

# Historical and recent aufeis, the Indigirka river basin (Russia)

\*Olga Makarieva<sup>1,2</sup>, Andrey Shikhov<sup>3</sup>, Nataliia Nesterova<sup>2,4</sup>, Andrey Ostashov<sup>2</sup>

<sup>1</sup>*Melnikov Permafrost Institute of RAS, Yakutsk*

<sup>2</sup>*St. Petersburg State University, St. Petersburg*

<sup>3</sup>*Perm State University, Perm*

<sup>4</sup>*State Hydrological institute, St. Petersburg*

*RUSSIA*

\*omakarieva@gmail.com

**Abstract:** A detailed spatial geodatabase of aufeis (or naleds in Russian) within the Indigirka River watershed (305 000 km<sup>2</sup>), Russia, was compiled from historical Russian publications (year 1958), topographic maps (years 1970–1980's), and Landsat images (year 2013-2017). Identification of aufeis by late-spring Landsat images was performed with a semi-automated approach according to Normalized Difference Snow Index (NDSI) and additional data. After this, a cross-reference index was set for each aufeis, to link and compare historical and satellite-based aufeis data sets.

The aufeis coverage varies from 0.26 to 1.15% in different sub-basins within the Indigirka River watershed. The digitized historical archive (Cadastre, 1958) contains the coordinates and characteristics of 896 aufeis with total area of 2064 km<sup>2</sup>. The Landsat-based dataset included 1213 aufeis with a total area of 1287 km<sup>2</sup>. Accordingly, the satellite-derived total aufeis area is 1.6 times less than the Cadastre (1958) dataset. However, more than 600 aufeis identified from Landsat images are missing in the Cadastre (1958) archive. It is therefore possible that the conditions for aufeis formation may have changed from the mid-20th century to the present.

Most present and historical aufeis are located in the elevation band of 1000 – 1200 m. About 60% of total aufeis area is represented by just 10% of the largest aufeis. Interannual variability of aufeis area for the period of 2001-2016 was assessed for the Bolshaya Momskaya aufeis and for a group of large aufeis (11 aufeis with a areas from 5 to 70 km<sup>2</sup>) in the basin of the Syuryuktyakh River. The results of this analysis indicate a tendency towards an area decrease in the Bolshaya Momskaya aufeis in recent years, while no reduction in Syuryuktyakh River aufeis area was observed.

The combined digital database of the aufeis is available at <https://doi.pangaea.de/10.1594/PANGAEA.891036>.

**Keywords:** aufeis, Indigirka, Landsat, NDSI, Cadastre, Cadastral map, Bolshaya Momskaya aufeis

## 1. Introduction

Aufeis (naleds in Russian, icings in English) are accumulations of ice that are formed by freezing underground and surface waters on the surface of the earth or ice along streams and river valleys in arctic and subarctic regions. They affect water exchange and economic activity (Alekseev, 1987). Aufeis are found in permafrost regions such as Alaska (Slaughter, 1982), Siberia (Alekseev, 1987), Canada (Pollard, 2005), Greenland (Yde and Knudsen, 2005) and others (Yoshikawa et al., 2007). Aufeis formation can result in significant economic expenses as aufeis may negatively affect infrastructure and therefore natural resource extraction (Aufeis of Siberia..., Nauka, 1981). Moreover, the springs that often feed aufeis may in some cases be the

47 only source of water for remote communities (Simakov, Shilnikovskaya, 1958). In Russia, aufeis  
48 are found in the North-East, Transbaikal region, Yakutia, and West Siberia. Sokolov (1975)  
49 estimated that the total aufeis water storage in Russia to be at least 50 km<sup>3</sup>, which approximately  
50 equals the Indigirka River total annual streamflow.

51 The main hydrological role of aufeis is the seasonal redistribution of the groundwater  
52 component of river runoff, where the winter groundwater discharge is released to summer  
53 streamflow through melting of aufeis (Surface water resources, 1972). In most cases, the share of  
54 the aufeis component in a river's annual streamflow accounts for 3-7%, reaching 25-30% in  
55 particular river basins with an extremely large proportion of aufeis (Reedyk et al., 1995; Kane &  
56 Slaughter, 1973; Sokolov, 1975). The most significant water inflow from aufeis melting takes  
57 place in May-June (Sokolov, 1975). For example, the share of the aufeis flow accounts for more  
58 than 11% of total annual streamflow at the Indigirka River (gauging station Yurty, 51 100 km<sup>2</sup>).  
59 In May, aufeis melt may represent 50% of monthly total streamflow, but decreases in June to  
60 35% (Sokolov, 1975).

61 It is important to understand how climate change may impact aufeis formation because  
62 warming has been observed in this region causing the transformation of permafrost  
63 (Romanovsky et al., 2007), glaciers reduction (Ananicheva, 2014) and hydrological regime  
64 changes (Bring et al., 2016; Makarieva et al., 2018). Aufeis are formed by a complex connection  
65 between river and groundwater. Many studies have reported the increase of minimum flow in  
66 Arctic rivers (Rennermalm and Wood, 2010; Tananaev et al., 2016), including those where  
67 aufeis are observed in abundance (Makarieva et al., 2018, in review). A widely accepted  
68 hypothesis for permafrost regions is that a warming climate increases the connection between  
69 surface- and groundwater that in turn leads to the increase of streamflow, both in cold seasons  
70 and in annual flow (Bense et al., 2012; Ge et al., 2011; Walvoord et al., 2012; Walvoord and  
71 Kurylyk, 2016). Variation and changes in aufeis extent can be assessed using remote sensing  
72 techniques, where aufeis dynamics can serve as an indicator of groundwater change that is  
73 otherwise difficult to observe (Topchiev, 2008; Yoshikawa et al., 2007).

74 The understanding of how aufeis respond to a warming climate varies. Alekseev (2016)  
75 suggests three to 11 year up and down cycles of aufeis maximum annual size, which may vary  
76 up to 25-30% in comparison with long-term average values. However, the same author  
77 (Alekseev, 2016) states a general tendency to the decrease of aufeis volume for the last 50-60  
78 years in some aufeis-affected areas of Russia such as the Baikal region, South Yakutia, Kolyma  
79 region, Eastern Sayan Mountains, following the increase of global and local air temperature.

80 Some authors suggest that degradation of permafrost in the discontinuous and sporadic  
81 permafrost regions will lead to the decrease of the number of aufeis and even an almost complete  
82 disappearance. Meanwhile, in the zone of continuous permafrost in North-East Siberia, a climate  
83 warming of 2-3 °C is not projected to lead to significant changes in permafrost extent, but will  
84 increase the number and size of both through- and open taliks by the end of the 21st century  
85 (Pomortsev et al., 2010). Such a scenario may result in the reduction of area of large aufeis and  
86 formation of new small aufeis (Pomortsev et al., 2010).

87 In Alaska as well, no significant changes were documented in the area and volume of aufeis  
88 over the past few decades or even a century (Yoshikawa et. al, 2007). They suggested that the  
89 formation and melting of ice is less dependent on climate and more so on the source (spring)  
90 water properties such as temperature and volume.

91 In 1958, Simakov and Shilnikovskaya (1958) compiled and published a map inventory of  
92 aufeis of the North-East USSR (scale 1:2 000 000). Since then, there has been no update on the  
93 information on aufeis in this region, apart from some specific studies. In 1980-1982, an  
94 inventory of aufeis in the zone of the Baikal-Amur Mainline was published (Catalog of  
95 Aufeis..., 1980, 1981, 1982). Markov et al. (2017) summarized the results of field studies on  
96 aufeis in the southern mountain taiga of Eastern Siberia from 1976 to 1983. Grosse and Jones  
97 (2011) compiled the spatial geodatabase of frost mounds (or pingos) for northern Asia from  
98 topographic maps. Further, the glacier science community has mapped past and recent glacier

99 cover across the globe (GLIMS and NSIDC, 2005, updated 2017). However, as far as the authors  
100 are aware, no electronic catalogue of aufeis exists.

101 The aim of this study is to update the inventory of aufeis in the North-East of Russia using  
102 Landsat images, as well as to develop an electronic catalogue, which will contain data on historic  
103 and current location and characteristics of aufeis. Here we present work that has been completed  
104 for the Indigirka River basin (down to the Vorontsovo gauging station, 305 000 km<sup>2</sup>).

105 The new database, which includes geographic information system (GIS) formatted files, is  
106 freely available (Makarieva et al., 2018) and can be used both for both scientific purposes and  
107 for solving practical problems such as engineering construction and water supply studies.

## 108 2. Study region

109 The study region is the Indigirka River basin, which is located in Northeastern Siberia and  
110 covers an area of 305 000 km<sup>2</sup> (Fig. 1). Most of the basin is represented by highlands with a  
111 number of mountain ranges (< 3 003 m) including the Cherskiy and Suntar-Khayata mountains.  
112 The lowland elevation reaches heights up to 350 m.

