



1 **Water-balance and hydrology research in a mountainous**
2 **permafrost watershed in upland streams of the Kolyma River,**
3 **Russia: a database from the Kolyma Water-Balance Station,**
4 **1948-1997**

5 Olga Makarieva^{1,2,3}, Nataliia Nesterova^{3,4}, Lyudmila Lebedeva², Sergey Sushansky^{5,*}

6 ¹*Gidrotehproekt Ltd, St. Petersburg, 199178, Russia*

7 ²*Melnikov Permafrost Institute, Merzlotnaya St., 36, Yakutsk 677010, Russia*

8 ³*Saint Petersburg State University, Institute of Earth Sciences, 7/9 Universitetskaya nab, St.*
9 *Petersburg, Russia 199034*

10 ⁴*State Hydrological Institute, Department of Experimental Hydrology and Mathematical*
11 *Modelling of Hydrological Processes, 23 2-ya liniya VO, St. Petersburg, Russia 199053*

12 ⁵*formerly at: the Kolyma water-balance station, Stokovoye, Magadan district, Russia*
13 **retired*

14
15 Contact author: omakarieva@gmail.com

16 **Abstract:** As of 2017, 70 years have passed since the beginning of work at the Kolyma water-
17 balance station (KWBS), a unique scientific research hydrological and permafrost catchment.
18 The volume and duration (50 continuous years) of hydrometeorological standard and
19 experimental data, characterizing the natural conditions and processes occurring in mountainous
20 permafrost conditions, significantly exceeds any counterparts elsewhere in the world. The data
21 are representative of the vast territory of the North-East of Russia. In 1997, the station was
22 terminated, thereby leaving Russia without operating research watersheds in the permafrost zone.
23 This paper describes the dataset containing the series of daily runoff from 10 watersheds with
24 area from 0.27 to 21.3 km², precipitation, meteorological observations, evaporation from soil and
25 snow, snow surveys, soil thaw and freeze depths, and soil temperature for the period 1948-1997.
26 It also highlights the main historical stages of the station's existence, its work and scientific
27 significance, and outlines the prospects for its future, where the Kolyma water-balance station
28 could be restored to the status of a scientific research watershed and become a valuable
29 international center for hydrological research in permafrost. The data is available at
30 <https://doi.pangaea.de/10.1594/PANGAEA.881731>.

31 **Keywords:** water-balance and hydrological research, continuous permafrost, Kolyma River,
32 Kolyma water-balance station, streamflow, thaw/freezing depth, precipitation, snow cover,
33 evaporation from soil and snow

34

35



36 1. Introduction

37 In 2017 we celebrate 70 years since work on organizing the Kolyma Water-Balance
38 station (KWBS) began. This hydrological and permafrost research catchment has accumulated
39 standard and experimental data unique both in terms of their amount and duration.

40 In the paper «Save northern high-latitude catchments» Laudon et al. (2017) recognize the
41 KWBS as a currently functioning scientific station, even though scientific research was
42 suspended here 20 years ago, and nowadays only standard observations at the meteorological site
43 and the runoff gauge are carried out.

44 Eurasia contributes 75 % of the total terrestrial runoff to the Arctic Ocean and three of
45 the four major Arctic rivers are located in Siberia (Shiklomanov et al. 2002). Peterson et al.
46 (2002) suggested that the net discharge from the six largest Eurasian rivers flowing into the
47 Arctic Ocean (Severnaya Dvina, Pechora, Ob', Yenisey, Lena, and Kolyma) increased by 7%
48 during the 20th century. As it is also mentioned by Laudon et al. (2017), the number of scientific
49 and hydrological research stations in the Northern regions of the world has decreased by 40%,
50 and it happened alongside the most significant climate change in the Arctic in recorded history.

51 The Kolyma Water Balance Station (KWBS) is located in the headwaters of the Kolyma
52 River, in a mountainous region of continuous permafrost (Fig. 1). Runoff formation conditions at
53 the station are representative for an immense territory of the upper Kolyma River basin and
54 mountainous regions of North-East Russia. Although there are large mountainous areas in other
55 cold regions of the world in Canada, the USA and Europe, the combination of extremely severe
56 climate and continuous permafrost creates unique conditions in North-East Russia.

57 To the best of our knowledge, the first systematic cold-region hydrology observations in
58 North America begin not earlier than the 1960s. The Caribou-Poker Creeks Research Watershed
59 was established in 1969 (Hinzman et al., 2002).

60 At the KWBS there were 10 hydrological gauges at catchments ranging between 0.27 and
61 21.3 km², two meteorological plots, 55 (in total) precipitation gauges, over 30 frost tubes
62 (cryopedometers), several groundwater wells, evaporation, water-balance and runoff plots. In
63 addition, regular snow surveys were conducted, as well as experimental investigations of specific
64 hydrological and permafrost processes.

65 During the period 1948 to 1997, the KWBS accumulated a huge amount of data on
66 hydro-meteorological and special observations of a unique duration (40-50 years) that
67 characterize the natural setting, which, on the one hand, are nearly ungauged, and on the other
68 hand, are representative of the vast territory of North-East Russia and to some extent of other
69 alpine cold-climate areas in Europe and North America. The data were published in 40 reports,
70 the first one covering the period 1948-1957. Following issues were published annually (KWBS
71 Observation's materials, 1959-1997).

72 After 1997, water balance observations at the KWBS were ceased. One weather station
73 and five runoff gauges functioned at the KWBS up to mid-June, 2013, when an extreme flash
74 flood destroyed four level gauges. Nowadays only standard observations are conducted at the
75 Nizhnyaya meteorological site and at the Kontaktovy (Nizhny) runoff gauge.

