1	Historical and recent aufeis, the Indigirka river basin
2	(Russia)
3	*Olga Makarieva <sup>1,2,</sup> , Andrey Shikhov <sup>3</sup> , Nataliia Nesterova <sup>2,4</sup> , Andrey Ostashov <sup>2</sup>
4	<sup>1</sup> Melnikov Permafrost Institute of RAS, Yakutsk
5	<sup>2</sup> St. Petersburg State University, St. Petersburg
6	<sup>3</sup> Perm State University, Perm
7	<sup>4</sup> State Hydrological institute, St. Petersburg
8	RUSSIA

Abstract: A detailed spatial geodatabase of aufeis (or naleds in Russian) within the Indigirka River watershed (305 000 km<sup>2</sup>), Russia, was compiled from historical Russian publications (year 1958), topographic maps (years 1970–1980's), and Landsat images (year 2013-2017). Identification of aufeis by late-spring Landsat images was performed with a semiautomated approach according to Normalized Difference Snow Index (NDSI) and additional data. After this, a cross-reference index was set for each aufeis, to link and compare historical and satellite-based aufeis data sets.

\*omakarieva@gmail.com

17 The aufeis coverage varies from 0.26 to 1.15% in different sub-basins within the 18 Indigirka River watershed. The digitized historical archive (Cadastre, 1958) contains the coordinates and characteristics of 896 aufeis with total area of 2064 km<sup>2</sup>. The Landsat-based 19 20 dataset included 1213 aufeis with a total area of 1287 km<sup>2</sup>. Accordingly, the satellite-derived 21 total aufeis area is 1.6 times less than the Cadastre (1958) dataset. However, more than 600 22 aufeis identified from Landsat images are missing in the Cadastre (1958) archive. It is 23 therefore possible that the conditions for aufeis formation may have changed from the mid-24 20th century to the present.

Most present and historical aufeis are located in the elevation band of 1000 – 1200 m. About 60% of total aufeis area is represented by just 10% of the largest aufeis. Interannual variability of aufeis area for the period of 2001-2016 was assessed for the Bolshaya Momskaya aufeis and for a group of large aufeis (11 aufeis with a areas from 5 to 70 km<sup>2</sup>) in the basin of the Syuryuktyakh River. The results of this analysis indicate a tendency towards an area decrease in the Bolshaya Momskaya aufeis in recent years, while no reduction in Syuryuktyakh River aufeis area was observed.

32 The combined digital database of the aufeis is available at 33 <u>https://doi.pangaea.de/10.1594/PANGAEA.891036</u>.

34

9

Keywords: aufeis, Indigirka, Landsat, NDSI, Cadastre, Cadastral map, Bolshaya Momskaya
 aufeis

37

#### 38 **1. Introduction**

39 Aufeis (naleds in Russian, icings in English) are accumulations of ice that are formed by 40 freezing underground and surface waters on the surface of the earth or ice along streams and 41 river valleys in arctic and subarctic regions. They affect water exchange and economic activity 42 (Alekseev, 1987). Aufeis are found in permafrost regions such as Alaska (Slaughter, 1982), 43 Siberia (Alekseev, 1987), Canada (Pollard, 2005), Greenland (Yde and Knudsen, 2005) and 44 others (Yoshikawa et al., 2007). Aufeis formation can result in significant economic expenses as 45 aufeis may negatively affect infrastructure and therefore natural resource extraction (Aufeis of 46 Siberia..., Nauka, 1981). Moreover, the springs that often feed aufeis may in some cases be the

only source of water for remote communities (Simakov, Shilnikovskaya, 1958). In Russia, aufeis
are found in the North-East, Transbaikal region, Yakutia, and West Siberia. Sokolov (1975)
estimated that the total aufeis water storage in Russia to be at least 50 km<sup>3</sup>, which approximately
equals the Indigirka River total annual streamflow.

51 The main hydrological role of aufeis is the seasonal redistribution of the groundwater 52 component of river runoff, where the winter groundwater discharge is released to summer 53 streamflow through melting of aufeis (Surface water resources, 1972). In most cases, the share of 54 the aufeis component in a river's annual streamflow accounts for 3-7%, reaching 25-30% in 55 particular river basins with an extremely large proportion of aufeis (Reedyk et al., 1995; Kane & 56 Slaughter, 1973; Sokolov, 1975). The most significant water inflow from aufeis melting takes 57 place in May-June (Sokolov, 1975). For example, the share of the aufeis flow accounts for more than 11% of total annual streamflow at the Indigirka River (gauging station Yurty, 51 100 km<sup>2</sup>). 58 59 In May, aufeis melt may represent 50% of monthly total streamflow, but decreases in June to 60 35% (Sokolov, 1975).

61 It is important to understand how climate change may impact aufeis formation because 62 warming has been observed in this region causing the transformation of permafrost (Romanovsky et al., 2007), glaciers reduction (Ananicheva, 2014) and hydrological regime 63 64 changes (Bring et al., 2016; Makarieva et al., 2018). Aufeis are formed by a complex connection 65 between river and groundwater. Many studies have reported the increase of minimum flow in Arctic rivers (Rennermalm and Wood, 2010; Tananaev et al., 2016), including those where 66 67 aufeis are observed in abundance (Makarieva et al., 2018, in review). A widely accepted hypothesis for permafrost regions is that a warming climate increases the connection between 68 69 surface- and groundwater that in turn leads to the increase of streamflow, both in cold seasons 70 and in annual flow (Bense et al., 2012; Ge et al., 2011; Walvoord et al., 2012; Walvoord and 71 Kurylyk, 2016). Variation and changes in aufeis extent can be assessed using remote sensing 72 techniques, where aufeis dynamics can serve as an indicator of groundwater change that is 73 otherwise difficult to observe (Topchiev, 2008; Yoshikawa et al., 2007).

The understanding of how aufeis respond to a warming climate varies. Alekseev (2016) suggests three to 11 year up and down cycles of aufeis maximum annual size, which may vary up to 25-30% in comparison with long-term average values. However, the same author (Alekseev, 2016) states a general tendency to the decrease of aufeis volume for the last 50-60 years in some aufeis-affected areas of Russia such as the Baikal region, South Yakutia, Kolyma region, Eastern Sayan Mountains, following the increase of global and local air temperature.

