



**WFDEI-GEM-CaPA: A 38-year High-Resolution Meteorological Forcing Data Set for Land  
Surface Modeling in North America**

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30 **Abstract:**

31 Cold regions hydrology is very sensitive to the impacts of climate warming. Future warming is expected to  
 32 increase the proportion of winter precipitation falling as rainfall. Snowpacks are expected to undergo less  
 33 sublimation, form later and melt earlier and possibly more slowly, leading to earlier spring peak  
 34 streamflow. More physically realistic and sophisticated hydrological models driven by reliable climate  
 35 forcing can provide the capability to assess hydrologic responses to climate change. However, hydrological  
 36 processes in cold regions involve complex phase changes and so are very sensitive to small biases in the  
 37 driving meteorology, particularly in temperature and precipitation. Cold regions often have sparse surface  
 38 observations, particularly at high elevations that generate a major amount of runoff. The effects of  
 39 mountain topography and high latitudes are not well reflected in the observational record. The best  
 40 available gridded data in Canada is from the high resolution forecasts of the Global Environmental  
 41 Multiscale (GEM) atmospheric model and the Canadian Precipitation Analysis (CaPA) but this dataset has  
 42 a short historical record. The EU WATCH ERA-Interim reanalysis (WFDEI) has a longer historical record,  
 43 but has often been found to be biased relative to observations over Canada. The aim of this study,  
 44 therefore, is to blend the strengths of both datasets (GEM-CaPA and WFDEI) to produce a less-biased long  
 45 record product (WFDEI-GEM-CaPA). First, a multivariate generalization of the quantile mapping technique  
 46 was implemented to bias-correct WFDEI against GEM-CaPA at  $3h \times 0.125^\circ$  resolution during the 2005-  
 47 2016 overlap period, followed by a hindcast of WFDEI-GEM-CaPA from 1979. The final product (WFDEI-  
 48 GEM-CaPA, 1979-2016) is freely available at the Federated Research Data Repository  
 49 (<http://dx.doi.org/10.20383/101.0111>).

50 **Subject Keywords:** cold regions processes, observations, bias correction, North America

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## 52    **1        Introduction**

53            Accurate and reliable weather and climate information at watershed-to-basin scale is in  
54    increasingly high demand by policy-makers, scientists, and other stakeholders for various purposes such  
55    as water resources management (Barnett et al., 2005), infrastructure planning (Brody et al., 2007), and  
56    ecosystem modelling (IPCC, 2013). Particularly, the potential impacts of a warming climate on water  
57    availability in snow-dominated high latitude regions continue to be a serious concern given that over the  
58    past several decades, these regions have experienced some of the most rapid warming on earth (Demaria  
59    et al., 2016; Diffenbaugh et al., 2012; Islam et al., 2017; Martin and Etchevers, 2005; Stocker et al., 2013).  
60    The on-going science suggests that these warming trends are resulting in the intensification of the  
61    hydrologic cycle, leading to substantial recent observed changes in the hydro-climatic regimes of major  
62    river basins in North America (Coopersmith et al., 2014; DeBeer et al., 2016; Dumanski et al., 2015).  
63    Changes in the timing and magnitude of river discharge (Dibike et al., 2016), shifts in extreme temperature  
64    and precipitation regimes (Asong et al., 2016b; Vincent et al., 2015) and changes in snow, ice, and  
65    permafrost regimes are anticipated (IPCC, 2013). Substantial evidence also indicates that the long-held  
66    notion of stationarity of hydrological processes is becoming invalid in a changing climate. As pointed out  
67    by Milly et al. (2008), this loss of stationarity means that there will be an increase in the likelihood and  
68    frequency of extreme weather and climate events, including floods, droughts, and heat and cold waves.

69            Water resources in most land areas north of 30° N are heavily dependent on natural water storage  
70    provided by snowpacks and glaciers, with water accumulated in the solid phase during the cold season  
71    and released in the liquid phase during warm events and the warm season. Particularly, the Rocky  
72    Mountains, the hydrologic apex of North America with headwater streams flowing to the Arctic, Atlantic  
73    and Pacific oceans, constitute an integral part of the global hydrologic cycle (Fang et al., 2013). Flows in  
74    these high elevation headwaters depend heavily on meltwater from snowpacks and glaciers. However,  
75    given that it is characterized by a highly varying cold region hydro-climate, studies indicate that it is in



76 these high elevation regions where climate variability and change is expected to be most pronounced in  
77 terms of its impacts on water supply (Beniston, 2003; Kane et al., 1991; Prowse and Beltaos, 2002; Woo  
78 and Pomeroy, 2011). More physically realistic and sophisticated hydrological models driven by reliable  
79 climate forcing information can enhance our ability to assess short- and long-term regional hydrologic  
80 responses to increasing variability and uncertainty in hydro-climatic conditions in a changing climate.  
81 Nonetheless, hydrological processes in cold regions involve complex phase changes and so are very  
82 sensitive to small biases in the driving meteorology, particularly in temperature and precipitation.

