dots, maxI_{30n}} \quad (2)$$

139 where  $k$  is the number of times (years) a record for a given month exists in the data set, and  $maxI_{30}$  is the maximum 30-minute  
140 intensity (mm hr<sup>-1</sup>) for each monthly record (Yu 2005). Since maximum 30-minute intensity is most accurately determined  
141 from data with as high frequency of measurement as possible, deriving values from data with lower resolutions results in  
142 underestimation bias, therefore necessitating use of the temporal downscaling model for  $MX.5P$ . The downscaling model took  
143 GHCN-Daily data to estimate the  $MX.5P$  value that would be expected if derived from the 1-min data. The downscaling model  
144 is a machine learning regression using Gradient Boosting trained with 609 ASOS stations (Fullhart et al., 2020b). The model  
145 requires 11 predictor variables shown in Table 3, which are statistics that may be determined from daily data and geographic  
146 information, some of which are already CLIGEN inputs. While  $MX.5P$  from 1-min resolution was estimated by the model, the  
147 predictor variable with the single most predictive power was  $MX.5P$  derived from daily data, which was calculated based on  
148 an assumption that intensity was constant for the duration of daily intervals (and was therefore grossly underestimated).  $MEAN$   
149  $P$  and  $SDEV P$  were also important predictors. The  $MX.5P$  values estimated by the model were found to have an RMSE of  
150 0.148 inches (3.76 mm) (Fullhart et al., 2020b).

152 The second intensity parameter,  $Time-Pk, TimePk$ , represents values at 12 equal intervals along the probability  
153 density cumulative distribution function of normalized time-to-peak intensity for events recorded at a given station ( $Time$   
154  $Pk, TimePk$  is the only input parameter that does not represent monthly values, though there are 12 values per station, each  
155 representing quantiles of the PDF/CDF). For a given  $Time-Pk, TimePk$  interval, the definition is as follows:

156

$$157 \text{ TimePk}(i) = \frac{N_{tp(i)}}{N_{tot}} \quad (3)$$

158

$$159 \text{ TimePk}(i) = \frac{N_{tp(i)}}{N_{tot}} \quad (3)$$

160

161 where  $\text{TimePk}(i)$  is the  $\text{TimePk}$  value at interval  $i$ ;  $tp$  is time-to-peak intensity normalized to the event duration;  
162  $N_{tp(i)}$  is the number of events where  $tp \leq i$ ; and  $N_{tot}$  is the total number of events. Interval,  $i$ , ranges between 1/12 and 12/12,  
163 and varies by increments of 1/12. (Yu 2005). Events were separated by  $\geq 6$  hours of no precipitation.

164

165 In Fullhart et al. (2020a), it was shown that using climate average  $\text{TimePk}$  values for the Köppen-Geiger  
166 climate classification of a given station resulted in  $<10\%$  error relative to true  $\text{TimePk}$  values, suggesting little variation  
167 of  $\text{TimePk}$  within climate classifications. In this previous study, a different weather station network was used—the  
168 U.S. Climate Reference Network (USCRN) at 5-min resolution (Diamond et al., 2013). For the new dataset of CLIGEN inputs,  
169 the analysis was repeated for the climate classifications represented by the 1-min ASOS network, though in some cases, climate  
170 classifications exclusive to the USCRN were used. Table A1 shows the assumed  $\text{TimePk}$  values for each climate  
171 classification. Of the 30 highest-order climate classifications, 19 were represented by ASOS and USCRN. The remaining 11  
172 classifications were assumed to be the averages of the other  $\text{TimePk}$  values within respective first-order groups (of  
173 which there are 5, where A is tropical, B is arid, C is temperate, D is cold, and E is polar). As such, the climate classification  
174 of each station was used to index the assumed  $\text{TimePk}$  values used in the CLIGEN input files. The climate classification  
175 of each station was determined based on the Köppen-Geiger climate map of Beck et al. (2018) representing the 1980-2016  
time period at  $0.083^\circ$  resolution.

## 176 2.4 Temperature

177 The 5 temperature-related parameters,  $TMAX AV$ ,  $TMIN AV$ ,  $SD TMAX$ ,  $SD TMIN$  and  $DEW PT$ ; (refer to Table 2 for  
178 definitions), have straight-forward calculations. However, the required data were only available for a subset of GHCN-Daily  
179 stations. To avoid limiting the analysis to this subset of stations, these data were instead derived from the model outputs of the  
180 ERA5 global meteorological/climate analysis (“ECMWF ReAnalysis”, with ERA5 being the fifth major global reanalysis).  
181 The ERA5 analysis was created by The European Centre for Medium-Range Weather Forecasts and the Copernicus Climate  
182 Change Service (Balsamo et al., 2018). ~~ERA5 provides climate and land surface outputs at various temporal resolutions,~~  
183 ~~including daily and monthly;~~ Hersbach et al., 2020). Google Earth Engine was used to download maximum and minimum  
184 temperatures at daily resolution, and average dew point temperatures at monthly resolution; ~~from a grid with  $0.25^\circ \times 0.25^\circ$~~   
185 ~~spatial resolution. Values obtained from the grid were unchanged, without any weighting based on proximity to neighbouring~~  
186 ~~cells or other forms of interpolation. The monthly dew point temperature was a convenient aggregation of data equivalent to~~  
187 ~~the  $DEW PT$  CLIGEN parameter, while daily resolution was needed for the remaining CLIGEN temperature parameters to~~  
188 ~~determine both the average and standard deviation of daily max/min temperatures.~~ Use of the ERA 5 model also allowed

Formatted: Font: Italic

189 continuous time-series to be obtained without gaps for the 30-year, 20-year and 10-year datasets (from 1990 through 2019,  
190 2000 through 2019, and 2010 through 2019, respectively).