113 The climate of the study area is distinctly continental with annual average and lowest  
114 monthly air temperature varying from -16.1 and -47.1 °C, respectively, at the Oymyakon  
115 meteorological station (726 m, 1930-2012) to -13.1 and -33.8 °C, respectively, at the  
116 Vostochnaya station (1 288 m, 1942-2012). Most precipitation (over 60%) occurs in the summer  
117 season. Average annual precipitation at the Oymyakon weather station is 180 mm and at the  
118 Vostochnaya station 278 mm.

119 The Indigirka River basin is located in the zone of continuous permafrost. Permafrost  
120 depth can reach 450 m in the mountains, up to 180 m in river valleys and intermountain areas,  
121 with taliks found in river beds and fractured deposits. The hydrogeological regime is affected by  
122 the active layer, which varies from 0.3 m to over 2 m (Explanatory note ..., 1991). The river  
123 runoff regime is characterized by high snowmelt freshet, summer-autumn rainfall floods, and  
124 low winter flow. In winter, small- and medium-sized rivers completely freeze. Freshet starts in  
125 May-June and lasts for approximately 1.5 months. Melt waters from aufeis, glaciers, and snow  
126 patches add to the river discharge in summer.

127 In total, about 10 000 aufeis with a total combined area of about 14 000 km<sup>2</sup> (Sokolov,  
128 1975) are known in North-East Russia. The watershed area covered by aufeis varies from 0.4 to  
129 1.3%, reaching 4% in some river basins (Tolstikhin, 1974). Most aufeis are of ground water  
130 origin; significantly less often they are formed out of river waters or are of a mixed type  
131 (Tolstikhin, 1974).

## 133 3. Materials and methods

### 134 3.1 The database of aufeis based on the Cadastre (1958) and topographic maps

135 The inventory map (scale 1:2 000 000) and the Cadastre of aufeis of the North-East of the  
136 USSR (Simakov, Shilnikovskaya, 1958), hereinafter referred to as the Cadastral Map and the  
137 Cadastre, became the first summarizing quantitative work on aufeis within the territory. The  
138 effort was carried out in the framework of the Central complex thematic expedition of the North-  
139 East Geological Survey of the USSR.

140 The Cadastre contains data on 7 448 aufeis of different size and over 2 000 boolgunyakhs  
141 (frost mounds). Of the total number of aufeis, 7 006 are plotted based on air-photo interpretation  
142 data, and another 442 on geological reports from field data. It should be noted that aufeis were  
143 identified based on geomorphologic features, meaning that in some cases only the areas or river  
144 valleys with aufeis were identified but not aufeis themselves.

145 In the Cadastre (1958) and our digitalization, the following characteristics of the aufeis are  
146 presented: location (the name of the river, the distance from the mouth or source), size  
147 (maximum length, average width, and area) and the dates of ice recording in aerial images  
148 (ranging from 08.06.1944 to 27.09.1945). Areas of the aufeis were evaluated via planimetry.

149 Only very large aufeis ( $> 3.3 \text{ km}^2$ ) were plotted on the Cadastral Map (1958), while the  
150 others are shown as point locations. Each aufeis on the Cadastral Map (1958) has its  
151 corresponding number, whose identifier and corresponding information can be found in the  
152 Cadastre (1958). As noted by Simakov and Shilnikovskaya (1958), some very small aufeis  
153 ( $< 0.01 \text{ km}^2$ ) could have been missed due to their indecipherability on aerial images, or they  
154 might have already melted by the time of the aerial photography. The example of the Cadastral  
155 Map's sheet (1958) for the Indigirka River upper reaches is presented in fig. 2.

156 Here, we developed the GIS database of aufeis in the Indigirka River basin up to the cross-  
157 section at the Vorontsovo gauging station based on the Cadastre (1958) and topographic maps.  
158 Our compilation contains data on 896 aufeis. The aufeis are presented as point objects in our  
159 database. The areas are specified for only 808 aufeis. The total area of all the aufeis with  
160 specified area accounts for  $2063.6 \text{ km}^2$  and the areas of individual aufeis vary from  $0.01$  to  $82$   
161  $\text{km}^2$ .

162 In the Cadastre, the dates of ice recording for 592 aufeis (66%) are presented, based on  
163 aerial images within the study area. The average seasonal date of recording is August 2, ranging  
164 from June 8 to September, 27. The dates of ice recording for the remaining 34% of the aufeis  
165 were not described, meaning that aufeis detection could be carried out based not on the visible  
166 ice presence at the aerial images but on geomorphological features of river valleys. Therefore,  
167 the Cadastre might contain data on old aufeis glades, where the aufeis themselves were absent.

168 Spatial positioning of the Cadastral Map of aufeis was conducted using the location  
169 description by Russian topographic maps with the scale of 1:200 000. Grosse and Jones (2011)  
170 used the same set of maps for compiling the dataset of pingos (frost mounds) in northern Asia  
171 and described those maps in details therein. The maps at 1:200 000 scale were based on more  
172 detailed maps of 1:50 000 and 1:100 000 scale, which were derived from aerial photography  
173 acquired in the 1970–1980's. The use of 1: 200 000 scale guarantees the position assessment  
174 precision to within 100 m. Each map sheet was visually searched for aufeis and identified aufeis  
175 were marked with an area polygon in a GIS layer. The locations of 330 aufeis (area  $358 \text{ km}^2$ )  
176 were determined based on topographic maps. When digitized, a point was plotted in the middle  
177 of an aufeis at a topographic map.

178 The locations of the remaining aufeis were determined with the positioned map of the  
179 Cadastre. Additionally, 11 aufeis were found, which were absent in the Cadastre, but present in  
180 the topographic maps. Aufeis areas were estimated by digitalization of the maps. Areas of the  
181 remaining aufeis were estimated with the Cadastre. It was not possible to estimate the area of 88  
182 aufeis, as they were not drawn on the topographic maps and only their location, but not area, was  
183 stated in the Cadastre.

184 Table 1 contains the structure of the GIS dataset of aufeis according to the Cadastre.

### 185 **3.2 Identification of aufeis based on Landsat data**

186 Aufeis location and area are relatively easy to determine using Landsat and/or Sentinel-2  
187 images, received immediately after snow cover melting. Snow and ice are known to be  
188 characterized by relatively high reflectance in the visible and near infrared spectral bands and its  
189 significant decrease in mid infrared band. Normalized Difference Snow Index (NDSI) is based  
190 on this pattern and is calculated according to the formula (Hall et al., 1995):

$$191 \quad NDSI = (GREEN - SWIR1) / (GREEN + SWIR1)$$

192 where SWIR1 is reflectance in mid infra-red band ( $1.56 - 1.66 \mu\text{m}$  for the Landsat-8 images),  
193 and GREEN is reflectance in the green band ( $0.525 - 0.6 \mu\text{m}$  for the Landsat-8 images).  
194 Following Hall et al. (1995), the threshold value for snow and ice is set at 0.4. Apart from using  
195 NDSI, other indices have been suggested to detect aufeis by Landsat images (but not used here).  
196 These are Normalized Difference Glacier Index (NDGI) and Maximum Difference Ice Index  
197 (MDII). Their advantages and disadvantages are discussed by Morse and Wolfe (2015).

198 Landsat-based detection of aufeis required some additional data to exclude other surface  
199 types with similar spectral characteristics, such as snow-covered areas, turbid water, etc. It is

200 problematic to separate floodplain lakes from aufeis by late-spring satellite images, because  
201 many of these lakes are still ice-covered in May-June. Morse and Wolfe (2015) recommended  
202 creating a mask of water surface by mid-summer images (when all water bodies are already not  
203 covered by ice), to exclude them from further analysis.

204 Aufeis detection in the Indigirka River basin was carried out based on the Landsat-8 OLI  
205 satellite images, 2013-2017, downloaded from the United States Geological Survey web-service  
206 (<https://earthexplorer.usgs.gov>). We used Landsat 8 collection 1 level-one terrain-corrected  
207 product (L1T) with radiometric and geometric corrections. In total, 33 images completely  
208 covering the Indigirka river basin were processed. We selected late-spring images (between 15  
209 May and 18 June), to detect the maximum possible number of aufeis, since in June they melt  
210 intensively. There was between 1-20% of cloudiness in some images.

211 Preprocessing of the images was performed with the use of Semi-Automatic Classification  
212 Plugin module (QGIS 2.18). It includes the calculation of surface reflectance and atmospheric  
213 correction by Dark Object Subtraction (DOS1) image-based algorithm, described by (Chavez,  
214 1996).

215 The Aufeis detection algorithm was realized in ArcGIS with the help of the ModelBuilder  
216 application. Apart from the Landsat images, the digital terrain model (DTM) GMTED2010  
217 (Danielson and Gesch, 2011) with a spatial resolution of 250 m was used to build a network of  
218 thalwegs within the study basin. This is essential for semi-automated separation of the aufeis  
219 from snow-covered areas in late-spring Landsat images. Indeed, almost all aufeis are located  
220 either at streams or thalwegs, or in immediate proximity to them. On the contrary, the snow  
221 cover in late spring mainly remains on mountain ridges and other elevated locations, i.e.  
222 relatively far from thalwegs. Based on the preliminary analysis of aufeis location in relation to  
223 the created network of thalwegs, we found that a 1.5 km wide buffer zone around the thalwegs  
224 covers almost all aufeis. So, snow and ice covered areas, which are located outside this buffer,  
225 are excluded from further analysis.