76 Observation results were reflected in more than 250 publications, dedicated to different
77 aspects of runoff formation in the continuous permafrost region, active layer dynamics,
78 underlying surface structure and its influence on hydrological processes. Based on the KWBS
79 materials, the following research was carried out:

- 80 • water balance formation (Boyarintsev, 1980; Boyarintsev, Gopchenko, 1992; Kuznetsov
81 et al., 1969; Suchansky, 2002; Zhuravin, 2004; Lebedeva et al., 2017),
- 82 • peak and spring flood runoff in small rivers in the permafrost zone (Boyarintsev, 1988),
- 83 • base flow (Boyarintsev, Nikolaev, 1986; Glotov, 2002),
- 84 • principles of runoff cryo-regulation (Alekseev et al., 2011),
- 85 • ice content dynamics of rocky talus deposits (Bantsekina, 2001),
- 86 • processes of intra-ground condensation (Reinyuk, 1959; Bantsekina et al., 2009;
87 Boyarintsev et al., 1991),



88 • floodplain taliks in continuous permafrost (Mikhaylov, 2013) and many others.
89 Collected data were also used for development and testing different geoscience models of:
90 • runoff formation (Gusev et al., 2006; Kuchment et al., 2000; Vinogradov et al., 2015;
91 Lebedeva et al., 2015; Semenova et al., 2013),
92 • climatic aspects (Shmakin, 1998),
93 • land surface and vegetation dynamics (Tikhmenev, 2008).
94 In this paper, we present a hydrometeorological and permafrost related dataset for 50
95 years from 1948 – 1997 for the Kolyma Water-Balance Station (KWBS), the Kontaktovy Creek
96 watershed, which is representative for vast mountainous territories of continuous permafrost
97 zone of Eastern Siberia and the North-East of Russia. This dataset is unique in terms of its
98 volume and duration of hydrometeorological standard and experimental data. It may be used in
99 many research tasks, but is of particular importance in studying runoff formation processes and
100 model development in permafrost regions.

101 2. Site description

102 The Kolyma water-balance station is located in the Tenkinsky district of the Magadan
103 region of Russia within the Upper-Kolyma highland. The station's territory – the Kontaktovy
104 Creek catchment with area 21.3 km² – is a part of the Pravy Itrikan River which flows into the
105 Kulu River basin, which is the right tributary of the Kolyma River. The station is located 16 km
106 from the Kulu village settlement. It is characterized by a mountain landscape, typical for the
107 upper reaches of the Kolyma River. The territory of the basin is severely cut up with creek
108 valleys. These valleys are narrow, with steep slopes, and watershed lines are mostly well
109 delineated. Absolute elevations of the basin range between 823 m a.s.l. near the Kontaktovy
110 Creek outlet and 1700 m a.s.l. at watershed divides. The length of the creek is 8.9 km. The
111 catchment is extended in the latitudinal direction and has an asymmetric shape. The slopes of the
112 catchment area have mainly southern exposure (53% of the slope area), the slopes of the
113 northern and eastern exposure have a 24% share, the western – 23%. The density of river
114 network in the basin is 2.5 km per square km. The main river canal is meandering. The steepness
115 of the slopes ranges from 200 to 800‰ (Fig. 1).

116 The station is located in the continuous permafrost zone. Permafrost thickness varies
117 from 120 to 210 m in valleys and can reach 300-400 m in highlands, following the relief.
118 Seasonal soil thaw depth depends on slope exposition, altitude and landscape and changes from
119 0.2-0.8 m on north-facing slopes to 1.5-3.0 m on south-facing ones.

120 KWBS is situated in the transitional zone between forest-tundra and coniferous taiga. Soil
121 types vary from stony-rock debris to clayey podzol with partially decayed organic material
122 underlain by frozen soil and bedrock. Most of the KWBS area is covered by rocky talus,
123 practically without vegetation (34%). Dwarf cedar and alder shrubs are common at south-facing
124 slopes and cover about 27% of the territory. Larch sparse woodland with moss-lichen cover is
125 typical for steep north-facing slopes (12%). Open terrain larch wood (15%) and swampy sparse
126 growth forest with minimal permafrost thaw depth, constant excessive stagnant moisturizing,
127 tussock or knobby microrelief (12%) characterize creek valleys. The estimates of landscape
128 distribution are given here after Korolev (1984).

129

130 2.1 History of KWBS

131 The Kolyma water balance station (KWBS) was established on October 15, 1947 and
132 was initially known as the Itrikanskaya runoff station of the Dalstroy (Far North Construction
133 Trust organized in 1931) Hydrometeorological Service. In 1948-1956 and 1957-1969 it was
134 called the Kulinskaya and the Kolyma runoff station respectively. The primary goal of this
135 station was studying runoff formation processes in small river catchments in mountain
136 permafrost landscapes, typical for northeast USSR.

137 As soon as May, 1948, the first runoff observations at the Kontaktovy Creeks and
138 Vstrecha brook were launched, as well as regular observations at the Nizhnyaya weather station



139 (850 m a.s.l.). A few months later, on September 1, 1948, observations at the Verkhnyaya
140 weather station (1220 m a.s.l.) were started. In 1948, stage gauges Sredny, Nizhny and Vstrecha
141 were equipped with automatic water level recorders, gauging footbridges and flumes.

142 During the period 1949-1957, at the Vstrecha brook catchment, a rain-gauge network was
143 organized. Runoff gauges at the Severny, Dozhdemerny, Vstrecha brooks were equipped with
144 various hydrometric facilities. Observations on soil, water and snow evaporation, soil freezing
145 and thawing commenced, as well as experimental observations at a runoff plot.

146 At the end of the 1940s and early 1950s, technical staff of the station were mainly former
147 convicts. During the first few years, the workers of the station built houses for themselves,
148 collected firewood and organized the household. The winter of 1955-56 appeared to be
149 especially severe for the staff, since due to the deep snow cover it was difficult to move around
150 the territory of KWBS, there was no transport connection with the Tenkinskaya highway,
151 delivering of firewood, needed for heating houses and service buildings, was also difficult. When
152 it was impossible to get to the highway by car, bread and mail were delivered from the Kulu
153 village settlement utilizing horses once every 7-10 days.

