



1 Glider data collected during the Algerian Basin Circulation 2 Unmanned Survey from 2014 to 2016

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23

24 **Abstract**

25

26 We present data collected in the framework of the Algerian Basin Circulation Unmanned Survey -
27 ABACUS project. ABACUS main objective is the monitoring of the basin circulation and of the
28 surface and intermediate water masses physical and biological properties in a key region of the
29 Mediterranean Sea circulation. Data have been collected through deep glider cruises in the
30 Western Mediterranean Sea during the autumns of 2014, 2015 and 2016. Glider missions were
31 realized in the Algerian Basin, between the Island of Mallorca and the Algerian Coast. Across the
32 three glider missions, eight repeated transects were obtained which enabled us to investigate the
33 basin scale circulation and the presence of mesoscale structures utilising both the adaptive
34 sampling capabilities of the gliders and the higher resolution of the data.

35 After collection, all data passed a quality control procedure and were then made available through
36 an unrestricted repository host by the SOCIB Data Centre at <https://doi.org/10.25704/b200-3vf5>.

37 The actual dataset spans three fall seasons, providing an important contribution to the data
38 collection in the chronically undersampled Algerian Basin.

39 Temperature and salinity data collected in the first 975 m of the water column allowed us to
40 identify the main water masses and describe their characteristics improving the understanding of
41 the dynamics of the region. On the time scale of the project, data show a large variability at the
42 surface layer and reduced variability at the intermediate and deep layers.

43 Our measurements have been successfully compared to data previously collected in the area from
44 1909 to 2011. Results showed similar overall distribution, ranges and variability as the historical
45 data, with no outliers in the surface or deep layers.

46

47



48 1 Introduction

49

50 The southwestern Mediterranean Sea is an important transit region characterized by the presence
51 of both fresh surface waters coming from the Atlantic (Atlantic Water, hereafter AW) and more
52 saline waters which typically reside in the Mediterranean region (Mediterranean Water, MW). At
53 intermediate levels (400 – 1000 m) there is the Levantine Intermediate Water (LIW), typified by
54 subsurface temperature and salinity maxima, while deeper layers (>1000 m) are occupied by the
55 Western Mediterranean Deep Water (WMDW) (Millot, 1999; Millot et al., 2006). Most of the
56 Western Mediterranean is occupied by the Algerian Basin (AB), a wide and deep basin comprised
57 between the Balearic Islands, the Algerian Coast and the Sardinia Channel, where an intense
58 inflow/outflow regime exists and complex circulation patterns take place (e.g., Pascual et al., 2013;
59 Cotroneo et al., 2016; Aulicino et al., 2018). Typically, AW entering through the Strait of Gibraltar
60 flows eastward at the AB surface, mainly inside the Algerian Current (AC), while more saline MW,
61 formed in the eastern and northern parts of the Mediterranean, flows westward at the
62 intermediate and deep layers (Millot, 1985; Testor et al., 2005). As previously demonstrated by
63 several studies, AW and MW interact at different scales, from basin-scale to mesoscale and sub-
64 mesoscale (Robinson and Golnaraghi, 1994; Fusco et al., 2003; Vidal-Vijande et al., 2011), allowing
65 a high seasonal and interannual variability in the basin. This aspect is favoured by the presence of
66 the AC. After leaving the Alboran Sea (Tintoré et al., 1991), this 30-50 Km wide along-slope current
67 flows eastward along the Algerian coast, carrying the AW eastwards (Testor et al., 2005). Typically,
68 the AC becomes unstable along its path due to complex hydrodynamic processes, and forms
69 several meanders which frequently evolves to isolated cyclonic and anticyclonic mesoscale eddies
70 (e.g., Millot, 1985; Moran et al., 2001; Ruiz et al., 2002; Font et al., 2004; Escudier et al., 2016;
71 Cotroneo et al., 2016; Pessini et al., 2018) promoting an intense mesoscale activity all over the AB.
72 These structures present high levels of kinetic energy (Pascual et al., 2013; Escudier, 2016) and
73 impact the distribution of physical and chemical properties of water masses, especially at surface
74 and intermediate depths (Taupier-Letage et al., 2003; Olita et al., 2011).

75 In the last two decades, both satellites and numerical simulations data have been largely used to
76 study mesoscale processes, partially balancing the scarcity of in situ observations. This improved
77 our knowledge of the large-scale surface features (Vignudelli et al., 2003; Isern-Fontanet et al.,
78 2016), but a complete understanding of mesoscale and submesoscale processes in the basin is still
79 needed. To this aim, frequent dedicated observations at higher horizontal and vertical resolution,
80 along the water column, are essential (Pascual et al., 2017). Multi-platform monitoring strategies,
81 which enable the integration of data from satellites, Autonomous Underwater Vehicles (AUV) and
82 numerical models, have already demonstrated their capabilities in the assessment of
83 oceanographic processes in different regions of the global ocean, such as the Atlantic Ocean
84 (Shcherbina et al., 2015) and the Mediterranean Sea (Carret et al., 2018; Troupin et al., 2018;
85 Pascual et al., 2017; Aulicino et al., 2016). The value of these is improved when implemented along
86 repeated monitoring lines coincident with the satellite groundtracks (Aulicino et al., 2018). Among
87 other AUVs, measurements at small sampling intervals collected through gliders (<5 Km spatial
88 resolution) have contributed to several ocean studies (Rudnick, 2016; Heslop et al., 2012), from
89 dynamics (e.g., Ruiz et al., 2009; Bosse et al., 2016; Cotroneo et al., 2016; Thomsen et al., 2016) to
90 physical and biogeochemical exchanges (e.g., Ruiz et al., 2012; Bouffard et al., 2010; Cotroneo et
91 al., 2016; Olita et al., 2017) or assessment of altimetry data (Heslop et al., 2017). However, gliders
92 generally have a limited speed (e.g. 25 cm/s in the horizontal) that can give synopticity problems
93 when monitoring mesoscale phenomena (Rudnick and Cole, 2011; Alvarez and Mourre, 2012;
94 Aulicino et al., 2016; Cotroneo et al., 2016) or wide areas (Liblik et al., 2016). Thus, their
95 combination with reliable satellite products, numerical simulations and, when available, other



96 platforms such as drifters, ARGO floats and ship-borne CTDs or a multiple glider missions
97 represent the good strategy for implementing observatories for marine science research which
98 aim to study basin scale and/or coastal processes (Aulicino et al., 2018) where issues of
99 synchronicity are a concern.

100 Furthermore, this integrated glider-based approach is also important for validating new satellite
101 products, such as the now fully-operational Sentinel-3 mission (Drinkwater and Rebhan, 2007;
102 Heslop et al., 2017) and the upcoming Surface Water and Ocean Topography (SWOT) wide-swath
103 radar interferometer (Fu and Ferrari, 2008), where multi-platform data are expected to help
104 distinguishing noise from small dynamical structures, as a result of cross calibrations between the
105 different sensors (Bouffard et al., 2010; Pascual et al., 2013).

106 In this study, we present data collected through three SLOCUM G2 (Figure 1) glider missions
107 carried out in the AB during fall 2014, 2015 and 2016. These cruises were developed along a
108 repeated monitoring line set across the AB and allowed the collection of a huge dataset of physical
109 and biological ocean parameters in the first 975 m of the water column. For each mission, glider
110 main track was chosen to lie under overflying altimeter satellites tracks, in order to optimize the
111 inter-comparison of the data from the two platforms. Additionally, during the 2014 and 2016
112 glider cruises, the sampling track was changed in order to sample specific mesoscale circulation
113 structures in the area. Even though this study is restricted to a particular area, i.e. the AB, we
114 strongly believe that the scientific community could use these data to enlarge the knowledge of
115 the western Mediterranean Sea and to refine the implementation of glider missions and
116 integrated multi-platform observatories in other regions. Glider characteristics, sensors details,
117 paths and mission strategies are described in Section 2; data format and quality control
118 procedures are presented in Section 3. In section 4, a comparison between ABACUS observations
119 and historical data, as well as sample ABACUS transects, are presented. Finally, Section 5 reports
120 main conclusions.