Some authors suggest that degradation of permafrost in the discontinuous and sporadic permafrost regions will lead to the decrease of the number of aufeis and even an almost complete disappearance. Meanwhile, in the zone of continuous permafrost in North-East Siberia, a climate warming of 2-3 °C is not projected to lead to significant changes in permafrost extent, but will increase the number and size of both through- and open taliks by the end of the 21th century (Pomortsev et al., 2010). Such a scenario may result in the reduction of area of large aufeis and formation of new small aufeis (Pomortsev et al., 2010).

In Alaska as well, no significant changes were documented in the area and volume of aufeis over the past few decades or even a century (Yoshikawa et. al, 2007). They suggested that the formation and melting of ice is less dependent on climate and more so on the source (spring) water properties such as temperature and volume.

91 In 1958, Simakov and Shilnikovskaya (1958) compiled and published a map inventory of 92 aufeis of the North-East USSR (scale 1:2 000 000). Since then, there has been no update on the 93 information on aufeis in this region, apart from some specific studies. In 1980-1982, an 94 inventory of aufeis in the zone of the Baikal-Amur Mainline was published (Catalog of 95 Aufeis..., 1980, 1981, 1982). Markov et al. (2017) summarized the results of field studies on 96 aufeis in the southern mountain taiga of Eastern Siberia from 1976 to 1983. Grosse and Jones 97 (2011) compiled the spatial geodatabase of frost mounds (or pingos) for northern Asia from 98 topographic maps. Further, the glacier science community has mapped past and recent glacier

99 cover across the globe (GLIMS and NSIDC, 2005, updated 2017). However, as far as the authors100 are aware, no electronic catalogue of aufeis exists.

101 The aim of this study is to update the inventory of aufeis in the North-East of Russia using 102 Landsat images, as well as to develop an electronic catalogue, which will contain data on historic 103 and current location and characteristics of aufeis. Here we present work that has been completed 104 for the Indigirka River basin (down to the Vorontsovo gauging station, 305 000 km<sup>2</sup>).

105 The new database, which includes geographic information system (GIS) formatted files, is 106 freely available (Makarieva et al., 2018) and can be used both for both scientific purposes and 107 for solving practical problems such as engineering construction and water supply studies.

### 108 **2.** Study region

109 The study region is the Indigirka River basin, which is located in Northeastern Siberia and 110 covers an area of 305 000 km<sup>2</sup> (Fig. 1). Most of the basin is represented by highlands with a 111 number of mountain ranges (< 3 003 m) including the Cherskiy and Suntar-Khayata mountains. 112 The lowland elevation reaches heights up to 350 m.

113 The climate of the study area is distinctly continental with annual average and lowest 114 monthly air temperature varying from -16.1 and -47.1 °C, respectively, at the Oymyakon 115 meteorological station (726 m, 1930-2012) to -13.1 and -33.8 °C, respectively, at the 116 Vostochnaya station (1 288 m, 1942-2012). Most precipitation (over 60%) occurs in the summer 117 season. Average annual precipitation at the Oymyakon weather station is 180 mm and at the 118 Vostochnaya station 278 mm.

119 The Indigirka River basin is located in the zone of continuous permafrost. Permafrost 120 depth can reach 450 m in the mountains, up to 180 m in river valleys and intermountain areas, 121 with talks found in river beds and fractured deposits. The hydrogeological regime is affected by 122 the active layer, which varies from 0.3 m to over 2 m (Explanatory note ..., 1991). The river 123 runoff regime is characterized by high snowmelt freshet, summer-autumn rainfall floods, and 124 low winter flow. In winter, small- and medium-sized rivers completely freeze. Freshet starts in 125 May-June and lasts for approximately 1.5 months. Melt waters from aufeis, glaciers, and snow 126 patches add to the river discharge in summer.

In total, about 10 000 aufeis with a total combined area of about 14 000 km<sup>2</sup> (Sokolov, 128 1975) are known in North-East Russia. The watershed area covered by aufeis varies from 0.4 to 129 1.3%, reaching 4% in some river basins (Tolstikhin, 1974). Most aufeis are of ground water 130 origin; significantly less often they are formed out of river waters or are of a mixed type 131 (Tolstikhin, 1974).

132

## 133 **3. Materials and methods**

134

# 3.1 The database of aufeis based on the Cadastre (1958) and topographic maps

The inventory map (scale 1:2 000 000) and the Cadastre of aufeis of the North-East of the USSR (Simakov, Shilnikovskaya, 1958), hereinafter referred to as the Cadastral Map and the Cadastre, became the first summarizing quantitative work on aufeis within the territory. The effort was carried out in the framework of the Central complex thematic expedition of the North-East Geological Survey of the USSR.

The Cadastre contains data on 7 448 aufeis of different size and over 2 000 boolgunyakhs (frost mounds). Of the total number of aufeis, 7 006 are plotted based on air-photo interpretation data, and another 442 on geological reports from field data. It should be noted that aufeis were identified based on geomorphologic features, meaning that in some cases only the areas or river valleys with aufeis were identified but not aufeis themselves.

In the Cadastre (1958) and our digitalization, the following characteristics of the aufeis are presented: location (the name of the river, the distance from the mouth or source), size (maximum length, average width, and area) and the dates of ice recording in aerial images (ranging from 08.06.1944 to 27.09.1945). Areas of the aufeis were evaluated via planimetering. Only very large aufeis (>  $3.3 \text{ km}^2$ ) were plotted on the Cadastral Map (1958), while the others are shown as point locations. Each aufeis on the Cadastral Map (1958) has its corresponding number, whose identifier and corresponding information can be found in the Cadastre (1958). As noted by Simakov and Shilnikovskaya (1958), some very small aufeis (<0.01 km<sup>2</sup>) could have been missed due to their indecipherability on aerial images, or they might have already melted by the time of the aerial photography. The example of the Cadastral Map's sheet (1958) for the Indigirka River upper reaches is presented in fig. 2.

Here, we developed the GIS database of aufeis in the Indigirka River basin up to the crosssection at the Vorontsovo gauging station based on the Cadastre (1958) and topographic maps. Our compilation contains data on 896 aufeis. The aufeis are presented as point objects in our database. The areas are specified for only 808 aufeis. The total area of all the aufeis with specified area accounts for 2063.6 km<sup>2</sup> and the areas of individual aufeis vary from 0.01 to 82 km<sup>2</sup>.