83 Cold regions often have sparse surface observations, particularly at the high elevations that  
84 generate a major amount of runoff. The effects of mountain topography and high latitudes are currently  
85 not well reflected in the observational record. Ground-based measurements (e.g. gauges) are limited  
86 especially over the Rocky Mountains, and suffer from gross inaccuracies associated with cold climate  
87 processes (Asong et al., 2017; Wang and Lin, 2015; Wong et al., 2017). The advent and use of weather  
88 radar systems have addressed some of the short-comings of gauge coverage, at least where radar exists.  
89 Unfortunately, in Canada, for example, the spatial coverage of weather radar is limited to the southern  
90 (south of 55° N) part of the country (Fortin et al., 2015b). Recently, improved satellite products have  
91 emerged such as the Global Precipitation Measurement (GPM) mission that provides meteorological  
92 information at fine spatiotemporal resolutions and regular intervals. But, the GPM is still at its early stage  
93 and only covers the region south of 60° N (Asong et al., 2017; Hou et al., 2014).

94 The capability of the current generation of Earth System Models (ESMs) to represent  
95 meteorological variables is therefore of major interest for hydrological climate change impact studies in  
96 cold regions watersheds. Despite substantial progress being made, raw outputs from regional and global  
97 ESMs still differ largely from observational reference meteorology due partly to spatial scale mismatches  
98 and systematic biases (Taylor et al., 2012). Therefore, ESM outputs are often downscaled and biases are  
99 adjusted statistically before being used in hydrological simulations (Asong et al., 2016b; Chen et al., 2013;



100 Chen et al., 2018; Gudmundsson et al., 2012). Apart from uncertainty due to the many empirical statistical  
101 techniques which have been developed to post-process ESM outputs (Maraun, 2016), the quality and  
102 length of the reference observational data set for bias correction remains a major issue (Reiter et al., 2016;  
103 Schoetter et al., 2012; Sippel et al., 2016). In Canada and other regions of North America, regional gridded  
104 data sets such as the combined Global Environmental Multiscale (GEM) atmospheric model forecasts (Yeh  
105 et al., 2002) and the Canadian Precipitation Analysis—CaPA (Mahfouf et al., 2007) have been found to  
106 perform comparably to ground observations, both statistically and hydrologically (Alavi et al., 2016;  
107 Boluwade et al., 2018; Eum et al., 2014; Fortin et al., 2015a; Gbambie et al., 2017; Wong et al., 2017).  
108 However, GEM-CaPA is too short to be used to directly correct ESM climate due to unsynchronized  
109 internal variability—the recommended minimum record length for bias correction is 30 years (Maraun,  
110 2016; Maraun et al., 2017). Other gridded products such as the EU WATCH ERA-Interim reanalysis—WFDEI  
111 (Weedon et al., 2014) and Princeton (Sheffield et al., 2006) have a longer historical record, but have been  
112 found to be biased relative to observations over Canada (Wong et al., 2017) and the United States (Behnke  
113 et al., 2016; Sapiano and Arkin, 2009). However, the WFDEI has been found to outperform other long-  
114 record gridded products (Chadburn et al., 2015; Park et al., 2016; Wong et al., 2017).

115 The aim of this study, therefore, is to combine the strengths of both datasets (GEM-CaPA and  
116 WFDEI) to produce a less-biased long record product (WFDEI-GEM-CaPA) using a multi-stage bias  
117 correction framework. First, a multivariate generalization of the quantile mapping technique was  
118 implemented to bias-correct WFDEI against GEM-CaPA at  $3\text{h} \times 0.125^\circ$  resolution during the 2005-2016  
119 period, followed by a hindcast of WFDEI-GEM-CaPA from 1979.

## 120 2 Methodology

### 121 2.1 Data sources

122 Hourly archived forecast data from the GEM model were acquired from Environment and Climate  
123 Change Canada ([http://collaboration.cmc.ec.gc.ca/cmc/cmoi/product\\_guide/submenus/rdps\\_e.html](http://collaboration.cmc.ec.gc.ca/cmc/cmoi/product_guide/submenus/rdps_e.html),



124 last access: 28 September 2018). The fields include downward incoming solar radiation, downward  
125 incoming longwave radiation and pressure at the surface, as well as specific humidity, air temperature,  
126 and wind speed at approximately 40 m above ground surface. The 40 m level was used because surface  
127 variables (2 m temperature, 2 m specific humidity, and 10 m wind speed) are only available from 2010 in  
128 the archive. The GEM data are approximately 24 km resolution from October 2001, approximately 15 km  
129 from June 2004, and approximately 10 km resolution from November 2012, and are provided on a rotated  
130 latitude/longitude grid in Environment and Climate Change Canada—ECCC ‘standard file’ format. The  
131 archived data are of former operational forecasts, and contain model outputs from versions of GEM prior  
132 to 2.0.0 through 5.0.0. A field for total precipitation (6-hourly) was acquired from the complementary  
133 CaPA product ([http://collaboration.cmc.ec.gc.ca/cmc/cmoi/product\\_guide/submenus/capa\\_e.html](http://collaboration.cmc.ec.gc.ca/cmc/cmoi/product_guide/submenus/capa_e.html), last  
134 access: 28 September 2018), which incorporates observed precipitation from meteorological weather  
135 stations, and more recently from radar, into the precipitation field from GEM. The CaPA data are  
136 approximately 10 km resolution from January 2002, also on a rotated latitude/longitude grid in ECCC  
137 ‘standard file’ format. The data contain reanalysis outputs from CaPA 2.4b8 from 2002–2012, and of  
138 former operational analyses from versions of CaPA 2.3.0 through 4.0.0 from November 2012 onward. The  
139 fields from GEM and CaPA were spatially interpolated and re-projected to a regular latitude/longitude  
140 grid at 0.125° resolution. From GEM, they were interpolated using a bilinear algorithm, while CaPA was  
141 interpolated using nearest neighbor (Schulzweida et al., 2004). Where necessary, GEM fields were  
142 converted to SI units and CaPA was converted to a precipitation rate in SI units for better compatibility  
143 with certain simulation models.