## 191 2.5 Solar Radiation

192 Incoming shortwave radiation is represented in CLIGEN by the *SOL.RAD* and *SD RAD* parameters ~~which (refer to~~  
193 ~~Table 2 for definitions) that~~ require ~~that daily~~ solar radiation ~~is known~~ with units of langley/d where 1 langley = 41,840 J/m<sup>2</sup>.  
194 These parameters were calculated with relatively high frequency (3-hr) ~~measurement estimates~~ that captured daily and day-  
195 to-day variability of radiation. ~~This data came taken~~ from the Global Land Data Assimilation System model (GLDAS)  
196 produced by NASA ~~at averaged 3-hr intervals~~ (Fang et al., 2009) ~~at 0.25° x 0.25° resolution~~. The outputs of the reprocessed  
197 GLDAS 2.0 and GLDAS 2.1 versions were used and ~~download downloaded~~ from Google Earth Engine. ~~(again, no weighting~~  
198 ~~of values was done based on proximity to neighbouring cells)~~. The most recent data available was used to create continuous  
199 time-series with temporal ranges being the same as those for the temperature parameters. For an individual day, incoming solar  
200 radiation was modeled by fitting a gaussian curve through the 3-hr time-averaged data points. Doing this avoided  
201 underestimation caused by time-averaging, which would have occurred by considering the 3-hr datapoints alone. Also, if the  
202 3-hr intervals did not coincide with the time of peak intensity, comparison to ground observations from Ameriflux data  
203 (discussed more later) showed that the gaussian curve tended to better approximate peak radiation than the greatest 3-hr  
204 datapoint.

205 A number of stations that existed on coasts or on small islands, particularly in the Pacific Ocean, did not have solar  
206 radiation data coverage for their locations because the GLDAS product covers only locations beyond a certain coastal  
207 proximity. In total, 390 stations had this problem. For these stations, data from the nearest station with existing data was used.  
208 300 of the stations with missing data were within 100 km of a station with data. Some proximities, however, were much further,  
209 with islands in the south Pacific being examples. Similarly, some locations in the existing U.S. CLIGEN input dataset used for  
210 validation ~~(created by Srivastava et al., (2019))~~ did not have observed solar radiation, and their parameter values were taken  
211 from the nearest station with available data, which in some cases were at considerable distances, potentially leading to poor  
212 validation in Sect. 3.

213 To ensure locations are matched for validation, a separate validation from that of Sect. 3 was done for solar radiation  
214 parameters. In this, GLDAS output was compared to 10 ground-based Ameriflux stations that monitor ecosystem fluxes  
215 including solar radiation (Hargrove et al., 2003). The Ameriflux network has stations distributed across the North and South  
216 American continents, and the 10 stations were selected from a range of latitudes and climates as a representation of global  
217 variability. From these stations, a single year was selected that had the fewest data gaps. Comparison to corresponding GLDAS  
218 outputs showed reasonable agreement with an RMSE of 36.6 langley/d and with GLDAS being overestimated by <1% for  
219 monthly values of *SOL.RAD*. Error was more evident for *SD RAD* suggesting that GLDAS was not optimum for capturing the  
220 day-to-day variability of radiation. The RMSE for *SD RAD* was 38.6 langley/d with GLDAS being underestimated by 24.1%.

## 221 2.5 Wind

222 — Very few applications of CLIGEN have used wind data in the past, perhaps the only one being the blowing  
223 snow component in WEPP (Nicks et al., 1989). CLIGEN inputs require high-frequency measurement of wind speed (m/s) and  
224 azimuthal wind direction. This includes mean, standard deviation, and skewness of daily wind speed on a monthly basis; and  
225 determinations of the average daily percentage of time with wind directions coming from the 4 cardinal directions, 4  
226 intercardinal directions, and the 8 sub-divisions of these (e.g. NNE, ENE), on a monthly basis. However, wind data was not  
227 obtainable for the locations corresponding to the GHCN-Daily stations with the level of detail needed for creating CLIGEN  
228 input files. The solution to this was to use the “International Conversion Programs” tool (availability given in Sect. 4), which  
229 takes the known daily precipitation accumulation and temperature parameters from an international station of interest and finds  
230 the existing station in the U.S. CLIGEN dataset with the most similar climate, allowing its wind parameters to be used (and  
231 other remaining parameters, if needed). Information regarding the locations from where wind parameters were taken from are  
232 given at the bottom of each input file.

## 233 3 Validation

234 — Each parameter except for the wind parameters were compared to an existing dataset for the U.S. and its  
235 territories created in 2015 using NOAA NCDC DSI-3260 data at 15-min resolution and consisting of 40-year records for 2,648  
236 stations (Srivastava et al., 2019). This limited the validation to only stations for the U.S., and from those, only the new stations  
237 within 10 km of an existing CLIGEN station were accepted. This resulted in the validation of 61 stations for the 30-year  
238 dataset, 53 stations for the 20-year dataset, and 204 stations for the 10-year dataset. For each of the validated parameters,  
239 RMSE, percent bias, and percent error were determined, where it was assumed that values from the existing U.S. dataset were  
240 the true values (performance metric definitions are given in Table A2). A summary of the validation is seen in Table 4.  
241 Inconsistencies between the two datasets were attributed to: differences of data sources, differences in temporal resolution of  
242 data used, differences in record lengths, and whether data was interpolated or taken from nearby stations.