154 Twenty to twenty five staff members were accommodated in three small huts, hardly
155 suitable for living. That winter they mainly had to collect and prepare firewood in the afternoon;
156 in the morning everybody had to go (despite their rank or position) in deep snow and at -50°C
157 to the nearest small river valley looking for firewood, then they pulled it back home, where they
158 were firing furnaces. Only by the time it got dark, it became warm enough to stay in the work-
159 room and they could start observation data processing.

160 The working day lasted till 10 or 11 p.m. Since there was no electricity, they used
161 kerosene lamps filled with a mixture of petrol and salt. In summer 1956 there were only 13
162 people left at the station, some of them were taken to help with haymaking to prepare hay for
163 their subsidiary holding that consisted of two cows and a horse (as recollected by the Chief of the
164 station V.G. Osipov, the hydrologist-technician A.I. Ipatieva, Informational letter..., 1988).

165 In 1957 the station was handed over to the jurisdiction of the Kolyma Hydro-
166 meteorological service administration, and in 1958 it was partially connected to electricity. At
167 that time there were active steps taken toward fitting out the station with new types of devices
168 and equipment, engaging new specialists in hydro-meteorology, and building accommodation
169 facilities.

170 In 1960 runoff observations at the Yuzhny brook were begun, rationalization of the
171 precipitation network was continued, and radio rain gauges were installed.

172 In 1963 two new water-balance sites (##2 and 3) were organized.

173 In 1968 runoff measurements were started at the unique research object, at the Morozova
174 brook catchment, which has no vegetation cover and is composed of rocky talus.

175 In 1969 the Kolyma runoff station was renamed into the Kolyma Water Balance station
176 (KWBS). In these years there was a transition to broad experimental water balance observations
177 of all of its elements and to an enhanced technical level of research.

178 Since 1970, the KWBS carried out snowpack observations at avalanche sites of the
179 Tenkinskaya road, as well as stratigraphy, temperature and physical and mechanical properties of
180 snow at four sites. Since 1980 there were introduced additional observations on dynamics of
181 icing formation at the Kontaktovy creek. In 1982 observations on soil moisture were started at 3
182 agro-hydrological sites at the fields of the «Kulu» state farm.

183 In 1976 the station hosted a delegation of USA scientists. They highly praised the
184 professional and personal qualities of the station's staff members, their commitment, on which
185 extensive field studies and theoretical works were based, despite the equipment being rather
186 simple and living conditions extreme. According to Slaughter and Bilello (1977), the data
187 recorded at the KWBS, were unique and unprecedented for world practice.

188 Since the beginning of the 1990s, the research program at the station has been gradually
189 cutting back. After 1997, water balance observations at the KWBS were ceased. One weather
190 station and five runoff gauges functioned at the KWBS up to mid of June, 2013, when an



191 extreme flash flood destroyed four level gauges. Nowadays only standard observations are
192 conducted at the Nizhnyaya meteorological site and at the Kontaktovy (Nizhny) runoff gauge.

193 3. Data description

194 3.1 Meteorological observations

195 The observations of meteorological elements were carried out at three meteorological
196 stations in different periods (Fig. 2). The database includes daily values of air temperature, water
197 vapour pressure, vapour pressure deficit, atmospheric pressure, wind speed, low and total cloud
198 amount, and surface temperature (Table 1).

199 The meteorological station Verkhnyaya (1220 m a.s.l., 1948-1972) was located in the
200 upper reaches of the Dozhemerny brook in the saddle between two hills. The horizon is closed
201 by the hills from the south and north which are at 30-40 m distance from the site. The horizon is
202 open from the east and west, strong winds are observed here. The surface at the station plot is
203 hummocky, covered with grassy vegetation with no woody vegetation around it. The nearest
204 building – the station house – was located 48 m away from the station plot. The depth of
205 seasonal thaw of permafrost reaches 1.5 m.

206 The meteorological station Nizhnyaya (850 m a.s.l., 1948-1997) is located on the edge of
207 a larch forest, on the terrace-watershed between the Kontaktovy creek and the Ugroza brook,
208 which has a slight slope to the SW. The nearest trees are located 50 m away, the buildings – 100
209 m from the station. The site is surrounded by mountains up to 1400 m a.s.l., the nearest of them
210 are at a distance of 200-500 m. The angle of the horizon is 11 degrees. The height of the weather
211 vane is 11.3 m. The surface at the station plot is covered by hummocks, with moss, peat and
212 individual bilberry and blueberry bushes. The area is surrounded by a sparse larch forest from
213 the north, east and south. The depth of permafrost seasonal thaw reaches 1.5 m.

214 The meteorological station Kulu (670 m a.s.l., 1981-1991) was located on the right cliff
215 of the broad valley of the Kulu river. The slope has a western exposure (4-6 degrees). The angle
216 of the horizon is 4 degrees. The height of the weather vane is 10.7 m. The area is surrounded by
217 larch trees of 6-8 m height from the west, north and east. The soils are loamy with the inclusion
218 of small gravel. The underlying surface consists of berry, grass and sphagnum mosses,
219 sometimes bare soil. The depth of seasonal thawing of permafrost reaches 1.8-2 m.

220 In September 1992, the Kulu station (635 m a.s.l., 1992-1997) was moved to the
221 residential building of the KWBS at the south-eastern part of the Kulu village. Residential and
222 technical buildings are located around the meteorological plot. There was a road to the south of
223 the station. The soil is marshy, covered with rubble and grass. The angle of the horizon is 4
224 degrees. This new location of the Kulu station is marked as Kulu2 station in the database.

225 The observations at all meteorological plots were conducted according to the program of
226 a meteorological station of the 2nd category.