144 We also used the gridded WFDEI meteorological forcing data which has a global 0.5° spatial  
145 resolution and 3-h time step covering the period 1979–2016 ([http://www.eu-watch.org/data\\_availability](http://www.eu-watch.org/data_availability),  
146 last access: 25 July 2018). Weedon et al. (2014) used the ERA-Interim surface meteorology data as baseline  
147 information to derive the WFDEI product. Firstly, ERA-Interim data were interpolated at half-degree



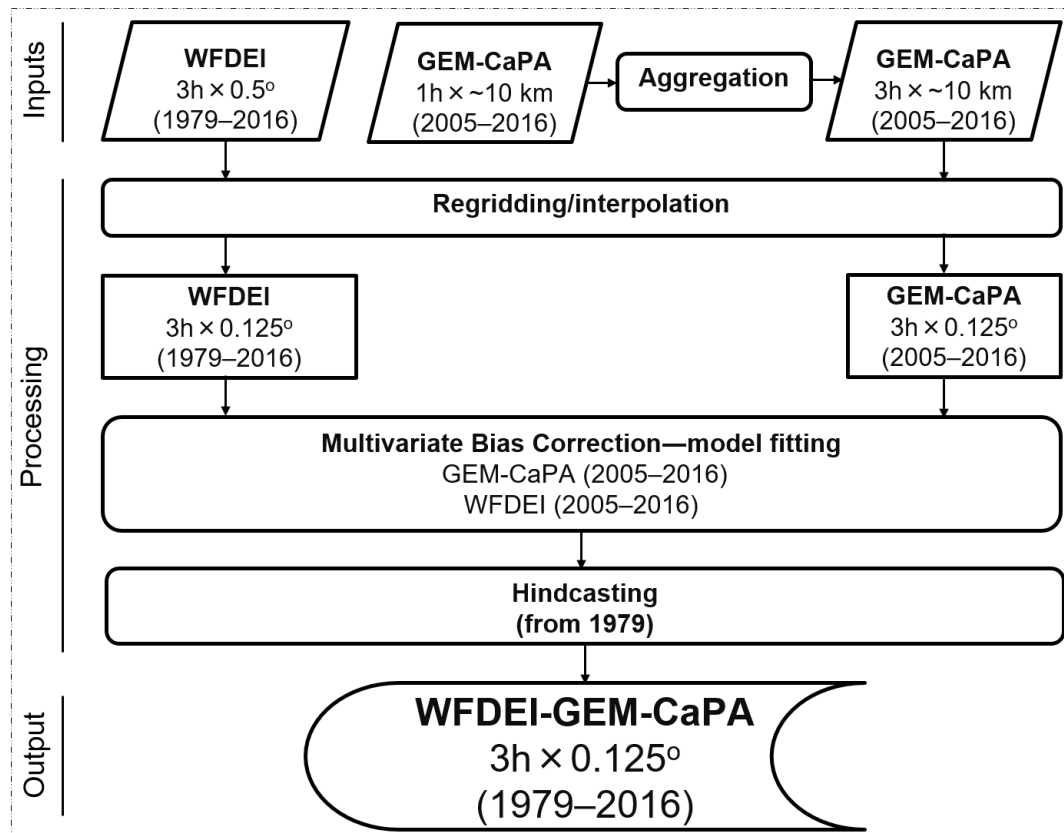
148 spatial resolution to match the land–sea mask defined by the Climatic Research Unit (CRU). Subsequently,  
149 corrections for elevation and monthly bias of climate trends in the ERA-Interim fields were applied to the  
150 interpolated data. The WFDEI data have two sets of precipitation data: the Global Precipitation  
151 Climatology Centre product (GPCC) and CRU Time Series version 3.1 (CRU TS3.1). Thus, two variants of the  
152 WFDEI product are available—WFDEI-GPCC and WFDEI-CRU. We used the WFDEI-CRU data set because it  
153 goes up to 2016 while the WFDEI-GPCC had only been updated until 2013 at the time of our analysis.

## 154 **2.2 Data processing and bias correction workflow**

155 The workflow for the multi-stage bias correction is shown in Fig.1. Bias correction was done after  
156 aggregating 1-h GEM-CaPA estimates to 3-h (the values at each time step represent the mean of the  
157 previous 3-h period, to make it consistent with WFDEI) and interpolating both WFDEI and GEM-CaPA to  
158 0.125° resolution. For bias correction, a multi-stage approach was implemented as follows. A multivariate  
159 generalization of the quantile mapping technique (Cannon, 2018) which combines quantile delta mapping  
160 (Cannon et al., 2015) and random orthogonal rotations to match the multivariate distributions of two data  
161 sets was implemented to bias-correct WFDEI against GEM-CaPA at 3-h\*0.125° resolution during the 2005-  
162 2016 period. Models were fitted to data for each calendar month while accounting for inter-variable  
163 dependence structure. Using the fitted models (2005-2016), a hindcast was made of WFDEI between  
164 1979-2004. Finally, the corrected WFDEI data derived from the fitted (2005-2016) and hindcast (1979-  
165 2004) periods were concatenated to obtain the bias-corrected WFDEI-GEM-CaPA product (1979-2016).