226 The process of aufeis detection by Landsat images consisted of the following steps:

- 227 • Detection of snow-ice bodies with the NDSI threshold of 0.4.
- 228 • Creation of a water mask with threshold values of the Normalized Difference Water  
229 Index (NDWI) (taken equal to 0.3), and reflectance in the near-infrared band (taken equal  
230 to 0.04).
- 231 • Extraction of the detected snow-ice bodies by the buffer zone around thalwegs (1.5 km  
232 wide).
- 233 • Conversion to vector format, area calculation and removal of objects smaller than 5  
234 Landsat pixels (0.45 ha).

235 The suggested algorithm allows successful aufeis detection if an image is predominantly  
236 snow-free. At the end of May/early June, many aufeis in mountain regions are still covered by  
237 snow. Their detection required later images, obtained in mid-June.

238 Morse and Wolfe (2015) suggested a new spectral index MDII for automatically  
239 distinguishing snow bodies from ice ones. However, here some of the high elevation aufeis were  
240 partially covered with snow at the image acquisition time. Instead of automatic processing, the  
241 outlining of high elevation aufeis was conducted manually when snow cover was present, with  
242 separation of aufeis from adjacent snow covered areas.

243 Further, during melt season, the aufeis often divide into several neighboring areas. When  
244 assessing the number of aufeis with satellite data, it is therefore necessary to aggregate the areas  
245 into one aufeis, if they are located at a distance <150 m (or five Landsat pixels) from each other,  
246 and within one aufeis glade.

247 As a result of semi-automated processing of Landsat images, aufeis with a total area of 1  
248 253.9 km<sup>2</sup> were detected. During the subsequent comparison with the Cadastre data (see section  
249 3.3 for more details), over 100 aufeis, with a total area of 33.5 km<sup>2</sup>, were delineated manually.  
250 The gaps were mainly due to the presence of snow cover and/or cloud coverage in the images.  
251 To reduce the number of gaps, two to three images of the same territory were used. The total

252 number of aufeis, identified with the Landsat images in the Indigirka River basin, was 1 213 and  
253 their total area 1 287.4 km<sup>2</sup>. Therefore, an omission error of automatic aufeis detection can be  
254 estimated as 2.7% of their total area.

255 The structure of the GIS dataset of aufeis according to Landsat images is presented in  
256 Table 2.

257

### 258 **3.3. Cross reference between historical and satellite-based aufeis data collection**

259 Cross-verification of aufeis data collections by the Cadastre (1958) and satellite imagery  
260 was performed in two steps. At the first step, we found the closest aufeis in the Landsat-derived  
261 dataset for each aufeis from the Cadastre data if the distance between them was less than 5000  
262 m. The determination of search radius was based on a preliminary analysis of the aufeis locations  
263 by the Cadastre in relation to Landsat-based dataset. As a result, the cross index (identifier of the  
264 closest aufeis in the Landsat-derived dataset) and minimum distance (m) to the closest aufeis  
265 were determined for aufeis from the Cadastre. For Landsat-based dataset, the cross index is the  
266 key field for the reference to the dataset from the Cadastre.

267 At the second step, a full manual verification was performed to find the mistakenly  
268 interrelated aufeis. For example, if the closest aufeis from the Cadastre and from the Landsat-  
269 based dataset were at a distance of less than 5000 m, but in different thalwegs, they were  
270 considered as different (unrelated) aufeis.

271 In total, 260 aufeis from the Cadastre were not verified by Landsat images. For them, the  
272 NoData value (−9999) was set in the Cross Index and Distance fields of attributive table (see  
273 Table 1 with the structure of GIS dataset from Cadastre).

274

## 275 **4. Results**

### 276 **4.1 Comparison of the historical and modern data collection**

277 The results of the comparison are presented in Table 3. In total, 634 aufeis from the  
278 Cadastre were found by the Landsat images. They correspond to 611 aufeis identified with the  
279 images, meaning that in 23 cases, one aufeis in an image corresponds to two aufeis in the  
280 Cadastre. But 262 aufeis from the Cadastre were not detected by the satellite images. Those are  
281 mainly small aufeis, which melt by the middle of June. However, among them there are also 43  
282 large aufeis over 1 km<sup>2</sup> (fig. 3-a). It is likely that since the mid-20th century, when the field  
283 observations were conducted and the Cadastre of aufeis was compiled, some aufeis could have  
284 disappeared.

285 A little over half of the aufeis detected by Landsat images are included in the Cadastre: a  
286 total of 602 aufeis detected (the total area of 250.4 km<sup>2</sup>) are not included in the Cadastre (fig. 3-  
287 b). Such a significant difference can be caused by the following reasons:

288 1. In some cases a single aufeis, according to the Cadastre, corresponds with two or more  
289 aufeis by satellite image;

290 2. Aufeis are characterized by significant interannual variability, which results in possible  
291 formation of new aufeis in areas where they previously were not observed (Alekseev, 2015;  
292 Pomortsev et al., 2010; Atlas of snow..., 1997).

293 Total aufeis area evaluated based on satellite images, appeared to be 1.6 times smaller than  
294 stated in the Cadastre (1958). First and foremost, such difference can be explained by the fact  
295 that it was not the area of the aufeis themselves, but instead the aufeis glades, that were reported  
296 in the Cadastre (1958) and this corresponds to the maximum aufeis area during one or several  
297 seasons. With the satellite data, the areas of the aufeis themselves were assessed and when mid-  
298 June images were used, the aufeis area was significantly smaller than the typical annual  
299 maximum.

300 Aufeis area distribution according to the Cadaster and satellite data is shown as Lorenz  
301 curves (fig. 4). In both cases, the shape of the curves signifies a high degree of irregularity which

302 is similar: 10% of the largest aufeis make up 61 and 57% of their total area according to the  
303 Landsat and the Cadastre data, respectively.

304 The cross-verification of the Cadastre and satellite data show that almost 60% of aufeis  
305 that are unconfirmed in the Landsat imagery and that are therefore only present in the Cadastre,  
306 have an individual aufeis area less than 0.25 km<sup>2</sup> (Fig. 5-a). The confirmed aufeis account for  
307 about 20% of the area stated in the Cadastre. Thus, it was mainly small aufeis that were not  
308 confirmed in the Landsat images. Conversely, Fig. 5-b shows that almost 60% of the aufeis  
309 detected in the Landsat images but not listed in the Cadastre have an area each of less than 0.25  
310 km<sup>2</sup>.

#### 311 **4.2. Aufeis distribution by elevation**

312 In general, aufeis distributions by elevation as assessed with the Cadastre and Landsat data  
313 are quite similar, although there are some differences that are elevation-specific (fig. 6). Most  
314 aufeis are located in the elevation band of 1 000 – 1 200 m. At lower elevations (up to 800 m)  
315 the number of aufeis according to Landsat data is higher than stated in the Cadastre. At the  
316 elevations of 1 400-2 000 m, more aufeis are identified in the Cadastre data than by the satellite  
317 images. This can be explained by the fact that many aufeis located at high altitudes often have a  
318 small area, so they could have been missed during the analysis of the satellite data. Further, they  
319 could have been covered with snow at the image acquisition time, which would increase the  
320 possibility of them being missed.

321 The elevation band of 200-300 m is characterized by the location of large aufeis. Though  
322 less than 2.5% and 5.0% of aufeis by the Cadastre and Landsat images are situated here, they  
323 represent about 11 and 13% of aufeis area from the datasets respectively (fig. 7).  
324

#### 325 **4.3 Aufeis distribution by river basins**

326 In the Indigirka River basin, there are several zones with a high density of aufeis: in the  
327 southern part (the Sutar and Kuidusun Rivers basins), as well as in the central part (Chersky  
328 Range slopes) (fig. 8). The largest aufeis identified with satellite images are located in the  
329 Syuryuktyakh River basin on the north-east slopes of the Chersky Range. Meanwhile, aufeis are  
330 almost absent in the northernmost (lowland) part of the Indigirka basin.

331 We analyzed the aufeis coverage for six river basins with available streamflow data. The  
332 headwater part of the Indigirka River, with the gauge near the Yurty village (area 51 100 km<sup>2</sup>), is  
333 the basin with the largest aufeis coverage (Table 4). Correlation between average elevation of the  
334 basins and their aufeis coverage (expressed as a percentage) is statistically significant. Among 6  
335 basins, the Spearman rank correlation coefficients between the basin average elevation and  
336 aufeis percentage are 0.71 and 0.77 by the Cadastre and satellite data, respectively.

#### 337 **4.4 Aufeis area interannual variability**

338 The assessment of aufeis area interannual variability was conducted in two areas: for the  
339 Bolshaya Momskaya aufeis, which is located in the Moma River channel (area in the Cadastre is  
340 82 km<sup>2</sup>), and for a group of large aufeis (total area in the Cadastre is 287.8 km<sup>2</sup>) in the  
341 Syuryuktyakh River basin, which is the left-bank tributary of the Indigirka River.

