

1 **AIMERG: a new Asian precipitation dataset (0.1°/half-hourly, 2000-2015) by calibrating GPM**  
2 **IMERG at daily scale using APHRODITE**

3 Ziqiang Ma<sup>1</sup>, Jintao Xu<sup>2</sup>, Siyu Zhu<sup>1</sup>, Jun Yang<sup>3</sup>, Guoqiang Tang<sup>4,5</sup>, Yuanjian Yang<sup>6</sup>, Zhou Shi<sup>2</sup>, Yang  
4 Hong<sup>1,7</sup>

5 <sup>1</sup>*Institute of Remote Sensing and Geographical Information Systems, School of Earth and Space Sciences, Peking*  
6 *University, Beijing, 100871, China*

7 <sup>2</sup>*Institute of Agricultural Remote Sensing and Information Technology Application, College of Environmental and*  
8 *Resource Sciences, Zhejiang University, Hangzhou, 310058, China*

9 <sup>3</sup>*National Satellite Meteorological Centre, China Meteorological Administration, Beijing, 100081, China*

10 <sup>4</sup>*University of Saskatchewan Coldwater Lab, Canmore, Alberta, Canada, T1W 3G1*

11 <sup>5</sup>*Centre for Hydrology, University of Saskatchewan, Saskatoon, Saskatchewan, Canada, S7N 1K2*

12 <sup>6</sup>*School of Atmospheric Physics, Nanjing University of Information Science and Technology, Nanjing 210044, China*

13 <sup>7</sup>*School of Civil Engineering and Environmental Science, University of Oklahoma, Norman, OK, 73019, United States*

14 Correspondences: Ziqiang Ma ([ziqma@pku.edu.cn](mailto:ziqma@pku.edu.cn)); Prof. Yang Hong ([yanghong@ou.edu](mailto:yanghong@ou.edu))

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16 **AIMERG: a new Asian precipitation dataset (0.1°/half-hourly, 2000-2015) by calibrating GPM**  
17 **IMERG at daily scale using APHRODITE**

18

19 **Highlights**

20 ● **A new effective daily calibration approach, DSTDCA, for improving GPM IMERG**

21 ● **A new AIMERG precipitation data (0.1°/half-hourly, 2000-2015, Asia) is provided**

22 ● **Bias of AIMERG is significantly improved compared with that of IMERG**

23 ● **APHRODITE is more suitable than GPCC in anchoring IMERG over the Asia**

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31 **Abstract**

32       Precipitation estimates with fine quality and spatio-temporal resolutions play significant roles in  
33 understanding the global and regional cycles of water, carbon and energy. Satellite-based precipitation  
34 products are capable of detecting spatial patterns and temporal variations of precipitation at fine  
35 resolutions, which is particularly useful over poorly gauged regions. However, satellite-based  
36 precipitation products are the indirect estimates of precipitation, inherently containing regional and  
37 seasonal systematic biases and random errors. In this study, focusing on the potential drawbacks in  
38 generating Integrated Multi-satellitE Retrievals for Global Precipitation Measurement (IMERG) and its  
39 recently updated retrospective IMERG in the Tropical Rainfall Measuring Mission (TRMM) era (finished  
40 in July, 2019), which were only calibrated at monthly scale using ground observations, Global  
41 Precipitation Climatology Centre (GPCC, 1.0°/Monthly), we aim to propose a new calibration algorithm  
42 for IMERG at daily scale, and to provide a new AIMERG precipitation dataset (0.1°/ half-hourly, 2000-  
43 2015, Asia) with better quality, calibrated by Asian Precipitation Highly Resolved Observational Data  
44 Integration (APHRODITE, 0.25°/Daily) at the daily scale for the Asian applications. And the main  
45 conclusions include but not limited to: (1) the proposed daily calibration algorithm (Daily Spatio-  
46 Temporal Disaggregation Calibration Algorithm, DSTDCA) is effective in considering the advantages  
47 from both satellite-based precipitation estimates and the ground observations; (2) AIMERG performs  
48 better than IMERG at different spatio-temporal scales, in terms of both systematic biases and random  
49 errors, over the China Mainland; and (3) APHRODITE demonstrates significant advantages than GPCC

50 in calibrating the IMERG, especially over the mountainous regions with complex terrain, e.g., the Tibetan  
51 Plateau. Additionally, results of this study suggest that it is a promising and applicable daily calibration  
52 algorithm for GPM in generating the future IMERG in either operational scheme or retrospective manner.

53 The AIMERG data record (0.1°/half-hourly, 2000-2015, Asia) is freely available at [http://argi-  
54 basic.hihanlin.com:8000/d/d925fecf60/](http://argi-basic.hihanlin.com:8000/d/d925fecf60/). Additionally, the AIMERG data is also freely accessible at  
55 <https://doi.org/10.5281/zenodo.3609352> (for the period from 2000 to 2008) (Ma et al., 2020a) and  
56 <http://doi.org/10.5281/zenodo.3609507> (for the period from 2009 to 2015) (Ma et al., 2020b).

57 **Keywords: Precipitation; IMERG; APHRODITE; Calibration; Daily scale; Asia;**

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## 59 **1. Introduction**

60 Precipitation is among the most essential hydroclimatic factors, and also most difficult to estimate  
61 due to its great small-scale variabilities (Yatagai et al., 2012; Huffman et al., 2019a). High spatio-  
62 temporal resolution precipitation dataset with fine quality is essential for various scientific and  
63 operational applications, including but not limited to driving the hydrological models, and supporting  
64 the predictions of droughts and floods (Beck et al., 2017, 2018). There are mainly two principal  
65 approaches for measuring the global precipitation: ground-based gauge observing, and satellite-based  
66 remote sensing, which resulting in three mainstreams of global precipitation products, namely gauge

67 analysis precipitation data, satellite-based only precipitation estimates, and satellite-gauge combined  
68 precipitation products, based on the consideration that ground-based gauge data are clearly important  
69 for anchoring the satellite estimates (Huffman et al., 2007, 2019a).

70 In recent years, a large number of quasi-global satellite precipitation products with various  
71 temporal and spatial resolutions have been developed and released to the public, such as the PMW-based  
72 CPC Morphing technique (CMORPH) (hereafter, for Acronyms, see the Appendix) (Joyce et al., 2004),  
73 and IR-based PERSIANN (Sorooshian et al., 2000) and PERSIANN-CCS (Hong et al., 2004). As the  
74 milestone in the satellite-based precipitation measurement process, the TRMM and its successor GPM  
75 developed a flexible framework for generating the most popular precipitation products, TMPA (1998-  
76 present, 0.25°/3 hourly) and IMERG (2014-present, 0.1°/half-hourly), as well as the retrospective  
77 IMERG (2000-present, 0.1°/half-hourly) from GPM era to TRMM era, which aimed at intercalibrating,  
78 merging, and interpolating all MW estimates of the GPM constellation, IR estimates, and gauge  
79 observations (Huffman et al., 2019b). The “Final run” version of IMERG (hereafter refer to IMERG),  
80 incorporating the monthly gauge analysis, provides the state-of-the-art precipitation estimate with finest  
81 spatio-temporal resolutions so far, while it still contains large uncertainties, e.g., greatly overestimating  
82 the precipitation, at daily and hourly scales from regions to regions, especially over the mountainous  
83 areas, such as the Tibetan Plateau, China (Tang et al., 2016; Lu et al., 2019; Xu et al., 2019), which is  
84 greatly potentially resulted by the calibration procedures in the process of generating the IMERG.  
85 Currently, the IMERG product (following the gauge correction method of TMPA approach) (Huffman

86 et al., 2007) has been produced by anchoring the multi-satellite-only precipitation estimates using the  
87 monthly analysis Satellite-Gauge product (1.0°/monthly, 1979 to the present, delayed by about 3 months)  
88 from the GPCC (Adler et al., 2003, 2018), therefore, the IMERG performed better at monthly and annual  
89 scales than those at finer temporal scales (e.g., daily, hourly).

90 Satellite-based precipitation products have significant advantages in detecting the variations of  
91 precipitation at fine spatio-temporal resolutions, especially over the poorly gauged regions. However, as  
92 the indirect estimates of precipitation, satellite-based precipitation products are inherently containing  
93 regional, seasonal, and diurnal systematic biases and random errors (Ebert et al., 2007), which could be  
94 effectively alleviated by anchoring the satellite-only precipitation products using gauge-based  
95 observations (Huffman et al., 2007). Therefore, great efforts have been taken on exploring the calibrations  
96 on the satellite-only precipitation estimates using gauge analysis. Historically, GPCP has provided the  
97 lion's share of the early efforts in the process of developing calibration algorithms for the satellite-only  
98 precipitation estimates in generating SG products (2.5°/monthly). For instance, to correct the bias of the  
99 multi-satellite only estimates (mainly based on PMW and IR data) on a regional scale, the multi-satellite  
100 estimate was firstly multiplied by the ratio of the large-scale (with moving window size  $5 \times 5$ ) average  
101 gauge analysis to the large-scale average of the multi-satellite estimate, and then the SG estimate was  
102 finally derived by combining the gauge-adjusted multi-satellite estimate and the gauge analysis with  
103 inverse-error-variance weighting (Huffman et al., 1997; Adler et 2003; Adler 2018). Recently, a two-step  
104 strategy was proposed to remove the bias inherent in the multi-satellite only precipitation estimates using

105 PDF matching method and to combine the bias-corrected estimates with the gauge analyses using OI  
106 algorithm (Xie and Xiong, 2011; Shen et al., 2014). And a similar improved PDF algorithm was applied  
107 to generate the GSMaP data, which was adjusted at the daily scale by the gauge analysis (0.5°/daily) from  
108 the CPC (Mega et al., 2014). While GPM IMERG adjusted the multi-satellite precipitation estimates  
109 (0.1°/half hourly) at the monthly scale using the ratios between the original monthly multi-satellite-only  
110 and the monthly satellite-gauge data, in combination with the original monthly multi-satellite-only and  
111 GPCC (1.0°), in the month (Huffman et al., 2019a). There is still much room for exploring the improved  
112 algorithms for calibrating the multi-satellite-only precipitation estimates at finer spatiotemporal scales, e.  
113 g, 0.25°/daily, which is also one of the next vital focuses by the GPM (Huffman et al., 2019a).

114 As for anchoring the satellite precipitation estimates, the quality and spatio-temporal resolutions of  
115 the gauge analysis precipitation data are the key factors. Though the GPCC has developed a series of  
116 gauge-based precipitation analysis datasets with the quality and spatio-temporal resolutions continually  
117 improved, accurate estimations of precipitation over the land are still greatly difficult with limited  
118 networks of rain gauges. In Asia, great efforts also have been mainly paid on generating gauge-analysis  
119 precipitation products at the monthly scale (Chen et al., 2002; Mitchell and Jones 2005; Matsuura and  
120 Willmott 2009; Schneider et al. 2008), and limited explorations at the daily scale, e.g., Rajeevan and  
121 Bhate (2009) explored daily grid precipitation data over India with data from more than 2,500 rain  
122 gauges. Meanwhile, significant differences among those products had been reported by Yatagai et al  
123 (2005, 2012). To more accurately monitor and predict the Asian hydro-meteorological environment, the

124 APHRODITE project (starting in 2006) aimed at developing the state-of-the-art gridded precipitation  
125 datasets at the resolutions of 0.25°/daily covering the entire Asia based on the largest numbers of ground  
126 observations from multi-sources. Since the release of APHRODITE products (1951-2015, 0.25°/daily,  
127 Last update October 5, 2018), APHRODITE daily grid precipitation data sets have been widely used,  
128 and it distinguished from other gauge analysis data by considering the different interpolation schemes  
129 and climatology characteristics, especially over the mountainous regions with complex terrain, e.g., the  
130 Tibetan Plateau (Yatagai et al., 2012).

131 The aim of this study is to explore the calibration approach at daily scale on the retrospective  
132 IMERG data using APHRODITE product, in both TRMM and GPM eras, from 2000 to 2015.  
133 Meanwhile, a new calibration approach, Daily Spatio-Temporal Disaggregation Calibration Algorithm  
134 (DSTDCA), is proposed and suggested for the GPM in their future algorithms; and a new AIMERG  
135 precipitation dataset (0.1°/ half-hourly, 2000-2015, Asia) (Ma et al., 2020a, b) with better quality is to  
136 be provided publicly for the Asian applications.

## 137 **2. Data**

### 138 **2.1 IMERG**

139 To generate the IMERG product, IMERG focused on intercalibrating, merging, and interpolating  
140 “all” satellite MW-based precipitation estimates, together with MW-calibrated IR-based precipitation  
141 estimates, precipitation gauge analyses, and potentially other precipitation estimators at fine spatio-

142 temporal scales for the both TRMM and GPM eras over the entire globe. Currently, IMERG is at its  
143 Version 06 stage ([https://pmm.nasa.gov/sites/default/files/document\\_files/IMERG\\_ATBD\\_V06.pdf](https://pmm.nasa.gov/sites/default/files/document_files/IMERG_ATBD_V06.pdf)),  
144 based on which IMERG has been retrospect to the TRMM era at the end of September, 2019, and  
145 IMERG is now available back to June 2000 (half-hourly/0.1°) ([https://pmm.nasa.gov/data-  
146 access/downloads/gpm](https://pmm.nasa.gov/data-access/downloads/gpm)). The “Final run” of IMERG combines the GPCC Monitoring product, the V8  
147 Full Data Analysis for the majority of the time (currently 1998-2016), and the V6 Monitoring Product  
148 from 2017 to the then-present. The Monitoring Product is posted about two months after the month of  
149 observations from ~7,000-8,000 stations world-wide, which is relative sparse, especially over the Asia  
150 (Schneider et al. 2014, 2018).