121



122

123 *Figure 1. The SLOCUM G2 glider during the ABACUS pre-mission test at sea. Credit M. Torner.*

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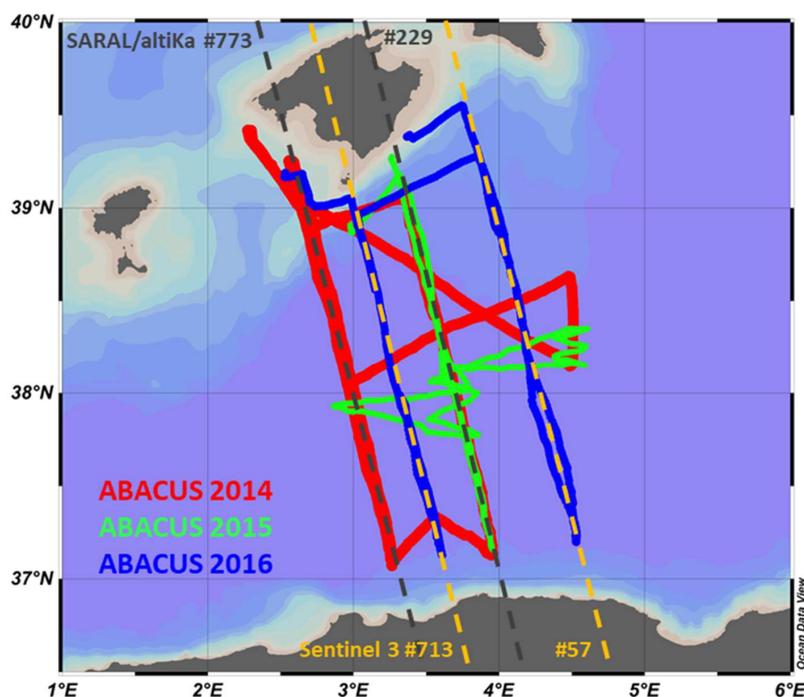


127 2 Glider field activities and technical details

128 Gliders are AUV able to provide high resolution hydrographic and bio-chemical measurements.
129 These vehicles control their buoyancy to allow vertical motion in the water column and make use
130 of their hydrodynamic shape and small fins to make horizontal motions. Changing the position of
131 the center of gravity with respect to the center of buoyancy helps controlling both pitch and roll
132 (Webb et al., 2001). This results in a typical saw-toothed navigation pathway to a maximum depth
133 of 1000 m.

134 During the autumns from 2014 to 2016, three deep SLOCUM G2 glider missions were carried out
135 in the AB (Figure 2) in the framework of the SOCIB (Balearic Islands Coastal Observing and
136 Forecasting System) Glider Facility Open Access Programme and supported by the Joint European
137 Research Infrastructure network for Coastal Observatories (JERICO) Trans National Access (TNA).
138 Funding agreements were developed under the JERCIO TNA third call (grant No 262584) for
139 ABACUS 1, the SOCIB external access (ABACUS 2) for ABACUS 2 and JERICO-NEXT TNA first call
140 (grant No. 654410) for ABACUS 3.

141 Up to 2016, a total of 8 repeated transects were realized between the Island of Mallorca and the
142 Algerian Coast. In 2014 and 2015, after the realization of the defined transects, the glider track
143 was changed in order to sample specific mesoscale structures in the study area (Figure 2).
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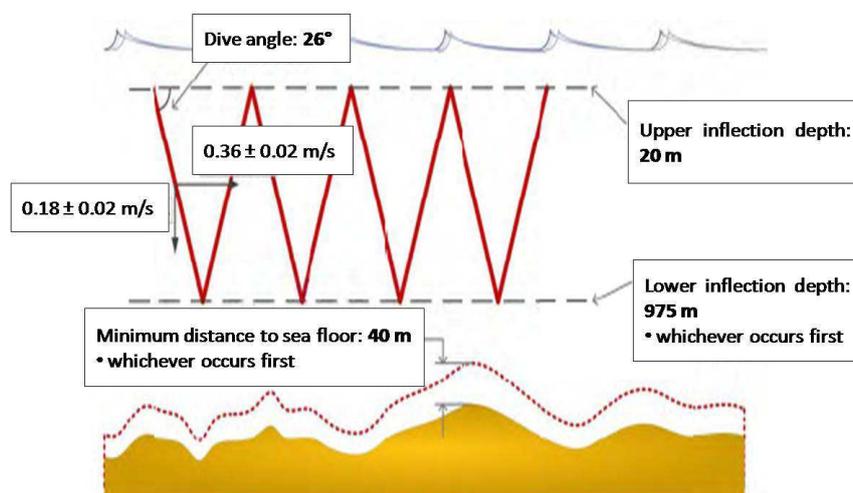


145

146 *Figure 2. Glider tracks during the missions of ABACUS 2014 (red dots), 2015 (green dots) and 2016*
147 *(blue dots). The deviations from the monitoring line were undertaken to sample some mesoscale*
148 *structures identified through near real time satellite altimetry and SST maps. The groundtracks of*
149 *the SARAL/altika (grey dashed lines) and Sentinel 3 (yellow dashed lines) satellites over the study*
150 *area are also showed.*



151 During these cruises ABACUS gliders collected Temperature, Salinity, Turbidity, Oxygen and
152 Chlorophyll (CHLA) concentration data in the first 975 m of the water column. Each mission had an
153 average duration of about 40 days and was always performed between September and December.
154 In situ data collection was supported by remotely sensed data from different platforms over the
155 Western Mediterranean Sea. In particular, gridded altimetry data provided by the Archiving,
156 Validation and Interpretation of Satellite Oceanographic data (AVISO), Sea Surface Temperature
157 (SST) and CHLA concentration information from MODerate resolution Imaging Spectroradiometer
158 (MODIS) data acquired by NASA were used to provide a large scale description of the dynamics
159 and surface water masses. ABACUS field activities were performed in collaboration with SOCIB and
160 the Mediterranean Institute for Advanced Studies (IMEDEA) using a SLOCUM G2 glider for deep
161 water (1000 m maximum depth) with a vertical speed of 0.18 ± 0.02 m/s, resulting in a horizontal
162 velocity of about 0.36 m/s. During all missions, the data acquisition design was set to dive with a
163 descending angle of 26° between 20 m and maximum depth (Figure 3).
164



165
166 *Figure 3. Glider navigation scheme during the ABACUS missions. The upper inflection depth was*
167 *changed after the ABACUS 2014 mission in order to allow the glider reaching the surface at every*
168 *second profile.*

169
170 Several physical and optical biochemical sensors were carried aboard the glider for measuring the
171 ocean temperature, salinity, oxygen, turbidity and CHLA concentration, at different rates
172 according to depth. In particular, ABACUS gliders were equipped with a glider-customized CTD by
173 Seabird-Scientific (<http://www.seabird.com/glider-payload-ctd>) measuring temperature, salinity
174 (derived from conductivity) and depth, and a two-channel combo Fluorometer-Turbidity sensor by
175 Wetlabs-Instruments (<http://www.seabird.com/eco-puck>), both of them embedded inside the
176 central hull segment of the glider commonly known as the “science-bay”. Additionally, an oxygen
177 optode to measure absolute oxygen concentration and saturation (%) by AADI
178 (<https://www.aanderaa.com/productsdetail.php?Oxygen-Optodes-2>) was externally mounted on
179 the aft, and wet, section of the glider.