342 Cloudless images from Landsat-5 (TM), Landsat 7 (ETM+) and Landsat-8 (OLI) were  
343 used with the acquisition dates between May 1 and June 30. In the USGS archives, there are no  
344 Landsat-5 images for the study territory for the 1984-2007 period. This limits the duration of  
345 satellite observations on aufeis to the period since 1999 (when the Landsat-7 satellite was  
346 launched). Also, the clouds complicate the acquisition of representative data. The list of the  
347 acquisition dates and assessed aufeis area values are presented in Table. 5.

348 Both areas are located at low elevations (Bolshaya Momskaya 430 to 500 m and  
349 Syuryuktyakh 200 to 500 m), which contributes to the relatively early and intensive aufeis melt  
350 in spring. The aufeis reach their maximum area by the beginning of May. Using the available

351 satellite images it is impossible to make a reliable conclusion on aufeis area increase or decline,  
352 because the acquisition dates vary significantly from year to year. However, it is possible to  
353 make some conclusions based on the available data:

354 1. In 2002-2017 the Bolshaya Momskaya aufeis did not reach the maximum area stated in  
355 the Cadastre (82 km<sup>2</sup>), even though the satellite image was acquired during the first week of May  
356 (2005) when aufeis melting had not yet started. Comparing two images, taken in similar  
357 conditions (08.05.2005 and 15.05.2013), it was found that aufeis area in 2013 was smaller by  
358 18.1 km<sup>2</sup> than in 2005. Accordingly, the Bolshaya Momskaya aufeis may have seen a decreasing  
359 trend over time in its maximum coverage.

360 2. The area of the largest aufeis in the Syuryuktyakh River basin in May 2014 was 78.0  
361 km<sup>2</sup>, which is 8 km<sup>2</sup> larger than stated in the Cadastre. One may note also that the maximum  
362 aufeis areas in the Syuryuktyakh River basin were detected by the images received at the end of  
363 the period (2014-2017), including mid-June (18.06.2015). Therefore, it can be suggested that the  
364 aufeis areas within Syuryuktyakh River basin have not decreased since 2002.

365

## 366 5. Discussion

367 The most important uncertainty in the obtained results relates to our ability to draw a  
368 conclusion on the long-term trend of total aufeis area comparing the historical and satellite-  
369 derived datasets. The total area of aufeis estimated by Landsat images is 38% less than according  
370 to the Cadastre. Is it possible to confirm that such a significant reduction in the aufeis area really  
371 occurred? Considering this issue, it is important to emphasize some limitations of the  
372 methodology and the created datasets.

373 The main limitation of the historical aufeis dataset is that the Cadastre provides an area of  
374 aufeis glades, but not the aufeis themselves. Simakov and Shilnikovskaya (1958) noted that the  
375 areas of aufeis glades match the average annual maximum of the ice-covered area. Alekseev  
376 (2005) states that the assessment of the stages and patterns of the development of aufeis glades  
377 based on the analysis of their landscape and geomorphological features is difficult due to the lack  
378 of research on temporal aspects of mutual transitions of landscape facies and their factorial  
379 dependencies. However, studying the aufeis landscapes in the central part of Eastern Sayan  
380 Mountains, Alekseev (2005) assumes that the vegetation community which is a typical indicator  
381 of aufeis development may persist for 200-300 years after the beginning of aufeis processes  
382 attenuation.

383 The satellite-derived assessment of the aufeis area has the following main source of  
384 uncertainty. It is often impossible to determine the maximum area of aufeis by satellite images,  
385 since it is observed at the beginning of snow melt season, when aufeis are still covered with  
386 snow. In late spring and beginning of summer, the area of aufeis may already been significantly  
387 reduced in comparison with the maximum values, due to melting and mechanical destruction.

388 Maximum intensity of aufeis melt in the studied region is observed in June when spring  
389 flood river streams actively erode the aufeis surface. Sokolov (1975) reported the results of the  
390 observations at the Anmyngynda aufeis carried out in 1962-1965. This aufeis is located in the  
391 upstream area of the Kolyma river basin (723 m a.s.l.) and may be used as being representative  
392 of the mountainous part of the studied region. In 1962-1965, the aufeis area changed from 5.1 to  
393 6.2 km<sup>2</sup> with mean maximum area of 5.7 km<sup>2</sup>. Aufeis melt has been observed to begin on  
394 average on the 10<sup>th</sup> of May. During May, the aufeis area decreased by 15% of the total area on  
395 average. At the end of June, the remaining area was 34% of the maximum, i.e. during this month  
396 more than 50% of the aufeis area has been destroyed. In the period from July to September, the  
397 melting slowed down: in July the aufeis decreased by 22%, in August by 8%, in September by  
398 3%. The area of aufeis at lower absolute elevations decreases faster at first half of the summer,  
399 and in the upstream areas – in the second one (Sokolov, 1975).



400 Some aufeis in the mountainous regions could be missed by satellite images, since they  
401 can be covered with snow until the end of June. However, their contribution to the total area is  
402 non-significant.

403 Taking into account all the above-described limitations, and also that more than 600  
404 aufeis that were missing in the Cadastre were found by Landsat images, we conclude that it is  
405 not correct to make a conclusion about long-term trends of aufeis area based on the entire created  
406 dataset. Following Pavelsky and Zarnetske (2017), we decided to examine only several of the  
407 largest aufeis deposits in order to identify the long-term trend.

408 We selected the 38 largest aufeis with an area  $\geq 10$  km<sup>2</sup> according to the Cadastre dataset,  
409 confirmed by satellite data. Their total area decreased from 858.1 km<sup>2</sup> according to the Cadastre  
410 to 356.3 km<sup>2</sup> according to recent Landsat images. Conversely, we also selected the largest aufeis  
411 according to satellite data (18 aufeis with satellite-estimated area  $\geq 10$  km<sup>2</sup>). Their total area also  
412 decreased significantly (from 428.6 km<sup>2</sup> according to the Cadastre to 343.5 km<sup>2</sup> according to  
413 Landsat images). We also analyzed 8 giant aufeis with areas  $\geq 35$  km<sup>2</sup> according to the Cadastre  
414 dataset. They all were confirmed by the satellite images; however seven of the eight had a  
415 significantly smaller area (from 2 to 21 km<sup>2</sup>) with the decrease being 2-10 times. Only one giant  
416 aufeis in the Syuryuktyakh River basin has the area by Landsat larger than by Cadastre, at 72 and  
417 64 km<sup>2</sup> accordingly. It should be noted that the formation of new (mainly small) aufeis can  
418 slightly reduce the rate of the aufeis area decrease.

419

## 420 **6. Conclusion**

421 The research conducted here is the first step of the study aimed at the development of a  
422 GIS database of the aufeis of North-East Russia. Historical data of the Cadastre (1958) and  
423 topographic maps were used to create a geodatabase of aufeis in the Indigirka River basin (up to  
424 the Vorontsovo gauge, with the area of 305 000 km<sup>2</sup>). It contains historical data on 896 aufeis  
425 with total area of 2063.6 km<sup>2</sup>. Aufeis detection was conducted for the 2013-2017 period using  
426 Landsat imagery with 1213 aufeis identified having a total area of 1287.4 km<sup>2</sup>. The historical  
427 dataset from the Cadastre (1958) and more recent satellite-based dataset were compared and  
428 combined in the joint Catalogue of aufeis within the Indigirka River basin, available at the  
429 PANGAEA repository (<https://doi.pangaea.de/10.1594/PANGAEA.891036>).

430 Recent total aufeis area is 1.6 times smaller than stated in the Cadastre (1958). The more  
431 significant changes occurred to 38 large and giant aufeis (area  $\geq 10$  km<sup>2</sup>) with total decrease of  
432 area by 501.8 km<sup>2</sup> (or 66% of the total reduction). Simultaneously, the historical Cadastre  
433 archive is lacking data on over 600 aufeis that were identified using satellite images. This  
434 suggests that the Cadastre data is incomplete, while there may also have been significant change  
435 in aufeis formation conditions in the last half century.

436 The analysis of large and giant aufeis seems to indicate that there has been a significant  
437 decrease in aufeis area over the period of last 70 years. Additional analysis of historical aerial  
438 photography data could help to clarify the issue of aufeis area decline trend since the middle of  
439 the 20th century to the present. One of the further study goals will be to find out the extent to  
440 which these changes are climate-derived and to identify their impact on river streamflow.

441

442 **Acknowledgements.** The authors are grateful to David Post, Anna Liljedahl and an  
443 anonymous reviewer for valuable comments and assistance with English.