243 Overall, reasonable agreement was found, with PERROR being below 20% for the majority of parameters. As  
244 expected, record length is a factor in the comparison to the 40-year U.S. dataset. Percent error increased slightly on average  
245 (~5%) with decreasing record length, going from the 30-year to 10-year dataset. Though a small increase, this difference likely  
246 reflected the potential for capturing short-term climate dynamics by the 20-year and 10-year datasets. For the 5 parameters  
247 related to daily accumulation, the parameter with the highest error was *SKEW P*, with error up to 30%. The sign of PBIAS for  
248 *SKEW P* was consistently positive suggesting that the GHCN-Daily data showed less skewness towards high daily  
249 accumulation.

250 Error was also considerable for the two parameters related to precipitation intensity, *MX.5P* and *Time-PkTimePk*. The  
251 discrepancies were due to multiple issues including the fact that the DSI-3260 dataset uses 15-min resolution compared to the  
252 1-min resolution that the *MX.5P* downscaling model and *Time-PkTimePk* distributions were based on. As mentioned, the

253 downscaling model was previously shown to produce an average error of 0.148 inches (3.76 mm) (Fullhart et al., 2020b). In  
254 the comparison to the DSI-3260 dataset, downscaled *MX.5P* values resulted in discrepancy of up to 37% error for *MX.5P*.  
255 Interval values for *Time-PkTimePk* distributions were generally smaller in magnitude and approached unity later in the  
256 distribution, meaning that the peak intensity of storms generally happened later in their duration than in the DSI-3260 data.  
257 This may be expected given the relatively coarse 15-min resolution of DSI-3260, and particularly when considering shorter  
258 storms, such as convective storms, the apparent peak intensity may have considerable uncertainty.

259 Temperature parameters were generally in agreement with no consistent estimation bias, except for *DEW PT*, which  
260 was slightly underestimated on average by up to 6%. Errors for *SOL.RAD* were up to 6%, with a slight overestimation bias of  
261 up to 3%. While *SOL.RAD* was in good agreement, *SD SOL* indicated up to 193% more day-to-day variability of solar radiation.  
262 The GLDAS data for solar radiation generally agreed better with the variability of the Ameriflux network that was discussed  
263 in Sect. 2.5, with GLDAS showing 24% less variability than Ameriflux. Given the reasonable agreement between GLDAS  
264 and Ameriflux, and good agreement of *SOL.RAD* with the DSI-3260 data, the substantial underestimation bias of *SD SOL* may  
265 be the result of errors in the existing U.S. inputs.

266 While the U.S. represents a wide range of climate types, limitation of the validation to only the U.S. is a hinderance  
267 to quality assurance of the new dataset. However, each of the source data have their own quality assurances prior to going to  
268 product. Particularly for the ERA5 and GLDAS global products, biases are documented and are known to happen on regional  
269 and continental spatial scales, and may relate to extremes in temperature, moisture, geographic location, etc. (Zhou et al., 2013;  
270 Ji et al., 2015; Urraca et al., 2018; Wang et al., 2019). Therefore, the uncertainty of each CLIGEN parameter also depends on  
271 the particular source data.

#### 272 **4 Data Availability**

273 The new international CLIGEN [input](#) dataset is available at the National Agriculture Library [websiteOnline](#)  
274 [Repository](#)—Ag Data Commons—at [https://data.nal.usda.gov/dataset/international-climate-benchmarks-and-input-](https://data.nal.usda.gov/dataset/international-climate-benchmarks-and-input-parameters-stochastic-weather-generator-cligen)  
275 [parameters-stochastic-weather-generator-cligen](https://data.nal.usda.gov/dataset/international-climate-benchmarks-and-input-parameters-stochastic-weather-generator-cligen) (Fullhart et al., 2020c; DOI: <https://doi.org/10.15482/USDA.ADC/1518706>)  
276 and is separated into three datasets according to 30-year, 20-year and 10-year record lengths. To run the CLIGEN inputs,  
277 CLIGEN may be downloaded at [https://www.ars.usda.gov/midwest-area/west-lafayette-in/national-soil-erosion-](https://www.ars.usda.gov/midwest-area/west-lafayette-in/national-soil-erosion-research/docs/wepp/cligen/)  
278 [research/docs/wepp/cligen/](https://www.ars.usda.gov/midwest-area/west-lafayette-in/national-soil-erosion-research/docs/wepp/cligen/). Additional resources and materials are available at this website including the “International  
279 Conversion Programs” tool. The international CLIGEN dataset will also be added to the web interface for running the hillslope-  
280 scale erosion and runoff model, RHEM, available at <https://apps.tucson.ars.ag.gov/rhem/>. The station of interest will be  
281 selectable in the input parameters panel under “Climate Station” and under “International”.