227 The climate of the study area is severely continental with harsh long winters and short but
228 warm summers. Average annual temperature at the Nizhnyaya meteorological plot during 1949-
229 1996 is -11.3 °C. Mean monthly temperature in January was -33.6 °C, in July +13.2 °C (Fig. 3-
230 4). The absolute minimum daily temperature of -53.0 °C was registered in 1982 and the absolute
231 maximum daily temperature was +22.8 °C (1988). The period of negative air temperatures lasts
232 from October to April, freeze-free period is, on average, 130 days long.

233 Air temperature inversions are observed at the KWBS. In December air temperature
234 gradient reaches +2.0, in May it accounts for -0.5°C (100 m)⁻¹ respectively. Mean wind velocity
235 is more than twice higher at Verkhnyaya station in comparison with Nizhnyaya station (3.0 and
236 1.3 m s⁻¹ accordingly).

237

238 3.2 Precipitation

239 In total, the precipitation was observed at 47 gauges within KWBS territory during
240 different periods. Continuous daily all-year around precipitation data is available for the period
241 1948-1997 for the gauge (#12) at meteorological station Nizhnyaya and for the gauge (#54) at



242 meteorological station Kulu for 1981-1997. Four gauges have the data of daily totals during
243 warm season for the period for more than 30 years and another 18 gauges for different shorter
244 periods. Usually the start of daily observations at those gauges was initiated by the beginning of
245 snowmelt period and lasted until the end of September. Monthly sums of precipitation were
246 measured at 30 gauges, 10-days and 5-days sums – at 21 and 18 gauges respectively.

247 Rain-gauge stations for measuring daily precipitation totals in 1948 were equipped with
248 the Nipher-shielded rain gauges and later – with the Tretyakov-shielded precipitation gauges. In
249 1948-1958 the observations were carried out with both devices in parallel, after 1959 only
250 Tretyakov gauges were used.

251 There were three main types of precipitation gauges which were used at the KWBS – the
252 Tretyakov, the Kosarev and pit rain gauge (Fig.5). In 1960-1963 there was an attempt to register
253 precipitation with automatic radio-precipitation gauges, but due to improper performance of the
254 devices those observations were stopped.

255 In 1988, precipitation observations at the KWBS were carried out with 36 precipitation
256 and rain gauges, distributed relatively evenly throughout the area and altitudinal zones. Average
257 density of the precipitation network at that time accounted for 1.6 units per 1 km².

258 Precipitation at Nizhnyaya meteorological plot in the 1949-1997 period varied from 229
259 (1991) to 474 (1990) mm per year with mean value of 342 mm. Maximum and minimum amount
260 of precipitation is observed in July and March and account for 71 and 7 mm respectively (Fig. 4-
261 5).

262

263 3.3 Snow surveys

264 Snow cover at KWBS is formed in the first weeks of October, and melts in the third week
265 of May. Snow cover observations were started in 1950 and initially conducted at two
266 catchments, at two meteorological plots and four typical squares. In 1959-1960 the number of
267 catchments with snow surveys reached five. Up to 1971, snow surveys were conducted once per
268 month starting in November and finishing in May at small catchments (the Severny, Yuzhny,
269 Dogdemerny and Vstrecha) and once before spring snowmelt at the Kontaktovy – Nizhny. Since
270 1972, the observations were reduced to one survey per year (usually at the end of April) for all
271 watersheds. Table 2 shows the number of snow routes, their total length and number of
272 measurement points, including their distribution among different landscapes of the catchments
273 (Fig. 6). Snow depth was measured every 10 m, snow density – every 50 m.

274 Mean, maximum and minimum observed snow water equivalent (SWE) before spring
275 freshet at the Kontaktovy – Nizhny accounted for 121, 213 (1985) and 59 (1964) mm
276 respectively in the 1960-1997 period. In general, rocky talus and tundra bush landscape are
277 characterized by lower SWE due to wind blowing. Much snow is accumulated in the forest
278 landscape. However, at the Morozova brook watershed which is fully covered by rocky talus
279 landscape, mean SWE before snowmelt was estimated as 161 mm with the maximum value of
280 298 mm observed in 1985 reaching 0.99 m snow height (Table 3, Table 4, Fig. 7).

281

282 3.4 Soil evaporation

283 From 1950 to 1953, five soil evaporation plots were opened at the sites with diverse
284 underlying surfaces and different expositions and altitudes. The plots were equipped with the
285 Rykachev and Gorshenin evaporimeters with evaporation area 1000 cm². The observations until
286 1958 are considered to be approximate due to the absence of accompanying rain gauges and
287 scales of required accuracy. From 1958 to 1966, the measurements were conducted at the soil
288 evaporation plot, located near the Nizhnyaya weather station. The observations of evaporation
289 were carried out with two evaporimeters GGI-500 and the Rykachev evaporimeter, precipitation
290 – with a rain pit-gauge. Three soil evaporation sites were established in different landscapes – in
291 the Vstrecha (1967), Morozova (1971) and Yuzhny (1977) brooks basins. The measurements
292 were carried out with standard weighing evaporimeters GGI-500-50, which, due to the physical



293 proximity of permafrost, were changed to GGI-500-30, meaning that their height was decreased
294 to 30 cm.

295 Average values of annual soil evaporation were previously estimated by Semenova et al.
296 (2013) and Lebedeva et al. (2017) based on partial KWBS data set as the following: 140 mm for
297 larch and swampy sparse growth forest, 110 mm in dwarf cedar and alder shrubs of tundra belt,
298 and about 70 mm for rocky talus. This database presents full data of observations which allow
299 for correction of the values as described below.

300 The highest values of soil evaporation during the summer period were observed at the
301 larch forest (site 9) and reached 136 mm. At a similar landscape (site 1), this value is lower, at
302 119 mm, which indicates the influence of local factors. The lowest values of soil evaporation are
303 104 mm at the plot located at dwarf cedar tree bush (site 7). In July, soil evaporation values
304 range from 33 to 40 mm, depending on the landscape. In September, the contribution of
305 evaporation decreases to 14-24 mm (Table 5).