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167



**Figure 1.** A schematic representation of inputs and bias correction procedure used to produce the WFDEI-GEM-CaPA meteorological forcing data set.

### 3 Results and discussion

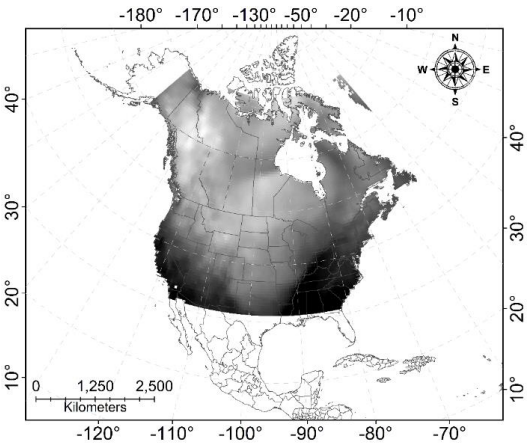
Table 1 presents an overview of the seven variables processed in this study. Note that the GEM 40 m variables are used directly to correct WFDEI surface variables (2 m temperature, 2 m specific humidity, and 10 m wind speed). Therefore, the corrected WFDEI-GEM-CaPA data reflect 40 m variables. The spatial coverage of the WFDEI-GEM-CaPA data is depicted in Fig. 2. It spans the land region between longitude 50.0625° W to 149.9375° W and latitude 31.0625° N to 71.9375° N.





**Table 1:** List variables processed in this study with heights and units in each dataset.

	WFDEI		GEM-CaPA		WFDEI-GEM-CaPA	
Variable	Height	Unit	Height	Unit	Height	Unit
Precipitation	Surface	$\text{kg m}^{-2} \text{s}^{-1}$	surface	$\text{kg m}^{-2} \text{s}^{-1}$	surface	$\text{kg m}^{-2} \text{s}^{-1}$
Air Temperature	2 m	K	40 m	K	40 m	K
Specific Humidity	2 m	$\text{kg kg}^{-1}$	40 m	$\text{kg kg}^{-1}$	40 m	$\text{kg kg}^{-1}$
Wind Speed	10 m	$\text{m s}^{-1}$	40 m	$\text{m s}^{-1}$	40 m	$\text{m s}^{-1}$
Surface Pressure	Surface	Pa	Surface	Pa	Surface	Pa
Surface Downwelling	Surface	$\text{W m}^{-2}$	Surface	$\text{W m}^{-2}$	Surface	$\text{W m}^{-2}$
Shortwave Radiation	Surface	$\text{W m}^{-2}$	Surface	$\text{W m}^{-2}$	Surface	$\text{W m}^{-2}$
Surface Downwelling	Surface	$\text{W m}^{-2}$	Surface	$\text{W m}^{-2}$	Surface	$\text{W m}^{-2}$
Longwave Radiation	Surface	$\text{W m}^{-2}$	Surface	$\text{W m}^{-2}$	Surface	$\text{W m}^{-2}$

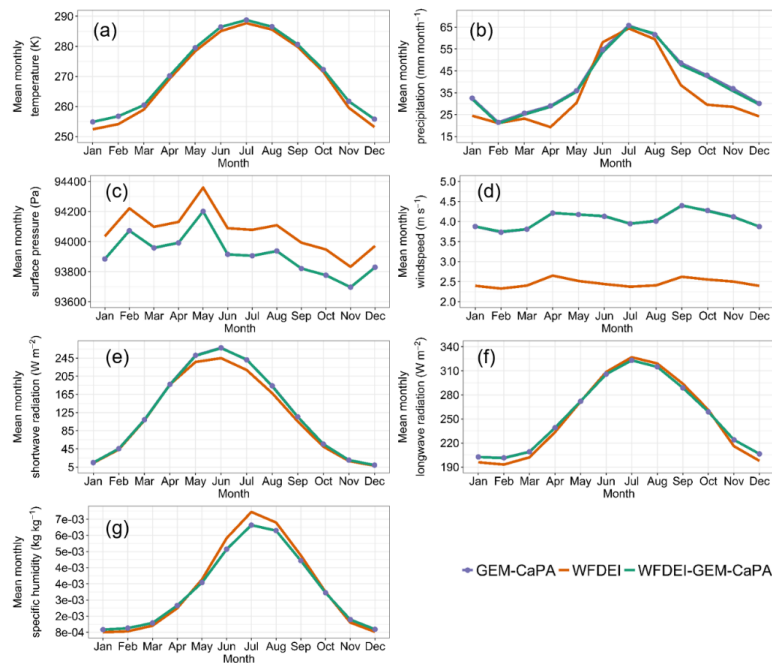


**Figure 2:** Spatial domain of the WFDEI-GEM-CaPA dataset spanning the region between longitude 50.0625° W to 149.9375° W and latitude 31.0625° N to 71.9375° N