## 151 **2.2 APHRODITE**

152 Since the release of the APHRODITE product (0.25°/Daily, 1951-2007), it has been widely used as  
153 one of state-of-the-art daily grid precipitation datasets over the Asia, for hydro-climatological related  
154 studies (Yatagai et al., 2012; Menegoz et al., 2013; Sunilkumar et al., 2019). APHRODITE has been  
155 demonstrated to replicate ‘ground truth’ observations very well (Duncan and Bigg, 2012) and represents  
156 the optimal dataset for analyzing historical precipitation variability and change. Recently, the  
157 APHRODITE data has been updated from the former period 1951-2007 to a longer period 1951-2015, in  
158 September, 2018, with continuous efforts of quality control (QC) flagging some data (Hamada et al.,

159 2011). The APHRODITE data could be available through the website ([http://aphrodite.st.hirosaki-](http://aphrodite.st.hirosaki-u.ac.jp/download/)  
160 [u.ac.jp/download/](http://aphrodite.st.hirosaki-u.ac.jp/download/)).

### 161 **2.3 CMPA**

162 The China Merged Precipitation Analysis (CMPA, 0.1°/hourly, 2008-2015) were generated by using  
163 hourly rain gauge data at more than 30, 000 automatic weather stations in China, with the combination of  
164 the CMORPH precipitation product, and provided by the Chinese Meteorological Administration  
165 (<http://data.cma.cn>) (Shen et al., 2014). The OI method was adopted to estimate the areal precipitation  
166 distribution based on the gauge observations (Yong et al., 2010), but uncertainty still exists in the  
167 interpolated precipitation field particularly over West China with relatively sparse gauge networks. For  
168 grid boxes with gauges, the observed precipitation values are exactly the gauge observation or the  
169 averaged observation when more than one gauge locates in a grid.

### 170 **2.4 Point-based rain gauge data from meteorological stations**

171 The hourly rain gauge datasets from 57, 835 national ground stations used in this study, in 2015,  
172 were collected from the National Meteorological Information Center of CMA (<http://data.cma.cn>). All  
173 the gauge data have undergone strict quality control in three levels, which includes (1) the extreme values'  
174 check, (2) internal consistency check, and (3) spatial consistency check (Shen et al., 2010). Most gauges  
175 are located over the eastern and southern parts of the Mainland China, and relatively sparse gauge  
176 networks are located across the northern and western parts, especially over the Tibetan Plateau. The

177 limited number of gauges could be a source of error in evaluation of satellite precipitation products in  
178 such areas (Shen et al., 2014).

## 179 **2.5 Point-based rain gauge data from hydrological stations**

180 The hourly ground precipitation observations from around 500 hydrological stations (the number of  
181 station varied from year to year) used in this study were collected from Hydrology Bureau of Zhejiang  
182 Province, southeastern China (<http://data.cma.cn/>). The quality control follows two steps: (1) the datasets  
183 are filtered by threshold value after being collected from rain gauges; (2) the outliers are identified through  
184 manual processing. With careful data quality control, the rain gauge datasets have satisfying performances  
185 on the accuracy and validity.

186 There are five datasets used in this study (refer to Table 1 for a summary of the datasets). IMERG  
187 and APHRODITE were used for generating the AIMERG data, and the others were used for evaluating  
188 and comparing the IMERG and AIMERG at different scales.

189

190 **Table 1.** List of satellite-based, gauge-based, and satellite-gauge combination precipitation products used  
 191 in this study.

Short name	Full name	Spatial and temporal sampling	Time period	References
IMERG	Integrated Multi-satellitE Retrievals for Global Precipitation Measurement	0.1°/half-hourly	2000-present	Huffman et al. (2019b) <a href="https://pmm.nasa.gov/data-access/downloads/gpm">https://pmm.nasa.gov/data-access/downloads/gpm</a> (last access: 17 January 2020)
APHRODITE	Asian Precipitation Highly Resolved Observational Data Integration Towards Evaluation of Water Resources	0.25°/daily	1951-2015	Yatagai et al. (2012) <a href="http://aphrodite.st.hirosaki-u.ac.jp/download/">http://aphrodite.st.hirosaki-u.ac.jp/download/</a> (last access: 17 January 2020)
CMPA	China Merged Precipitation Analysis	0.1°/hourly	2008-present	Shen et al. (2014) <a href="http://data.cma.cn">http://data.cma.cn</a> (last access: 17 January 2020)
	Point-based rain gauge data	hourly	2010-present	Shen et al. (2010) <a href="http://data.cma.cn">http://data.cma.cn</a> (last access: 17 January 2020)

192

### 193 3. Methodology

#### 194 3.1 Calibration Procedure of the Daily Spatio-Temporal Disaggregation Calibration 195 Algorithm, DSTDCA

196 According to previous evaluations on IMERG (Lu et al., 2019; Xu et al., 2019), there are at least two  
197 characteristics resulting its significant overestimations: (1) the amplitude of hourly or half-hourly  
198 estimated rainfall rates are significantly amplified by IMERG compared with ground observations, which  
199 might be caused by the benchmark of GPCC and GPCP SG data for calibrations, and (2) the IMERG  
200 algorithm is generally over detecting precipitation events, resulting a large fraction of false alarm but  
201 unreal precipitation events. Therefore, this study selects the APHRODITE data as the benchmark for  
202 calibrating IMERG at daily scale, based on the proposed approach, DSTDCA, and the main steps of the  
203 DSTDCA are shown as follows:

204 (1) IMERG data ( $0.1^\circ$ /half-hourly) are accumulated to IMERG data at the daily scale ( $0.1^\circ$ ), which  
205 are used to generate the spatial disaggregation weights. As the spatial resolution of APHRODITE data is  
206  $0.25^\circ$ , the moving window size of 3 by 3 is selected, and the daily spatial disaggregation weights ( $0.1^\circ$ )  
207 based on IMERG is obtained by calculating the ratios between the daily rainfall accumulations at the  
208 central grid and the average daily rainfall accumulations in the corresponding  $3 \times 3$  window. The daily  
209 spatial disaggregation weights consider the relative spatial patterns of the precipitation captured by the  
210 IMERG;

211 (2) Based on the daily precipitation accumulations of IMERG, the half-hourly temporal  
212 disaggregation weights ( $0.1^\circ$ ) are derived by calculating the ratios between the each half-hourly

213 precipitation estimates and the corresponding daily precipitation estimates. If the daily accumulation  
214 estimate is equal to zero, then each half-hourly temporal disaggregation weight is set as zero;

215 (3) As there is a small fraction of grids in APHRODITE with no data at daily scale, the no data grids  
216 in APHRODITE data are firstly filled with the data according to its nearest neighbor with effective value;

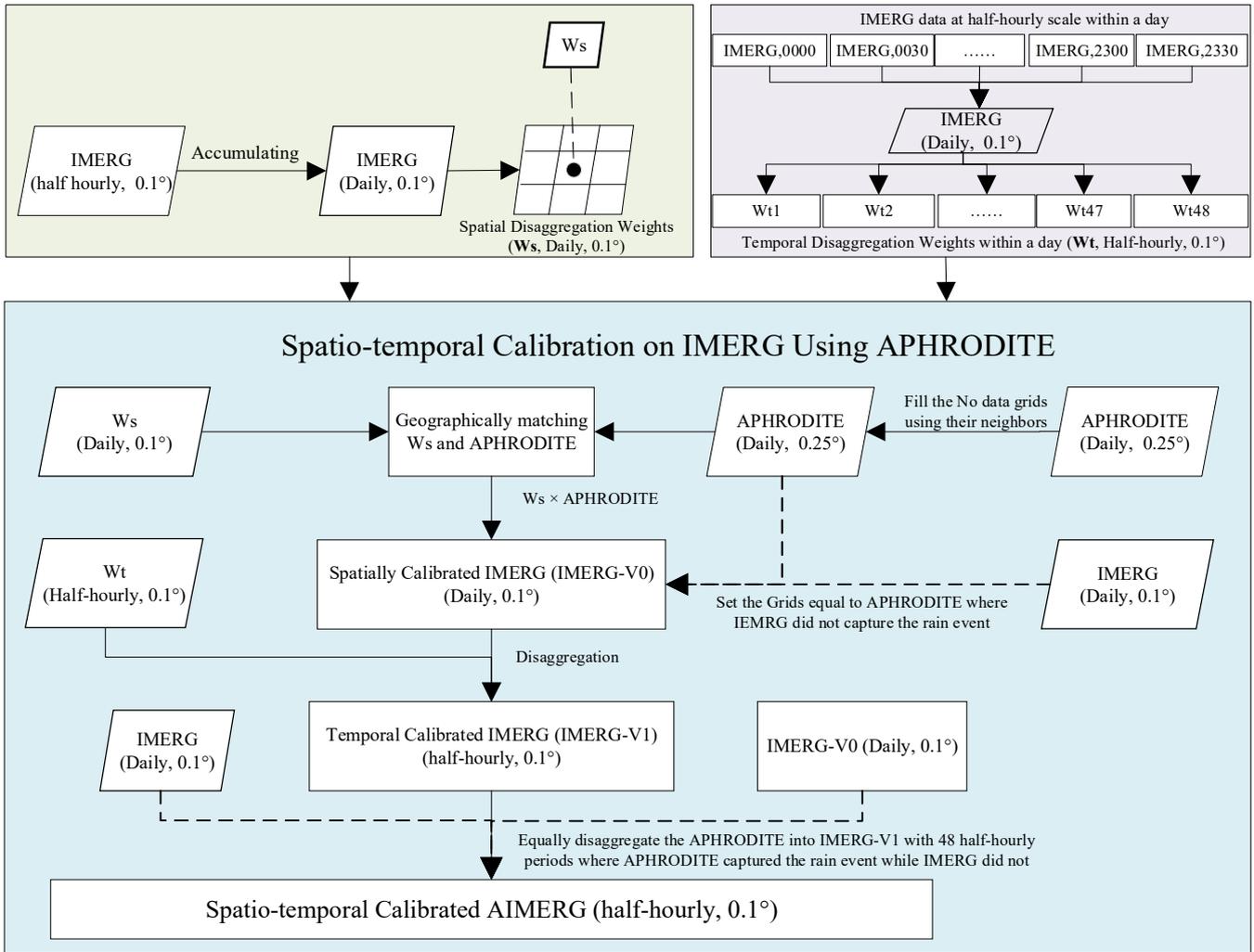
217 (4) Spatial calibrations: the daily calibrated IMERG using APHRODITE data are obtained by  
218 multiplying the spatial disaggregation weights based on IMERG ( $0.1^\circ$ /daily) from step (1) by daily  
219 APHRODITE data ( $0.25^\circ$ / daily) from step (3). In this step, to match the IMERG ( $0.1^\circ$ ) and APHRODITE  
220 ( $0.25^\circ$ ), the numbers and weights of the APHRODITE grids corresponding to each IMERG pixel are  
221 determined, according to the relative spatial locations and coverage relationships between the each pixel  
222 of IMERG ( $0.1^\circ$ ) and the corresponding pixels of APHRODITE ( $0.25^\circ$ );

223 (5) Temporal calibrations: the half-hourly calibrated IMERG are obtained by multiplying the half-  
224 hourly temporal disaggregation weights ( $0.1^\circ$ /half-hourly) from step (2) by the daily calibrated IMERG  
225 from step (4);

226 (6) By considering the situations that APHRODITE data captured the precipitation while the IMERG  
227 did not, the half-hourly calibrated IMERG is further processed by equally disaggregating the value from  
228 the daily APHRODITE data at the corresponding grid into 48 half-hourly periods, which are regarded as  
229 the half-hourly calibrated IMERG values in the corresponding day;

230 (7) By considering the situations that IMERG data captured the precipitation while the APHRODITE  
231 did not, the 48 half-hourly calibrated IMERG values in corresponding days and locations are all set as  
232 zero, to meet the ground truth observations. And this consideration has been already conducted in the  
233 fourth step;

234 After all the above-mentioned procedures, the final calibrated AIMERG (0.1° half-hourly) data are  
235 obtained by considering both the total precipitation controls and the effective precipitation events  
236 measured by the “ground truth” observations by APHRODITE data over the Asia.



239 Figure 1. The flowchart of the Daily Spatio-Temporal Disaggregation Calibration Algorithm,

240 DSTDCA, to generate the AIMERG dataset over the Asia, 2000-2015

241 **3.2 Evaluation Metrics**

242 To evaluate the IMERG and its calibrations comprehensively, seven metrics (CC, MAE, BIAS,  
 243 RMSE, POD, FAR, CSI) were selected (Tang et al., 2016). Generally, CC is used to describe the  
 244 agreements between satellite estimates and gauge observations; MAE, RMSE, and BIAS are used to  
 245 indicate the error and bias of satellite estimates compared with gauge observations; and the POD, FAR,  
 246 and CSI are used to demonstrate the capabilities to correctly capture the precipitation events of satellite  
 247 precipitation estimates against the ground observations. The detailed information of these evaluation  
 248 metrics are listed in Table 2.

249 **Table 2** Formulas and perfect values of the evaluation metrics used in this study<sup>a</sup>.