282 **5 Conclusions**

283 Validation of CLIGEN inputs in the new international dataset showed reasonable agreement with parameter values  
284 for existing U.S. CLIGEN inputs. The 30-year, 20-year and 10-year datasets are generally in close agreement, ~~although some~~  
285 ~~uncertainty exists and in some cases, the methods used to create this dataset may offer an improvement over existing CLIGEN~~  
286 ~~input files. However, issues arise~~ due to the assumptions ~~that were~~ taken for addressing ~~pervasive~~ data gaps ~~and the degree to~~  
287 ~~which short-term climate dynamics have a role in influencing~~ NOAA-GHCN records. Validation of the climate benchmarks:  
288 ~~In some cases, use of~~ by comparison to other records is complicated by use of discontinuous time-series, and uncertainty is  
289 higher ~~resolution climate data for parameterization may offer an improvement over existing CLIGEN input files in places with~~  
290 ~~non-stationary climates or long-term cycles.~~

**Formatted:** Font: Not Bold

291 The new dataset of CLIGEN inputs allows the CLIGEN weather generator to be more readily applied to its  
292 various applications. The input files also serve to represent climate benchmarks for a selection of variables that are generally  
293 unobtainable from a single source. The coverage of stations is particularly dense in Europe, Australia, and North America, and  
294 offers the potential to improve the spatial analysis of processes in different fields that require climate records. For a number of  
295 CLIGEN's applications, the production of climate data is a secondary concern, but is often a labor-intensive task. The use of  
296 this dataset may allow researchers to put more effort and resources towards their primary study or area of focus without needing  
297 to address the production of climate inputs.

**Formatted:** Font: Bold

298

299

300

301

302

303

304

305

306

307

308

309

310

311

313 Table A1: *TimePk* distribution interval values for global Köppen-Geiger climate classifications.

Interval	1/12	2/12	3/12	4/12	5/12	6/12	7/12	8/12	9/12	10/12	11/12	12/12
Af	0.22	0.30	0.36	0.44	0.50	0.58	0.63	0.70	0.77	0.83	0.90	1.00
Am	0.25	0.36	0.43	0.51	0.58	0.66	0.73	0.79	0.84	0.90	0.94	1.00
Aw	0.27	0.39	0.48	0.56	0.63	0.71	0.77	0.81	0.86	0.90	0.95	1.00
Bwh	0.16	0.26	0.35	0.43	0.52	0.61	0.69	0.76	0.84	0.90	0.95	1.00
Bwk	0.15	0.26	0.36	0.45	0.53	0.62	0.69	0.76	0.83	0.89	0.96	1.00
BSh	0.16	0.27	0.36	0.46	0.54	0.64	0.71	0.77	0.83	0.89	0.95	1.00
BSk	0.12	0.22	0.32	0.40	0.48	0.57	0.65	0.74	0.82	0.89	0.96	1.00
Csa	0.07	0.17	0.26	0.36	0.45	0.54	0.62	0.70	0.78	0.86	0.94	1.00
Csb	0.07	0.17	0.25	0.34	0.43	0.52	0.61	0.69	0.77	0.85	0.94	1.00
Csc	0.07	0.17	0.26	0.35	0.44	0.53	0.61	0.70	0.78	0.86	0.94	1.00
Cwa	0.10	0.20	0.29	0.38	0.46	0.55	0.64	0.72	0.80	0.87	0.94	1.00
Cwb	0.10	0.20	0.29	0.38	0.46	0.55	0.64	0.72	0.80	0.87	0.94	1.00
Cwc	0.10	0.20	0.29	0.38	0.46	0.55	0.64	0.72	0.80	0.87	0.94	1.00
Cfa	0.20	0.31	0.40	0.48	0.56	0.65	0.72	0.78	0.84	0.90	0.96	1.00
Cfb	0.07	0.15	0.24	0.32	0.40	0.51	0.60	0.69	0.78	0.86	0.94	1.00
Cfc	0.13	0.23	0.32	0.40	0.48	0.58	0.66	0.74	0.81	0.88	0.95	1.00
Dsa	0.17	0.27	0.37	0.45	0.53	0.61	0.68	0.75	0.82	0.88	0.94	1.00
Dsb	0.08	0.17	0.25	0.34	0.42	0.52	0.60	0.69	0.78	0.85	0.93	1.00
Dsc	0.27	0.38	0.48	0.56	0.64	0.70	0.76	0.81	0.87	0.91	0.95	1.00
Dsd	0.17	0.27	0.37	0.45	0.53	0.61	0.68	0.75	0.82	0.88	0.94	1.00
Dwa	0.16	0.29	0.40	0.49	0.58	0.67	0.74	0.80	0.86	0.91	0.96	1.00
Dwb	0.16	0.27	0.37	0.46	0.55	0.63	0.70	0.78	0.83	0.90	0.95	1.00
Dwc	0.16	0.28	0.38	0.48	0.56	0.65	0.72	0.79	0.85	0.91	0.96	1.00
Dwd	0.16	0.28	0.38	0.48	0.56	0.65	0.72	0.79	0.85	0.91	0.96	1.00
Dfa	0.15	0.26	0.35	0.45	0.53	0.62	0.70	0.77	0.84	0.90	0.96	1.00
Dfb	0.13	0.23	0.32	0.41	0.50	0.59	0.67	0.75	0.83	0.89	0.95	1.00
Dfc	0.25	0.36	0.45	0.53	0.60	0.67	0.72	0.79	0.85	0.90	0.95	1.00
Dfd	0.18	0.28	0.37	0.46	0.54	0.63	0.70	0.77	0.84	0.90	0.95	1.00
ET	0.28	0.41	0.51	0.58	0.66	0.74	0.78	0.82	0.87	0.91	0.94	1.00

Formatted: Position: Horizontal: Left, Relative to: Column  
Vertical: In line, Relative to: Margin, Horizontal: 0", Wrap  
Around