The suitability of the bias correction algorithm to reproduce the observed annual cycle and inter-annual variability of the variables was assessed for the fitting (2005-2016) and hindcast (1979-2004) periods. Data extracted over the entire Mackenzie River basin is used to demonstrate the quality of the bias correction exercise and uniqueness of the resulting output. Fig. 3 shows the annual cycle for GEM-CaPA, WFDEI and WFDEI-GEM-CaPA during the fitting period. Overall, the monthly distributions show that the bias was removed for all variables resulting in the very close distributions between GEM-CaPA and WFDEI-GEM-CaPA. The bias was particularly large for wind speed, an important variable for both



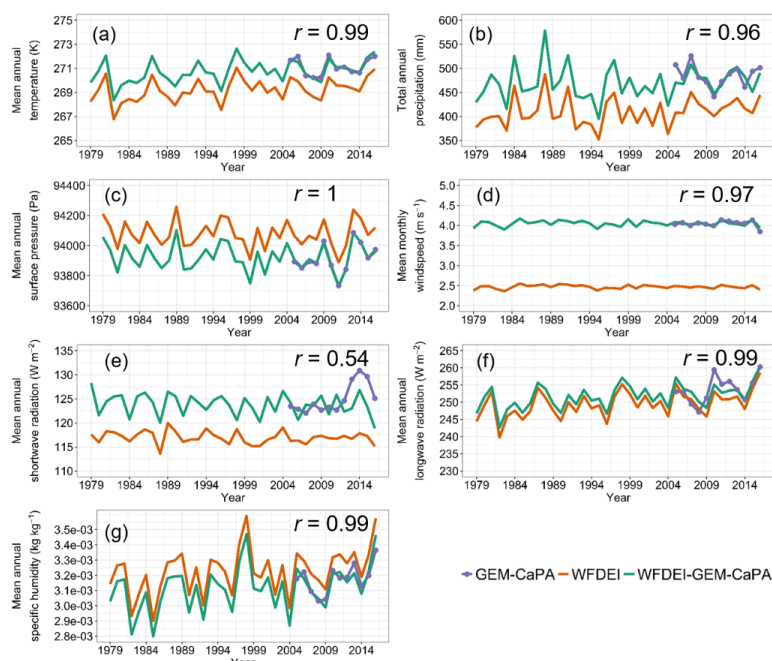
191 mountainous and prairie hydrological processes, but was successfully removed. Fig. 4 shows the mean  
 192 annual time series of the seven variables over the 1979-2016 period. It is noticeable that the bias is  
 193 corrected while the inter-annual variability is well preserved between WFDEI and WFDEI-GEM-CAPA,  
 194 except for shortwave radiation where the inter-annual variability is not fully preserved as shown by the  
 195 correlation between the WFDEI and WFDEI-GEM-CAPA annual series. However, this should not be a major  
 196 issue when impact models are forced with these data.



197  
 198 **Figure 3:** Annual cycle of GEM-CaPA (dark slate blue), WFDEI (orange) and bias corrected data—WFDEI-  
 199 GEM-CaPA (green) for air temperature (a), precipitation (b), surface pressure (c), wind speed (d),  
 200 shortwave radiation (e), longwave radiation (f), and specific humidity (g) during the fitting period (2005-  
 201 2016).

202

203



**Figure 4:** Time series of GEM-CaPA (dark slate blue), WFDEI (orange) and bias corrected data—WFDEI-GEM-CaPA (green) for air temperature (a), precipitation (b), surface pressure (c), wind speed (d), shortwave radiation (e), longwave radiation (f), and specific humidity (g) during the periods 2005-2016 (GEM-CaPA) and 1979-2016 (WFDEI and WFDEI-GEM-CaPA). The correlation ( $r$ ) between the WFDEI and WFDEI-GEM-CaPA annual series is indicated for each variable.

The foregoing analyses have shown that the bias in the WFDEI data was removed for both the fitting and hindcast periods. However, some potential limitations remain—for example, WFDEI was interpolated directly from  $0.5^\circ$  to  $0.125^\circ$  and bias-corrected against GEM-CaPA at  $0.125^\circ$ . The interpolation does not add any event-scale spatial variability for a variable like precipitation which is very variable across different scales. These issues have been reviewed extensively by (Cannon, 2018; Maraun, 2013; Maraun et al., 2010; Storch, 1999).



## 216 4 Conclusions

217 Cold regions hydrology is very sensitive to the impacts of climate warming. More physically  
218 realistic hydrological models driven by reliable climate forcing can provide the capability to assess  
219 hydrologic responses to climate variability and change. However, cold regions often have sparse surface  
220 observations, particularly at high elevations that generate a significant amount of runoff. By making  
221 available this long-term dataset, we hope it can be used to better understand and represent the  
222 seasonal/inter-annual variability of hydrological fluxes and the timing of runoff, and their long-term  
223 trends. This unique data set will also prove valuable for bias correction of climate model projections to  
224 assess potential impacts of future climate change on the hydrology and water resources of North America.

## 225 5 Data availability

226 The latest dataset is available at the Federated Research Data Repository  
227 (<http://dx.doi.org/10.20383/101.0111>).

## 228 Author contribution

229 Z.E., H.W., J.P., A.P., and M.E. conceived of and designed the experiment. D.P. preprocessed the  
230 GEM-CaPA data, A.C. developed the bias correction model code and guided the computing procedures  
231 while Z.E. performed the simulations. M.E. extracted the sample data used in generating Fig.3 and 4. Z.E.  
232 prepared the manuscript with contributions from all co-authors.