180 The accuracy of the measurements, their vertical resolution and depth range vary according to the
181 specific instrument and measured variable as reported in Table 1.

182 Temperature, salinity and oxygen data were sampled to full diving depth (0-975 m depth) while
183 the acquisition of the other optical parameters ceased at 300 m depth.

184

Parameter	Instrument	Sampling rate (Hz)	Vertical resolution (m)	Depth range (m)	Accuracy	Resolution
Temperature (T), Conductivity (C), Depth (D)	Seabird GPCTD Glider payload pumped CTD	1/2	0.4	-5 to -975	T \pm 0.002 °C C \pm 0.0003 S/m D \pm 0.1% fsr*	T 0.001 °C C 0.00001 S/m D 0.002% fsr*
Oxygen	AADI Optode 5013	1/4	0.8	-5 to -975	<8 μ M or 5%	<1 μ M
Fluorescence (F), Turbidity (Tu)	WetlabsFLNTUsk	1/8	1.6	-5 to -150	Sensitivity F 0.015 \pm 0.123 μ g/L Tu 0.005 \pm 0.123 NTU	
		1/16	3.2	-150 to -300		

185 *Table 1. Sampling rate and vertical resolution of ABACUS glider data (adapted from Cotroneo et al.,*
186 *2016; Aulicino et al., 2018). * Full Scale Range*

187

188 A specific sampling rate was defined for each instrument (Table 1). Physical parameters
189 (temperature and salinity) were sampled at 1/2 Hz, resulting in a vertical resolution of 0.4 m along
190 the water column. Oxygen concentration was acquired at 1/4 Hz (vertical resolution 0.8 m) whilst
191 turbidity and chlorophyll were sampled at 1/8 Hz until 150 m depth and at 1/16 Hz from that level
192 until 300 m depth, with a vertical resolution of 1.6 m and 3.2 m respectively.

193 Calibration processes and regular maintenance were carried out before and after every glider
194 mission and guarantee the quality of the measurements. While temperature and salinity sensors
195 were regularly calibrated, unfortunately the optical sensors used during the ABACUS missions
196 were not calibrated before the cruises. This results into the impossibility to use the derived-
197 variable values as absolute values. Nonetheless, gradients along space, depth and time can
198 successfully be observed and discussed.

199 Generally, real time data transmission from the glider can be adaptively configured before and
200 during the mission. For the ABACUS missions, it was set to occur at about every 6 hours, in
201 correspondence with every second upcast for most of the glider transects. Real time data (having
202 a file size ranging between 20 and 90 Kb) were transmitted using Iridium satellite link and
203 contained a subset of all measurements, populated taking one out of every three samples for each
204 parameter. This strategy permitted the retrieval of a first overview of the data collected, as well as
205 the eventual transmission of new sampling and navigation directives to the glider. The full
206 resolution dataset remained stored locally inside the glider until the vehicle was physically brought
207 back to SOCIB facilities once the mission was completed. Then, data were transferred to SOCIB
208 Data Center where pre-processing, processing and validation were carried out and NetCDF files
209 created.

210 The structure and content of the ABACUS NetCDF file is described in section 3.

211 The pre-mission activities were carried out at the SOCIB glider facility (Tintoré et al., 2013) and
212 included all ballasting, pressure tests, compass validation, and adjustment operations needed to
213 assure the glider capability to reach the surface. This capability is provided by adjusting the overall
214 weight of the vehicle and by distributing part of that weight so that inclination and roll are neutral



215 when the glider sets its mechanical actuators (pump, mass shifter, air bladder and fin) in neutral
216 mode (0 cc, 0 inch, no inflation and 0 rad respectively). That tuning is driven by the hydrographical
217 characteristics of the target waters to be navigated, i.e. their minimum density, since they
218 condition the execution of the mentioned operations. Within this scope, the climatological
219 maximum value of temperature and minimum value of salinity for the studied area and period
220 have been analysed. These data were used as extreme hydrographic characteristics of the target
221 waters and allowed us to derive the minimum density (1024.0683 Kg/m^3) needed to precisely tune
222 the glider for the target waters. This tune up is required in order to make sure that the vehicle will
223 always be able to break into the surface and to raise its tail above the water to have clear and
224 robust communications.

225 As regards compass, Merckelbach et al (2008) developed a method to assess the glider compass
226 error in Slocum gliders. This approach has been followed for gliders used during the later ABACUS
227 experiments. The methodology should take place at any location away from sources of hard or
228 soft electromagnetic material (e.g. sport ground), thus it was performed in a forest cleaning to
229 minimize electromagnetic interference effects. Using a wooden platform, the glider was placed
230 horizontally and a plastic-made compass stand aligned with the true North was used allowing
231 precise and repetitive heading gradients of 15 degrees. The measurement error was then
232 estimated by calculating the difference between the heading measured by the glider and the
233 heading that was physically imposed using the stand. A maximum error of less than 10 degrees
234 was observed during the three tests so that compass re-calibration was not considered necessary.
235 All these operations, as part of the pre-missions tests, contributed in assessing the quality of the
236 collected data.

237 The glider tracks were designed to cross the basin from the island of Mallorca (Spain) to about 20
238 miles off the Algerian coast (Figure 2). The timing of the missions was accurately planned in order
239 to provide synoptic in situ data with respect to the satellite SARAL/ALtiKa and Sentinel-3A
240 passages, being also comparable among the different ABACUS missions (Aulicino et al., 2018).

241 In 2014, during the ABACUS 1 mission, three glider transects were completed along the
242 neighbouring SARAL/AltiKa groundtracks 229 and 773. The satellite overpassed the glider on 17
243 September and 26 November 2014 (track 773) and on 12 December 2014 (satellite track 229).
244 During ABACUS 2 mission, the glider was overflown by the SARAL/AltiKa satellite on 23 October
245 2015, along the groundtrack 229 with the glider halfway along the transect. In fact, analysis of the
246 data collected during the 2014 mission, highlighted the need of a higher degree of synopticity
247 between the glider and satellite data. A specific deployment plan was then designed for the
248 following ABACUS missions in order to have the glider approximately in the middle of the transect
249 during the satellite passages. The improvement obtained in synopticity are extensively discussed
250 in (Cotroneo et al., 2016; Aulicino et al., 2018).

251 During October 2014, a deviation from the planned sampling track was realized in order to
252 investigate the presence and sample a mesoscale eddy (see red butterfly shape track in Figure 2).
253 The eddy presence was first identified on the basis of near real time satellite altimetry and SST,
254 then the glider track was modified during one of the glider surface communication periods. The
255 adaptive sampling capabilities of the glider were then successfully used to perform two full depth
256 transects across the mesoscale eddy (Cotroneo et al., 2016). After that, the scheduled ABACUS
257 sampling track was restarted. The same strategy was applied in 2015 when the glider deviated
258 twice from the planned track to monitor the edge of a possible mesoscale eddy (saw-tooth green
259 track in figure 2).



260 Finally, during the ABACUS 3 mission, two new glider transects were performed along the Sentinel-
261 3A groundtracks 57 and 713, the latter being located between the SARAL/AltiKa groundtracks 229
262 and 773 (see Figure 2). During this mission the glider was overflown on 12 November 2016 (track
263 57) and 5 December 2016 (track 713).