Formatted Table

Formatted: Position: Horizontal: Left, Relative to: Column  
Vertical: In line, Relative to: Margin, Horizontal: 0", Wrap  
Around

Formatted: Position: Horizontal: Left, Relative to: Column  
Vertical: In line, Relative to: Margin, Horizontal: 0", Wrap  
Around

Formatted: Position: Horizontal: Left, Relative to: Column  
Vertical: In line, Relative to: Margin, Horizontal: 0", Wrap  
Around

Formatted: Position: Horizontal: Left, Relative to: Column  
Vertical: In line, Relative to: Margin, Horizontal: 0", Wrap  
Around

Formatted: Position: Horizontal: Left, Relative to: Column  
Vertical: In line, Relative to: Margin, Horizontal: 0", Wrap  
Around

Formatted

EF	0.28	0.41	0.51	0.58	0.66	0.74	0.78	0.82	0.87	0.91	0.94	1.00
----	------	------	------	------	------	------	------	------	------	------	------	------

**Formatted:** Position: Horizontal: Left, Relative to: Column  
Vertical: In line, Relative to: Margin, Horizontal: 0", Wrap  
Around

**Formatted:** Font: 9 pt, Bold

**Table A2: Statistical measures of performance. Observed (O) and predicted (P) values are compared by each metric.**

Performance metric	Abbreviation	Equation
Root mean square error	RMSE	$\sqrt{\frac{1}{n} \sum (O - P)^2}$
Percent Bias	PBIAS	$\left[ \frac{\sum O - P}{\sum O} \right] \left[ \frac{\sum (O - P)}{\sum O} \right] \times 100$
Percent Error	PERROR	$\frac{1}{n} \left[ \sum \frac{abs(O - P)}{O} \right] \times 100$

### Author Contributions

AF calculated input parameters, GA provided expertise on data management and integration with the RHEM web interface, MN and MW gave their expertise on project guidance, and all authors were involved in writing the manuscript.

### Competing Interests

The authors declare that they have no conflict of interest.

### Acknowledgements

The authors wish to express their appreciation for everyone involved in creating and maintaining the various climate networks that were used. Funding for this project was given through the Agricultural Research Service Headquarters Grant, and the Southwest Watershed Research Center.

### References

Baffaut, C., Nearing, M. A., and Nicks, A. D.: Impact of CLIGEN parameters on WEPP-predicted average annual soil loss, Trans. ASAE 39, 447-457, 1996.

333  
334 Balsamo, G., Dutra, E., Albergel, C., Munier, S., Calvet, J. C., Munoz-Sabater, J., and de Rosnay, P.: ERA-5 and ERA-Interim  
335 driven ISBA land surface model simulations: which one performs better?, *Hydr. and Earth Sys. Sci.*, 22, 3515-3532,  
336 <https://doi.org/10.5194/hess-22-3515-2018>, 2018.

337  
338 Beck, H. E., Zimmermann, N. E., McVicar, T. R., Vergopolan, N., Berg, A., and Wood, E. F.: Present and future Köppen-  
339 Geiger climate classification maps at 1-km resolution, *Sci. Data*, 5, 180214, <https://doi.org/10.1038/sdata.2018.214>, 2018.

340  
341 Bojinski, S., Verstraete, M., Peterson, T. C., Richter, C., Simmons, A., and Zemp, M.: The concept of essential climate  
342 variables in support of climate research, applications, and policy, *Bull. Am. Meteor. Soc.*, 95, 1431-1443,  
343 <https://doi.org/10.1175/BAMS-D-13-00047.1>, 2014.

344  
345 Diamond, H. J., Karl, T. R., Palecki, M. A., Baker, C. B., Bell, J. E., Leeper, R. D., ... and Goodge, G.: US Climate Reference  
346 Network after one decade of operations: Status and assessment, *Bull. Am. Meteor. Soc.*, 94, 485-498,  
347 <https://doi.org/10.1175/BAMS-D-12-00170.1>, 2013.

348  
349 Doesken, N. J., McKee, T. B., and Davey, C.: Climate data continuity—what have we learned from the ASOS automated surface  
350 observing system, *Proc. 13th Conf. App. Clim.*, 2002.

351  
352 Durre, I., Menne, M. J., Gleason, B. E., Houston, T. G., and Vose, R. S.: Comprehensive automated quality assurance of daily  
353 surface observations, *J. App. Meteor. Clim.*, 49, 1615-1633, <https://doi.org/10.1175/2010JAMC2375.1>, 2010.

354  
355 Fang, H., Beaudoin, H. K., Teng, W. L., and Vollmer, B. E.: Global Land Data Assimilation System (GLDAS) products,  
356 services and application from NASA hydrology data and information services center (HDISC), *ASPRS Ann. Conf.*, 2009.

357  
358 Fullhart, A.T., Nearing, M.A., and Weltz, M.A.: Temporally downscaling precipitation intensity factors for Köppen climate  
359 regions in the U.S., *J. So. Wat. Con.*, in press, 2020a

360  
361 Fullhart, A.T., Nearing, M. A., McGehee, R. P., and Weltz, M.A.: Temporally downscaling a precipitation intensity factor for  
362 soil erosion modeling using the NOAA-ASOS weather station network, *Catena*, ~~in—press~~194, 104709,  
363 <https://doi.org/10.1016/j.catena.2020.104709>, 2020b.