In the Cadastre, the dates of ice recording for 592 aufeis (66%) are presented, based on aerial images within the study area. The average seasonal date of recording is August 2, ranging from June 8 to September, 27. The dates of ice recording for the remaining 34% of the aufeis were not described, meaning that aufeis detection could be carried out based not on the visible ice presence at the aerial images but on geomorphological features of river valleys. Therefore, the Cadastre might contain data on old aufeis glades, where the aufeis themselves were absent.

168 Spatial positioning of the Cadastral Map of aufeis was conducted using the location 169 description by Russian topographic maps with the scale of 1:200 000. Grosse and Jones (2011) 170 used the same set of maps for compiling the dataset of pingos (frost mounds) in northern Asia 171 and described those maps in details therein. The maps at 1:200 000 scale were based on more 172 detailed maps of 1:50 000 and 1:100 000 scale, which were derived from aerial photography 173 acquired in the 1970–1980's. The use of 1: 200 000 scale guarantees the position assessment 174 precision to within 100 m. Each map sheet was visually searched for aufeis and identified aufeis were marked with an area polygon in a GIS layer. The locations of 330 aufeis (area 358 km<sup>2</sup>) 175 176 were determined based on topographic maps. When digitized, a point was plotted in the middle 177 of an aufeis at a topographic map.

The locations of the remaining aufeis were determined with the positioned map of the Cadastre. Additionally, 11 aufeis were found, which were absent in the Cadastre, but present in the topographic maps. Aufeis areas were estimated by digitalization of the maps. Areas of the remaining aufeis were estimated with the Cadastre. It was not possible to estimate the area of 88 aufeis, as they were not drawn on the topographic maps and only their location, but not area, was stated in the Cadastre.

184

Table 1 contains the structure of the GIS dataset of aufeis according to the Cadastre.

185

## 3.2 Identification of aufeis based on Landsat data

Aufeis location and area are relatively easy to determine using Landsat and/or Sentinel-2 images, received immediately after snow cover melting. Snow and ice are known to be characterized by relatively high reflectance in the visible and near infrared spectral bands and its significant decrease in mid infrared band. Normalized Difference Snow Index (NDSI) is based on this pattern and is calculated according to the formula (Hall et al., 1995):

191

$$NDSI = (GREEN - SWIR1) / (GREEN + SWIR1)$$

where SWIR1 is reflectance in mid infra-red band  $(1.56 - 1.66 \,\mu\text{m}$  for the Landsat-8 images), and GREEN is reflectance in the green band  $(0.525 - 0.6 \,\mu\text{m}$  for the Landsat-8 images). Following Hall et al. (1995), the threshold value for snow and ice is set at 0.4. Apart from using NDSI, other indices have been suggested to detect aufeis by Landsat images (but not used here). These are Normalized Difference Glacier Index (NDGI) and Maximum Difference Ice Index (MDII). Their advantages and disadvantages are discussed by Morse and Wolfe (2015).

198 Landsat-based detection of aufeis required some additional data to exclude other surface 199 types with similar spectral characteristics, such as snow-covered areas, turbid water, etc. It is 200 problematic to separate floodplain lakes from aufeis by late-spring satellite images, because 201 many of these lakes are still ice-covered in May-June. Morse and Wolfe (2015) recommended 202 creating a mask of water surface by mid-summer images (when all water bodies are already not 203 covered by ice), to exclude them from further analysis.

Aufeis detection in the Indigirka River basin was carried out based on the Landsat-8 OLI satellite images, 2013-2017, downloaded from the United States Geological Survey web-service (https://earthexplorer.usgs.gov). We used Landsat 8 collection 1 level-one terrain-corrected product (L1T) with radiometric and geometric corrections. In total, 33 images completely covering the Indigirka river basin were processed. We selected late-spring images (between 15 May and 18 June), to detect the maximum possible number of aufeis, since in June they melt intensively. There was between 1-20% of cloudiness in some images.

Preprocessing of the images was performed with the use of Semi-Automatic Classification Plugin module (QGIS 2.18). It includes the calculation of surface reflectance and atmospheric correction by Dark Object Subtraction (DOS1) image-based algorithm, described by (Chavez, 1996).

215 The Aufeis detection algorithm was realized in ArcGIS with the help of the ModelBuilder 216 application. Apart from the Landsat images, the digital terrain model (DTM) GMTED2010 217 (Danielson and Gesch, 2011) with a spatial resolution of 250 m was used to build a network of 218 thalwegs within the study basin. This is essential for semi-automated separation of the aufeis 219 from snow-covered areas in late-spring Landsat images. Indeed, almost all aufeis are located 220 either at streams or thalwegs, or in immediate proximity to them. On the contrary, the snow 221 cover in late spring mainly remains on mountain ridges and other elevated locations, i.e. 222 relatively far from thalwegs. Based on the preliminary analysis of aufeis location in relation to 223 the created network of thalwegs, we found that a 1.5 km wide buffer zone around the thalwegs 224 covers almost all aufeis. So, snow and ice covered areas, which are located outside this buffer, 225 are excluded from further analysis.

The process of aufeis detection by Landsat images consisted of the following steps:

• Detection of snow-ice bodies with the NDSI threshold of 0.4.

226

227

231

232

233

234

- Creation of a water mask with threshold values of the Normalized Difference Water Index (NDWI) (taken equal to 0.3), and reflectance in the near-infrared band (taken equal to 0.04).
  - Extraction of the detected snow-ice bodies by the buffer zone around thalwegs (1.5 km wide).
  - Conversion to vector format, area calculation and removal of objects smaller than 5 Landsat pixels (0.45 ha).

The suggested algorithm allows successful aufeis detection if an image is predominantly snow-free. At the end of May/early June, many aufeis in mountain regions are still covered by snow. Their detection required later images, obtained in mid-June.

Morse and Wolfe (2015) suggested a new spectral index MDII for automatically distinguishing snow bodies from ice ones. However, here some of the high elevation aufeis were partially covered with snow at the image acquisition time. Instead of automatic processing, the outlining of high elevation aufeis was conducted manually when snow cover was present, with separation of aufeis from adjacent snow covered areas.