264 Resolution of sampling was defined according to the energetic constraints of the platform and to
265 the scientific aims of the missions, which required high resolution in both horizontal and vertical
266 directions to monitor large scale, as well as, mesoscale processes. During the first ABACUS mission,
267 the glider was programmed to sample only during downcasts with a final along-track resolution of
268 almost 4 km once the oblique profiles are projected into the vertical. As a result of the first data
269 analysis, the sampling plan for the following missions was modified. During ABACUS 2 (2015) and
270 ABACUS 3 (2016) cruises, both downcasts and upcasts data were collected, thus obtaining an
271 improved spatial along-track resolution of about 2 km, while diving angle and speed remained
272 unchanged. Apart from increasing horizontal sampling resolution, this new strategy also has an
273 effect on data quality, enabling the application of the thermal lag correction developed by Garau
274 et al. (2011) which requires consecutive up and down profiles of temperature and salinity.
275

276 **3 Data quality control**

277 After each mission, data were transferred from the internal glider memory to the SOCIB Data
278 Center, where data processing was carried out and production of delayed time NetCDF files at
279 different elaboration levels (i.e., level 0, level 1, level 2) occurred (Troupin et al., 2016; Cusi et al.,
280 2013) before web dissemination of the data.

281 Each NetCDF file comprises the main cruise information, data and a short abstract. Level 0
282 contains the raw data collected by the glider without any elaboration or correction and organized
283 in vectors. Level 1 includes data regularized, corrected and/or derived from raw glider data. Each
284 variable is stored in a single vector containing all the observations acquired during the cruise and
285 is described and commented individually. Level 2 dataset includes the regularly sampled vertical
286 profiles generated from depth binning of already processed (level 1) glider data. In this level, each
287 variable is stored in matrix. Depth increases with the number of rows, with each row
288 corresponding to about 1 m depth, while time and consequently the number of profiles changes
289 along the columns.

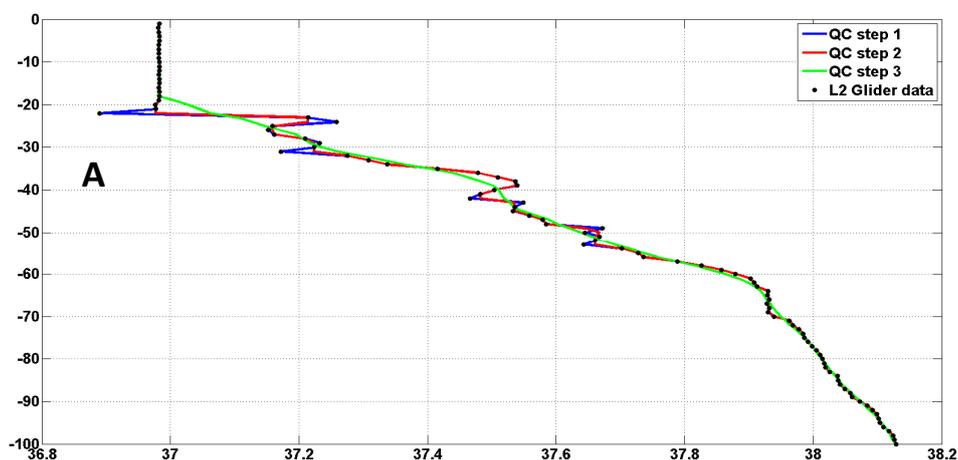
290 Delayed mode data processing for level 1 and 2 included thermal lag correction, filtering and 1 m
291 bin vertical averaging (Troupin et al., 2016). Nevertheless, some suspicious data were still present
292 in the surface layer after this standard quality control protocol. The presence of these doubtful
293 values at the shallower turning depth of the glider supports the hypothesis of a possible effect of
294 glider navigation phase change on the sampling capability.

295 For this reason, during the 2015 and 2016 missions, the glider was programmed to reach the
296 surface after every profile in order to avoid this issue, improve the data collection in the very
297 surface layer (depth < 20 m) and provide a more suitable dataset for the comparison with satellite
298 data.

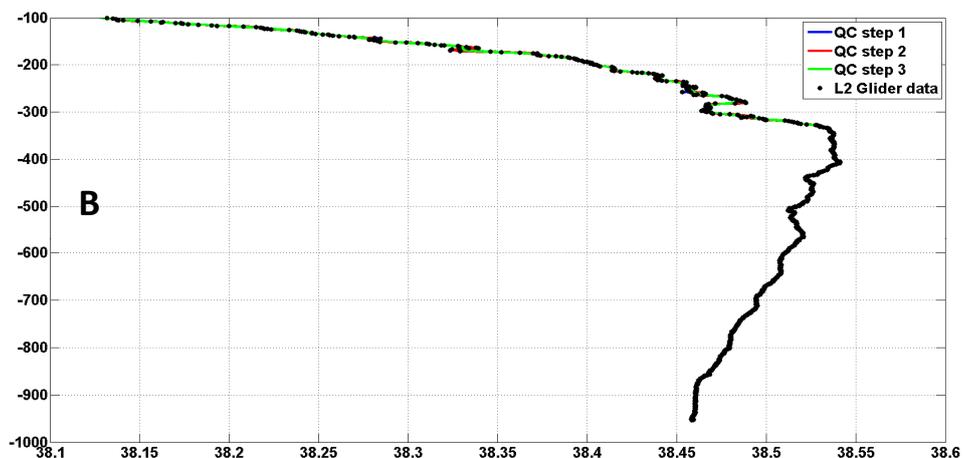
299 Nevertheless, an additional quality control procedure was developed and regularly performed at
300 University of Naples “Parthenope” on all ABACUS data. In particular, this procedure includes an
301 additional single-point spike control, the interpolation of single missing data along the profiles, the
302 application of a dedicated median filter, a 5-point running mean along the depth and, finally, an
303 iterative comparison between adjacent profiles. This tool allowed us to identify and discard bad



304 data and artefacts that were still present after the standard quality control. This final product is
 305 indicated as level 3 of the ABACUS dataset. An example of the level 3 elaboration steps is given in
 306 Figure 4. This additional procedure was routinely adopted for the quality control of all ABACUS
 307 glider missions for each geophysical parameter.
 308 All dataset elaboration levels from 0 to 2 are available at <https://doi.org/10.25704/b200-3vf5>,
 309 while level 3 data are freely available upon direct request to the authors. Data samples presented
 310 in this manuscript are obtained from level 3 data.
 311



312



313

314 *Figure 4. Effect of the additional quality control used to generate level 3 data. An example is shown*
 315 *for a salinity profile of the ABACUS 2014 mission. The surface layer is shown in panel A (0-100m*
 316 *depth) while the intermediate and deep water layers are shown in panel B (from 100 to 1000 m*
 317 *depth). Black dots represent the original level 2 data after the standard quality control procedures,*
 318 *blue line shows the single missing data interpolation, the median filter effect is shown through the*
 319 *red line, finally the green line represents the effect of the 5-point running mean on the data.*