364  
365 Fullhart, A. T., Nearing, M. A., Armendariz, G., Weltz, M. A.: International climate benchmarks and input parameters for a  
366 stochastic weather generator, CLIGEN (dataset), *Ag Data Commons*, <https://doi.org/10.15482/USDA.ADC/1518706>, 2020c.

Formatted: English (United States)

367  
368 Hargrove, W. W., Hoffman, F. M., and Law, B. E.: New analysis reveals representativeness of the AmeriFlux network, *Trans.*  
369 *Am. Geophys. Union*, **84**, 529-535, <https://doi.org/10.1029/2003EO480001>, 2003.

370  
371 [Hersbach, H., Bell, B., Berrisford, P., Hirahara, S., Horányi, A., Muñoz-Sabater, J., ... and Simmons, A. The ERA5 global](#)  
372 [reanalysis, \*J. Roy. Meteor. Soc.\*, \*\*730\*\*, 1999-2049. <https://doi.org/10.1002/qj.3803>, 2020.](#)

373  
374 Hollmann, R., Merchant, C. J., Saunders, R., Downy, C., Buchwitz, M., Cazenave, A., ... and Holzer-Popp, T.: The ESA  
375 climate change initiative: Satellite data records for essential climate variables, *Bull. Am. Meteor. Soc.*, **94**, 1541-1552,  
376 <https://doi.org/10.1175/BAMS-D-11-00254.1>, 2013.

377  
378 IPCC: Climate Change 2013: The Physical Science Basis, Summary for Policymakers. Contribution of Working Group I to  
379 the Fifth Assessment Report of the Intergovernmental Panel on Climate Change, Cambridge University Press, Cambridge,  
380 United Kingdom and New York, NY, USA, 2013.

381  
382 Ji, L., Senay, G. B., and Verdin, J. P.: Evaluation of the Global Land Data Assimilation System (GLDAS) air temperature data  
383 products, *J. Hydrometeorol.*, **16**, 2463-2480, <https://doi.org/10.1175/JHM-D-14-0230.1>, 2015.

384  
385 Kinnell, P. I. A.: CLIGEN as a weather generator for RUSLE2. *Catena*, **172**, 877-880,  
386 <https://doi.org/10.1016/j.catena.2018.09.016>, 2019.

387  
388 Lobo, G. P., Frankenberger, J. R., Flanagan, D. C., and Bonilla, C. A.: Evaluation and improvement of the CLIGEN model for  
389 storm and rainfall erosivity generation in Central Chile, *Catena*, **127**, 206-213, <https://doi.org/10.1016/j.catena.2015.01.002>,  
390 2015.

391  
392 Mehan, S., Guo, T., Gitau, M. W., and Flanagan, D. C.: Comparative study of different stochastic weather generators for long-  
393 term climate data simulation, *Climate*, **5**, 26, <https://doi.org/10.3390/cli5020026>, 2017.

394  
395 Menne, M. J., Durre, I., Vose, R. S., Gleason, B. E., and Houston, T. G.: An overview of the Global Historical Climatology  
396 Network-Daily Database, *J. of Atmos. Oc. Tech.*, **29**, 897-910, <https://doi.org/10.1175/JTECH-D-11-00103.1>, 2012.

397  
398 Nearing, M. A., Jetten, V., Baffaut, C., Cerdan, O., Couturier, A., Hernandez, M., ... and Souchère, V.: Modeling response of  
399 soil erosion and runoff to changes in precipitation and cover, *Catena*, **61**, 131-154,  
400 <https://doi.org/10.1016/j.catena.2005.03.007>, 2005.

401  
402 Nicks, A. D., Lane, L. J., Nearing, M. A., and Stone, J. J.: WEPP hillslope profile erosion model user summary. USDA—  
403 Water Erosion Prediction Project: Hillslope Profile Model Documentation, NSERL Report, 2, 1989.  
404  
405 Panagos, P., Borrelli, P., Meusburger, K., Yu, B., Klik, A., Lim, K. J., and Sadeghi, S. H.: Global rainfall erosivity assessment  
406 based on high-temporal resolution rainfall records, *Sci. Reports*, 7, 1-12, <https://doi.org/10.1038/s41598-017-04282-8>, 2017.  
407  
408 Pruski, F. F. and Nearing, M. A.: Runoff and soil-loss responses to changes in precipitation: A computer simulation study, *J.*  
409 *So. Wat. Con.*, 57, 7-16, 2002.  
410  
411 Rodell, M., Houser, P. R., Jambor, U. E. A., Gottschalck, J., Mitchell, K., Meng, C. J., ... and Entin, J. K.: The global land data  
412 assimilation system, *Bull. Am. Meteor. Soc.*, 85, 381-394, <https://doi.org/10.1175/BAMS-85-3-381>, 2004.  
413  
414 Srivastava, A., Flanagan, D. C., Frankenberger, J. R., and Engel, B. A.: Updated climate database and impacts on WEPP model  
415 predictions. *J. So. Wat. Con.*, 74, 334-349, <https://doi.org/10.2489/jswc.74.4.334>, 2019.  
416  
417 Urraca, R., Huld, T., Gracia-Amillo, A., Martínez-de-Pison, F. J., Kaspar, F., and Sanz-Garcia, A.: Evaluation of global  
418 horizontal irradiance estimates from ERA5 and COSMO-REA6 reanalyses using ground and satellite-based data, *Sol.*  
419 *En.*, 164, 339-354, <https://doi.org/10.1016/j.solener.2018.02.059>, 2018.  
420  
421 Vaghefi, P. and Yu, B.: Use of CLIGEN to simulate decreasing precipitation trends in the southwest of Western  
422 Australia, *Trans. ASABE*, 59, 49-61, <https://doi.org/10.13031/trans.59.10829>, 2016.  
423  
424 Vaghefi, P. and Yu, B.: Validation of CLIGEN parameter adjustment methods for Southeastern Australia and Southwestern  
425 Western Australia, *J. Hydrometeorol.*, 18, 2011-2028, <https://doi.org/10.1175/JHM-D-16-0237.1>, 2017.  
426  
427 Wang, C., Graham, R. M., Wang, K., Gerland, S., and Granskog, M. A.: Comparison of ERA5 and ERA-Interim near-surface  
428 air temperature, snowfall and precipitation over Arctic sea ice: Effects on sea ice thermodynamics and evolution,  
429 *Cryosphere*, 13, 1661-1679, <https://doi.org/10.5194/tc-13-1661-2019>, 2019.  
430  
431 Wischmeier, W. H. and Smith, D. D.: Predicting rainfall erosion losses: A guide to conservation planning (No. 537),  
432 Department of Agriculture, Science and Education Administration, 1978.  
433