Further, during melt season, the aufeis often divide into several neighboring areas. When assessing the number of aufeis with satellite data, it is therefore necessary to aggregate the areas into one aufeis, if they are located at a distance <150 m (or five Landsat pixels) from each other, and within one aufeis glade.

As a result of semi-automated processing of Landsat images, aufeis with a total area of 1 248 253.9 km<sup>2</sup> were detected. During the subsequent comparison with the Cadastre data (see section 249 3.3 for more details), over 100 aufeis, with a total area of 33.5 km<sup>2</sup>, were delineated manually. 250 The gaps were mainly due to the presence of snow cover and/or cloud coverage in the images. 251 To reduce the number of gaps, two to three images of the same territory were used. The total number of aufeis, identified with the Landsat images in the Indigirka River basin, was 1 213 and
 their total area 1 287.4 km<sup>2</sup>. Therefore, an omission error of automatic aufeis detection can be
 estimated as 2.7% of their total area.

The structure of the GIS dataset of aufeis according to Landsat images is presented in Table 2.

257 258

### 3.3. Cross reference between historical and satellite-based aufeis data collection

259 Cross-verification of aufeis data collections by the Cadastre (1958) and satellite imagery 260 was performed in two steps. At the first step, we found the closest aufeis in the Landsat-derived dataset for each aufeis from the Cadastre data if the distance between them was less than 5000 261 262 m. The determination of search radius was based on a preliminary analysis of the aufeis locations 263 by the Cadastre in relation to Landsat-based dataset. As a result, the cross index (identifier of the 264 closest aufeis in the Landsat-derived dataset) and minimum distance (m) to the closest aufeis 265 were determined for aufeis from the Cadastre. For Landsat-based dataset, the cross index is the 266 key field for the reference to the dataset from the Cadastre.

At the second step, a full manual verification was performed to find the mistakenly interrelated aufeis. For example, if the closest aufeis from the Cadastre and from the Landsatbased dataset were at a distance of less than 5000 m, but in different thalwegs, they were considered as different (unrelated) aufeis.

In total, 260 aufeis from the Cadastre were not verified by Landsat images. For them, the
NoData value (-9999) was set in the Cross Index and Distance fields of attributive table (see
Table 1 with the structure of GIS dataset from Cadastre).

#### 275 **4. Results**

276

#### 4.1 Comparison of the historical and modern data collection

277 The results of the comparison are presented in Table 3. In total, 634 aufeis from the 278 Cadastre were found by the Landsat images. They correspond to 611 aufeis identified with the 279 images, meaning that in 23 cases, one aufeis in an image corresponds to two aufeis in the 280 Cadastre. But 262 aufeis from the Cadastre were not detected by the satellite images. Those are 281 mainly small aufeis, which melt by the middle of June. However, among them there are also 43 282 large aufeis over 1 km<sup>2</sup> (fig. 3-a). It is likely that since the mid-20th century, when the field 283 observations were conducted and the Cadastre of aufeis was compiled, some aufeis could have 284 disappeared.

A little over half of the aufeis detected by Landsat images are included in the Cadastre: a total of 602 aufeis detected (the total area of 250.4 km<sup>2</sup>) are not included in the Cadastre (fig. 3b). Such a significant difference can be caused by the following reasons:

1. In some cases a single aufeis, according to the Cadastre, corresponds with two or moreaufeis by satellite image;

290 2. Aufeis are characterized by significant interannual variability, which results in possible
291 formation of new aufeis in areas where they previously were not observed (Alekseev, 2015;
292 Pomortsev et al., 2010; Atlas of snow..., 1997).

Total aufeis area evaluated based on satellite images, appeared to be 1.6 times smaller than stated in the Cadastre (1958). First and foremost, such difference can be explained by the fact that it was not the area of the aufeis themselves, but instead the aufeis glades, that were reported in the Cadastre (1958) and this corresponds to the maximum aufeis area during one or several seasons. With the satellite data, the areas of the aufeis themselves were assessed and when mid-June images were used, the aufeis area was significantly smaller than the typical annual maximum.

Aufeis area distribution according to the Cadaster and satellite data is shown as Lorenz curves (fig. 4). In both cases, the shape of the curves signifies a high degree of irregularity which is similar: 10% of the largest aufeis make up 61 and 57% of their total area according to theLandsat and the Cadastre data, respectively.

The cross-verification of the Cadastre and satellite data show that almost 60% of aufeis that are unconfirmed in the Landsat imagery and that are therefore only present in the Cadastre, have an individual aufeis area less than 0.25 km<sup>2</sup> (Fig. 5-a). The confirmed aufeis account for about 20% of the area stated in the Cadastre. Thus, it was mainly small aufeis that were not confirmed in the Landsat images. Conversely, Fig. 5-b shows that almost 60% of the aufeis detected in the Landsat images but not listed in the Cadastre have an area each of less than 0.25 km<sup>2</sup>.

### 311 **4.2. Aufeis distribution by elevation**

312 In general, aufeis distributions by elevation as assessed with the Cadastre and Landsat data 313 are quite similar, although there are some differences that are elevation-specific (fig. 6). Most 314 aufeis are located in the elevation band of  $1\ 000 - 1\ 200\ m$ . At lower elevations (up to 800 m) the number of aufeis according to Landsat data is higher than stated in the Cadastre. At the 315 316 elevations of 1 400-2 000 m, more aufeis are identified in the Cadastre data than by the satellite images. This can be explained by the fact that many aufeis located at high altitudes often have a 317 318 small area, so they could have been missed during the analysis of the satellite data. Further, they 319 could have been covered with snow at the image acquisition time, which would increase the 320 possibility of them being missed.

The elevation band of 200-300 m is characterized by the location of large aufeis. Though less than 2.5% and 5.0% of aufeis by the Cadastre and Landsat images are situated here, they represent about 11 and 13% of aufeis area from the datasets respectively (fig. 7).