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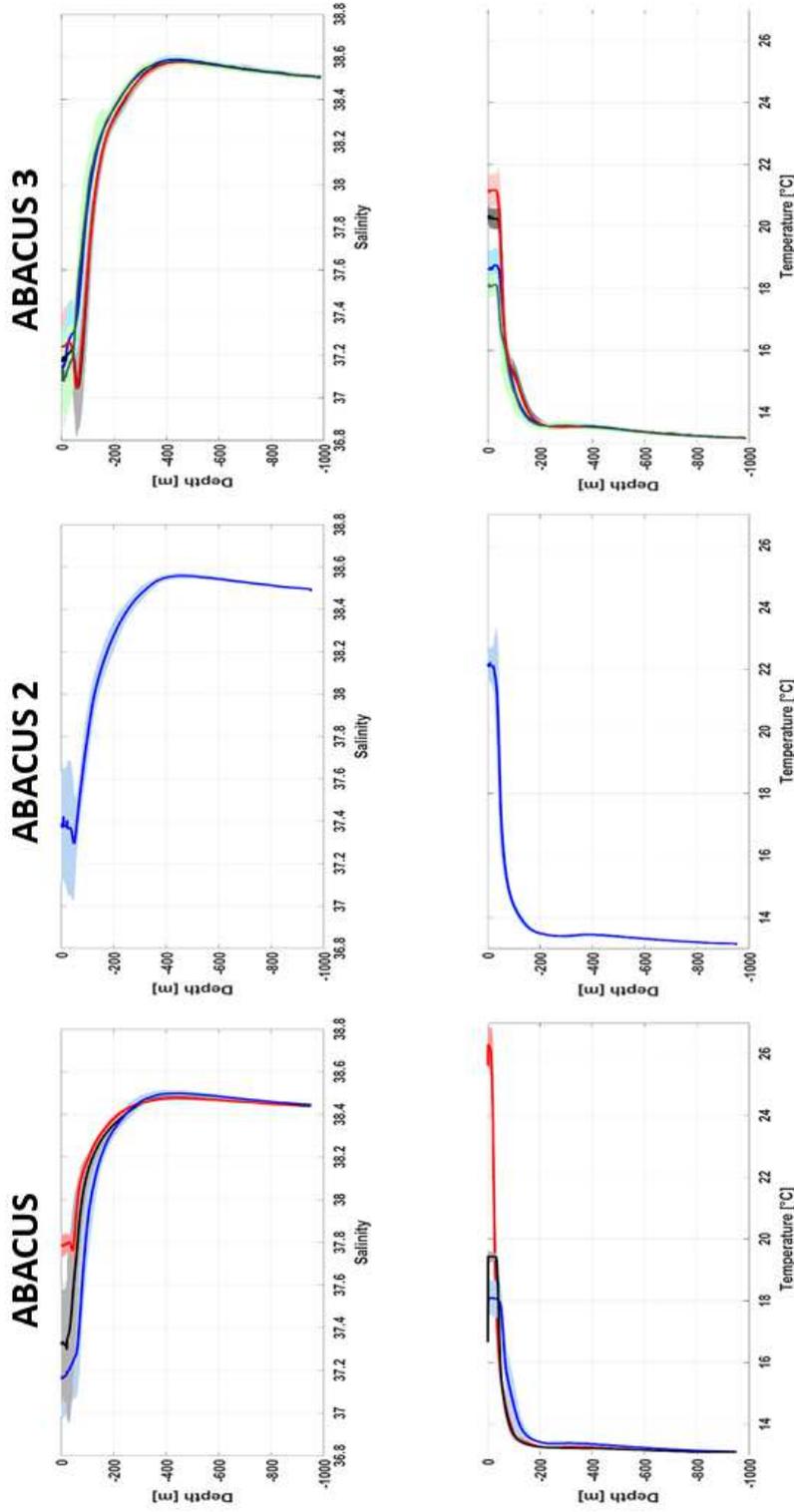


321

322 Figure 5 shows the mean temperature and salinity profile calculated for each transect completed
323 during the ABACUS cruises from 2014 to 2016. For each mission, a mean profile was calculated
324 using data of a single Mallorca-Algerian coast (or vice versa) transect. In this figure, the first
325 available mean profile for each mission is shown in blue, the second in black, the third in red and
326 the fourth in green. Standard deviation along depth during each transect was also computed for
327 both temperature and salinity (shaded areas in Figure 5).

328 The ABACUS profiles are characterized by a high level of variability in the surface layer that
329 considerably decreases in the intermediate and deeper layers, in agreement with existing
330 literature in the area (Manca et al., 2004; Fusco et al., 2008). Mean standard deviation values
331 range between 0.01 and 0.04 for salinity and between 0.05 °C and 0.15 °C for temperature; these
332 values are generally larger than the precision of the glider sensors previously described (Table 1).

333 This supports the assumption that glider measurements are accurate enough to represent the
334 variability of the basin. Ship-based CTD profiles collected at the same time and location of the
335 ABACUS casts and corrected to bottle samples could provide material for an interesting
336 comparison between glider and CTD probe data, but unfortunately, no ship-based CTD data were
337 collected in correspondence of the ABACUS missions up to 2016.



339
340 Figure 5. Mean salinity (upper panels) and temperature (lower panels) profiles (thick lines) for each ABACUS transect. From left to right, ABACUS 1,
341 ABACUS 2 and ABACUS 3 data are represented. For each mission, the first available mean profile is shown in blue, the second in black, the third in
342 red and the fourth in green. Shaded area represents the standard deviation calculated at each depth for each mean profile.

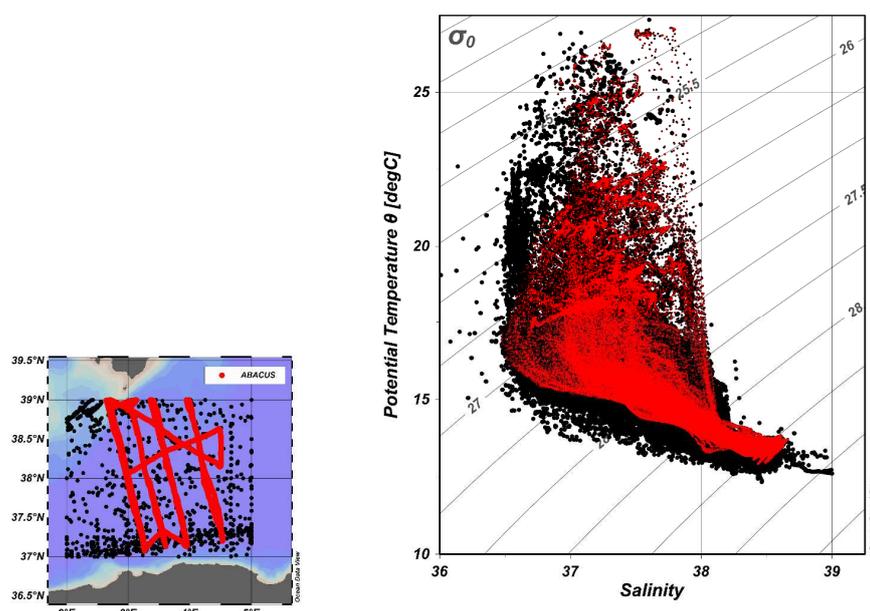


343 4 Data comparison and transect samples

344 In order to test the reliability of ABACUS dataset, a comparison was performed between glider
345 data and a set of historical oceanographic measurements collected in the study area. In particular,
346 a composite dataset including temperature and salinity data along the water column was realized
347 merging data from the MedarMedatlas II project (MEDAR group, 2002), from the Coriolis CORA-
348 3.4 Dataset (Cabanes et al., 2013) and from the World Ocean Hydrographic Profiles (WOHP - V1.0
349 database, in agreement with Viktor Gouretski). The resulting dataset spans from 1909 to 2011
350 with a regular distribution of the data across the different seasons. After the application of
351 standard quality control procedures, the calculation of potential temperature (Θ) was carried out
352 and the Θ/S couples from this dataset and the ABACUS observations were compared.

353 Figure 6 shows the comparison between ABACUS Θ/S data and all the available observations in the
354 study area. ABACUS data have the same distribution of the historical data, with no outliers in the
355 surface or deep layers. These results confirm that glider ABACUS data captured, and correctly
356 describe, the main thermohaline properties of the AB water masses and their variability.