444

## 445 **7. References**

446 Alekseev, V.R.: Naledi. Novosibirsk, Nauka, Moscow, 1987 (in Russian).

- 447           Alekseev, V.R.: Long-term variability of the spring taryn-aufeises. *Ice and Snow*, 56(1),  
448 73-92, 2016. doi:10.15356/2076-6734-2016-1-73-92.
- 449           Alekseyev, V.R.: Cryogenesis and geodynamics of icing valleys. *Geodynamics &*  
450 *Tectonophysics* 6 (2), 171–224, 2015. doi:10.5800/GT-2015-6-2-0177.
- 451           Alekseev, V.R.: Landscape indication of aufeis phenomena. Novosibirsk, Nauka, 2005,  
452 364 p. (in Russian).
- 453           Ananicheva, M.D.: Estimation of the areas, volumes and heights of the boundary of the  
454 feeding of glacier systems of the Northeast of Russia from the space images of the beginning of  
455 the 21st century. *Ice and Snow*, 1 (125), 35-48, 2014.
- 456           Atlas of snow and ice resources of the world, 11 (2). Institute of geography RAS, Moscow,  
457 270 pp., 1997 (in Russian).
- 458           Aufeis of Siberia and the Far East. Nauka, Novosibirsk, 244 pp. 1981. (in Russian).
- 459           Bense, V. F., Kooi, H., Ferguson, G., and Read, T.: Permafrost degradation as a control on  
460 hydrogeological regime shifts in a warming climate. *Journal of Geophysical*  
461 *Research*, 117(F03036), 2012. doi: 10.1029/2011JF002143.
- 462           Bring, A., Fedorova, I., Dibike, Y., Hinzman, L., Mård, J., Mernild, S. H., Prowse, T.,  
463 Semenova, O., Stuefer, S. L., and Woo, M.-K.: Arctic terrestrial hydrology: Asynthesis of  
464 processes, regional effects, and research challenges, *J. Geophys. Res. Biogeosci.*, 121, 621–649,  
465 doi:10.1002/2015JG003131, 2016.
- 466           Cadastre to the map of the north-east of the USSR Scale 1: 2000000. Shilnikovskaya Z.G.  
467 Central complex thematic expedition of the North-Eastern geological department. Magadan, 398  
468 pp., 1958. (in Russian).
- 469           Catalogue of Aufeis in the Baikal-Amur Railroad Zone, 1. Aufeis in the Upper Part of the  
470 Chara River Basin. *Gidrometeoizdat*, Leningrad, 63 pp., 1980 (in Russian) .
- 471           Catalogue of Aufeis in the Baikal-Amur Railroad Zone, 2. Aufeis in the Muya River Basin.  
472 *Gidrometeoizdat*, Leningrad, 84 pp., 1981 (in Russian) .
- 473           Catalogue of Aufeis in the Baikal-Amur Railroad Zone, 3. Aufeis in the Upper Angara  
474 River Basin. *Gidrometeoizdat*, Leningrad. 96 pp., 1982 (in Russian) Chavez Jr. P.S. Image-based  
475 atmospheric corrections – revisited and improved. *Photogram. Eng. and Remote Sens*, 1996,  
476 62(9), 1025–1036.
- 477           Danielson, J.J., and Gesch, D.B.: Global multi-resolution terrain elevation data 2010  
478 (GMTED2010): U.S. Geological Survey Open-File Report 2011–1073, 26 pp., 2011.
- 479           Explanatory note to the geocryological map of the USSR, scale 1: 2 500 000, 125 pp., 1991  
480 (in Russian).
- 481           Ge, S., J. McKenzie, C. Voss, and Wu, Q.: Exchange of groundwater and surface-water  
482 mediated by permafrost response to seasonal and long term air temperature variation, *Geophys.*  
483 *Res. Lett.*, 38, L14402, 2011. doi:10.1029/2011GL047911.
- 484           GLIMS and NSIDC: Global Land Ice Measurements from Space glacier database.  
485 Compiled and made available by the international GLIMS community and the National Snow  
486 and Ice Data Center, Boulder CO, U.S.A., doi:10.7265/N5V98602, 2005, updated 2017.

- 487 Grosse, G. and Jones, B. M.: Spatial distribution of pingos in northern Asia, *The*  
488 *Cryosphere*, 5, 13-33, 2011. <https://doi.org/10.5194/tc-5-13-2011>.
- 489 Hall, D K., Riggs, G.A. and Salomonson, V.V.: Development of methods for mapping  
490 global snow cover using Moderate Resolution Imaging Spectroradiometer (MODIS) data.  
491 *Remote Sensing of Environment*, 54, 127-140, 1995.
- 492 Kane, D. L. and Slaughter C. W.: Recharge of a central Alaska lake by subpermafrost  
493 groundwater. *Nat. Acad. Sci*, 38, 458-462, 1973.
- 494 Makarieva, O., Nesterova, N., Post, D. A., Sherstyukov, A., and Lebedeva, L.: Warming  
495 temperatures are impacting the hydrometeorological regime of Russian rivers in the zone of  
496 continuous permafrost, *The Cryosphere Discuss.*, <https://doi.org/10.5194/tc-2018-157>, in review,  
497 2018.
- 498 Makarieva, O., Shikhov, A., Ostashov, A., Nesterova, N.: Aufeises (naleds) of the North-  
499 East of Russia: GIS catalogue for the Indigirka River basin (Russia), PANGAEA, 2018.  
500 available at: <https://doi.pangaea.de/10.1594/PANGAEA.891036>.
- 501 Map of the icings of the North-East of the USSR Scale 1: 2000000. Simakov, A.S.,  
502 Shilnikovskaya, Z.G. North-Eastern Geological Administration of the Main Directorate of  
503 Geology and Subsoil Protection. Central complex thematic expedition. Magadan. 1958 (in  
504 Russian).
- 505 Markov, M. L., Vasilenko, N. G. and Gurevich, E. V.: Icing fields of the BAM zone:  
506 expeditionary investigations. SPb.: Nestor-History, 320 pp., 2016 (in Russian).
- 507 Morse, P.D. and Wolfe, S.A.: Geological and meteorological controls on icing (aufeis)  
508 dynamics (1985 to 2014) in subarctic Canada. *Journal of Geophysical Research: Earth Surface*.  
509 120, 1670–1686, 2015.
- 510 Pavelsky, T.M. and Zarnetske, J.P.: Rapid decline in river icings detected in Arctic Alaska:  
511 Implications for a changing hydrologic cycle and river ecosystems. *Geophysical Research*  
512 *Letters*. 44(7), 3228–3235, 2017.
- 513 Pollard, W. H.: Icing processes associated with high Arctic perennial springs, Axel  
514 Heiberg Island, Nunavut, Canada, *Permafrost Periglacial Processes*, 16, 51 – 68, 2005.
- 515 Pomortsev, O.A., Kashkarov, E.P. and Popov, V.F.: Aufeis: global warming and processes  
516 of ice formation (rhythmic basis of long-term prognosis). *Bulletin of Yakutsk State University*, 7  
517 (2). 40-48, 2010 (in Russian).
- 518 Reedyk, S., Woo, M.K. and Prowse, T.D.: Contribution of icing ablation to streamflow in a  
519 discontinuous permafrost area. *Canadian Journal of Earth Science*, 32: 13-20, 1995.
- 520 Rennermalm, A. K., Wood, E. F., and Troy, T. J.: Observed changes in pan-arctic cold-  
521 season minimum monthly river discharge, *Clim. Dyn.*, 35(6), 923–939, 2010.
- 522 Romanovsky, V. E., Sazonova, T. S., Balobaev, V. T., Shender, N. I., and Sergueev, D. O.:  
523 Past and recent changes in air and permafrost temperatures in eastern Siberia, *Global Planet.*  
524 *Change*, 56, 399–413, doi:10.1016/j.gloplacha.2006.07.022, 2007.
- 525 Simakov, A.S. and Shilnikovskaya, Z.G.: The map of the north-east of the USSR. A Brief  
526 Explanatory Note. The North-Eastern Geological Administration of the Main Directorate of  
527 Geology and Subsoil Protection, Magadan, 40 pp., 1958 (in Russian).