Statistic metrics	Equation	Perfect value	Value ranges
Correlation Coefficient (CC)	$CC = \frac{\frac{1}{N} \sum_{n=1}^N (S_n - \bar{s})(G_n - \bar{G})}{\sigma_S \sigma_G}$	1	[-1, 1]
Mean Error (ME)	$ME = \sum_{n=1}^N (S_n - G_n)$	0	$(-\infty, +\infty)$
Relative Bias (BIAS)	$BIAS = \frac{\sum_{n=1}^N (S_n - G_n)}{\sum_{i=1}^n G_n} \times 100\%$	0	$(-\infty, +\infty)$
Root Mean Square Error (RMSE)	$RMSE = \sqrt{\frac{1}{N} \sum_{n=1}^N (S_n - G_n)^2}$	0	$[0, +\infty)$
Probability of Detection (POD)	$POD = \frac{n_{11}}{n_{11} + n_{01}}$	1	[0, 1]
False Alarm Ratio (FAR)	$FAR = \frac{n_{10}}{n_{11} + n_{10}}$	0	[0, 1]
Critical Success Index (CSI)	$CSI = \frac{n_{11}}{n_{11} + n_{10} + n_{01}}$	1	[0, 1]

250 <sup>a</sup>Notation: n is the sample numbers; S<sub>n</sub> is satellite precipitation estimate; G<sub>n</sub> is gauge-based precipitation; σ<sub>G</sub> is the standard deviations of  
 251 gauge-based precipitation; σ<sub>S</sub> is the standard deviations of satellite-based precipitation estimate. n<sub>11</sub> is the precipitation event detected by  
 252 both gauge and satellite simultaneously; n<sub>10</sub> is the precipitation event detected by the satellite but not detected by the gauge; n<sub>01</sub> is contrary  
 253 to n<sub>10</sub>; n<sub>00</sub> is the precipitation events detected neither by the gauge nor the satellite.

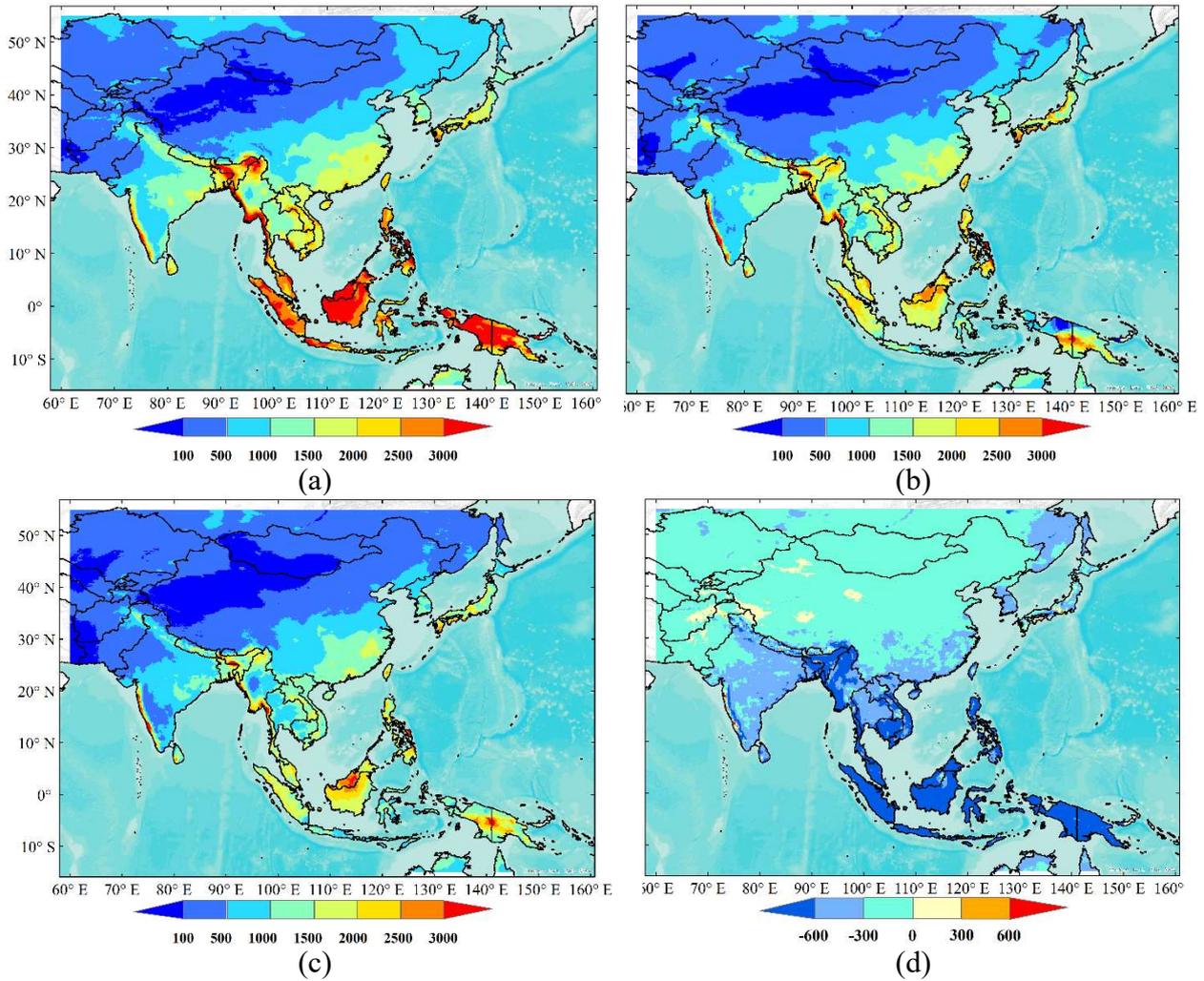
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## 255 4. Results

### 256 4.1 AIMERG Product

257 Generally, both IMERG and APHRODITE share similar spatial patterns with precipitation volumes  
258 decreasing from southeast to northwest in Asia, while compared with APHRODITE data (Fig. 2b),  
259 IMERG greatly overestimates the precipitation over Arunachal Pradesh, coastal Indochina and Western  
260 Ghats, and the Indonesia (Fig. 2a). Corrected by APHRODITE, the spatial patterns and volumes of  
261 AIMERG are much more similar to those of APHRODITE, especially along the Himalayas, coastal  
262 Indochina and Western Ghats, and the Indonesia (Fig. 2c). Compared with APHRODITE, AIMERG  
263 seems floating up and down in terms of the volumes, for instance, AIMERG is larger and smaller than  
264 APHRODITE in eastern Indonesia and northeastern Asia, respectively. Though AIMERG is smaller  
265 than IMERG over most regions, there are still some areas where the volumes of AIMERG are larger than  
266 those of IMERG, e.g., in western Tibetan Plateau, Middle East, and along the western coast of India (Fig.  
267 2d).

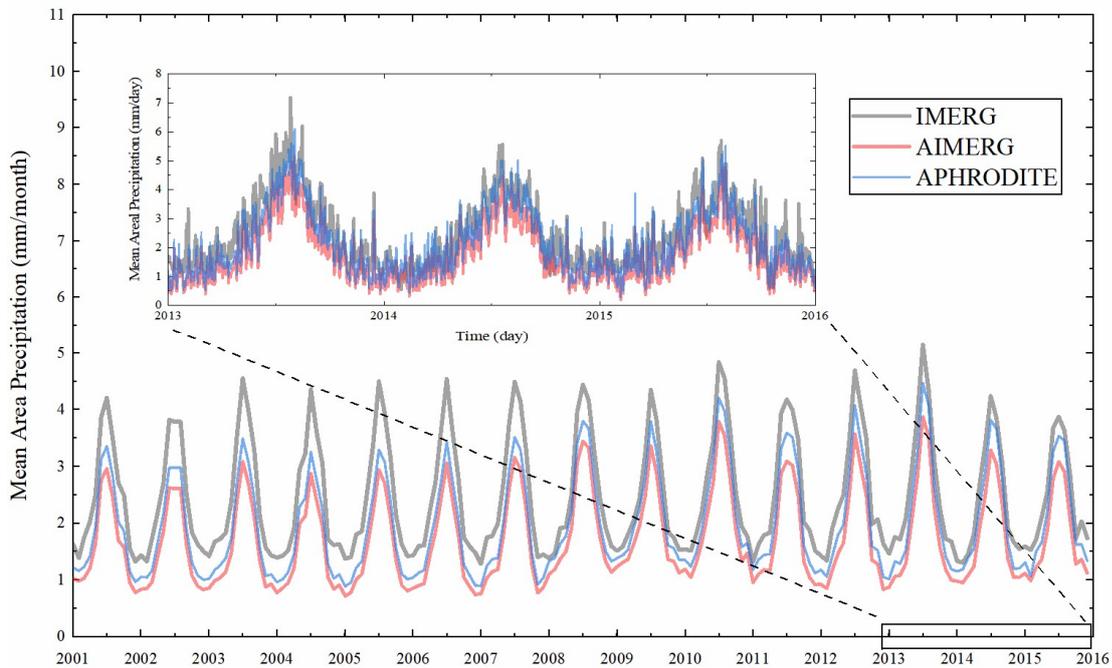
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269 Figure 2. Spatial patterns of Asian mean annual gridded precipitation products of (a) IMERG, 0.1°, (b)  
 270 APHRODITE, 0.25°, and (c) AIMERG, 0.1°, and (d) AIMERG-IMERG, 0.1°, respectively, during the  
 271 period of 2001-2015. The background map used in this study was provided by Esri, USGS and NOAA  
 272 ([http://goto.arcgisonline.com/maps/World\\_Terrain\\_Base](http://goto.arcgisonline.com/maps/World_Terrain_Base), last access: 17 January 2020).

273

274 The temporal patterns of the mean areal precipitation over the Monsoon Asia of the three products  
275 demonstrate that the systematic bias of IMERG is significantly reduced in both dry and wet seasons,  
276 shown in Fig. 3. IMERG is around 1.5 times larger than APHRODITE at monthly scale. Though much  
277 more close to the APHRODITE, AIMERG is still a little smaller than the APHRODITE, which means  
278 the calibration algorithm proposed by this study tends to underestimate the precipitation compared with  
279 calibration benchmark, APHRODITE. At daily scale, IMERG is generally larger than APHRODITE,  
280 while at some special days, APHRODITE is larger than IMERG, which might result the AIMERG may  
281 be also larger than IMERG.



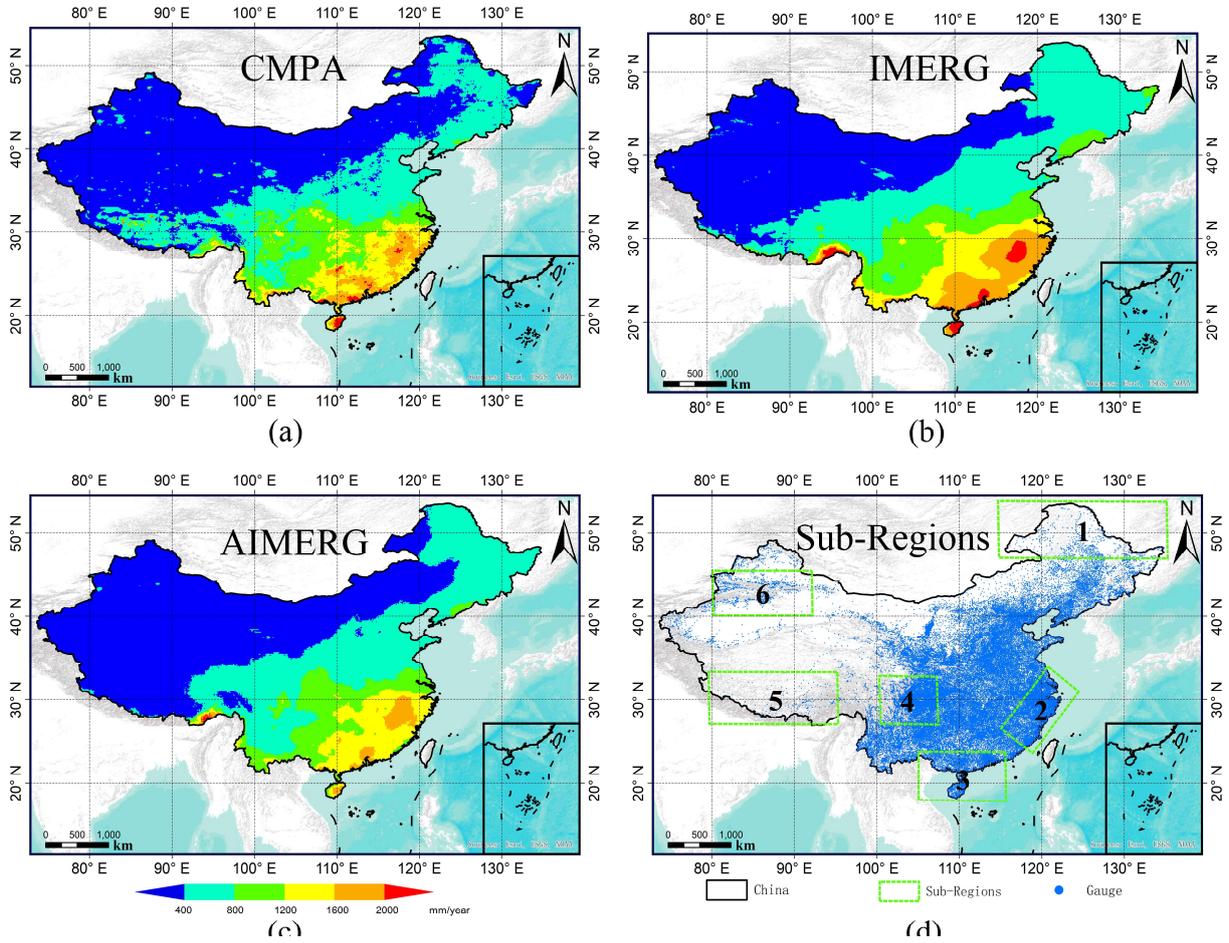
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283 Figure 3. The temporal variations of mean Asian gridded precipitation products of IMERG, APHRODITE,  
284 and AIMERG, respectively, during the period of 2001-2015.