434 Yu, B.: Adjustment of CLIGEN parameters to generate precipitation change scenarios in southeastern Australia, *Catena*, 61,  
435 196-209, <https://doi.org/10.1016/j.catena.2005.03.004>, 2005.

436

437 Zhang, X. C.: Generating correlative storm variables for CLIGEN using a distribution-free approach, *Trans. ASAE*, 48, 567-  
438 575, <https://doi.org/10.13031/2013.18331>, 2005.

439

440 Zhang, Y., Liu, B., Wang, Z., and Zhu, Q.: Evaluation of CLIGEN for storm generation on the semiarid Loess Plateau in  
441 China, *Catena*, 73, 1-9, <https://doi.org/10.1016/j.catena.2007.08.001>, 2008.

442

443 Zhang, X. C. J.: Adjusting skewness and maximum 0.5 hour intensity in CLIGEN to improve extreme event and sub-daily  
444 intensity generation for assessing climate change impacts, *Trans. ASABE*, 56, 1703-1713,  
445 <https://doi.org/10.13031/trans.56.10004>, 2013.

446

447 Zhou, X., Zhang, Y., Yang, Y., and Han, S.: Evaluation of anomalies in GLDAS-1996 dataset, *Wat. Sci. Tech.*, 67, 1718-1727,  
448 <https://doi.org/10.2166/wst.2013.043>, 2013.

449

450

451

452

453

454

455

456

457

458

459

460

461

462

463

464

Formatted: Font: 10 pt

465 **Table 1: Station counts for continent/region and each of the 30-year, 20-year and 10-year datasets. Oceania is the region represented**  
466 **by south Pacific islands and extending north to Hawaii.**

Station Counts	North America	South America	Europe	Africa	Asia	Australia	Oceania	Antarctica	Total
30-year	1,860	170	2,089	9	118	3,423	4	0	7,673
20-year	996	112	374	7	11	834	2	0	2,336
10-year	1,332	8	413	6	52	864	19	0	2,694
Total	4,188	290	2,876	22	181	5,121	25	0	12,703

467  
468  
469  
470  
471  
472  
473  
474  
475  
476  
477  
478  
479  
480  
481  
482  
483  
484  
485  
486  
487





510 **Table 4: Summary of the validation of parameters to the 2015 U.S. CLIGEN dataset. created by Srivastava et al. (2019).**

	30-year dataset			20-year dataset			10-year dataset		
	RMSE	PBIAS	PERROR	RMSE	PBIAS	PERROR	RMSE	PBIAS	PERROR
MEAN P	0.08	-12.16	19.95	0.07	1.18	14.76	0.08	1.13	21.17
S DEV P	0.10	-2.70	15.06	0.10	2.92	16.45	0.14	1.08	24.17
SKEW P	1.35	8.05	20.15	1.11	7.13	22.93	1.29	15.98	30.36
P(W/W)	0.07	2.48	10.35	0.06	-1.35	10.3332	0.09	-3.6870	16.6866
P(W/D)									
		-			-			-	
	0.05	11.7980	19.20	0.0406	6.3090	14.2025	0.0506	12.8314	23.0729
TMAX AV	3.49	3.18	3.97	5.43	-0.41	6.77	3.75	0.66	4.28
TMIN AV	4.56	-8.55	15.79	6.23	-10.62	13.67	4.76	-7.93	11.33
SD TMAX	1.07	7.93	9.01	1.37	11.56	13.28	1.30	9.62	11.85
SD TMIN	1.53	6.87	11.34	1.22	7.80	13.01	1.04	4.45	10.98
SOL.RAD	22.55	-1.08	5.85	29.10	-2.90	5.87	26.91	-2.75	5.65
SD SOL	51.85	-135.54	146.33	68.09	-193.42	202.42	63.04	-173.21	181.51
MX .5 P	0.23	24.9391	29.9591	0.27	28.3736	31.9390	0.31	33.2625	37.3028
DEW PT	3.66	5.62	8.94	2.00	0.45	5.14	2.56	0.48	5.85
Time Pk	0.33	30.92	33.43	0.30	28.33	31.08	0.30	28.77	31.66

Formatted: Position: Horizontal: Left, Relative to: Column  
Vertical: In line, Relative to: Margin, Horizontal: 0", Wrap  
Around