324325

#### 4.3 Aufeis distribution by river basins

In the Indigirka River basin, there are several zones with a high density of aufeis: in the southern part (the Suntar and Kuidusun Rivers basins), as well as in the central part (Chersky Range slopes) (fig. 8). The largest aufeis identified with satellite images are located in the Syuryuktyakh River basin on the north-east slopes of the Chersky Range. Meanwhile, aufeis are almost absent in the northernmost (lowland) part of the Indigirka basin.

We analyzed the aufeis coverage for six river basins with available streamflow data. The headwater part of the Indigirka River, with the gauge near the Yurty village (area 51 100 km<sup>2</sup>), is the basin with the largest aufeis coverage (Table 4). Correlation between average elevation of the basins and their aufeis coverage (expressed as a percentage) is statistically significant. Among 6 basins, the Spearman rank correlation coefficients between the basin average elevation and aufeis percentage are 0.71 and 0.77 by the Cadastre and satellite data, respectively.

337

### 4.4 Aufeis area interannual variability

The assessment of aufeis area interannual variability was conducted in two areas: for the Bolshaya Momskaya aufeis, which is located in the Moma River channel (area in the Cadastre is 82 km<sup>2</sup>), and for a group of large aufeis (total area in the Cadastre is 287.8 km<sup>2</sup>) in the Syuryuktyakh River basin, which is the left-bank tributary of the Indigirka River.

Cloudless images from Landsat-5 (TM), Landsat 7 (ETM+) and Landsat-8 (OLI) were used with the acquisition dates between May 1 and June 30. In the USGS archives, there are no Landsat-5 images for the study territory for the 1984-2007 period. This limits the duration of satellite observations on aufeis to the period since 1999 (when the Landsat-7 satellite was launched). Also, the clouds complicate the acquisition of representative data. The list of the acquisition dates and assessed aufeis area values are presented in Table. 5.

Both areas are located at low elevations (Bolshaya Momskaya 430 to 500 m and Syuryuktyakh 200 to 500 m), which contributes to the relatively early and intensive aufeis melt in spring. The aufeis reach their maximum area by the beginning of May. Using the available satellite images it is impossible to make a reliable conclusion on aufeis area increase or decline,
 because the acquisition dates vary significantly from year to year. However, it is possible to
 make some conclusions based on the available data:

1. In 2002-2017 the Bolshaya Momskaya aufeis did not reach the maximum area stated in the Cadastre (82 km<sup>2</sup>), even though the satellite image was acquired during the first week of May (2005) when aufeis melting had not yet started. Comparing two images, taken in similar conditions (08.05.2005 and 15.05.2013), it was found that aufeis area in 2013 was smaller by 18.1 km<sup>2</sup> than in 2005. Accordingly, the Bolshaya Momskaya aufeis may have seen a decreasing trend over time in its maximum coverage.

2. The area of the largest aufeis in the Syuryuktyakh River basin in May 2014 was 78.0 km<sup>2</sup>, which is 8 km<sup>2</sup> larger than stated in the Cadastre. One may note also that the maximum aufeis areas in the Syuryuktyakh River basin were detected by the images received at the end of the period (2014-2017), including mid-June (18.06.2015). Therefore, it can be suggested that the aufeis areas within Syuryuktyakh River basin have not decreased since 2002.

#### 5. Discussion

366

The most important uncertainty in the obtained results relates to our ability to draw a conclusion on the long-term trend of total aufeis area comparing the historical and satellitederived datasets. The total area of aufeis estimated by Landsat images is 38% less than according to the Cadastre. Is it possible to confirm that such a significant reduction in the aufeis area really occurred? Considering this issue, it is important to emphasize some limitations of the methodology and the created datasets.

373 The main limitation of the historical aufeis dataset is that the Cadastre provides an area of 374 aufeis glades, but not the aufeis themselves. Simakov and Shilnikovskaya (1958) noted that the 375 areas of aufeis glades match the average annual maximum of the ice-covered area. Alekseev 376 (2005) states that the assessment of the stages and patterns of the development of aufeis glades 377 based on the analysis of their landscape and geomorphological features is difficult due to the lack 378 of research on temporal aspects of mutual transitions of landscape facies and their factorial 379 dependencies. However, studying the aufeis landscapes in the central part of Eastern Sayan 380 Mountains, Alekseev (2005) assumes that the vegetation community which is a typical indicator of aufeis development may persist for 200-300 years after the beginning of aufeis processes 381 382 attenuation.

The satellite-derived assessment of the aufeis area has the following main source of uncertainty. It is often impossible to determine the maximum area of aufeis by satellite images, since it is observed at the beginning of snow melt season, when aufeis are still covered with snow. In late spring and beginning of summer, the area of aufeis may already been significantly reduced in comparison with the maximum values, due to melting and mechanical destruction.

388 Maximum intensity of aufeis melt in the studied region is observed in June when spring 389 flood river streams actively erode the aufeis surface. Sokolov (1975) reported the results of the 390 observations at the Anmyngynda aufeis carried out in 1962-1965. This aufeis is located in the 391 upstream area of the Kolyma river basin (723 m a.s.l.) and may be used as being representative 392 of the mountainous part of the studied region. In 1962-1965, the aufeis area changed from 5.1 to  $6.2 \text{ km}^2$  with mean maximum area of  $5.7 \text{ km}^2$ . Aufeis melt has been observed to begin on 393 average on the 10<sup>th</sup> of May. During May, the aufeis area decreased by 15% of the total area on 394 395 average. At the end of June, the remaining area was 34% of the maximum, i.e. during this month 396 more than 50% of the aufeis area has been destroyed. In the period from July to September, the 397 melting slowed down: in July the aufeis decreased by 22%, in August by 8%, in September by 398 3%. The area of aufeis at lower absolute elevations decreases faster at first half of the summer, 399 and in the upstream areas – in the second one (Sokolov, 1975).

400 Some aufeis in the mountainous regions could be missed by satellite images, since they 401 can be covered with snow until the end of June. However, their contribution to the total area is 402 non-significant.

Taking into account all the above-described limitations, and also that more than 600 aufeis that were missing in the Cadastre were found by Landsat images, we conclude that it is not correct to make a conclusion about long-term trends of aufeis area based on the entire created dataset. Following Pavelsky and Zarnetske (2017), we decided to examine only several of the largest aufeis deposits in order to identify the long-term trend.