#### 285 **4.2 Assessments on IMERG and AIMERG at national and regional scales**

286 The spatial patterns of CMPA demonstrate much more similar to those of AIMERG, especially in  
287 the southeastern China where dense rain gauges are located, while both CMPA and IMERG overestimate  
288 the precipitation along the Himalayas where the meteorological gauges are sparse and mainly the satellite-  
289 based observations are applied (Fig. 4). Obviously, the IMERG significantly overestimates the  
290 precipitation in the southeast coast of China, where typhoons always visit (Fig. 4 b). For deciding the  
291 sub-regions (Fig. 4 d), we have mainly considered three aspects: the representative climatic zones in  
292 China, the local distributions of the gauge stations, and the complexity of the topography. For instances,  
293 Sub-Region 1 represents the high latitude plain in the most north-eastern region of China under a cold  
294 climate (left top: 115.0° E, 54.0°N; right bottom: 135.0° E, 47.0°N); Sub-Region 2 represents the south-  
295 eastern coastal area of China influenced greatly by the Asian Monsoons (left top: 115.0° E, 26.0°N; left  
296 bottom: 119.0° E, 24.0°N; right bottom: 124.0° E, 31.0°N; right top: 120.0° E, 34.0°N); Sub-Region 3  
297 represents the most southern region including the island Hainan in the tropical zone (left top: 105.0° E,  
298 24.0°N; right bottom: 115.0° E, 18.0°N); Sub-Region 4 represents the inner area of China covering the  
299 Yunnan-Kweichow Plateau and Sichuan Basin, under a humid inland climate (left top: 100.0° E, 33.0°N;  
300 right bottom: 107.0° E, 27.0°N); Sub-Region 5 represents the most southern Tibetan Plateau along the

301 Himalayas with complex terrains and high elevations above  $\sim 4000.0$  meters (left top:  $80.0^\circ$  E,  $33.0^\circ$ N;  
 302 right bottom:  $95.0^\circ$  E,  $27.0^\circ$ N); Sub-Region 6 represents the central Asia with complex terrains covering  
 303 the entire Tianshan Mountains in China under an arid inland climate (left top:  $80.0^\circ$  E,  $45.0^\circ$ N; right  
 304 bottom:  $92.0^\circ$  E,  $40.0^\circ$ N).



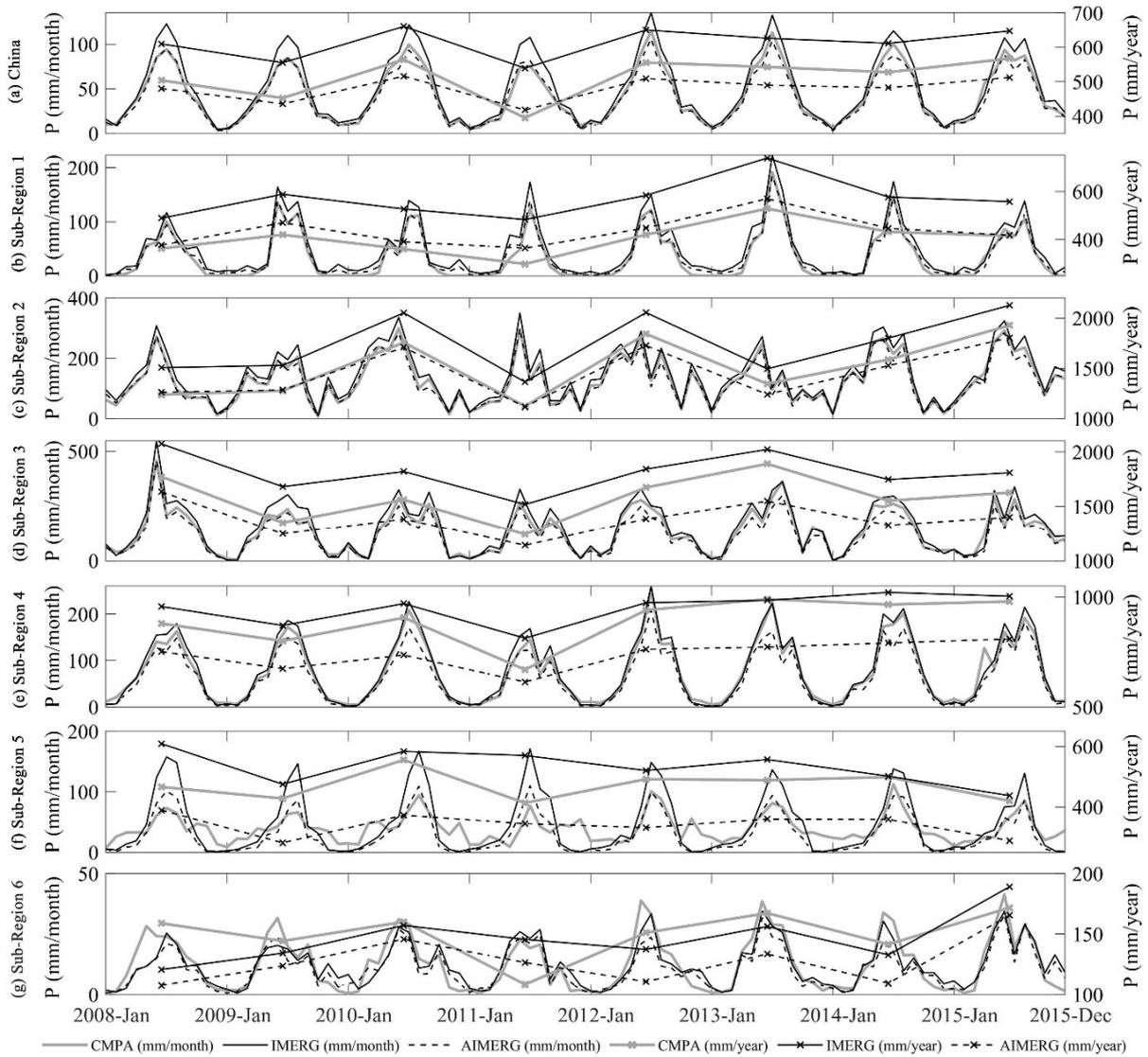
305

306 Figure 4 Spatial patterns of (a) CMPA, (b) IMERG, and (c) AIMERG over China Mainland From  
307 2008~2015, and (d) the spatial distributions of the ~ 50, 000 automatic meteorological stations in China  
308 Mainland. The accurate boundary information of the Sub-Regions: Sub-Region 1 (left top: 115.0° E,  
309 54.0°N; right bottom: 135.0° E, 47.0°N); Sub-Region 2 (left top: 115.0° E, 26.0°N; left bottom: 119.0°  
310 E, 24.0°N; right bottom: 124.0° E, 31.0°N; right top: 120.0° E, 34.0°N); Sub-Region 3 (left top: 105.0°  
311 E, 24.0°N; right bottom: 115.0° E, 18.0°N); Sub-Region 4 (left top: 100.0° E, 33.0°N; right bottom: 107.0°  
312 E, 27.0°N); Sub-Region 5 (left top: 80.0° E, 33.0°N; right bottom: 95.0° E, 27.0°N); Sub-Region 6 (left  
313 top: 80.0° E, 45.0°N; right bottom: 92.0° E, 40.0°N). The background map used in this study was provided  
314 by Esri, USGS and NOAA ([http://goto.arcgisonline.com/maps/World\\_Terrain\\_Base](http://goto.arcgisonline.com/maps/World_Terrain_Base), last access: 17  
315 January 2020).

316

317 The magnitudes of IMERG, AIMERG, and CMPA are compared at national and regional scale  
318 over the China Mainland from 2008 to 2015 (Fig. 5). Generally speaking, CMPA and AIMERG are almost  
319 same, and are significantly smaller than IMERG at both annual and monthly scales, additionally, CMPA  
320 is still a little larger than AIMERG over the China Mainland, which could be possibly resulted from the  
321 use of satellite observations in the CMPA and IMERG (Fig. 6a). The overall situations of the three  
322 product in sub-region 1 and 2 are similar with those over the China Mainland (Fig. 6 b-c), while both  
323 CMPA and IMERG are both significantly larger than AIMERG (Fig. 6 d-f). In sub-region 6, the Tianshan

324 Mountains, CMPA is almost even larger than IMERG, which indicates that large uncertainties should be  
 325 focused on sub-region 6 (Fig. 6 g).

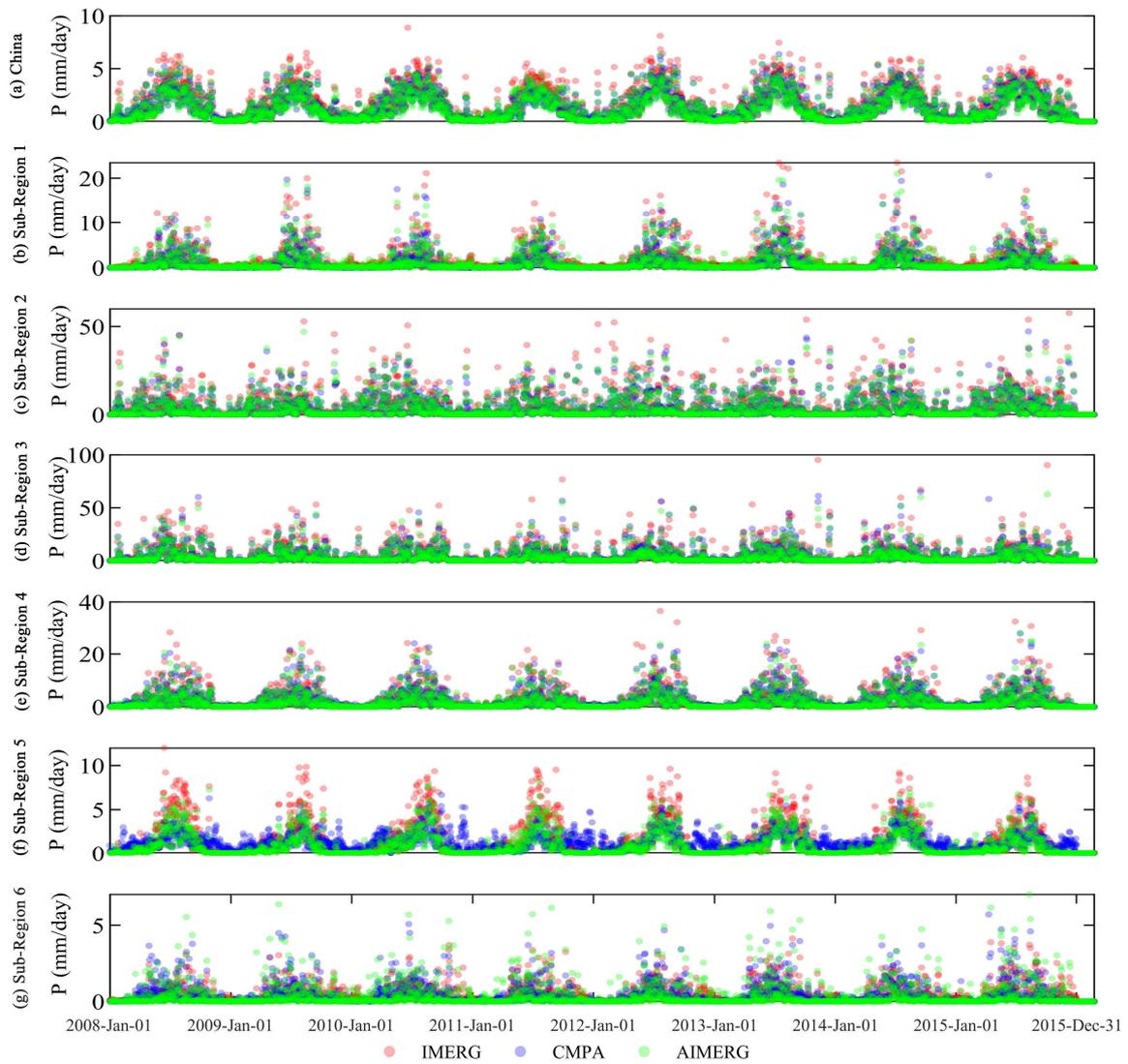


326

327 Figure 5 The temporal patterns of mean areal precipitation of the IMERG, CMPA, and AIMERG, over  
328 China Mainland and sub-regions from 2008 to 2015, at monthly and annual scales.