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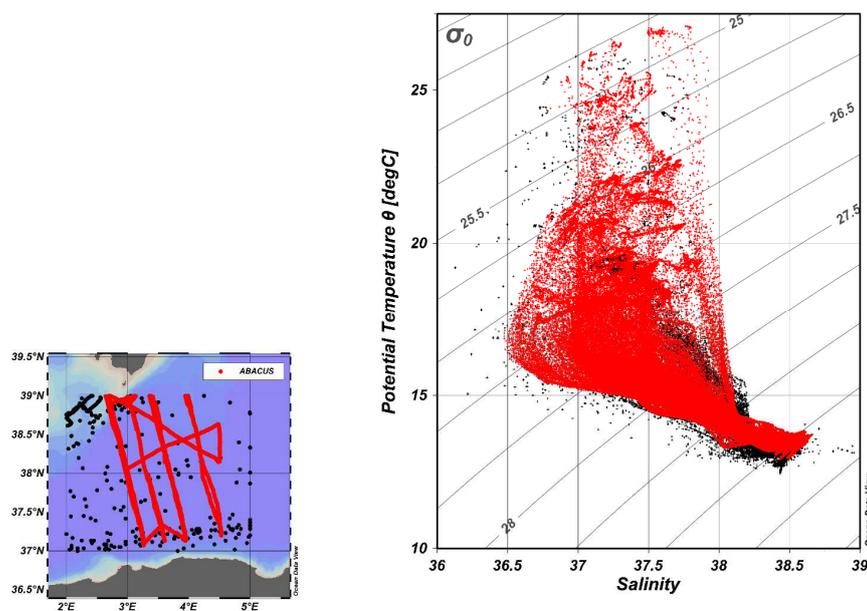
358

359 Figure 6. Θ/S diagram comparing historical oceanographic observations from 1909 to 2011 (black
360 dots) to ABACUS data (red dots). The associated map shows the spatial distribution of the data.

361

362 In order to perform an additional and more accurate test on the data reliability, we selected the
363 data collected during the fall season (September – December) from the merged historical dataset.
364 This subset was then used for a second comparison with the ABACUS data. Results are showed in
365 the Θ/S diagram in Figure 7. Again, the ABACUS data successfully represent the hydrographic
366 variability of the area, even when analysing data from a selected season. Furthermore, the
367 reduced number of historical data in the fall season and their sparse distribution are a clear
368 indication of the relative importance of ABACUS dataset for studying the AB.

369



370

371 Figure 7. θ/S diagram comparing historical oceanographic observations (black dots) during the fall
372 season (September – December from 1909 to 2011) to ABACUS data (red dots). The associated
373 map shows the spatial distribution of the data.

374

375 Several typical Mediterranean water masses have been identified in the $\theta-S$ diagrams derived from
376 the glider mounted CTD during the three ABACUS missions at sea. A summary overview is
377 provided in Table 2. The surface layer (0 to 50 m depth) was occupied by AW whose properties
378 vary greatly according to different stages of mixing, precise geographical position and residence
379 time in the Mediterranean Sea. Potential temperature, for example, ranges between 14.4°C and
380 27.0°C with colder waters always identified in the southern part of the AB. More mixed and
381 modified waters are present in its northern sector due to the influence of Balearic waters
382 (Cotroneo et al., 2016). As expected, the effect of seasonal cooling can be detected too; in fact,
383 December missions presents lower mean surface water temperatures (below 19°C).

384 As for salinity, its values in the surface layer range between 36.51 and 38.04, with higher values
385 measured next to the Mallorca Channel, characterized by shallow bottom and easier intrusion of
386 saltier waters (Aulicino et al., 2018).

387 The intermediate layers were typically occupied by LIW, characterized by a potential temperature
388 generally lower than 13.5 °C, a salinity of about 38.5 and a low oxygen concentration (see an
389 example in Figure 8c). Its presence was mostly identified between 300 and 550 m depth.

390 The deepest observed layers (between 700 and 1000 m depth) were usually characterized by the
391 presence of WMDW with typical θ values ranging between 12.9°C and 13.2°C and salinity ranging
392 between 38.44 and 38.49. These waters were saltier (about 0.05) during fall 2015 and 2016 than in
393 the 2014 glider mission.

394



395

396

Glider mission	AW	LIW	WMDW
ABACUS 1.1 September 2014	$14.44 \leq \theta \leq 27.01$ °C $36.56 \leq S \leq 37.98$	$13.21 \leq \theta \leq 13.35$ °C $38.48 \leq S \leq 38.52$	$12.91 \leq \theta \leq 13.17$ °C $38.44 \leq S \leq 38.49$
ABACUS 1.2 Nov - Dec 2014	$15.82 \leq \theta \leq 18.99$ °C $36.73 \leq S \leq 37.34$	$13.29 \leq \theta \leq 13.35$ °C $38.49 \leq S \leq 38.53$	$12.92 \leq \theta \leq 13.17$ °C $38.44 \leq S \leq 38.49$
ABACUS 2 October 2015	$16.11 \leq \theta \leq 23.88$ °C $36.52 \leq S \leq 38.04$	$13.32 \leq \theta \leq 13.51$ °C $38.53 \leq S \leq 38.59$	$12.99 \leq \theta \leq 13.17$ °C $38.49 \leq S \leq 38.54$
ABACUS 3 Nov - Dec 2016	$15.18 \leq \theta \leq 20.64$ °C $36.51 \leq S \leq 37.84$	$13.29 \leq \theta \leq 13.51$ °C $38.49 \leq S \leq 38.63$	$13.09 \leq \theta \leq 13.23$ °C $38.49 \leq S \leq 38.55$

397

398 Table 2. Temperature and salinity range values measured for Atlantic Water (AW), Levantine
399 Intermediate Water (LIW) and West Mediterranean Deep Water (WMDW) during ABACUS 1,
400 ABACUS 2 and ABACUS 3 missions.

401

402 The main properties of the water masses, as well as their spatial and vertical variability and
403 distribution can be successfully observed through the analysis of vertical transects.