We selected the 38 largest aufeis with an area  $\geq 10 \text{ km}^2$  according to the Cadastre dataset, 408 409 confirmed by satellite data. Their total area decreased from 858.1 km<sup>2</sup> according to the Cadastre to 356.3 km<sup>2</sup> according to recent Landsat images. Conversely, we also selected the largest aufeis 410 according to satellite data (18 aufeis with satellite-estimated area >10 km<sup>2</sup>). Their total area also 411 412 decreased significantly (from 428.6 km<sup>2</sup> according to the Cadastre to 343.5 km<sup>2</sup> according to Landsat images). We also analyzed 8 giant aufeis with areas  $\geq 35 \text{ km}^2$  according to the Cadastre 413 414 dataset. They all were confirmed by the satellite images; however seven of the eight had a 415 significantly smaller area (from 2 to 21  $\text{km}^2$ ) with the decrease being 2-10 times. Only one giant 416 aufeis in the Syuryuktyakh River basin has the area by Landsat larger than by Cadastre, at 72 and 64 km<sup>2</sup> accordingly. It should be noted that the formation of new (mainly small) aufeis can 417 slightly reduce the rate of the aufeis area decrease. 418

419 420

# 6. Conclusion

421 The research conducted here is the first step of the study aimed at the development of a 422 GIS database of the aufeis of North-East Russia. Historical data of the Cadastre (1958) and 423 topographic maps were used to create a geodatabase of aufeis in the Indigirka River basin (up to the Vorontsovo gauge, with the area of 305 000 km<sup>2</sup>). It contains historical data on 896 aufeis 424 with total area of 2063.6 km<sup>2</sup>. Aufeis detection was conducted for the 2013-2017 period using 425 Landsat imagery with 1213 aufeis identified having a total area of 1287.4 km<sup>2</sup>. The historical 426 427 dataset from the Cadastre (1958) and more recent satellite-based dataset were compared and 428 combined in the joint Catalogue of aufeis within the Indigirka River basin, available at the 429 PANGAEA repository (https://doi.pangaea.de/10.1594/PANGAEA.891036).

Recent total aufeis area is 1.6 times smaller than stated in the Cadastre (1958). The more significant changes occurred to 38 large and giant aufeis (area  $\ge 10 \text{ km}^2$ ) with total decrease of area by 501.8 km<sup>2</sup> (or 66% of the total reduction). Simultaneously, the historical Cadastre archive is lacking data on over 600 aufeis that were identified using satellite images. This suggests that the Cadastre data is incomplete, while there may also have been significant change in aufeis formation conditions in the last half century.

The analysis of large and giant aufeis seems to indicate that there has been a significant decrease in aufeis area over the period of last 70 years. Additional analysis of historical aerial photography data could help to clarify the issue of aufeis area decline trend since the middle of the 20th century to the present. One of the further study goals will be to find out the extent to which these changes are climate-derived and to identify their impact on river streamflow.

441

442 Acknowledgements. The authors are grateful to David Post, Anna Liljedahl and an
 443 anonymous reviewer for valuable comments and assistance with English.
 444

## 445 **7. References**

446

Alekseev, V.R.: Naledi. Novosibirsk, Nauka, Moscow, 1987 (in Russian).

- 447 Alekseev, V.R.: Long-term variability of the spring taryn-aufeises. Ice and Snow, 56(1), 448 73-92, 2016. doi:10.15356/2076-6734-2016-1-73-92.
- Alekseyev, V.R.: Cryogenesis and geodynamics of icing valleys. Geodynamics &
  Tectonophysics 6 (2), 171–224, 2015. doi:10.5800/GT-2015-6-2-0177.
- 451 Alekseev, V.R.: Landscape indication of aufeis phenomena. Novosibirsk, Nauka, 2005,
  452 364 p. (in Russian).

453 Ananicheva, M.D.: Estimation of the areas, volumes and heights of the boundary of the 454 feeding of glacier systems of the Northeast of Russia from the space images of the beginning of 455 the 21st century. Ice and Snow, 1 (125), 35-48, 2014.

- 456 Atlas of snow and ice resources of the world, 11 (2). Institute of geography RAS, Moscow,
  457 270 pp., 1997 (in Russian).
- 458 Aufeis of Siberia and the Far East.Nauka, Novosibirsk, 244 pp. 1981. (in Russian).

Bense, V. F., Kooi, H., Ferguson, G., and Read, T.: Permafrost degradation as a control on
hydrogeological regime shifts in a warming climate. Journal of Geophysical
Research, 117(F03036), 2012. doi: 10.1029/2011JF002143.

- Bring, A., Fedorova, I., Dibike, Y., Hinzman, L., Mård, J., Mernild, S. H., Prowse, T.,
  Semenova, O., Stuefer, S. L., and Woo, M.-K.: Arctic terrestrial hydrology: Asynthesis of
  processes, regional effects, and research challenges, J. Geophys. Res.Biogeosci., 121, 621–649,
  doi:10.1002/2015JG003131, 2016.
- 466 Cadastre to the map of the north-east of the USSR Scale 1: 2000000. Shilnikovskaya Z.G.
  467 Central complex thematic expedition of the North-Eastern geological department. Magadan, 398
  468 pp., 1958. (in Russian).
- 469 Catalogue of Aufeis in the Baikal-Amur Railroad Zone, 1. Aufeis in the Upper Part of the
  470 Chara River Basin. Gidrometeoizdat, Leningrad, 63 pp., 1980 (in Russian).
- 471 Catalogue of Aufeis in the Baikal-Amur Railroad Zone, 2. Aufeis in the Muya River Basin.
  472 Gidrometeoizdat, Leningrad, 84 pp., 1981 (in Russian).

473 Catalogue of Aufeis in the Baikal-Amur Railroad Zone, 3. Aufeis in the Upper Angara
474 River Basin. Gidrometeoizdat, Leningrad. 96 pp., 1982 (in Russian)Chavez Jr.P.S. Image-based
475 atmospheric corrections – revisited and improved. Photogram. Eng. and Remote Sens, 1996,
476 62(9), 1025–1036.

- 477 Danielson, J.J., and Gesch, D.B.: Global multi-resolution terrain elevation data 2010
  478 (GMTED2010): U.S. Geological Survey Open-File Report 2011–1073, 26 pp., 2011.
- 479 Explanatory note to the geocryological map of the USSR, scale 1: 2 500 000, 125 pp., 1991480 (in Russian).