329 As this study aims to propose a new algorithm for calibrating the IMERG product at the daily scale,  
330 the daily spatial patterns of IMERG, CMPA, and AIMERG were explored, which generally agree with  
331 those of IMERG, CMPA, and AIMERG at monthly scale (Fig. 6). In mountainous region, along the  
332 Himalayas, with relatively small precipitation, CPMA is greatly larger and smaller than the other two  
333 products (both IMERG and AIMERG) in dry seasons and wet seasons respectively (Fig. 6 f). One  
334 phenomenon should be noted that the CPMA seems abnormal along the Himalayas, which might be  
335 resulted by the limited ground observations used in CMPA, shown in Fig 4d, while APHRODITE data  
336 integrate large numbers of ground observations from the neighbor countries, such as India, Nepal, Bhutan,  
337 providing valuable information for retrieving high quality precipitation product around the Tibetan  
338 Plateau (Yatagai, 2012). Calibrated by APHRODITE at daily scale, AIMERG is significantly smaller  
339 than IMERG and CMPA at both annual and monthly scale, while there are also some situations that  
340 AIMERG is larger than IMERG and CMPA at daily scale, for example in sub-region 6, over the Tianshan  
341 mountains.

342



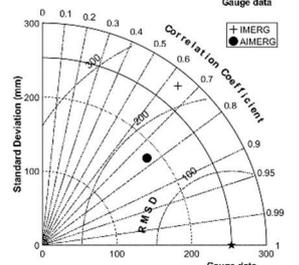
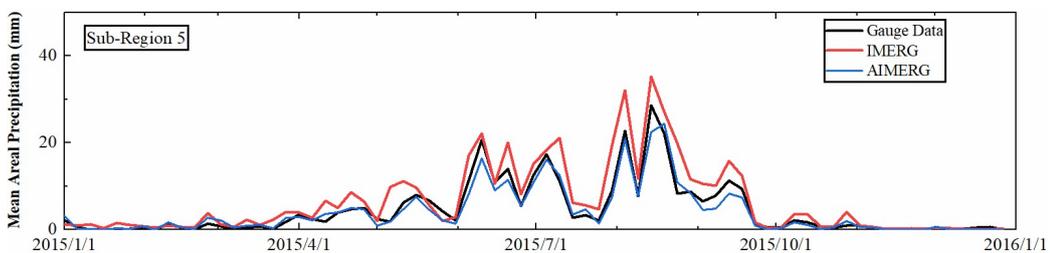
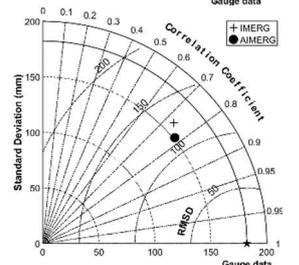
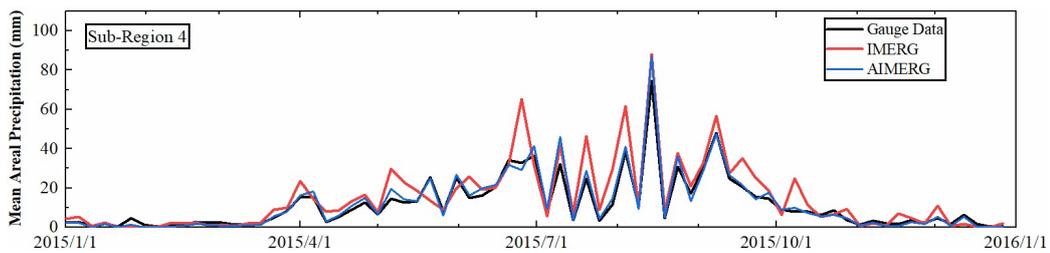
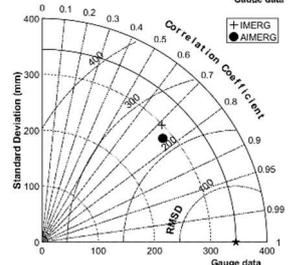
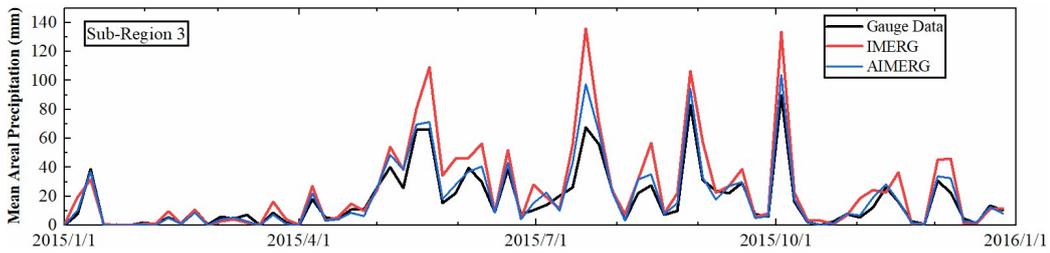
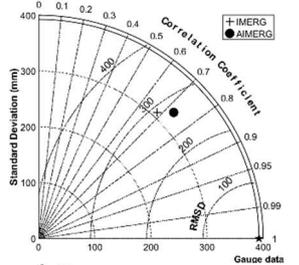
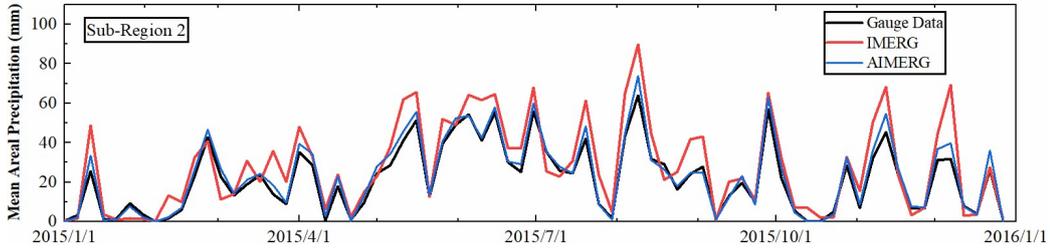
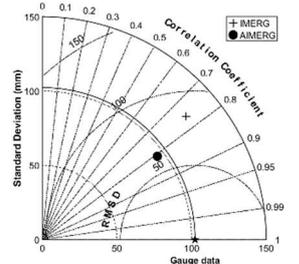
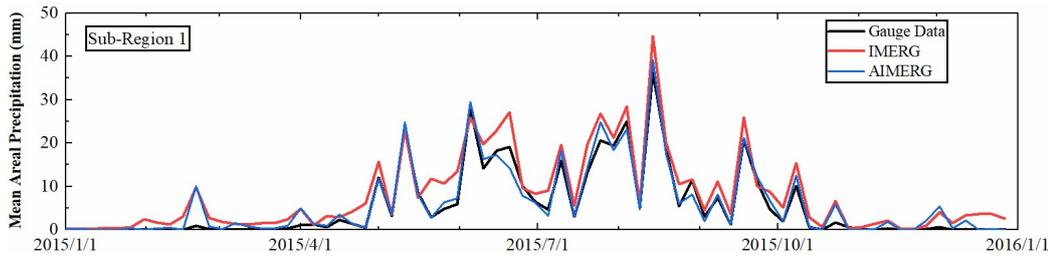
343

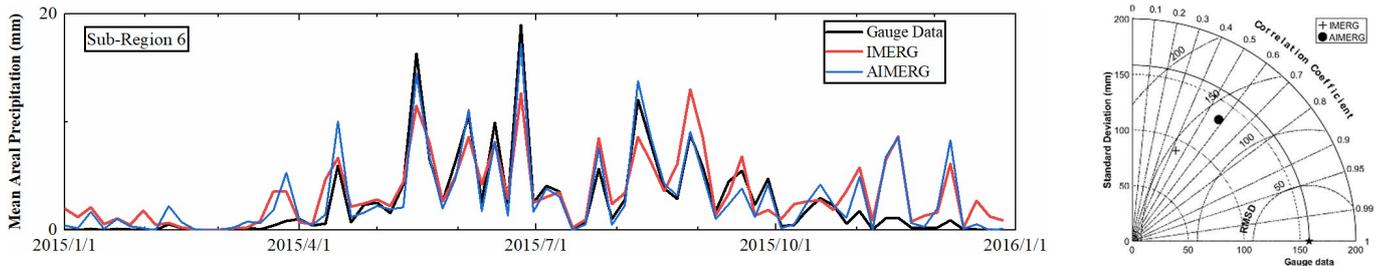
344 Figure 6. The temporal patterns of mean areal precipitation of the IMERG, CMPA, and AIMERG, over  
 345 China Mainland and sub-regions from 2008 to 2015, at daily scale.

346

347 Hourly ground observation data from more than 50,000 meteorological stations were used to assess  
348 the quality of the IMERG and its calibrations, AIMERG, over the six sub-regions, in 2015 (Fig. 7). The  
349 temporal patterns and volumes of mean areal precipitation by AIMERG and ground observations are  
350 almost same, while IMERG is generally larger than AIMERG and ground observations. Meanwhile, the  
351 IMERG still has the problems in overestimating and underestimating the precipitation in dry seasons  
352 (relatively large precipitation occurring) and wet seasons (relatively small precipitation happening),  
353 respectively, for example in sub-region 6, over the Tianshan Mountains. In terms of quantitative indices  
354 (Standard deviation, RMSD, and CC), AIMERG generally outperforms the IMERG against the ground  
355 observations, especially in sub-region 5, along the Himalayas, which indicates that the ground information  
356 from the neighbor countries integrated into the APHRODITE data greatly benefits the calibration results,  
357 AIMERG.

358

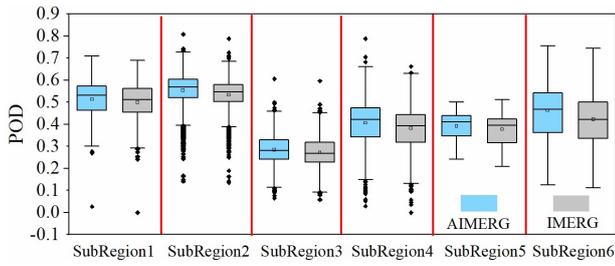




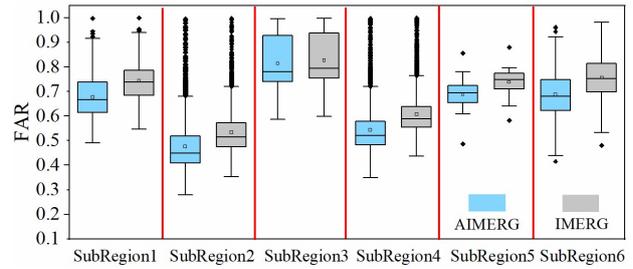
359 Figure 7. The temporal patterns and the volumes of IMERG, ground observations, and AIMERG, in six  
 360 sub-regions at daily scale; and the Taylor diagrams of performances on IMERG and AIMERG against  
 361 ground observations in terms of centered root-mean-square difference, correlation coefficient and  
 362 standard deviation in the six sub-regions at hourly scale, in 2015.

363

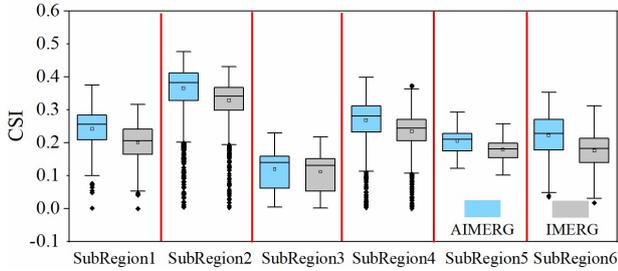
364 Figure 8 illustrates the numerical distributions of contingency statistics for IMERG and AIMERG,  
 365 at hourly scale, in six sub-regions, 2015. Generally, the POD values of AIMERG are larger than those of  
 366 IMERG (Fig. 8a), and FAR values of AIMERG are overall smaller than those of IMERG in each sub-  
 367 regions (Fig. 8b), which results the better performances of the comprehensive index, CSI, combining both  
 368 the characteristics of POD and FAR, in each sub-regions (Fig. 8c). Additionally, both the IMERG and  
 369 AIMERG perform best in sub-region 2, and worst in sub-region 3.



(a)



(b)



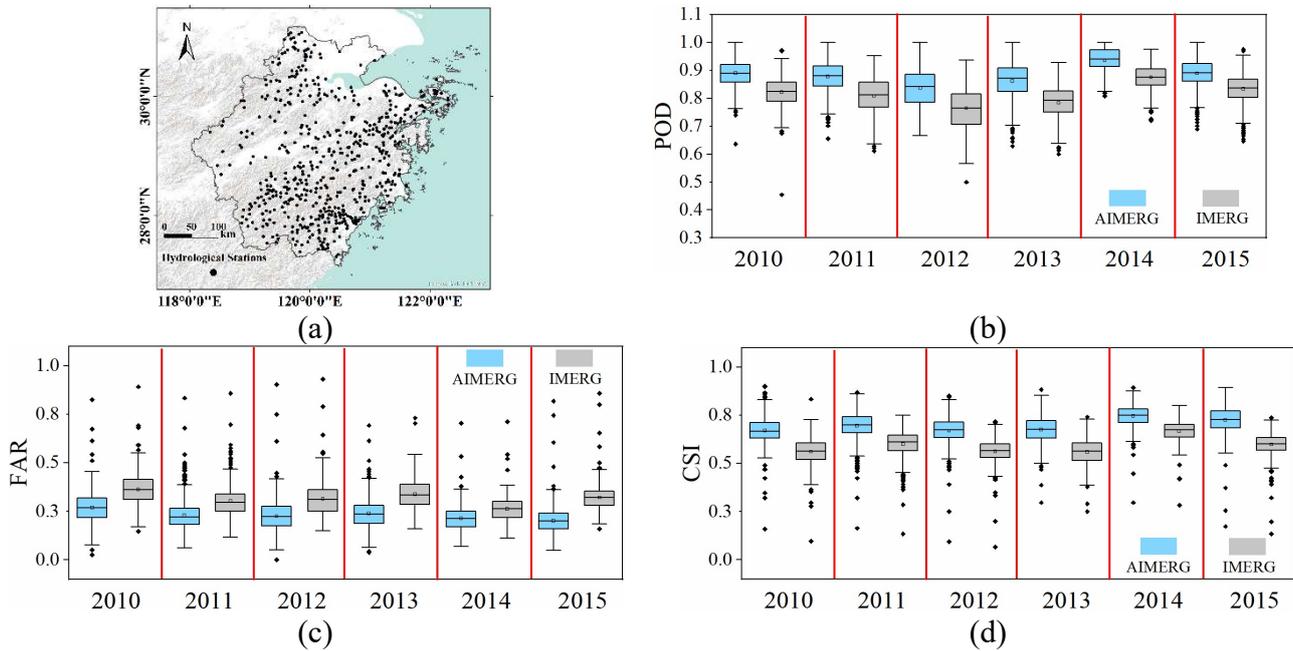
(c)

370 Figure 8. The boxplots demonstrate diagnose of IMERG and AIMERG against the ground observations  
 371 from the meteorological stations, at hourly scale, in six sub-regions, 2015.