306

307 **3.5 Snow evaporation**

308 Snow evaporation observations were conducted at the KWBS from 1951, but only the
309 data for the period 1968-1992 is considered to be consistent and reliable and is published in
310 described database. From 1968 to 1981, the observations were conducted with standard
311 evaporimeters GGI-500-6 at weather plot Nizhnyaya. In 1981 the snow evaporation observations
312 were transferred to the Kulu weather plot and lasted till 1992.

313 Observations of evaporation from snow were made mainly in the fall (September,
314 October) and in the spring (May - March). During winter months (January - February), the
315 observations were made only until 1973, because the amount of evaporation from snow proved
316 to be extremely insignificant for water balance. In the spring, during the intensive snowmelt,
317 additional weighing of the evaporimeters was carried out every 3-6 hours. In the database, the
318 evaporation values for night (20-8 hours) and daytime (8-20 hours) intervals are presented, only
319 those values that correspond to a full 12-hour period of observations are published. The accuracy
320 is 0.01 mm. The average values of evaporation from snow in mm per day are as follows:
321 January-February – -0.04; March – +0.09; April – +0.40; May – +0.74; September – +0.20;
322 October – +0.01. Typical values of evaporation from snow for 1976-1977 are presented at Fig. 8.

323

324 **3.6 Thaw/freeze depth**

325 Since 1952, the observations on permafrost seasonal thaw dynamics were conducted at
326 the KWBS. Danilin cryopedometers were installed at permafrost observation sites (Snyder et al.,
327 1971) which mostly were located in the approximate vicinity of Nizhnyaya and Verkhnyaya
328 weather stations at the slopes with different aspects and landscapes. During 1952-1997 38
329 cryopedometers were functioning in total. The longest observation period is 33 continuous years
330 (cryopedometer 17.5 located at the forest with bushes, maximum thawing is 130 cm, 1964-
331 1997). The deepest values of thawing were observed in rocky talus landscape and can reach
332 more than 240 cm. The shallowest values of thawing range from 60 to 70 cm at swampy forest.
333 Thawing of soils at the forest zone varies in large ranges and depends on the location of the
334 cryopedometer at a slope (Table 6, Fig. 9). Despite the fact that permafrost observation sites
335 were equipped with special bridges for observers to come close (Fig. 5), eventually surface
336 damage in the area where the device was installed began to influence thaw depth.

337

338 **3.7 Streamflow**

339 Runoff observations were carried out at 10 catchments: the creek Kontaktovy (the gauges
340 Verkhny, Sredny, Nizhny), brooks Morozova, Yuzhny, Vstrecha, Vstrecha (the mouth),
341 Dozhdemerny, Severny, Ugroza (Fig. 10). Key characteristics of the catchments are listed in
342 Table 7. All the water level gauges were equipped with «Valdai» water level recorders, as well
343 as needle and hook water level gauges. In spring and autumn, when recorders did not work
344 properly due to ice on the creeks, discharge was measured more frequently, every 4 hours. To



345 prevent the recorder floats from freezing, the wells were heated with electric bulbs. At small
346 Morozova and Yuzhny brooks runoff was measured by means of a V-notch weir, at the Severny
347 brook – with a flow measuring flume.

348 Flow at KWBS begins in May, most of it occurs in summer. For the summer period,
349 rainfall floods are typical (Fig. 11).

350 In October, small creeks freeze completely. Along the whole length of the Kontakovy
351 Creek, channel taliks can be found. They go all the way through the layer of alluvial sediments
352 and their depth reaches 15 m in the cross section of the Nizhny hydrological gauge (Mikhaylov,
353 2013). Taliks freeze in winter only partially.

354 Annual flow depth of the Kontakovy stream basin with area 21.3 km² (average altitude
355 1070 m) is 281 mm for the period 1948-1997, it increases with the elevation and at the
356 Morozova catchment (mean elevation 1370 m, basin area 0.63 km²) reaches 453 mm (1969-
357 1996). The flow from south-facing (Severny) and north-facing (Yuzhny) micro-watersheds with
358 area of 0.38 and 0.27 km² are 227 and 193 mm for the period 1960-1997 respectively.

359 In the database, mean daily values of streamflow are presented.

360 4. Data availability

361 All data presented in this paper are available from the “PANGAEA. Data Publisher for
362 Earth & Environmental Science” (see Makarieva et al., 2017,
363 <https://doi.pangaea.de/10.1594/PANGAEA.881731>). The directory includes 12 elements:

- 364 1. daily precipitation time series at different gauges within Kolyma Water-Balance Station
365 (KWBS), 1948-1997;
- 366 2. daily runoff time series at different gauges within Kolyma Water-Balance Station
367 (KWBS), 1948-1997;
- 368 3. evaporation time series at different sites within Kolyma Water-Balance Station (KWBS),
369 1950-1997;
- 370 4. meteorological observations at different sites within Kolyma Water-Balance Station
371 (KWBS), 1948-1997;
- 372 5. monthly precipitation time series at different gauges within Kolyma Water-Balance
373 Station (KWBS), 1948-1997;
- 374 6. precipitation (10 day sum) time series at different gauges within Kolyma Water-Balance
375 Station (KWBS), 1962-1997;
- 376 7. precipitation (5 day sum) time series at different gauges within Kolyma Water-Balance
377 Station (KWBS), 1966-1997;
- 378 8. snow survey line characteristics within Kolyma Water-Balance Station (KWBS), 1959-
379 1997;
- 380 9. snow survey time series at different sites within Kolyma Water-Balance Station (KWBS),
381 1950-1997;
- 382 10. soil temperature time series at the Nizhnyaya meteorological station within Kolyma
383 Water-Balance Station (KWBS), 1974-1981;
- 384 11. thaw depth and snow height time series at different sites within Kolyma Water-Balance
385 Station (KWBS), 1954-1997;
- 386 12. evaporation time series from snow within Kolyma Water-Balance Station (KWBS),
387 1968-1992.