481 Ge, S., J. McKenzie, C. Voss, and Wu, Q.: Exchange of groundwater and surface-water
482 mediated by permafrost response to seasonal and long term air temperature variation, Geophys.
483 Res. Lett., 38, L14402, 2011. doi:10.1029/2011GL047911.

GLIMS and NSIDC: Global Land Ice Measurements from Space glacier database.
Compiled and made available by the international GLIMS community and the National Snow
and Ice Data Center, Boulder CO, U.S.A., doi:10.7265/N5V98602, 2005, updated 2017.

487 Grosse, G. and Jones, B. M.: Spatial distribution of pingos in northern Asia, The 488 Cryosphere, 5, 13-33, 2011. https://doi.org/10.5194/tc-5-13-2011.

Hall, D K., Riggs, G.A. and Salomonson, V.V.: Development of methods for mapping
global snow cover using Moderate Resolution Imaging Spectroradiometer (MODIS) data.
Remote Sensing of Environment, 54, 127-140, 1995.

Kane, D. L. and Slaughter C. W.: Recharge of a central Alaska lake by subpermafrost
groundwater. Nat. Acad. Sci, 38, 458-462, 1973.

Makarieva, O., Nesterova, N., Post, D. A., Sherstyukov, A., and Lebedeva, L.: Warming
temperatures are impacting the hydrometeorological regime of Russian rivers in the zone of
continuous permafrost, The Cryosphere Discuss., https://doi.org/10.5194/tc-2018-157, in review,
2018.

Makarieva, O., Shikhov, A., Ostashov, A., Nesterova, N.: Aufeises (naleds) of the NorthEast of Russia: GIS catalogue for the Indigirka River basin (Russia), PANGAEA, 2018.
available at: https://doi.pangaea.de/10.1594/PANGAEA.891036.

501 Map of the icings of the North-East of the USSR Scale 1: 2000000. Simakov, A.S., 502 Shilnikovskaya, Z.G. North-Eastern Geological Administration of the Main Directorate of 503 Geology and Subsoil Protection. Central complex thematic expedition. Magadan. 1958 (in 504 Russian).

505 Markov, M. L., Vasilenko, N. G. and Gurevich, E. V.: Icing fields of the BAM zone: 506 expeditionary investigations. SPb.: Nestor-History, 320 pp., 2016 (in Russian).

Morse, P.D. and Wolfe, S.A.: Geological and meteorological controls on icing (aufeis)
dynamics (1985 to 2014) in subarctic Canada. Journal of Geophysical Research: Earth Surface.
120, 1670–1686, 2015.

Pavelsky, T.M. and Zarnetske, J.P.: Rapid decline in river icings detected in Arctic Alaska:
Implications for a changing hydrologic cycle and river ecosystems. Geophysical Research
Letters. 44(7), 3228–3235, 2017.

Pollard, W. H.: Icing processes associated with high Arctic perennial springs, Axel
Heiberg Island, Nunavut, Canada, Permafrost Periglacial Processes, 16, 51 – 68, 2005.

515 Pomortsev, O.A., Kashkarov, E.P. and Popov, V.F.: Aufeis: global warming and processes
516 of ice formation (rhythmic basis of long-term prognosis). Bulletin of Yakutsk State University, 7
517 (2). 40-48, 2010 (in Russian).

518 Reedyk, S., Woo, M.K. and Prowse, T.D.: Contribution of icing ablation to streamflow in a
519 discontinuous permafrost area. Canadian Journal of Earth Science, 32: 13-20, 1995.

Rennermalm, A. K., Wood, E. F., and Troy, T. J.: Observed changes in pan-arctic coldseason minimum monthly river discharge, Clim. Dyn., 35(6), 923–939, 2010.

Romanovsky, V. E., Sazonova, T. S., Balobaev, V. T., Shender, N. I., and Sergueev, D. O.:
Past and recent changes in air and permafrost temperatures in eastern Siberia, Global Planet.
Change, 56, 399–413, doi:10.1016/j.gloplacha.2006.07.022, 2007.

525 Simakov, A.S. and Shilnikovskaya, Z.G.: The map of the north-east of the USSR. A Brief
526 Explanatory Note. The North-Eastern Geological Administration of the Main Directorate of
527 Geology and Subsoil Protection, Magadan, 40 pp., 1958 (in Russian).

- Slaughter, C. W.: Occurrence of and recurrence of aufeis in an upland taiga catchment, in:
  Roger J.E. Brown Memorial Volume—Proceedings of the 4th Canadian Permafrost Conference,
  Natl. Res. Counc. of Can., Ottawa, Ont, 182 188, 1982.
- 531 Sokolov, B.L.: Naleds and river runoff. Gidrometeoizdat, Leninrgad, 190 pp., 1975 (in Sussian).
- 533 Surface waters resources of the USSR, 17. Leno-Indigirsky district. Gidrometeoizdat, 534 Leninrgad, 651 pp., 1972. (in Russian).
- 535 Tananaev, N. I., Makarieva, O. M., and Lebedeva, L. S.: Trends in annual and extreme 536 flows in the Lena River basin, Northern Eurasia. doi: 10.1002/2016GL070796, 2016.
- 537 Tolstikhin, O.N.: Naleds and ground waters in the North-East of the USSR. Science, 538 Novosibirsk, 164 pp., 1974 (in Russian).
- 539 Topchiev A.G.: Methods of aerospace monitoring of aufeis geosystems and underground 540 water of cryolithozone. Systems and means of informatics, 2008, 18 (3), pp. 304-327 (in 541 Russian)
- 542 US Geological Survey Server: <u>http://earthexplorer.usgs.gov</u>, last access: 1 August 2018.

543 Walvoord, M. A., Voss C. I. and Wellman T. P.: Influence of permafrost distribution on 544 groundwater flow in the context of climate-driven permafrost thaw: Example from Yukon Flats 545 Basin. Alaska, United States. Water Resour. W07524, Res., 48, 2012. 546 doi:10.1029/2011WR011595.