372

373 To assess the quality of the IMERG and AIMERG, entirely independent precipitation data from  
 374 around 500 hydrological stations, at hourly scale, from 2010 to 2015, were applied, which are relatively  
 375 even distributed in Zhejiang province (Fig. 9a). The POD values of AIMERG ( $\sim 0.9$ ) are general larger  
 376 than those of IMERG ( $\sim 0.8$ ), while the FAR values of AIMERG ( $\sim 0.3$ ) are significantly smaller than  
 377 those of IMERG ( $\sim 0.4$ ), which results in the overall capabilities of AIMERG to capture the precipitation  
 378 events are improved more than 10%, compared with IMERG, in terms of the CSI. The relative smaller  
 379 POD values and larger FAR values of IMERG in the Zhejiang province, southeastern coast of China,

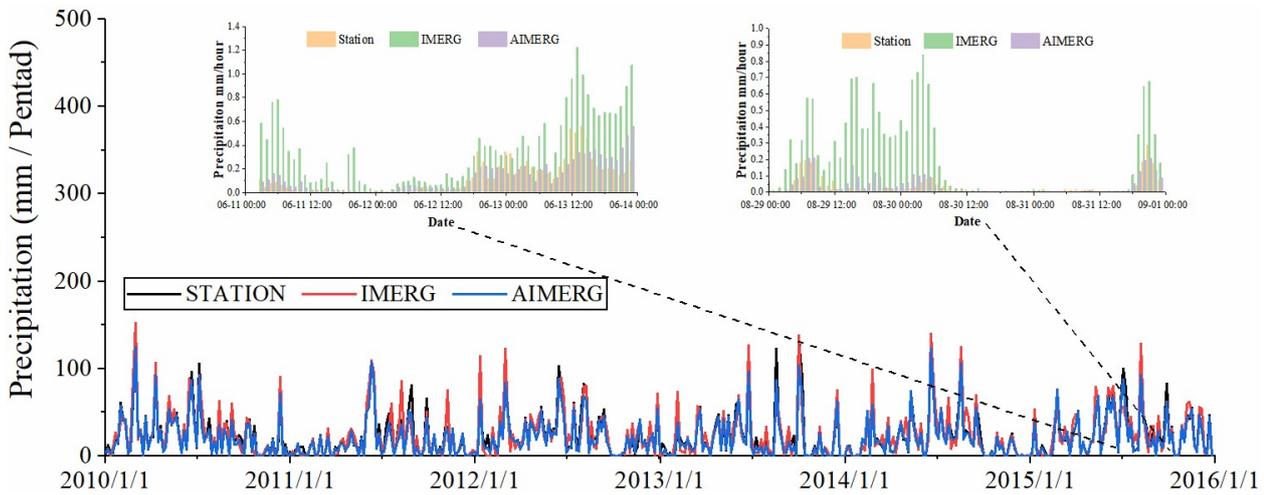
380 might be one of the potential drawbacks in accurately estimating the precipitation both qualitatively and  
 381 quantitatively.



382 Figure 9. The boxplots demonstrate diagnose of IMERG and AIMERG against the ground observations  
 383 from hydrological stations, respectively, at hourly scale, in Zhejiang province, 2010-2015. The  
 384 background map used in this study was provided by Esri, USGS and NOAA  
 385 ([http://goto.arcgisonline.com/maps/World\\_Terrain\\_Base](http://goto.arcgisonline.com/maps/World_Terrain_Base), last access: 17 January 2020).

386 From the temporal patterns of mean areal precipitation of IMERG, AIMERG, and ground  
 387 observations from hydrological stations, in Zhejiang province, 2010-2015 (Fig. 10), IMERG is general  
 388 larger than both AIMERG and ground observations. For instance, the IMERG significantly overestimates  
 389 the precipitation with up to ten times than that of AIMERG and ground observations, such as in the typical

390 periods, 0 a.m., June, 11 – 0 a.m., June, 14, 2015, and 0 a.m., Aug, 29 – 0 a.m., Sep, 1, 2015. Additionally,  
 391 both the temporal patterns and the magnitudes of AIMERG are almost same with those of ground  
 392 observations, compared with those of IMERG. Meanwhile, in some pentads with the heavy rain events,  
 393 both AIMERG and ground observations are larger than IMERG.



394

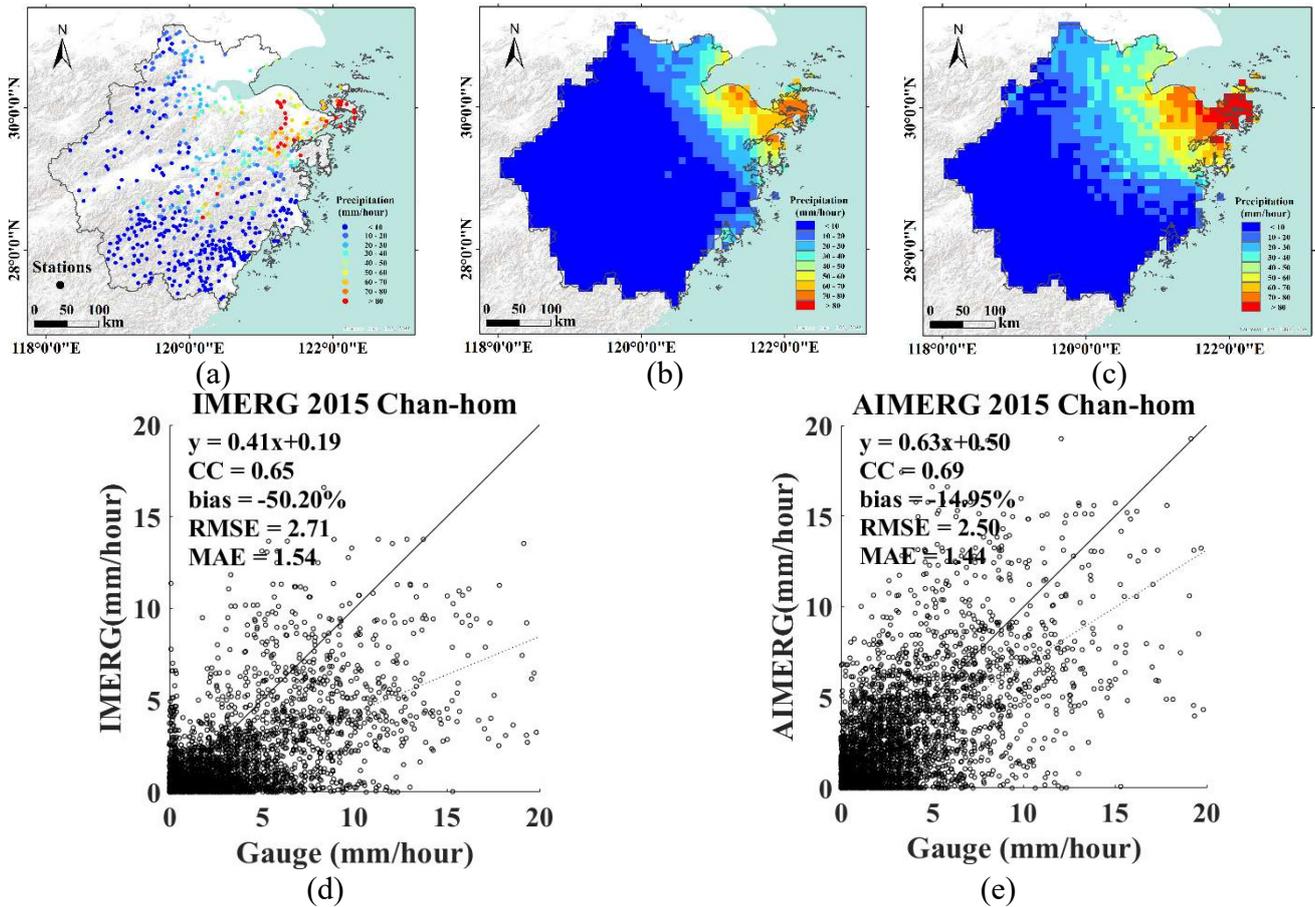
395 Figure 10. The temporal patterns of mean areal precipitation of IMERG, AIMERG, and the ground  
 396 observations from the independent hydrological stations, at daily/hourly scale, in Zhejiang province,  
 397 2010-2015.

398

399 One of the primary aims of the satellite-based precipitation estimates is to provide the high quality  
 400 precipitation information at hourly scale in the heavy rainfall events. Therefore, one typhoon event, Chan-  
 401 hom, is selected as an example for assessing the quality of the IMERG and AIMER in Zhejiang Province,

402 where is always threatened by the typhoons, shown in Fig. 11. Though the spatial patterns of IMERG and  
403 AIMERG are both similar to those of ground observations, IMERG still underestimates the precipitation,  
404 compared with AIMERG (Fig. 11 a-c). From the statistics, not only the systematic bias of IMERG (around  
405 -50%) is significantly improved, with bias of AIMERG around -10%, but also the random errors of  
406 IMERG (RMSE  $\sim$  2.7 mm/hour, MAE  $\sim$  1.5 mm/hour) are also reduced, compared with AIMERG (RMSE  
407  $\sim$  2.5 mm/hour, MAE  $\sim$  1.4 mm/hour), which meant the calibrations using APHRODITE on IMERG  
408 improved the abilities of original IMERG product to more accurately estimate the quantitative  
409 precipitation volumes, especially in heavy rainfall events (Fig. 11 c-d).

410



411 Figure 11. The typhoon, Chan-hom, is selected as an example for assessing the quality of the IMERG and  
 412 AIMER, occurred in the typical period 0 a.m., – 11 a.m., July, 11, 2015, in Zhejiang Province. The  
 413 background map used in this study was provided by Esri, USGS and NOAA  
 414 ([http://goto.arcgisonline.com/maps/World\\_Terrain\\_Base](http://goto.arcgisonline.com/maps/World_Terrain_Base), last access: 17 January 2020).

415

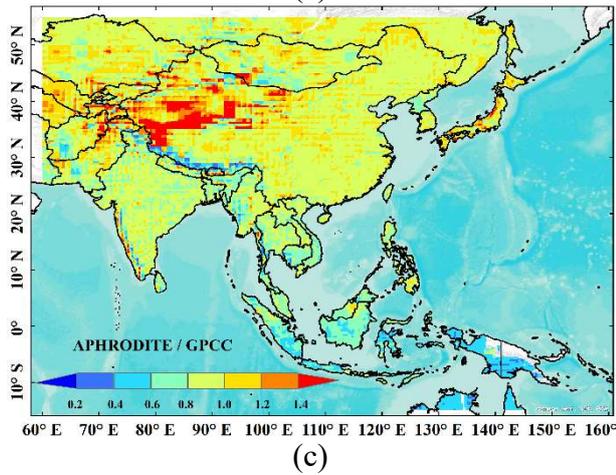
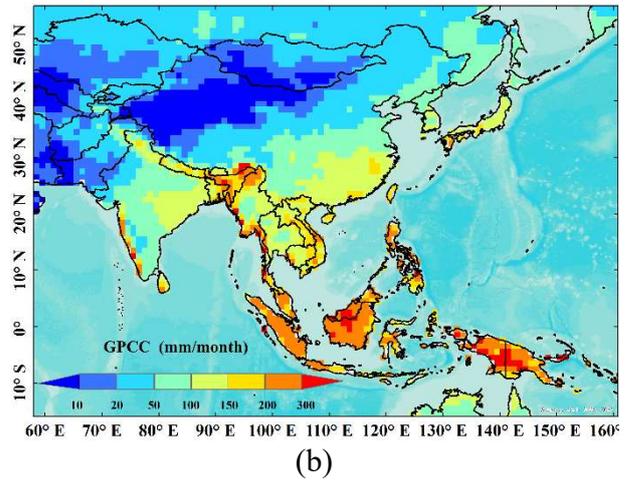
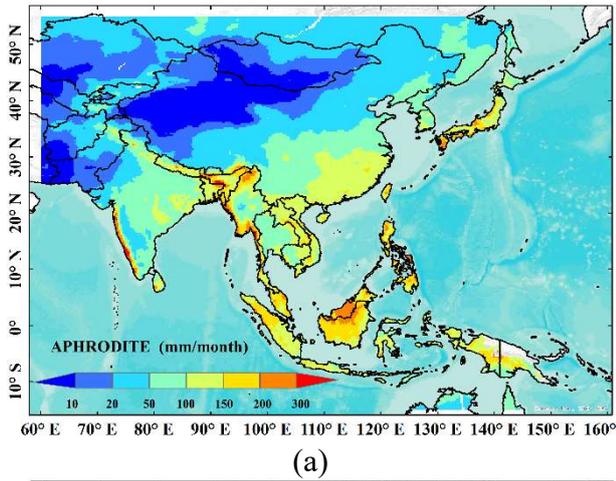
416 **5. Discussions**