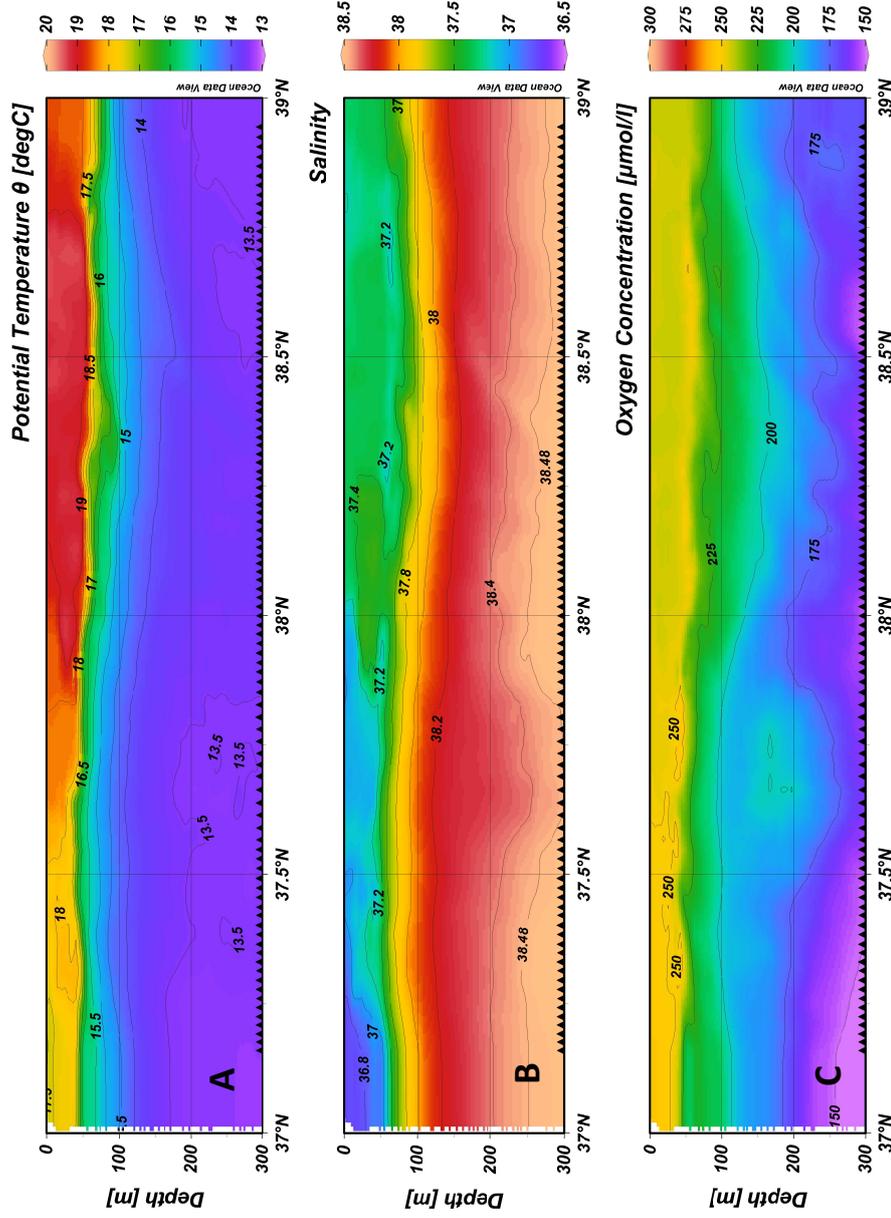
404 Figures 8 and 9 show the vertical sections along latitude of temperature, salinity, oxygen
405 concentration, CHLA concentration and turbidity collected along a transect of the ABACUS 3
406 mission. This North-South transect was realized from December 1st to December 9th 2016 along
407 the Sentinel 3 groundtrack number 713.

408 Temperature and salinity data collected in the surface layer (fig. 8 a,b) show a clear signature of
409 the AC presence. Lower temperatures and salinities are registered in the southern part of the
410 transect, highlighting the presence of AW recently entering the Mediterranean Sea. On the other
411 hand, the northern part of the transect is characterized by more saline and warm waters with
412 typical Mediterranean properties. Oxygen concentration in the surface layer presents the
413 expected distribution, without any significant latitude pattern (Fig. 8c)

414 The North-South pattern is again evident in the CHLA concentration section (fig. 8d). An increase
415 in the chlorophyll signal is registered in correspondence of the AC system at 50 m depth and, with
416 lower intensity, at the northern edge of the transect where terrestrial nutrient input from the
417 Mallorca island can be more important.

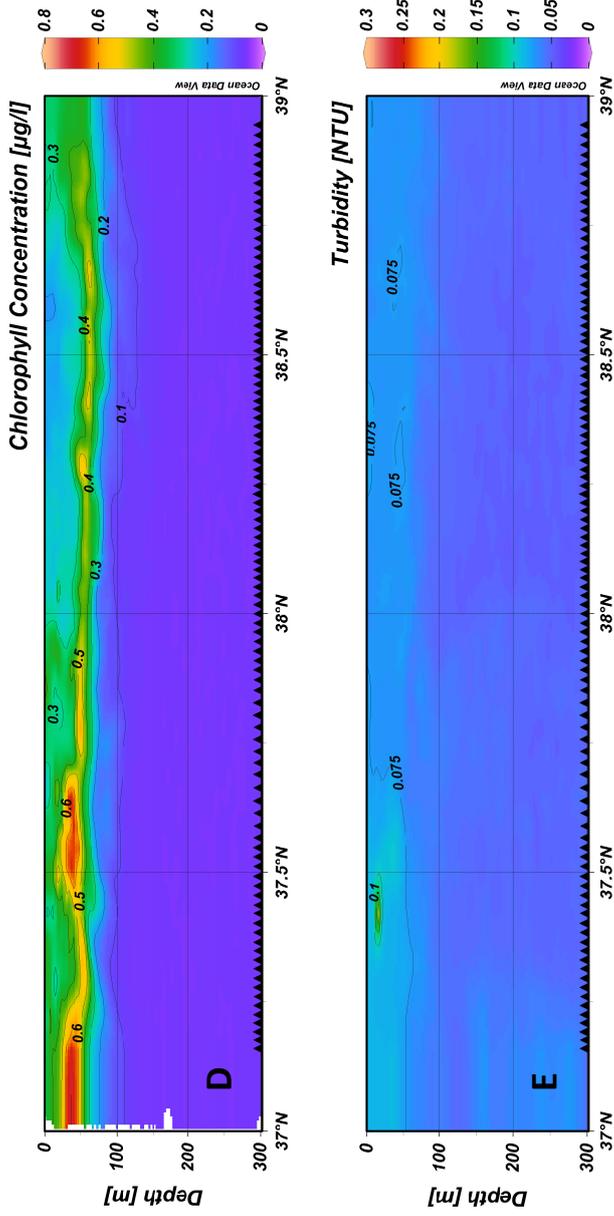
418 The CHLA concentration increase in the southern part of the transect shows the presence of a
419 lower concentration area at about 37.3°N. This signal may be associated to meanders or filaments
420 of the AC that deviate from the Eastward pattern of the current and impact on the biological
421 properties of the water masses.

422 The same signal can also be identified in the temperature and salinity sections, even if with
423 reduced intensity. The deeper layers (from 200 m to 1000 m depth) are mainly characterized by
424 the presence of the LIW. The relatively low oxygen concentration values and increased salinity
425 signals registered between 300 and 500 m depth along most of the transect are a clear signature
426 of the LIW presence (Fig. 9).