- 547 Walvoord, M. A. and Kurylyk, B. L.: Hydrologic Impacts of Thawing Permafrost—A 548 Review. Vadose Zone Journal, 15 (6), 2016. doi: https://doi.org/10.2136/vzj2016.01.001.
- 549 Yde, J. C., and Knudsen, N. T. Observations of debris-rich naled associated with a major
  550 glacier surge event, Disko Island, west Greenland, Permafrost Periglacial Processes, 16, 319 –
  551 325, 2005.
- 552 Yoshikawa, K., Hinzman, L.D., and Kane, D.L.: Spring and aufeis (icing) hydrology in 553 Brooks Range, Alaska, Journal of Geophysical Research: Biogeosciences, 112, 1–14, 2007.
- 554

Table 1 The structure of GIS database of aufeis by Cadastre (1958)

Field name	Field alias	Description
FID	FID	Index number (Object ID)
AufDataSrc	Aufeis data	Aufeis Cadastre data (1958) (for all objects)
	source	
Auf_area	Aufeis area	Aufeis area $(km^2)$ from the Cadastre (1958). If the data was
	Cadastre	missing, the area was calculated by topographic maps (1980) scale
	$(km^2)$	1: 200 000
Auf_index	Aufeis	Index of the aufeis in the Cadastre (1958) (it contains 0 if the
	index	aufeis was missing in the Cadastre, but found in the topographic
	Cadastre	map (1980) scale 1: 200 000)
Map_index	Cadastre	Index of the Cadastre (1958) map
	map index	
Auf_topo	Aufeis in	Presence of the aufeis at topographic map (0 – missing, 1 –
	topo	present)
Auf_in_map	Aufeis in	Presence of the aufeis in the Cadastre $(0 - missing, 1 - present)$
	map	
Toponumber	Торо	Nomenclature of the topographic map sheet
	number	
Date	Date	Date of fixing the presence of ice within the aufeis
Long	Long	Longitude, degree
Lat	Lat	Latitude, degree
Elevation	Elevation	Height above sea level (determined by Aster GDEM), m
Comment	Comment	Comments (mainly typos in the Cadastre map, or the method of
		determining aufeis area)
CrossIndex	Cross index	Cross index of aufeis derived from Landsat (if aufeis is not in
		Landsat, the value is missing)
Distance_m	Distance	Minimum distance between the aufeis from the Cadastre and the
	(m)	same aufeis from Landsat image (m)

Table 2 The structure of GIS database of aufeis by Landsat images (2013-2017)

Field name	Field alias	Description		
FID	FID	Index number (Object ID)		
AufDataSrc	Aufeis data	Landsat images (for all objects)		
	source			
WRS2_ID	Landsat	The Landsat scene identifier in the WRS2 graph of the US		
	WRS2_ID	Geological Survey (USGS). The first three digits indicate the		
		column number, and last three digits represent the line number.		
Image_Date	Landsat	The date of image		
	image date			
Comment	Comment	Additional information, for example, if the aufeis was partly covered by clouds and additional images were used to estimate the area		
CrossIndex	Cross index	Identifier of aufeis by Landsat images (key field for the reference to		
		the Cadastre data)		
Auf_Area	Aufeis area	Aufeis area by Landsat image, km <sup>2</sup>		
	(km <sup>2</sup> )			
Elevation	Average	Average elevation of aufeis, calculated by Aster GDEM digital		
	elevation	elevation model		

## Table 3 Data correlation of aufeis based on the Cadastre (1958) and the Landsat images

Data source	Matching aufeis	Not confirmed aufeis	
Data source	number and area (km <sup>2</sup> )	number and area (km <sup>2</sup> )	
Cadastre (1958)	634 (1905.0)	262 (158.6)	
Landsat	611 (1037.0)	602 (250.4)	

- 562 Table 4 Aufeis area coverage (percentage) in the sub-basins within the Indigirka River watershed
- 563 by the Cadastre and Landsat data

River	Area, km <sup>2</sup>	Average elevation, m a.s.l.	% aufeis coverage (Cadastre)	% aufeis coverage (Landsat)
Suntar River –Sakharinya River mouth	7680	1460	0.97	0.78
Elgi – 5 km upstream of the Artyk-Yuryakh River mouth	17600	1104	0.49	0.23
Nera – Ala-Chubuk	22300	1174	0.32	0.26
Indigirka – Yurty	51100	1256	1.15	0.80
Indigirka – Indigirskiy	83500	1185	0.82	0.56
Indigirka – Vorontsovo	305000	803	0.68	0.41

565 Table 5 Aufeis area changes, 2001-2017.

Bolshaya Mo	omskaya aufeis	The group of aufeis in t	he Syuryuktyakh River basin
Imagery date Aufeis area, km <sup>2</sup>		Imagery date	Aufeis area, km <sup>2</sup>
17.06.2002	29.2	26.06.2001	69.7
08.05.2005	66.2	29.06.2002	100.6
27.05.2006	57.9	04.06.2007	155.1
19.06.2009	39.5	17.06.2009	117.5
25.05.2011	61.7	22.06.2011	89.5
27.05.2012	49.6	21.05.2014	268
15.05.2013	48.1	18.06.2015	164.8
18.06.2017	21.9	04.06.2016	206.4



Fig. 1 Geographical location of the Indigirka river basin





Fig. 2 Subset of the Cadastral Map of the North-East of the USSR from 1958 (sheet 7, upper reaches of the Indigirka River – the basins of the rivers Suntar, Agayakan and Kuydusun).





Fig. 3 Difference between aufeis location according to the Cadastre and satellite data: a) – aufeis are
absent in the image but present in the Cadastre (Landsat-8 image of 18.06.2017); b) – aufeis are absent
(or their area is understated) in the Cadastre but present in the image (Landsat-8 image of 30.05.2016).











Fig. 5 Aufeis area distribution: a) - according to the Cadastre data, confirmed and not confirmed by

Landsat images, b) - according to Landsat images, confirmed and not confirmed by the Cadastre.





Fig. 6 Aufeis distribution by elevation within the Indigirka River basin.



Fig. 7 Aufeis area distribution by elevation within the Indigirka River basin.



Fig. 8 Aufeis in the Indigirka River basin according to the Cadastre and Landsat images. Black
 outlines with section lining represent the zones where aufeis area interannual variability was
 assessed