## 233 7 Competing interests

234 The authors declare that they have no conflict of interest.

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240 Global Institute for Water Security toward archiving the data at the Federated Research Data Repository.

## 241 9 References

- 242 Alavi, N., Bélair, S., Fortin, V., Zhang, S., Husain, S. Z., Carrera, M. L., and Abrahamowicz, M.:  
243 Warm Season Evaluation of Soil Moisture Prediction in the Soil, Vegetation, and Snow (SVS) Scheme,  
244 Journal of Hydrometeorology, 17, 2315-2332, 2016.
- 245 Asong, Z. E., Khaliq, M. N., and Wheeler, H. S.: Projected changes in precipitation and  
246 temperature over the Canadian Prairie Provinces using the Generalized Linear Model statistical  
247 downscaling approach, Journal of Hydrology, 539, 429-446, 2016b.
- 248 Asong, Z. E., Razavi, S., Wheeler, H. S., and Wong, J. S.: Evaluation of Integrated Multi-satellitE  
249 Retrievals for GPM (IMERG) over Southern Canada against Ground Precipitation Observations: A  
250 Preliminary Assessment, Journal of Hydrometeorology, 0, null, 2017.
- 251 Barnett, T. P., Adam, J. C., and Lettenmaier, D. P.: Potential impacts of a warming climate on  
252 water availability in snow-dominated regions, Nature, 438, 303, 2005.
- 253 Behnke, R., Vavrus, S., Allstadt, A., Albright, T., Thogmartin, W. E., and Radeloff, V. C.: Evaluation  
254 of downscaled, gridded climate data for the conterminous United States, Ecological Applications, 26,  
255 1338-1351, 2016.
- 256 Beniston, M.: Climatic change in mountain regions: a review of possible impacts. In: Climate  
257 variability and change in high elevation regions: Past, present & future, Springer, 2003.
- 258 Boluwade, A., Zhao, K. Y., Stadnyk, T. A., and Rasmussen, P.: Towards validation of the Canadian  
259 precipitation analysis (CaPA) for hydrologic modeling applications in the Canadian Prairies, Journal of  
260 Hydrology, 556, 1244-1255, 2018.
- 261 Brody, S. D., Zahran, S., Maghelal, P., Grover, H., and Highfield, W. E.: The Rising Costs of Floods:  
262 Examining the Impact of Planning and Development Decisions on Property Damage in Florida, Journal of  
263 the American Planning Association, 73, 330-345, 2007.
- 264 Cannon, A. J.: Multivariate quantile mapping bias correction: an N-dimensional probability  
265 density function transform for climate model simulations of multiple variables, Climate Dynamics, 50,  
266 31-49, 2018.
- 267 Cannon, A. J., Sobie, S. R., and Murdock, T. Q.: Bias Correction of GCM Precipitation by Quantile  
268 Mapping: How Well Do Methods Preserve Changes in Quantiles and Extremes?, Journal of Climate, 28,  
269 6938-6959, 2015.
- 270 Chadburn, S. E., Burke, E. J., Essery, R. L. H., Boike, J., Langer, M., Heikenfeld, M., Cox, P. M., and  
271 Friedlingstein, P.: Impact of model developments on present and future simulations of permafrost in a  
272 global land-surface model, The Cryosphere, 9, 1505-1521, 2015.
- 273 Chen, J., Brissette, F. P., Chaumont, D., and Braun, M.: Finding appropriate bias correction  
274 methods in downscaling precipitation for hydrologic impact studies over North America, Water  
275 Resources Research, 49, 4187-4205, 2013.
- 276 Chen, J., Li, C., Brissette, F. P., Chen, H., Wang, M., and Essou, G. R. C.: Impacts of correcting the  
277 inter-variable correlation of climate model outputs on hydrological modeling, Journal of Hydrology, 560,  
278 326-341, 2018.
- 279 Coopersmith, E. J., Minsker, B. S., and Sivapalan, M.: Patterns of regional hydroclimatic shifts: An  
280 analysis of changing hydrologic regimes, Water Resources Research, 50, 1960-1983, 2014.
- 281 DeBeer, C. M., Wheeler, H. S., Carey, S. K., and Chun, K. P.: Recent climatic, cryospheric, and  
282 hydrological changes over the interior of western Canada: a review and synthesis, Hydrol. Earth Syst.  
283 Sci., 20, 1573-1598, 2016.