- 528 Slaughter, C. W.: Occurrence of and recurrence of aufeis in an upland taiga catchment, in:  
529 Roger J.E. Brown Memorial Volume—Proceedings of the 4th Canadian Permafrost Conference,  
530 Natl. Res. Council of Can., Ottawa, Ont, 182 – 188, 1982.
- 531 Sokolov, B.L.: Naleds and river runoff. Gidrometeoizdat, Leninrgad, 190 pp., 1975 (in  
532 Russian).
- 533 Surface waters resources of the USSR, 17. Leno-Indigirsky district. Gidrometeoizdat,  
534 Leninrgad, 651 pp., 1972. (in Russian).
- 535 Tananaev, N. I., Makarieva, O. M., and Lebedeva, L. S.: Trends in annual and extreme  
536 flows in the Lena River basin, Northern Eurasia. doi: 10.1002/2016GL070796, 2016.
- 537 Tolstikhin, O.N.: Naleds and ground waters in the North-East of the USSR. Science,  
538 Novosibirsk, 164 pp., 1974 (in Russian).
- 539 Topchiev A.G.: Methods of aerospace monitoring of aufeis geosystems and underground  
540 water of cryolithozone. Systems and means of informatics, 2008, 18 (3), pp. 304-327 (in  
541 Russian)
- 542 US Geological Survey Server: <http://earthexplorer.usgs.gov>, last access: 1 August 2018.
- 543 Walvoord, M. A., Voss C. I. and Wellman T. P.: Influence of permafrost distribution on  
544 groundwater flow in the context of climate-driven permafrost thaw: Example from Yukon Flats  
545 Basin, Alaska, United States, Water Resour. Res., 48, W07524, 2012.  
546 doi:10.1029/2011WR011595.
- 547 Walvoord, M. A. and Kurylyk, B. L.: Hydrologic Impacts of Thawing Permafrost—A  
548 Review. Vadose Zone Journal, 15 (6), 2016. doi: <https://doi.org/10.2136/vzj2016.01.001>.
- 549 Yde, J. C., and Knudsen, N. T. Observations of debris-rich naled associated with a major  
550 glacier surge event, Disko Island, west Greenland, Permafrost Periglacial Processes, 16, 319 –  
551 325, 2005.
- 552 Yoshikawa, K., Hinzman, L.D., and Kane, D.L.: Spring and aufeis (icing) hydrology in  
553 Brooks Range, Alaska, Journal of Geophysical Research: Biogeosciences, 112, 1–14, 2007.
- 554

Table 1 The structure of GIS database of augeis by Cadastre (1958)

Field name	Field alias	Description
FID	FID	Index number (Object ID)
AufDataSrc	Augeis data source	Augeis Cadastre data (1958) (for all objects)
Auf_area	Augeis area Cadastre (km <sup>2</sup> )	Augeis area (km <sup>2</sup> ) from the Cadastre (1958). If the data was missing, the area was calculated by topographic maps (1980) scale 1: 200 000
Auf_index	Augeis index Cadastre	Index of the augeis in the Cadastre (1958) (it contains 0 if the augeis was missing in the Cadastre, but found in the topographic map (1980) scale 1: 200 000)
Map_index	Cadastre map index	Index of the Cadastre (1958) map
Auf_topo	Augeis in topo	Presence of the augeis at topographic map (0 – missing, 1 – present)
Auf_in_map	Augeis in map	Presence of the augeis in the Cadastre (0 – missing, 1 – present)
Toponumber	Topo number	Nomenclature of the topographic map sheet
Date	Date	Date of fixing the presence of ice within the augeis
Long	Long	Longitude, degree
Lat	Lat	Latitude, degree
Elevation	Elevation	Height above sea level (determined by Aster GDEM), m
Comment	Comment	Comments (mainly typos in the Cadastre map, or the method of determining augeis area)
CrossIndex	Cross index	Cross index of augeis derived from Landsat (if augeis is not in Landsat, the value is missing)
Distance_m	Distance (m)	Minimum distance between the augeis from the Cadastre and the same augeis from Landsat image (m)

556

557

Table 2 The structure of GIS database of augeis by Landsat images (2013-2017)

Field name	Field alias	Description
FID	FID	Index number (Object ID)
AufDataSrc	Augeis data source	Landsat images (for all objects)
WRS2_ID	Landsat WRS2_ID	The Landsat scene identifier in the WRS2 graph of the US Geological Survey (USGS). The first three digits indicate the column number, and last three digits represent the line number.
Image_Date	Landsat image date	The date of image
Comment	Comment	Additional information, for example, if the augeis was partly covered by clouds and additional images were used to estimate the area
CrossIndex	Cross index	Identifier of augeis by Landsat images (key field for the reference to the Cadastre data)
Auf_Area	Augeis area (km <sup>2</sup> )	Augeis area by Landsat image, km <sup>2</sup>
Elevation	Average elevation	Average elevation of augeis, calculated by Aster GDEM digital elevation model

558

559

Table 3 Data correlation of aufeis based on the Cadastre (1958) and the Landsat images

560

Data source	Matching aufeis number and area (km <sup>2</sup> )	Not confirmed aufeis number and area (km <sup>2</sup> )
Cadastre (1958)	634 (1905.0)	262 (158.6)
Landsat	611 (1037.0)	602 (250.4)

561

562 Table 4 Aufeis area coverage (percentage) in the sub-basins within the Indigirka River watershed  
563 by the Cadastre and Landsat data

River	Area, km <sup>2</sup>	Average elevation, m a.s.l.	% aufeis coverage (Cadastre)	% aufeis coverage (Landsat)
Suntar River –Sakharinya River mouth	7680	1460	0.97	0.78
Elgi – 5 km upstream of the Artyk-Yuryakh River mouth	17600	1104	0.49	0.23
Nera – Ala-Chubuk	22300	1174	0.32	0.26
Indigirka – Yurty	51100	1256	1.15	0.80
Indigirka – Indigirskiy	83500	1185	0.82	0.56
Indigirka – Vorontsovo	305000	803	0.68	0.41

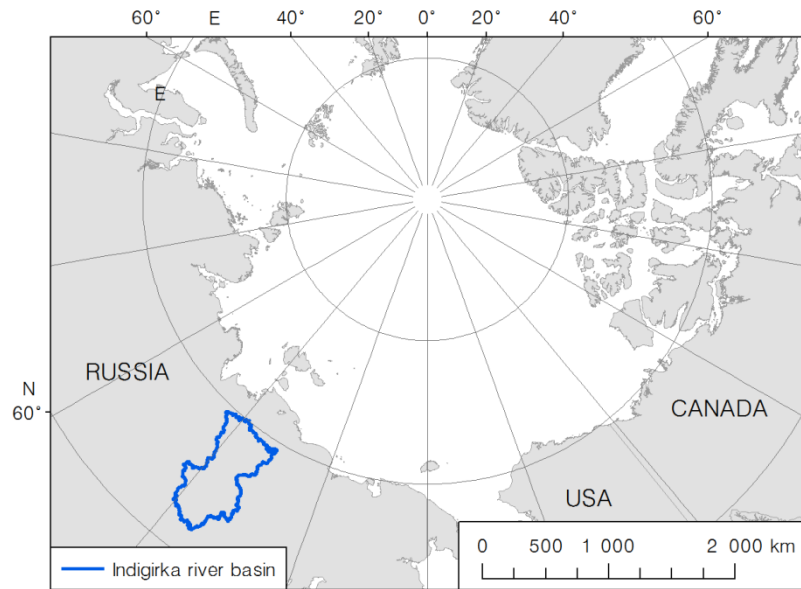
564

565 Table 5 Aufeis area changes, 2001-2017.

Bolshaya Morskaya aufeis		The group of aufeis in the Syuryuktyakh River basin	
Imagery date	Aufeis area, km <sup>2</sup>	Imagery date	Aufeis area, km <sup>2</sup>
17.06.2002	29.2	26.06.2001	69.7
08.05.2005	66.2	29.06.2002	100.6
27.05.2006	57.9	04.06.2007	155.1
19.06.2009	39.5	17.06.2009	117.5
25.05.2011	61.7	22.06.2011	89.5
27.05.2012	49.6	21.05.2014	268
15.05.2013	48.1	18.06.2015	164.8
18.06.2017	21.9	04.06.2016	206.4