Formatted Table

Formatted: Font: 8 pt

Formatted: Font: 8 pt

Formatted: Font: 8 pt

Formatted: Position: Horizontal: Left, Relative to: Column  
Vertical: In line, Relative to: Margin, Horizontal: 0", Wrap  
Around

Formatted: Position: Horizontal: Left, Relative to: Column  
Vertical: In line, Relative to: Margin, Horizontal: 0", Wrap  
Around

Formatted: Position: Horizontal: Left, Relative to: Column  
Vertical: In line, Relative to: Margin, Horizontal: 0", Wrap  
Around

Formatted: Position: Horizontal: Left, Relative to: Column  
Vertical: In line, Relative to: Margin, Horizontal: 0", Wrap  
Around

Formatted: Position: Horizontal: Left, Relative to: Column  
Vertical: In line, Relative to: Margin, Horizontal: 0", Wrap  
Around

Formatted: Position: Horizontal: Left, Relative to: Column  
Vertical: In line, Relative to: Margin, Horizontal: 0", Wrap  
Around

Formatted: Position: Horizontal: Left, Relative to: Column  
Vertical: In line, Relative to: Margin, Horizontal: 0", Wrap  
Around

Formatted: Position: Horizontal: Left, Relative to: Column  
Vertical: In line, Relative to: Margin, Horizontal: 0", Wrap  
Around

Formatted: Position: Horizontal: Left, Relative to: Column  
Vertical: In line, Relative to: Margin, Horizontal: 0", Wrap  
Around

Formatted: Position: Horizontal: Left, Relative to: Column  
Vertical: In line, Relative to: Margin, Horizontal: 0", Wrap  
Around

Formatted: Position: Horizontal: Left, Relative to: Column  
Vertical: In line, Relative to: Margin, Horizontal: 0", Wrap  
Around

Formatted: Position: Horizontal: Left, Relative to: Column  
Vertical: In line, Relative to: Margin, Horizontal: 0", Wrap  
Around

Formatted: Position: Horizontal: Left, Relative to: Column  
Vertical: In line, Relative to: Margin, Horizontal: 0", Wrap  
Around

Formatted: Position: Horizontal: Left, Relative to: Column  
Vertical: In line, Relative to: Margin, Horizontal: 0", Wrap  
Around

Formatted: Position: Horizontal: Left, Relative to: Column  
Vertical: In line, Relative to: Margin, Horizontal: 0", Wrap  
Around

Formatted

511

512

513

514

515

516

517

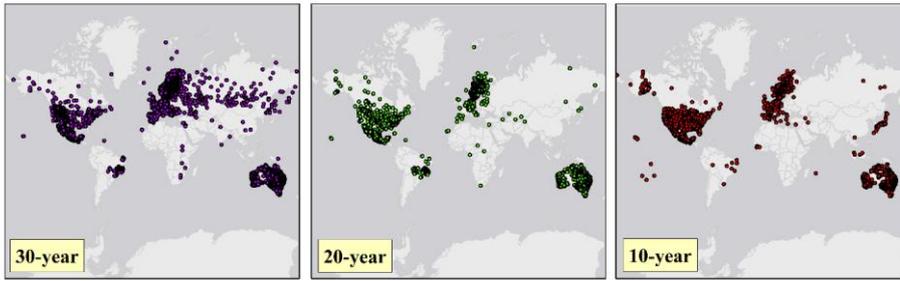
518

519

520

521

522



523

524 **Figure 1: Coverage of the three international CLIGEN input datasets according to the record length used to produce the monthly**  
525 **input parameters. The locations correspond to those of the GHCN-Daily stations accepted for use.**

526

527

528

529

530

531

532

533

534

535

536

537

538

539

540

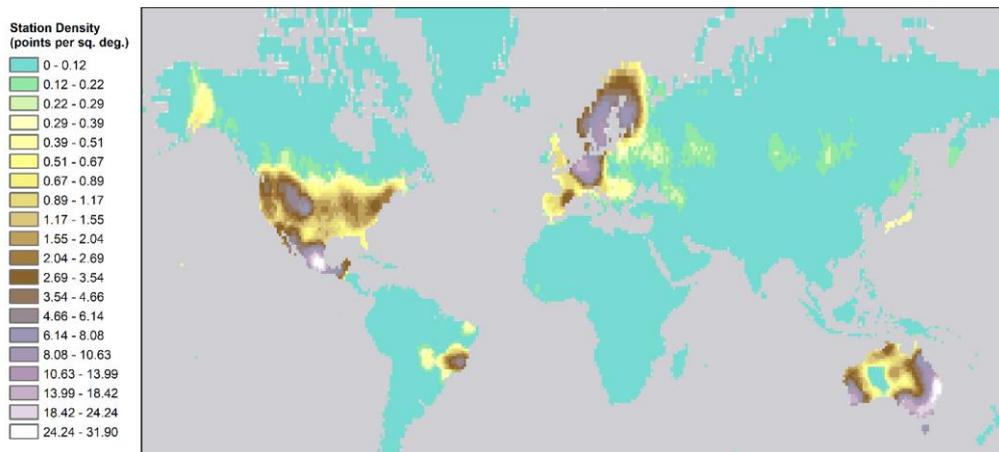
541

542

543

544

545



546

547 **Figure 2: Station density map representing all stations combined. The cell size is defined by lat./long. degree lines ( $1^\circ \times 1^\circ$ ). Densities**

548 **are calculated inside of circular neighbourhoods with radii of three degrees from the center of each cell.**

549