- 284 Demaria, E. M. C., Roundy, J. K., Wi, S., and Palmer, R. N.: The Effects of Climate Change on  
285 Seasonal Snowpack and the Hydrology of the Northeastern and Upper Midwest United States, *Journal of*  
286 *Climate*, 29, 6527-6541, 2016.
- 287 Dibike, Y., Prowse, T., Bonsal, B., and O'Neil, H.: Implications of future climate on water  
288 availability in the western Canadian river basins, *International Journal of Climatology*, doi:  
289 10.1002/joc.4912, 2016. n/a-n/a, 2016.
- 290 Diffenbaugh, N. S., Scherer, M., and Ashfaq, M.: Response of snow-dependent hydrologic  
291 extremes to continued global warming, *Nature Climate Change*, 3, 379, 2012.
- 292 Dumanski, S., Pomeroy, J. W., and Westbrook, C. J.: Hydrological regime changes in a Canadian  
293 Prairie basin, *Hydrological Processes*, 29, 3893-3904, 2015.
- 294 Eum, H.-I., Dibike, Y., Prowse, T., and Bonsal, B.: Inter-comparison of high-resolution gridded  
295 climate data sets and their implication on hydrological model simulation over the Athabasca Watershed,  
296 *Canada, Hydrological Processes*, 28, 4250-4271, 2014.
- 297 Fang, X., Pomeroy, J. W., Ellis, C. R., MacDonald, M. K., DeBeer, C. M., and Brown, T.: Multi-  
298 variable evaluation of hydrological model predictions for a headwater basin in the Canadian Rocky  
299 Mountains, *Hydrol. Earth Syst. Sci.*, 17, 1635-1659, 2013.
- 300 Fortin, V., Jean, M., Brown, R., and Payette, S.: Predicting snow depth in a forest–tundra  
301 landscape using a conceptual model allowing for snow redistribution and constrained by observations  
302 from a digital camera, *Atmosphere–Ocean*, 53, 200-211, 2015a.
- 303 Fortin, V., Roy, G., Donaldson, N., and Mahidjiba, A.: Assimilation of radar quantitative  
304 precipitation estimations in the Canadian Precipitation Analysis (CaPA), *Journal of Hydrology*, 531, Part  
305 2, 296-307, 2015b.
- 306 Gbambie, A. S. B., Poulin, A., Boucher, M.-A., and Arsenault, R.: Added Value of Alternative  
307 Information in Interpolated Precipitation Datasets for Hydrology, *Journal of Hydrometeorology*, 18, 247-  
308 264, 2017.
- 309 Gudmundsson, L., Bremnes, J. B., Haugen, J. E., and Engen-Skaugen, T.: Technical Note:  
310 Downscaling RCM precipitation to the station scale using statistical transformations &ndash; a  
311 comparison of methods, *Hydrol. Earth Syst. Sci.*, 16, 3383-3390, 2012.
- 312 Hou, A. Y., Kakar, R. A., Neeck, S., Azarbarzin, A. A., Kummerow, C. D., Kojima, M., Oki, R.,  
313 Nakamura, K., and Iguchi, T.: The Global Precipitation Measurement Mission, *Bulletin of the American*  
314 *Meteorological Society*, 95, 701-722, 2014.
- 315 IPCC: Climate Change 2013: The Physical Science Basis. Contribution of Working Group I to the  
316 Fifth Assessment Report of the Intergovernmental Panel on Climate Change, Cambridge University  
317 Press, Cambridge, United Kingdom and New York, NY, USA, 2013.
- 318 Islam, S. u., Déry, S. J., and Werner, A. T.: Future Climate Change Impacts on Snow and Water  
319 Resources of the Fraser River Basin, British Columbia, *Journal of Hydrometeorology*, 18, 473-496, 2017.
- 320 Kane, D. L., Hinzman, L. D., Woo, M.-k., and Everett, K. R.: Arctic hydrology and climate change.  
321 In: *Arctic ecosystems in a changing climate*, Elsevier, 1991.
- 322 Mahfouf, J. F., Brasnett, B., and Gagnon, S.: A Canadian precipitation analysis (CaPA) project:  
323 Description and preliminary results, *Atmosphere–Ocean*, 45, 1-17, 2007.
- 324 Maraun, D.: Bias Correcting Climate Change Simulations - a Critical Review, *Current Climate*  
325 *Change Reports*, 2, 211-220, 2016.
- 326 Maraun, D.: Bias Correction, Quantile Mapping, and Downscaling: Revisiting the Inflation Issue,  
327 *Journal of Climate*, 26, 2137-2143, 2013.
- 328 Maraun, D., Shepherd, T. G., Widmann, M., Zappa, G., Walton, D., Gutiérrez, J. M., Hagemann,  
329 S., Richter, I., Soares, P. M. M., Hall, A., and Mearns, L. O.: Towards process-informed bias correction of  
330 climate change simulations, *Nature Climate Change*, 7, 764, 2017.