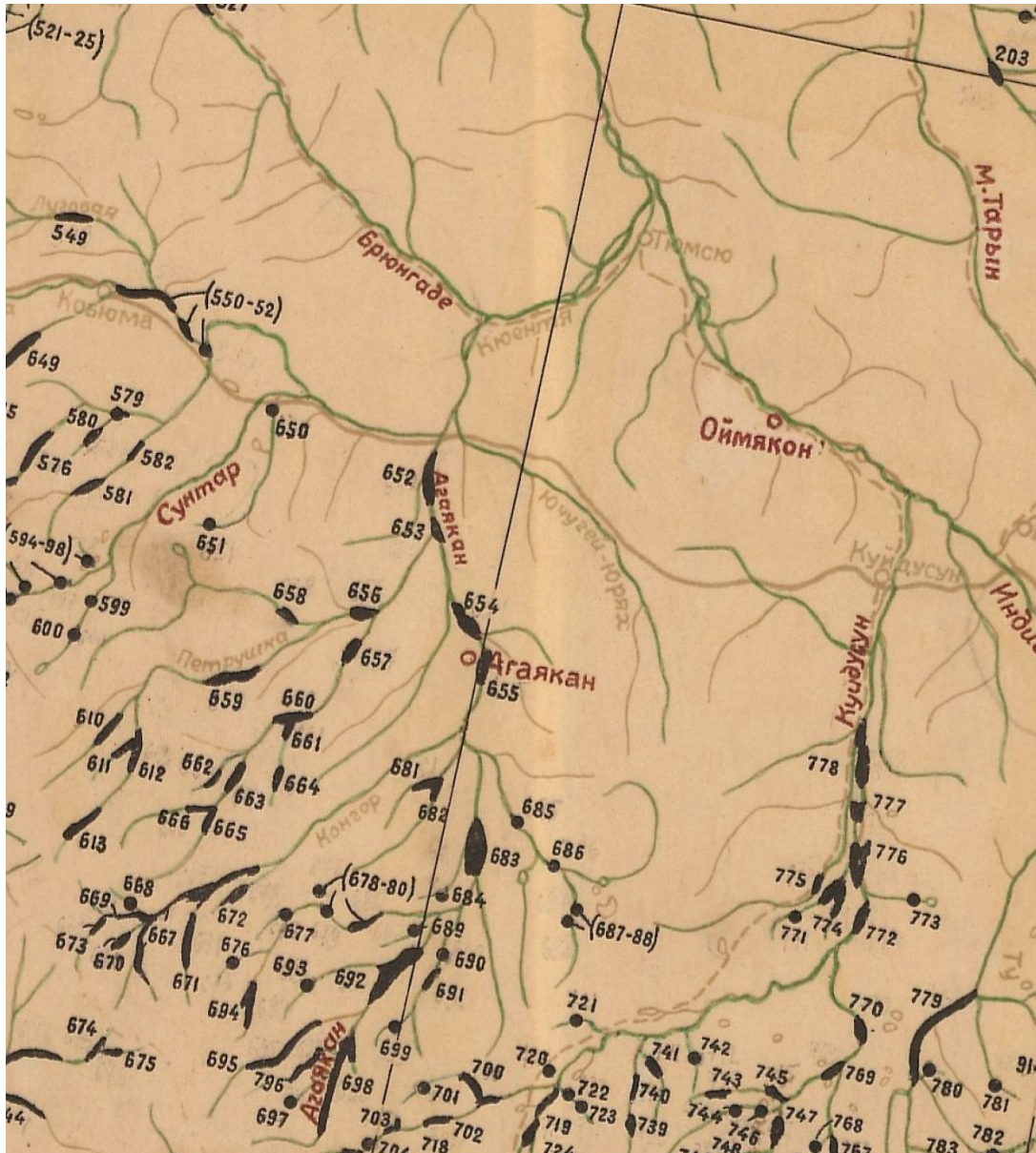
566

567



568  
569  
570  
571

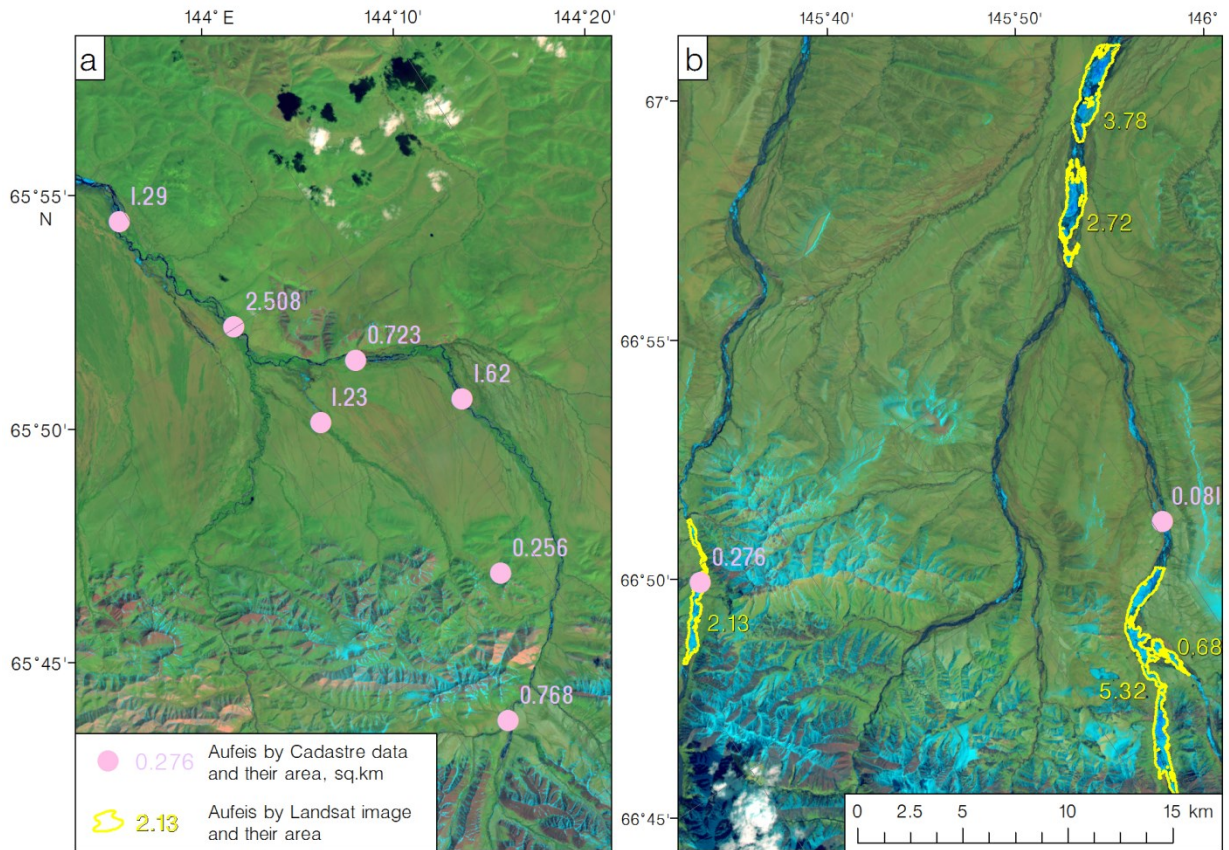
Fig. 1 Geographical location of the Indigirka river basin



572  
 573  
 574  
 575

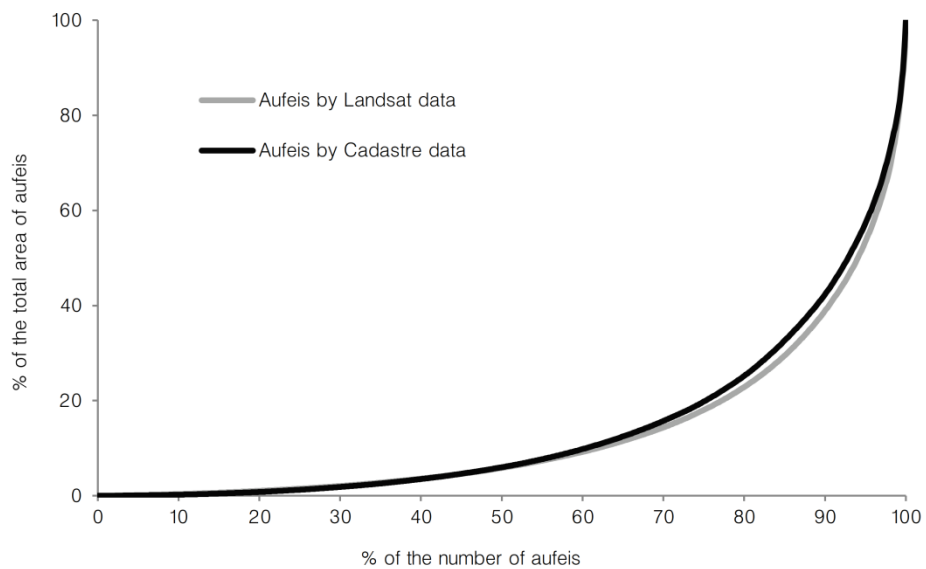
Fig. 2 Subset of the Cadastral Map of the North-East of the USSR from 1958 (sheet 7, upper reaches of the Indigirka River – the basins of the rivers Sунтар, Агаякан and Куйдусун).





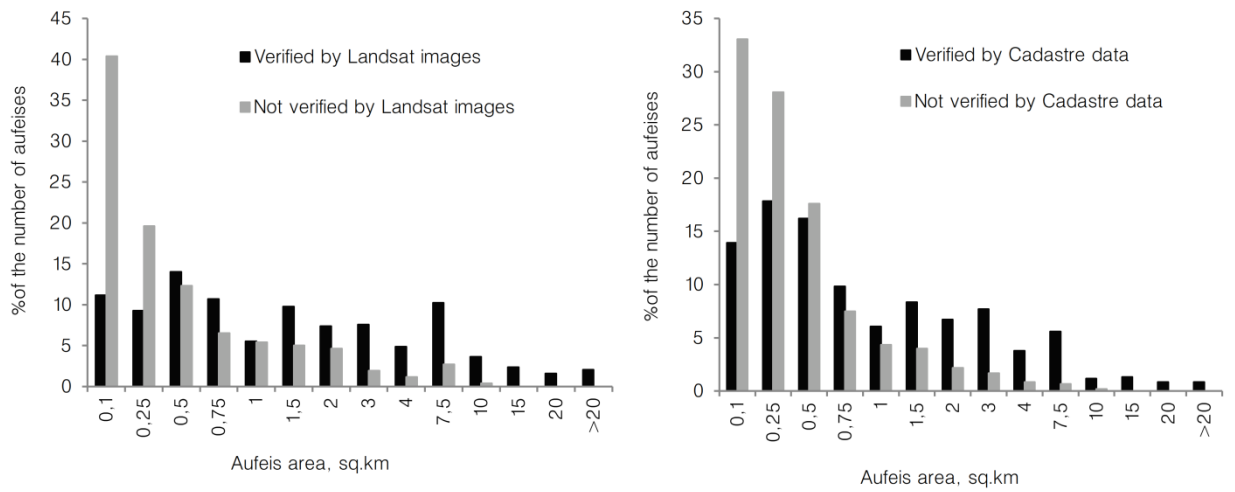
576  
577  
578  
579  
580  
581

Fig. 3 Difference between aufeis location according to the Cadastre and satellite data: a) – aufeis are absent in the image but present in the Cadastre (Landsat-8 image of 18.06.2017); b) – aufeis are absent (or their area is understated) in the Cadastre but present in the image (Landsat-8 image of 30.05.2016).