388

389 5. The future of the KWBS

390 In summer 2016, with the assistance of Melnikov Permafrost Institute of Siberian Branch
391 of Russian Academy of Science, a group of specialists, consisting of representatives of different
392 scientific institutions, conducted a reconnaissance survey of the KWBS in order to find out if it
393 was possible to carry out scientific research and stationary monitoring of permafrost and



394 hydrological processes at the station. Despite rather difficult logistic access to the KWBS, it was
395 considered possible to organize accommodation and provision of the station for the period of
396 summer expeditions. At first, the main goal of research resumption at the station would be a
397 renewal of regular observations of runoff, meteorological elements and active layer dynamics at
398 three small catchments (Morozova, Yuzhny, Severny) and the KWBS main-stream outlet
399 (Nizhny gauge) using advanced equipment with automatic data recording. As a result, some
400 unique runoff observations series – over 60 years long – will be continued, which will allow for
401 evaluation of climatic impact on permafrost and provide a scientifically based forecast on current
402 and future climate change impact on the hydrological regime.

403 During short 3-4 week field trips at the beginning and at the end of the warm (and
404 hydrological) season, it would be also possible to study specific processes of runoff formation
405 under permafrost conditions. Slope runoff occurs unevenly, and is concentrated in particular
406 areas, the drainage zones or preferential path flows. Reconnaissance surveys of the Kontakovy
407 creek catchment at the KWBS territory, 2016, revealed that there are several types of such zones
408 of slope runoff concentration. Another possible scientific task is to evaluate the role of cryogenic
409 redistribution of runoff, which regularly occurs due to ice freezing-melting in coarse-grained
410 slope deposits. Similar studies have already been carried out in mountain regions of permafrost,
411 including the KWBS (Sushansky, 1999; Bantsekina, 2001; Bantsekina, 2002; Boyarintsev et al.,
412 2006). Another research issue is the study of floodplain taliks (Mikhaylov, 2013) and aufeis
413 (Alexeev, 2016) and their impact on hydrological processes in the mountainous part of the
414 continuous permafrost zone. Field trips for a limited group of scientists could be covered with
415 relatively modest financial support through research grants. In the future, could the aim is for the
416 KWBS to get back its status of a research station, to receive state funding, obtain sponsor
417 support from gold mining companies of the Magadan region and become an international center
418 for complex studies in the field of permafrost hydrology.

419 The KWBS is situated in the region where monitoring of natural processes is extremely
420 sparse. From 1986 to 1999, the number of hydrological gauge stations in Far-East parts of
421 Siberia decreased by 73% (Shiklomanov et al., 2002). Resumption of water balance observations
422 and organization of complex research of permafrost, climate, and landscape, hydrological and
423 hydrogeological processes based on data collected at the KWBS would make it possible to get
424 new data, representative for the understudied territory of the Arctic in the context of
425 environmental changes. Considering insufficient knowledge about this territory and available
426 long-term data, the KWBS has the prospect to become a highly demanded complex international
427 center for testing natural process models at different scales – from point to regional, – validation
428 of remote sensing products and a place for multidisciplinary field research.

429 More than 20% of the Northern Hemisphere is covered by permafrost. Three of the four
430 largest rivers of the Arctic Ocean basin flow through Siberia. Many studies highlight ongoing
431 and intensifying changes of water, sediment and chemical fluxes at all spatial scales but
432 mechanisms of changes and future projections are highly uncertain. There are no research
433 centers that could conduct focused studies of hydrological processes at catchments in the
434 permafrost region in Russia. The KWBS incorporation into the international network for
435 monitoring natural processes in cold regions (Interact, SAON, CALM, GTN-P, etc.) could
436 significantly enhance international cooperation for better understanding of cold-region hydrology
437 for the last 70 years, present and future.

438 Nowadays, the resumption of continuous observations and research at the Kolyma station
439 appears to be a critical task due to increased interest in the natural processes of the Arctic region.
440 Present-day data, following the KWBS long-term observations series, could become a valuable
441 indicator of climate change and a basis for studying its impact on the state of the permafrost and
442 its associated hydrological regime. Currently, as the station infrastructure is still partly intact,
443 and some of the specialists who worked at the KWBS are still active and willing to help, it is
444 necessary to gain attention and support from the Russian and international scientific community
445 regarding the renewal of the KWBS before it is too late.



446 6. Conclusions

447 The presented dataset describes water balance, hydrometeorological and permafrost
448 related components at small research watershed in mountainous permafrost zone of North-East
449 Russia, the Kolyma water-Balance Station (KWBS). It includes 50 years of continuous daily
450 meteorological and streamflow data for main meteorological plot and runoff gauge of KWBS
451 and daily data of shorter periods for another two meteorological sites and 9 runoff gauges.
452 Meteorological data includes values of air temperature, water vapour pressure, vapour pressure
453 deficit, atmospheric pressure, wind speed, low and total cloud amount, and surface temperature.
454 The dataset also includes all-year daily, warm period daily, 5-, 10-days and monthly sums for 47
455 (in total) precipitation gauges within KWBS territory for different time spans over the period
456 1948-1997. It also contains soil evaporation data from different landscapes, snow evaporation
457 series from two sites; snow surveys results for different watersheds within KWBS, as well as
458 thaw/freeze depths at more than 30 observational sites.

459 The dataset is important because it characterizes the natural settings, which, on the one
460 hand, are nearly ungauged, and on the other hand, are representative for the vast mountainous
461 territory of Eastern Siberia and North-East Russia. It is unique because it combines water
462 balance, hydrological and permafrost data which allow for studying permafrost hydrology
463 interaction processes within the range of all scientific issues, from models development to
464 climate change impacts research.