## 417 **5.1. The potential drawbacks in processing the IMERG product**

418 From the document of “Algorithm Theoretical Basis Document (ATBD) Version 06” for  
419 generating the final IMERG product (Huffman et al., 2019a), we find that there are mainly two steps in  
420 the process: the first step is to derive the multi-satellite only precipitation inversion estimates, and the  
421 second step is to calibrate the satellite-based only precipitation estimates using the interpolated  
422 precipitation product based on ground observations, e.g., GPCC (1.0°/monthly). As lacking mature  
423 calibration algorithm for calibrating the multi-satellite-only precipitation estimates at daily scale, the  
424 current IMERG-Final product are only calibrated using the GPCC at monthly scale. The two aims of this  
425 study are to (1) provide a spatio-temporal calibration algorithm (DSTDCA) for anchoring the satellite-  
426 based precipitation estimates at daily scale, and (2) a new precipitation product with finer quality, namely  
427 AIMERG (half-hourly,  $0.1^\circ \times 0.1^\circ$ , 2000-2015, Asia) (Ma et al., 2020a, b), for Asian researcher. For  
428 anchoring the IMERG final product, we introduce the APHRODITE data (daily,  $0.25^\circ \times 0.25^\circ$ , 2000-  
429 2015, Asia), which were interpolated based on ground observations from the large numbers of rain gauges.  
430 Though the general spatial patterns of monthly mean precipitation estimates from both APHRODITE and  
431 GPCC, from 1951 to 2015, are similar, the volumes of them demonstrate significant differences,  
432 especially along the Himalayas, coastal Indochina and Western Ghats, and the Indonesia (Fig. 12 a-b).  
433 To much more clearly demonstrate the relative values of GPCC and APHRODITE, the spatial patterns of  
434 the ratio of monthly mean values of APHRODITE to those of GPCC are illustrated in Fig. 12 c, from

435 which we find that GPCC significantly overestimates the precipitation in the tropical rain range along the  
436 Indonesia, and along the southern Himalayas with complex terrain, while it significantly underestimates  
437 the precipitation in the north western Tibetan Plateau and Middle East, compared with the ground “truth”  
438 product, APHRODITE. Illustrated by Fig. 12, the GPCC plays vital roles for the final IMERG product,  
439 and the introduction of APHRODITE on calibrating the IMERG would be greatly benefiting the quality  
440 of the AIMERG.

441

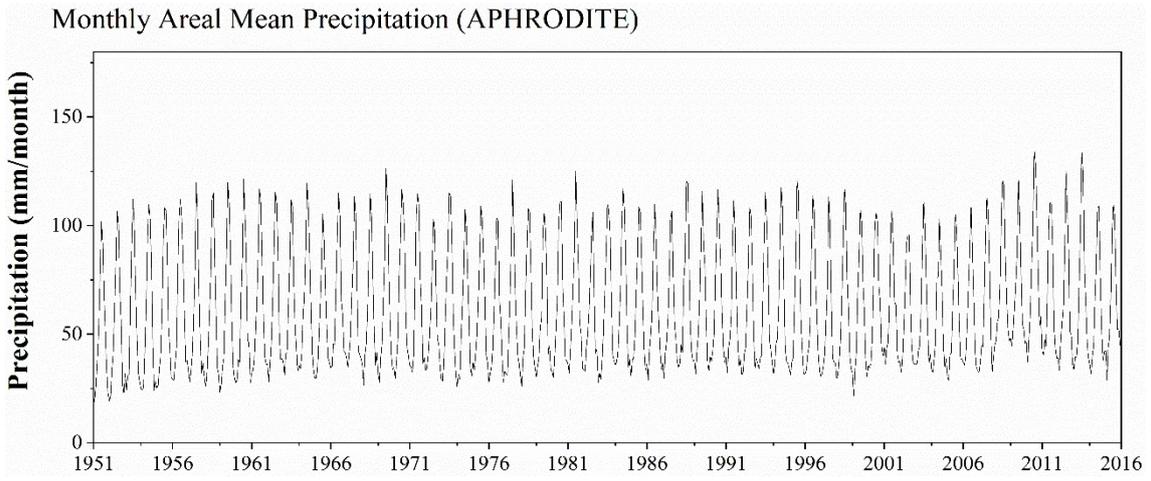


442 Figure 12. The spatial patterns of the monthly mean precipitation of (a) APHRODITE and (b) GPCC, and  
 443 (c) Ratios between monthly mean values of APHRODITE and GPCC, over the Asia in the period from  
 444 1951 to 2015. The background map used in this study was provided by Esri, USGS and NOAA  
 445 ([http://goto.arcgisonline.com/maps/World\\_Terrain\\_Base](http://goto.arcgisonline.com/maps/World_Terrain_Base), last access: 17 January 2020).

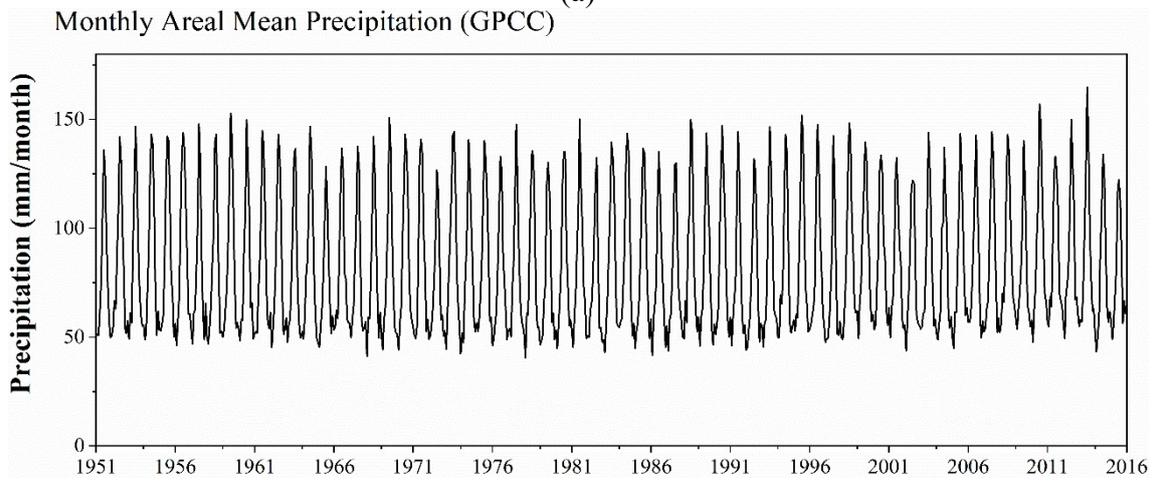
446 There are mainly two kinds of errors in the multi-satellite-only precipitation product, including  
 447 systematic bias and random errors (Shen et al., 2014). As seen in the above-mentioned results, the random

448 errors of the AIMERG are alleviated by using the APHRODITE data compared with IMERG (e.g., Fig.  
449 4-11). In terms of the systematic errors, we compared the monthly Asian mean precipitation estimates of  
450 both APHRODITE and GPCC, from 1951 to 2015 (Fig. 13). The monthly Asian mean precipitation of  
451 APHRODITE varies between  $\sim 25$  mm/month and  $\sim 100$  mm/month, while those of GPCC ranges from  
452  $\sim 50$  mm/month and  $\sim 150$  mm/month, which results the ratios of APHRODITE to GPCC fluctuate  
453 significantly from  $\sim 0.2$  to  $\sim 0.9$ , with average value  $\sim 0.7$ , which means that the GPCC at least  
454 overestimates the precipitation more than  $\sim 30\%$ , compared with the APHRODITE. Therefore, the  
455 introduction of APHRODITE data would greatly reduce the systematic errors of the IMERG final product,  
456 over the Asia.

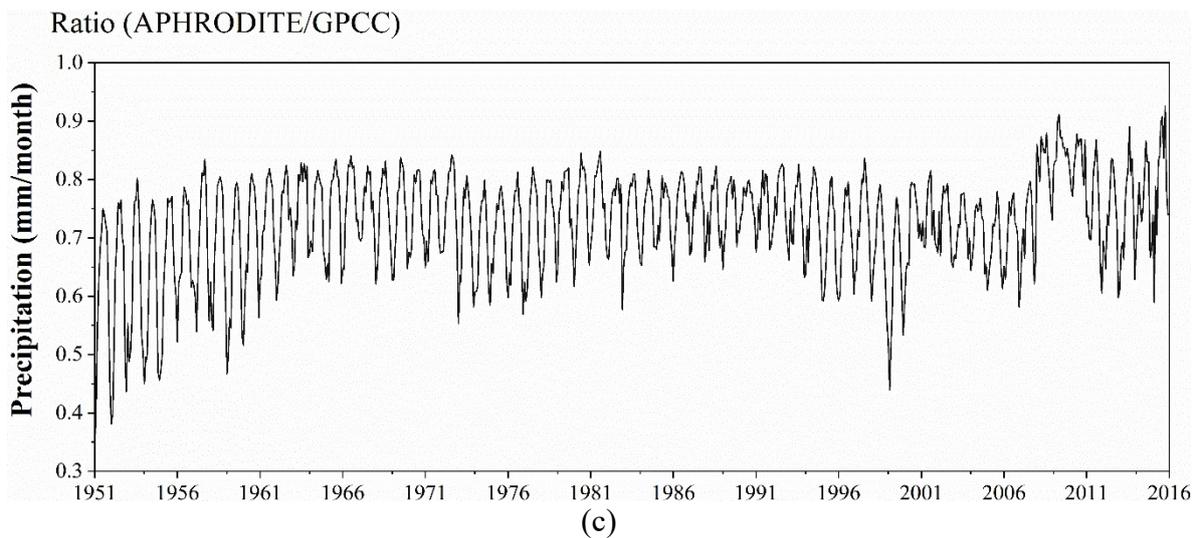
457



(a)



(b)



458 Figure 13. The temporal patterns of monthly areal mean precipitation of (a) APHRODITE, (b) GPCC,  
 459 and (c) APHRODITE/GPCC, 1951-2015.

460 **5.2. The controls on the range of the spatial weights based on IMERG**

461 As demonstrated in the document of the “ATBD”, gauge information is introduced into the original  
 462 multi-satellite-only half-hourly data to generate the final IMERG product. Firstly, the ratio between the  
 463 monthly accumulation of half-hourly multi-satellite-only field and the monthly satellite-gauge field is  
 464 calculated, then each half-hourly field of multi-satellite-only precipitation estimates in the corresponding  
 465 month is multiplied by the ratio field to generate the half-hourly calibrated IMERG. After various  
 466 experiments, the ratio values between the monthly satellite-gauge and the monthly accumulation of half-  
 467 hourly multi-satellite-only fields is limited to the range [0.2, 3] (Huffman et al., 2019a). The cap of 3 is  
 468 decided due to the value of 2 (used in TRMM V6) was too restrictive. Additionally, the cap of 3 was

469 finally applied because it performed better in matching the two accumulations than that of other larger  
470 values, for instance, the cap of 4 resulted in introducing unrealistic shifts to histogram of half-hourly  
471 precipitation rates for the month. Additionally, early in TRMM the lower bound of 0.5 was applied, which  
472 suggested a smaller value of the lower bound allows matching between the two accumulations without  
473 creating the egregious high snapshot values when the upper bound was expanded too far.

474         Inspired by the range of the ratio values between the monthly satellite-gauge and the monthly  
475 accumulation of half-hourly multi-satellite-only fields in generating IMERG, we consider the range [0,  
476 1.5] of the daily spatial disaggregation weights in this study is reasonable after careful checking the  
477 distributions of spatial disaggregation weights. The lower bound of 0 was selected based on the  
478 consideration if the IMERG did not capture the daily precipitation event, then the spatial disaggregation  
479 weight is still equal to zero, which agrees as most as possible to the original IMERG. While there are at  
480 least two reasons for setting the upper bound of the spatial disaggregation weights as 1.5: (1) most  
481 numerical values of spatial disaggregation weights are in the range [0, 1.5], and (2) there are obvious  
482 anomalies in the final calibrated AIMERG, especially along the coastal regions and edges of the specific  
483 precipitation event coverages, where the values of the spatial disaggregation weights are larger than 1.5.  
484 Though the range [0, 1.5] of spatial disaggregation weights was applied to obtain the final AIMERG in  
485 this study, we also consider that this is still an open-ended question.

486 **5.3. The advantages of APHRODITE data in anchoring the multi-satellite-only precipitation**  
487 **product**

488 It has been a great challenge to obtain precipitation estimates over the Tibetan Plateau and its  
489 surroundings, as there are very limited ground observations in this region, especially in its western parts  
490 (Ma et al., 2017). Incorporating a uniform precipitation gauge analysis is important and critical for  
491 controlling the bias that typifies the satellite precipitation estimates, e.g., using GPCC for TMPA and  
492 IMERG (Huffman et al., 2019a). Those projects (e.g., GPCC, TRMM, GPM) demonstrate that even  
493 monthly gauge analyses contribute significant improvements on the satellite-only precipitation estimates,  
494 at least for some regions in some seasons. Primarily explorations at CPC suggested substantial  
495 improvements in the bias corrections using daily gauge analysis, especially for regions, where there is a  
496 dense network of gauges (Mega et al., 2014). Foreseeably, GPM would try their best to calibrate the GPM  
497 multi-satellite only precipitation estimates at finer spatio-temporal scales (e.g., 0.25°/daily) worldwide.