- 331 Maraun, D., Wetterhall, F., Ireson, A. M., Chandler, R. E., Kendon, E. J., Widmann, M., Brien, S., Rust, H. W., Sauter, T., Themeßl, M., Venema, V. K. C., Chun, K. P., Goodess, C. M., Jones, R. G., Onof, C., Vrac, M., and Thiele-Eich, I.: Precipitation downscaling under climate change: Recent developments to bridge the gap between dynamical models and the end user, *Reviews of Geophysics*, 48, 2010.
- 332  
333  
334  
335 Martin, E. and Etchevers, P.: Impact of Climatic Changes on Snow Cover and Snow Hydrology in the French Alps. In: *Global Change and Mountain Regions: An Overview of Current Knowledge*, Huber, U. M., Bugmann, H. K. M., and Reasoner, M. A. (Eds.), Springer Netherlands, Dordrecht, 2005.
- 336  
337  
338 Milly, P. C. D., Betancourt, J., Falkenmark, M., Hirsch, R. M., Kundzewicz, Z. W., Lettenmaier, D. P., and Stouffer, R. J.: Stationarity Is Dead: Whither Water Management?, *Science*, 319, 573-574, 2008.
- 339  
340 Park, H., Yoshikawa, Y., Oshima, K., Kim, Y., Ngo-Duc, T., Kimball, J. S., and Yang, D.: Quantification of Warming Climate-Induced Changes in Terrestrial Arctic River Ice Thickness and Phenology, *Journal of Climate*, 29, 1733-1754, 2016.
- 341  
342  
343 Prowse, T. D. and Beltaos, S.: Climatic control of river-ice hydrology: a review, *Hydrological processes*, 16, 805-822, 2002.
- 344  
345 Reiter, P., Gutjahr, O., Schefczyk, L., Heinemann, G., and Casper, M.: Bias correction of ENSEMBLES precipitation data with focus on the effect of the length of the calibration period, *Meteorologische Zeitschrift*, 25, 85-96, 2016.
- 346  
347  
348 Sapiano, M. R. P. and Arkin, P. A.: An Intercomparison and Validation of High-Resolution Satellite Precipitation Estimates with 3-Hourly Gauge Data, *Journal of Hydrometeorology*, 10, 149-166, 2009.
- 349  
350 Schoetter, R., Hoffmann, P., Rechid, D., and Schlünzen, K. H.: Evaluation and Bias Correction of Regional Climate Model Results Using Model Evaluation Measures, *Journal of Applied Meteorology and Climatology*, 51, 1670-1684, 2012.
- 351  
352  
353 Schulzeida, U., Kornbluh, L., and Quast, R.: Climate data operators, Max-Planck-Institute for Meteorology, Hamburg and <http://www.mpimet.mpg.de/~cdo>, 2004. 2004.
- 354  
355 Sheffield, J., Goteti, G., and Wood, E. F.: Development of a 50-Year High-Resolution Global Dataset of Meteorological Forcings for Land Surface Modeling, *Journal of Climate*, 19, 3088-3111, 2006.
- 356  
357 Sippel, S., Otto, F. E. L., Forkel, M., Allen, M. R., Guillod, B. P., Heimann, M., Reichstein, M., Seneviratne, S. I., Thonicke, K., and Mahecha, M. D.: A novel bias correction methodology for climate impact simulations, *Earth Syst. Dynam.*, 7, 71-88, 2016.
- 358  
359  
360 Stocker, T. F., Qin, D., Plattner, G.-K., Alexander, L. V., Allen, S. K., Bindoff, N. L., Bréon, F.-M., Church, J. A., Cubasch, U., Emori, S., Forster, P., Friedlingstein, P., Gillett, N., Gregory, J. M., Hartmann, D. L., Jansen, E., Kirtman, B., Knutti, R., Krishna Kumar, K., Lemke, P., Marotzke, J., Masson-Delmotte, V., Meehl, G. A., Mokhov, I. I., Piao, S., Ramaswamy, V., Randall, D., Rhein, M., Rojas, M., Sabine, C., Shindell, D., Talley, L. D., Vaughan, D. G., and Xie, S.-P.: Technical Summary. In: *Climate Change 2013: The Physical Science Basis. Contribution of Working Group I to the Fifth Assessment Report of the Intergovernmental Panel on Climate Change*, Stocker, T. F., Qin, D., Plattner, G.-K., Tignor, M., Allen, S. K., Boschung, J., Nauels, A., Xia, Y., Bex, V., and Midgley, P. M. (Eds.), Cambridge University Press, Cambridge, United Kingdom and New York, NY, USA, 2013.
- 361  
362  
363  
364  
365  
366  
367  
368  
369 Storch, H. v.: On the Use of "Inflation" in Statistical Downscaling, *Journal of Climate*, 12, 3505-3506, 1999.
- 370  
371 Taylor, K. E., Stouffer, R. J., and Meehl, G. A.: An Overview of CMIP5 and the Experiment Design, *Bulletin of the American Meteorological Society*, 93, 485-498, 2012.
- 372  
373 Vincent, L. A., Zhang, X., Brown, R. D., Feng, Y., Mekis, E., Milewska, E. J., Wan, H., and Wang, X. L.: Observed Trends in Canada's Climate and Influence of Low-Frequency Variability Modes, *Journal of Climate*, 28, 4545-4560, 2015.
- 374  
375  
376 Wang, X. L. and Lin, A.: An algorithm for integrating satellite precipitation estimates with in situ precipitation data on a pentad time scale, *Journal of Geophysical Research: Atmospheres*, 120, 3728-3744, 2015.
- 377  
378





379 Weedon, G. P., Balsamo, G., Bellouin, N., Gomes, S., Best, M. J., and Viterbo, P.: The WFDEI  
380 meteorological forcing data set: WATCH Forcing Data methodology applied to ERA-Interim reanalysis  
381 data, *Water Resources Research*, 50, 7505-7514, 2014.

382 Wong, J. S., Razavi, S., Bonsal, B. R., Wheeler, H. S., and Asong, Z. E.: Inter-comparison of daily  
383 precipitation products for large-scale hydro-climatic applications over Canada, *Hydrol. Earth Syst. Sci.*,  
384 21, 2163-2185, 2017.

385 Woo, M.-K. and Pomeroy, J.: *Snow and Runoff: Processes, Sensitivity and Vulnerability*. In:  
386 *Changing Cold Environments*, John Wiley & Sons, Ltd, 2011.

387 Yeh, K.-S., Côté, J., Gravel, S., Méthot, A., Patoine, A., Roch, M., and Staniforth, A.: The CMC–  
388 MRB Global Environmental Multiscale (GEM) Model. Part III: Nonhydrostatic Formulation, *Monthly*  
389 *Weather Review*, 130, 339-356, 2002.

390