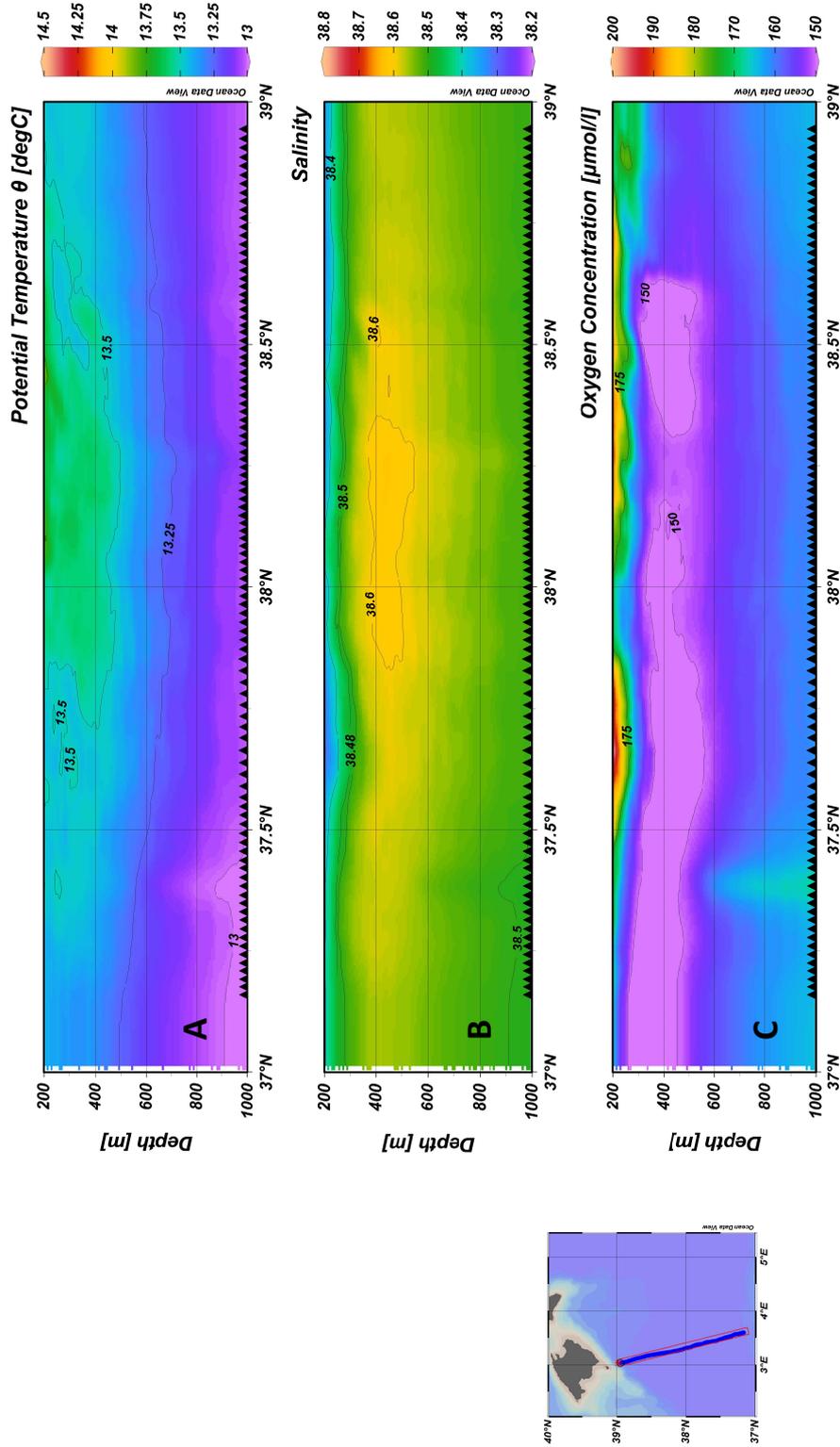


427 Figure 8. Surface layer (0-300 m depth) along track sections of potential temperature (a), salinity (b) and oxygen concentration (c) realized through
 428 ABACUS 3 glider data. Black triangles indicate the position of the single glider profiles.

429
 430
 431



432 Figure 8. Surface layer (0-300 m depth) along track sections of Chlorophyll concentration (d) and turbidity (e) realized through ABACUS 3 glider
433 data. Black triangles indicate the position of the single glider profiles.
434



435 Figure 9. Intermediate and deep layer (200-1000 m depth) along track sections of potential temperature (a), salinity (b) and oxygen concentration
 436 (c) realized through ABACUS 3 glider data. Black triangles indicate the position of the single glider profiles.
 437



438 The 800-1000 m depth layer, also shows the presence of a small scale structure at about 37.3°N.
439 In this layer, a lower temperature and salinity signal associated to increased oxygen concentration
440 can be found from 600 m to 1000 m depth. This structure may be associated to the signals
441 observed in the surface layer and surely deserves further investigations. All these data confirm the
442 glider ability to describe the main water mass properties at all scales from basin to mesoscale and
443 capture small scale structure along the water column and the investigated track. An example of
444 mesoscale variability analyzed through ABACUS glider data can be found in Cotroneo et al. (2016)
445 which present the vertical sections of multiparametric observations collected across a mesoscale
446 eddy.

447

448 **5 Conclusions**

449 The Mediterranean Sea is known to be particularly sensitive to changes in external forcings, thus
450 being one of the most responsive areas to climate change (Schroeder et al., 2017; Gualdi et al.,
451 2013) and its waters have already shown the presence of significant trends even in the deepest
452 layers (Fusco et al., 2008; Budillon et al., 2009). In this framework, the ABACUS dataset provides
453 about 1700 complete casts along the water column down to 975 m depth across the Algerian
454 Basin one of the key areas of the western Mediterranean where monitoring programmes should
455 be improved and intensified.

456 In particular, here we present data collected during a series of glider missions carried out from
457 2014 to 2016 in the framework of the ABACUS project. This effort allowed the collection of a large
458 dataset of physical (temperature and salinity) and biochemical (oxygen concentration, turbidity
459 and chlorophyll concentration) high resolution in situ observations. The reliability of these
460 measurements was tested and assessed through different quality control procedures, as well as
461 through comparisons with available historical datasets.

462 This dataset is available through an unrestricted repository at <https://doi.org/10.25704/b200-3vf5>,
463 where NetCDF files including different elaboration levels (from 0 to 2) and documentation are
464 easily accessible. This multiparametric dataset is expected to be particularly beneficial to
465 oceanographic studies focusing on the characterization of the hydrographic and biochemical
466 structure of the Western Mediterranean at all spatial scales. In fact, the presence of AW (at
467 different modification stages), LIW and WMDW, as well as their interannual variability, can be
468 observed and analysed, as reported by Aulicino et al. (2018). They included part of the presented
469 dataset in their multiplatform analyses, stressing the usefulness of glider repeated monitoring in
470 combination with altimetry and numerical simulations. Still, these observations already proved
471 their contribution to the analyses of the mesoscale and sub-mesoscale processes in the study
472 region (Cotroneo et al., 2016) whose study needs to be based on an appropriate high resolution in
473 situ dataset to possibly be coupled with satellite remote sensing data.

474 Then, we believe that the ABACUS glider dataset represents a valid unrestricted product which
475 could partially fill the lack of information in the AB, and a valuable tool for improving, together
476 with similar information collected in the framework of other AUV projects, our knowledge about
477 the dynamics of the Western Mediterranean and its physical and bio-chemical characteristics at
478 different spatial scales.

479 Moreover, we expect that this data could be used to improve and validate information derived
480 through numerical simulations, as well as to calibrate and validate present and future satellite
481 observations, especially those acquired through radar altimetry.

482 In a future perspective, cost efficient repeated glider cruises can then contribute to create a
483 network of endurance lines, monitoring both the short and long-term variability of the main
484 physical and biochemical parameters of the Mediterranean Sea. At a larger scale, the use of gliders



485 is crucial in the Global Ocean Observing System to fill gaps in transition regions between the open
486 ocean and shelf areas (Liblik et al., 2016) and to increase the array of observations in areas
487 traditionally lacking of in situ measurements, such as the Southern Mediterranean Sea. The
488 synergy between multiple gliders surveys and other data sources, i.e., satellites, models, coastal
489 radars, buoys and drifters, possibly including advisable occasional classical oceanographic cruises,
490 could represent the best strategy to be implemented.

491 As for the AB, the ABACUS monitoring activities are expected to be repeated and enhanced during
492 the coming years, taking into account the lesson learned through this study. From 2017 We
493 enlarged the ABACUS network and realized more glider missions per year in order to monitor both
494 seasonal and interannual variability along the Mallorca-Algerian coast transect.

495

496 **Data availability**

497 ABACUS glider data from 2014 to 2016 are available to the public in NetCDF format through an
498 unrestricted repository at <https://doi.org/10.25704/b200-3vf5>. A set of NetCDF files for data from
499 level 0 to level 2 is created for each glider deployment.

500 The SOCIB Data Centre hosts the data repository and offers a useful interface for data visualization
501 and download.

502

503

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505

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516 SOCIB Data Server hosts ABACUS data which are available at <https://doi.org/10.25704/b200-3vf5>.

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