465

466 7. Author contribution

467 O. Makarieva and N. Nesterova digitized and prepared the dataset for publication with
468 assistance from L. Lebedeva and S. Sushansky. The data were collected in 1948-1997 by
469 Hydrometeorological Service of USSR and Russia and published in Observation Reports (1948-
470 1997).

471

472 8. References

473 Alekseev V.P. (2016) Long-term variability of the spring taryn-aufeises. *Ice and Snow*,
474 56(1), 73-92. DOI:10.15356/2076-6734-2016-1-73-92

475 Alexeev V.P., Boyarintsev Ye.L., Gopchenko Ye.D., Serbov N.G., Zavalii N.V. (2011)
476 The mechanism of cryogenic runoff control at the formation of water balance of small mountain
477 rivers in the area of permafrost rocks. *Ukrainian hydrometeorological Journal*, 8, 182-194 (in
478 Russian).

479 Bantsekina T.V. 2003. Peculiarities of hydrothermal regime of seasonal thawing layer in
480 coarsely clastic rocks during spring summer period (with an example of Upper Kolyma
481 highland). PhD thesis, Melnikov Permafrost Institute, Yakutsk, 137pp (in Russian).

482 Bantsekina T.V. (2002) Temperature regime and dynamics of icing of coarse-grained
483 slope deposits without filling during spring-summer period (case study of the Kontaktovy creek).
484 *Kolyma*, 4, 9-13. (in Russian)

485 Bantsekina T.V. (2001) Dynamics of coarse-grained slope deposits icing during spring
486 thawing. *Kolyma*, 2, 28-31 (in Russian).

487



- 488 Bantsekina, TV, Mikhailov, V.M. (2009) To the assessment of the role of intra-soil
489 condensation of water vapor in the formation of the thermal and water regimes of large
490 sediments, Earth's Cryosphere, XIII, 1, 40-45
- 491 Boyarintsev, E.L (1980) Estimation of the losses of spring floods in the Upper Kolyma
492 basin, Meteorology, Climatology and Hydrology, 16, 19-24.
- 493 Boyarintsev E.L. (1988) Azonal factors of rainfall runoff formation in the territory of
494 Kolyma WBS. Proceedings DVNIGMI, 135, 67-93, (in Russian).
- 495 Boyarintsev E.L., Gopchenko E.D., Serbs N.G., Legostaev G.P. (1991) Concerning the
496 condensation of air vapors in the active layer of permafrost. M., Dep. in the IC VNIIGMI-WDC
497 16.01.91, 1046 GM 91. P. 17, (in Russian).
- 498 Boyarintsev E.L., Gopchenko E.D. et al. (1992) Summer period water balance of small
499 mountain catchments of the permafrost and its calculation. Meteorology, climatology and
500 hydrology, 27, 105-116.
- 501 Boyarintsev E.L., Nikolaev S.N. (1986) Groundwater runoff from small watersheds of
502 permafrost zone // Materials of scientific. Conf. on the problems of hydrology of the rivers of the
503 BAM zone and the Far East. L, Gidrometeoizdat, 297-307 (in Russian).
- 504 Boyarintsev E.L., Serbov N.G., Popova N.I. (2006) Formation of the water balance of the
505 spring floods of the small mountain watersheds of Upper Kolyma (based on materials from the
506 Kolyma Water-Balance Station) - Bulletin of the North-Eastern Scientific Center, Far-Eastern
507 Branch of the Russian Academy of Sciences, 4, 12-19 (in Russian).
- 508 Glotov V.E. (2002) Ground water of the Kontaktovy Creek watershed as a factor of
509 general drainage system formation. In: Glotov V, Ukhov N (eds.) Factors affecting the formation
510 of a general drainage system of minor mountain rivers in sub-arctic areas. SVKNII DVO RAN.
511 Magadan, 102-141 (in Russian).
- 512 Gusev E.M., Nasonova O.N., Dzhogan L.Ya. (2006) The Simulation of Runoff from
513 Small Catchments in the Permafrost Zone by the SWAP Model. Water Resources, 33, 2, 133-
514 145.
- 515 Hinzman, L.D., N. Ishikawa, K. Yoshikawa, W.R. Bolton, K.C. Petrone (2002)
516 Hydrologic Studies in Caribou Poker Creeks Research Watershed in Support of Long term
517 Ecological Research. Eurasian Journal of Forest Research. 5-2:67-71.
- 518 Informational letter #2 (1988) 40 years anniversary of the Kolyma Water Balance Station.
519 Kolymskiy Territorial Office on Hydrometeorology, Magadan (in Russian).
- 520 Korolev, Yu. B. (1984) Mapping of the vegetation cover in connection with an
521 assessment of its hydrological role (by the example of the Upper Kolyma). PhD thesis for
522 candidate of biological science, Magadan University, Yakutia, 231 p. (in Russian)
- 523 Kuchment L.S., Gelfan A.N., Demidov A.I. (2000) A Model of Runoff Formation on
524 Watersheds in the Permafrost Zone: Case Study of the Upper Kolyma River. Water Resources,
525 27, 4, 435-444.