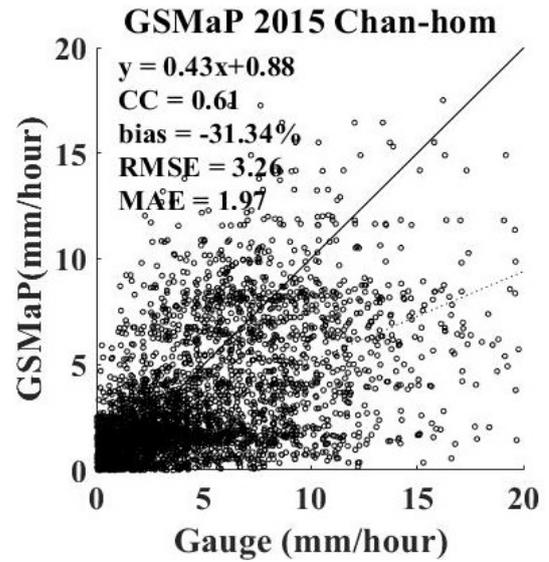
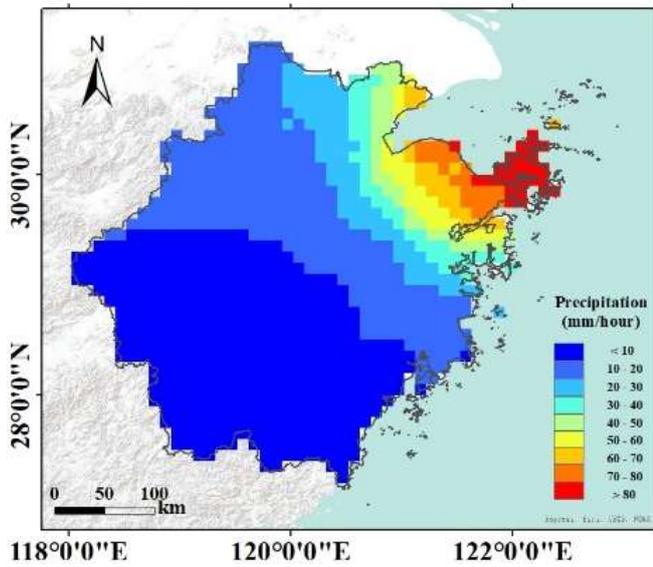
498 Currently, GPCC has been adopted to calibrate the TRMM TMPA and GPM IMERG at monthly  
499 scale. The Deutscher Wetterdienst (DWD) Global Precipitation Climatology Centre (GPCC) was  
500 established in 1989 to provide high-quality precipitation analyses over land based on conventional  
501 precipitation gauges from ~7,000-8,000 stations world-wide (Schneider et al. 2014, 2018). And two  
502 GPCC products were applied in the IMERG, the V8 Full Data Analysis for the majority of the time  
503 (currently 1998-2016), and the V6 Monitoring Product from 2017 to the then-present. Compared with

504 GPCC, APHRODITE has inherently advantages with significantly larger numbers of ground observations  
505 and finer spatio-temporal resolutions, over the Asia. APHRODITE projects aimed at collecting as most  
506 gauge information as possible from the Asian countries. There are mainly three kinds of gauge  
507 information sources used in APHRODITE analysis, the GTS-based data, data precompiled by other  
508 projects or organizations, and APHRODITE's own collection. More detailed information on the  
509 APHRODITE' data sources could be found at the website (<http://www.chikyu.ac.jp/precip/>) and the  
510 research of Yatagai (2012). Compared with the GPCC with the limited ground observations in and around  
511 the Tibetan Plateau in China, the neighboring countries provide plenty of ground observations in the  
512 APHRODITE data, in mountainous regions, and semi-arid and arid regions. Additionally, the spatio-  
513 temporal resolutions of APHRODITE (0.25°/daily) are finer than those of GPCC (1.0°/monthly).  
514 Therefore, APHRODITE has significant advantages in calibrating the IMERG data at daily scale.

#### 515 **5.4. Quantitatively and horizontally comparisons with other high resolution precipitation product**

516 Recently, Tang et al (2020) has conducted a comprehensive comparison of GPM IMERG with  
517 other nine state-of-the-art high resolution precipitation products, six satellite-based precipitation products  
518 (TRMM 3B42, 0.25°/3 hour; CMORPH, 0.25°/3 hour; PERSIANN-CDR, 0.25°/1 day; GSMaP 0.1°/1  
519 hour; CHIRPS, 0.05°/1 day; SM2RAIN, 0.25°/1 day) and three reanalysis datasets (ERA5, ~0.25°/1 hour;  
520 ERA-Interim, ~0.75°/3 hour; MERRA2~0.5° × 0.625°/1 hour) from 2000 to 2018, and found that the  
521 IMERG product generally outperformed other datasets, except the Global Satellite Mapping of

522 Precipitation (GSMaP), which was adjusted at the daily scale by the gauge analysis ( $0.5^\circ$ /daily) from the  
523 CPC (Mega et al., 2014). Therefore, we have compared the AIMERG with GSMaP, in case of the typhoon  
524 Chan-hom, which is coordinated with those in Figure 11. As shown in Fig. 14, though the spatial patterns  
525 of the GSMaP are similar with those of the AIMERG, the AIMERG provides much more details than  
526 GSMaP, especially over the northeastern Zhejiang Province. Meanwhile, AIMERG significantly  
527 overwhelms GSMaP in terms of both bias and random errors. For instance, GSMaP underestimates the  
528 precipitation (bias  $\sim -31\%$ ) twice as large as AIMERG (bias  $\sim -15\%$ ), and the random errors of GSMaP  
529 (MAE  $\sim 1.97$  mm/hour, RMSE  $\sim 3.26$  mm/hour) are also significantly larger than those of AIMERG  
530 (MAE  $\sim 1.44$  mm/hour, RMSE  $\sim 2.50$  mm/hour). Compared with the original IMERG in Figure 11,  
531 though the random errors of GSMaP are relatively larger, the bias of GSMaP ( $\sim -31\%$ ) is significantly  
532 smaller than that of the original IMERG ( $\sim -50\%$ ), which owes to the calibrations on the GSMaP at the  
533 daily scale. In future, we also encourage researchers to comprehensively evaluate and compare the  
534 AIMERG with other high resolution precipitation products at various spatio-temporal scales.



535 Figure 14. The typhoon, Chan-hom, is selected as an example for assessing the quality of the Gauge  
 536 adjusted GSMaP, occurred in the typical period 0 a.m., – 11 a.m., July, 11, 2015, in Zhejiang Province,  
 537 which is coordinated with those in Figure 11. The background map used in this study was provided by  
 538 Esri, USGS and NOAA ([http://goto.arcgisonline.com/maps/World\\_Terrain\\_Base](http://goto.arcgisonline.com/maps/World_Terrain_Base), last access: 17 January  
 539 2020).

540 The extent of the AIMERG could cover the Northern Eurasia, Middle East, Monsoon Asia, and  
 541 Japan. This study mainly evaluated the AIMERG in the China Mainland, which calls for Asia wide  
 542 evaluations in the future to assess both the algorithm and the corresponding precipitation product.

543 **6. Data Availability**

544 The AIMERG data record (0.1°/half-hourly, 2000-2015, Asia) is freely available at <http://argi->  
545 [basic.hihanlin.com:8000/d/d925fecf60/](http://basic.hihanlin.com:8000/d/d925fecf60/). Additionally, the AIMERG data is also freely accessible at  
546 <https://doi.org/10.5281/zenodo.3609352> (for the period from 2000 to 2008) (Ma et al., 2020a) and  
547 <http://doi.org/10.5281/zenodo.3609507> (for the period from 2009 to 2015) (Ma et al., 2020b).

## 549 7. Conclusions

550 As the milestone in the satellite-based precipitation measurement process, the TRMM and its  
551 successor GPM generate the most popular and the state-of-the-art satellite precipitation products for both  
552 water cycle related scientific researches and applications, TMPA (1998-present, 0.25°/3 hourly) and  
553 IMERG (2014-present, 0.1°/half-hourly), as well as the retrospective IMERG (2000-present, 0.1°/half-  
554 hourly) from GPM era to TRMM era. In this study, focusing on the potential drawbacks in generating  
555 IMERG and its recently updated retrospective IMERG (finished in July, 2019), which were only  
556 calibrated at monthly scale using limited ground observations, GPCC (1.0°/monthly), resulting the  
557 IMERG with large systematic bias and random errors, we introduce another daily gauge analysis product,  
558 APHRODITE (Last update October 5, 2018), to calibrate the IMERG at 0.25°/daily scale. Compared with  
559 GPCC, APHRODITE has inherently advantages with significantly larger numbers of ground observations  
560 and finer spatio-temporal resolutions (0.25°/daily), over the Asia.

561 We have proposed a new algorithm (Daily Spatio-Temporal Disaggregation Calibration Algorithm,  
562 DSTDCA) for calibrating IMERG at daily scale, and provided a new AIMERG precipitation dataset  
563 (0.1°/half-hourly, 2000-2015, Asia) (Ma et al., 2020a, b) with better quality, calibrated by APHRODITE  
564 at daily scale for the Asian applications. And the main conclusions include but not limited to: (1) the  
565 proposed daily calibration algorithm is effective in considering the advantages from both satellite-based  
566 precipitation estimates and the ground observations; (2) AIMERG performs better than IMERG at  
567 different spatio-temporal scales, in terms of both systematic biases and random errors, over the China  
568 Main land; and (3) APHRODITE demonstrates significant advantages than GPCP in calibrating the  
569 IMERG, especially over the mountainous regions with complex terrain, e.g., the Tibetan Plateau.  
570 Additionally, results of this study suggests that it is a promising and applicable daily calibration algorithm  
571 for GPM in generating the future IMERG in either operational scheme or retrospective manner.

572

### 573 **Author Contributions**

574 Dr. Ziqiang Ma designed and organized the manuscript. Drs. Jintao Xu, Siyu Zhu, Jun Yang and  
575 Yuanjian Yang prepared the related materials and run the models for generating AIMERG and the related  
576 assessments. Dr. Guoqiang Tang and Prof. Zhou Shi made contributions on the scientific framework of  
577 this study and discussed the interpretation of results. Prof. Yang Hong co-advised this study. All authors  
578 discussed the results and commented on the manuscript.

579

580 **Competing interests**

581 The authors declare they have no competing financial interests.

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591 Meteorological Data Sharing Service System (<http://cdc.nmic.cn/home.do>), the APHRODITE data  
592 provider (<http://aphrodite.st.hirosaki-u.ac.jp/download/>), and the IMERG data provider  
593 (<https://pmm.nasa.gov/data-access/downloads/gpm>).

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725 **Appendix A: Acronyms with definitions used in this study.**

AIMERG	Asian precipitation dataset by calibrating GPM IMERG at daily scale using APHRODITE
APHRODITE	Asian Precipitation Highly Resolved Observational Data Integration Towards Evaluation of Water Resources
ATBD	Algorithm Theoretical Basis Document
BIAS	Relative Bias
CC	Correlation Coefficient
CHIRPS	Climate Hazards group Infrared Precipitation with Stations
CLIMAT	Monthly Climatological Data
CMA	Chinese Meteorological Administration
CMORPH	Climate Prediction Center (CPC) MORPHing technique
CPC	Climate Prediction Center
CSI	Critical Success Index
DSTDCA	Daily Spatio-Temporal Disaggregation Calibration Algorithm
DWD	Deutscher Wetterdienst
ERA5	Fifth generation of ECMWF atmospheric reanalyses of the global climate
ERA-Interim	ECMWF ReAnalysis Interim
FAR	False Alarm Ratio
GEWEX	Global Energy and Water Exchange
GPCC	Global Precipitation Climatology Centre

GPM	Global Precipitation Measurement
GSMaP	Gauge-adjusted Global Satellite Mapping of Precipitation V7
GTS	Global Telecommunications System
IMERG	Integrated Multi-satellitE Retrievals for GPM
IR	Infrared
ME	Mean Error
MERRA2	The Modern-Era Retrospective Analysis for Research and Applications, Version 2
MW	Microwave
NHMs	National hydrological and meteorological services
NMIC	National Meteorological Information Center
OI	Optimal Interpolation
PDF	Probability Density Function
PERSIANN	Precipitation Estimation from Remotely Sensed Information using Artificial Neural Networks
PERSIANN- CCS	Precipitation Estimation from Remotely Sensed Information using Artificial Neural Networks-Cloud Classification System
PERSIANN- CDR	PERSIANN-Climate Data Record
PMW	Passive Microwave
POD	Probability of Detection

QC	Quality Control
RMSD	Root-mean-square Deviation
RMSE	Root Mean Square Error
SG	Satellite-Gauge
SM2RAIN	Soil Moisture to RAIN based on ESA Climate Change Initiative (CCI)
SYNOP	Synoptic Weather Report
TMPA	TRMM Multi-satellite Precipitation Analysis
TRMM 3B42	Tropical Rainfall Measuring Mission Multi-satellite Precipitation Analysis 3B42 V7