582  
583

Fig. 4 Lorenz curves illustrating aufeis area distribution according to the Cadastre and Landsat data



584

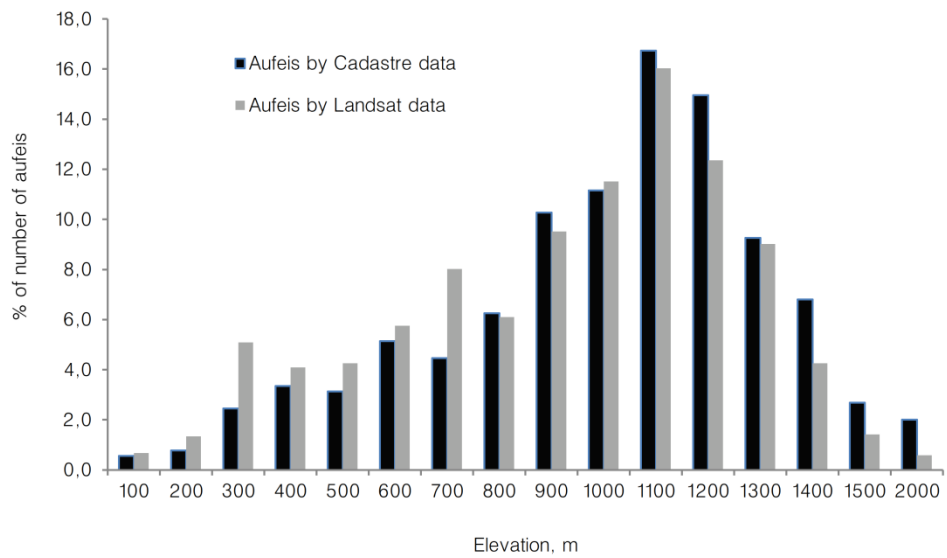
585

586

587

588

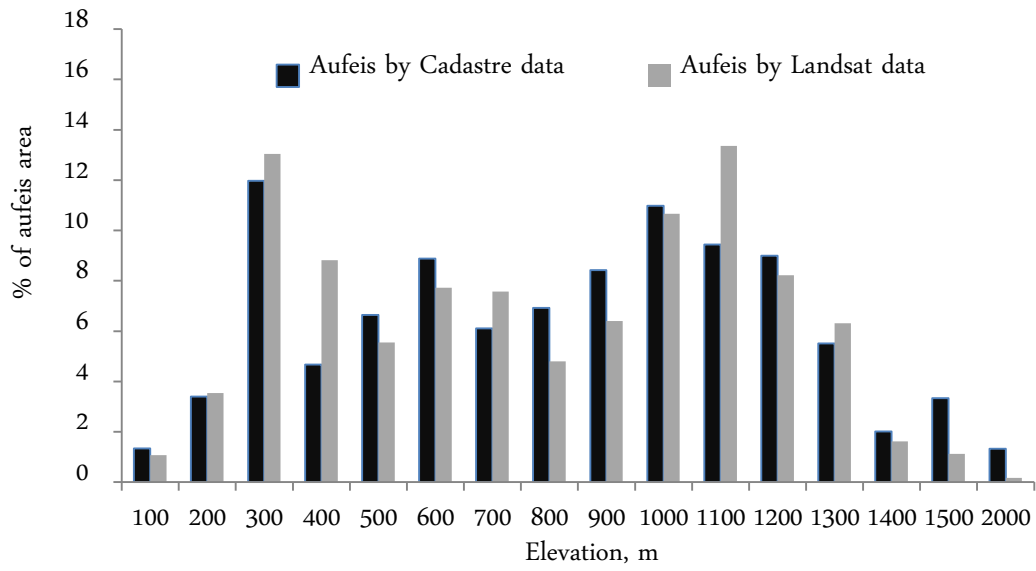
Fig. 5 Aufeis area distribution: a) – according to the Cadastre data, confirmed and not confirmed by Landsat images, b) – according to Landsat images, confirmed and not confirmed by the Cadastre.



589

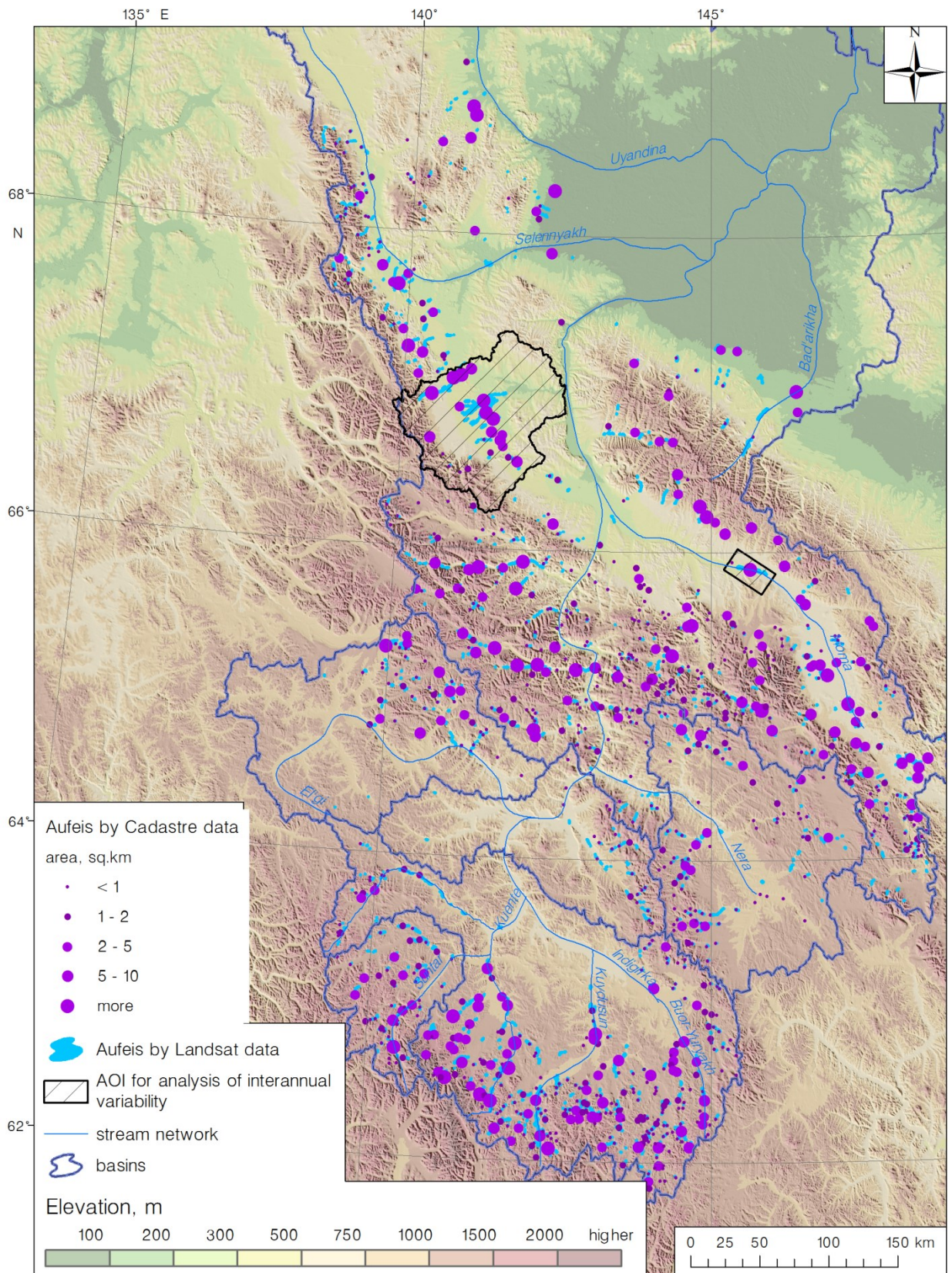
590

Fig. 6 Aufeis distribution by elevation within the Indigirka River basin.



591  
 592  
 593

Fig. 7 Aufeis area distribution by elevation within the Indigirka River basin.



594  
595  
596  
597

Fig. 8 Aufeis in the Indigirka River basin according to the Cadastre and Landsat images. Black outlines with section lining represent the zones where aufeis area interannual variability was assessed