- 526 Kuznetsov A.S., Nasibulin S.S., Ipatieva A.I. (1969) The first results of the study of
527 water balance on the rivers of the Upper Kolyma basin // Collection of works of the Magadan
528 Hydrometeorological Observatory, 2, Magadan, 98-121, (in Russian).
- 529 Laudon H, Spence C, Buttle J, Carey SK, McDonnell JJ, McNamara JP, Soulsby C,
530 Tetzlaff D. (2017) Saving northern high-latitude catchments. *Nature Geoscience*, 10, 324-325.
531 doi: 10.1038/ngeo2947.
- 532 Lebedeva L.S., Semenova O.M., Vinogradova T.A. (2015) Hydrological modeling:
533 seasonal thaw depths in different landscapes of the Kolyma Water Balance Station (Part 2).
534 *Earth's Cryosphere*, XIX, 2, 35-44.
- 535 Lebedeva L.S., Makarieva O.M., Vinogradova T.A. (2017) Spatial variability of the
536 water balance elements in mountain catchments in the North-East Russia (case study of the
537 Kolyma Water Balance Station). *Meteorology and Hydrology*, 4, 90-101 (in Russian).
- 538 Mikhaylov V.M. (2013) Floodplain taliks of North-East of Russia. Novosibirsk, "Geo",
539 244 p. (in Russian)
- 540 Observation Reports. Kolyma Water Balance Station, 1948–1997. Issues 1–40, 1959–
541 1998, Kolyma UGKS, Magadan (in Russian)
- 542 Peterson, B. J., R. M. Holmes, J. W. McClelland, C. J. Vörösmarty, R. B. Lammers, A. I.
543 Shiklomanov, I. A. Shiklomanov, and S. Rahmstorf (2002) Increasing river discharge to the
544 Arctic Ocean, *Science*, 298, 2171–2173.
- 545 Reynuk I.T. (1959) Condensation in active layer of permafrost. Proceedings of All Union
546 Scientific and Research Institute of Gold and Rare Metals. Issue 13 (permafrost studies), 1-24.
- 547 Semenova, O., Lebedeva, L., Vinogradov, Yu., (2013) Simulation of subsurface heat and
548 water dynamics, and runoff generation in mountainous permafrost conditions, in the Upper
549 Kolyma River basin, Russia. *Hydrogeol. J.* 21 (1), 107–119, doi:10.1007/s10040-012-0936-1
- 550 Shiklomanov, A. I., R. B. Lammers, and C. J. Vörösmarty (2002) Widespread decline in
551 hydrological monitoring threatens Pan-Arctic Research, *Eos Trans. AGU*, 83(2), 13–17,
552 doi:10.1029/2002EO000007.
- 553 Shmakin, A.B. (1998) The updated version of SPONSOR land surface scheme: PILPS
554 influenced improvements. *Glob. Plan. Change*, 19, 49-62.
- 555 Slaughter C.W., Billelo M.A. (1977) Kolyma Water Balance Station, Magadan oblast,
556 Northeast U.S.S.R.: United Station - Soviet Scientific Exchange Visit, Special Report 77-155,
557 Army Gold Regions Research and Engineering Laboratory. Hanover, 66.
- 558 Snyder F, Sokolov A, Szesztay K. (1971) Flood Studies: An International Guide for
559 Collection and Processing of Data. Unesco: Paris; 52 pp.
- 560 Sushansky S.I. (1999) Peculiarities of water balance elements in the Morozova Creek
561 catchment. *Kolyma*, 1, 33-40 (in Russian).



562 Sushansky S.I. (2002) History of creation, methods, objects and some results of studies in
563 the Kolyma water balance station. In: Glotov V, Ukhov N (eds.) Factors affecting the formation
564 of a general drainage system of minor mountain rivers in sub-arctic areas. SVKNII DVO RAN:
565 Magadan, 18-35.

566 Tikhmenev P.Ye. (2008) Peculiarities of succession process in disturbed lands of the
567 Kolyma river basin. Natural-resources potential, ecology and sustainable development of
568 Russian regions: Collected works VI International scientific-practical Conf. Penza: PGAU, 273-
569 275.

570 Vinogradov, Yu.B., Semenova, O.M., Vinogradova, T.A. (2015) Hydrological modeling:
571 heat dynamics in a soil profile (Part 1). Earth's Cryosphere (Kriosfera Zemli) XIX (1), 11–21.

572 Zhuravin S. (2004) Features of water balance for small mountainous basins in East
573 Siberia: Kolyma Water Balance Station case study. IAHS Publ 290, IAHS, Wallingford, UK,
574 28–40.

575

576 Table 1 List of meteorological observation data

Station	Nizhnyaya	Verkhnyaya	Kulu	Kulu2
Latitude	61.85	61.86	61.88	61.88
Longitude	147.67	147.61	147.43	147.44
Elevation, m	850	1220	670	635
Air temperature	1948-1997	1948-1972	1981-1991 1992-1997	
Water vapour pressure				
Vapour pressure deficit	1974-1997	n/a		
Atmospheric pressure	1951-1997	1951-1972		
Low cloud amount	1948-1997	1948-1972		
Total cloud amount				
Wind speed				
Surface temperature				

577

578 Table 2 Number of snow routes, their total length and number of measurements points –
579 maximum and minimum values within the whole period of observations

Watershed	Period	N	L	MP
Yuzhny	1960-1997	4	1400-1540	144-154
Severny	1950-1997	4/10	1950/2130	23/207
Morozova	1968-1997	2/5	960/2645	98/534
Ugroza	1983-1997	1	1200	120
Dozhdemerny	1959-1971	3	3240/5720	327/575
Vstrecha	1950-1997	1/17	2110/10850	119/1091
Kontaktovy	1960-1997	2/4	4830/13100	485/1314

580 N – amount of snow routes; L – total length of the route, m; MP – amount of snow thickness
581 measurement points.



582

583 Table 3 Mean, maximum and minimum observed snow water equivalent (SWE) (mm) before
584 spring freshet at different landscapes of the Kontaktovy creek watershed, 1960-1997

Landscape	SWE		
	mean	max (1985)	min (1964)
Forest	144	265	79
Dwarf cedar tree bush	127	247	39
Rocky talus	100	182	46
Boulders	66	127 (1974)	2
Kontaktovy Creek	121	213	59

585

586

587 Table 4 Mean, maximum and minimum snow water equivalent (SWE) (mm) before spring
588 freshet at different watersheds within KWBS

Watershed	Period	SWE		
		mean	max	min
Yuzhny	1960-1997	121	166	70
Severny	1950-1997	126	232	62
Morozova	1968-1997	161	298	71
Ugroza	1983-1994	133	200	93
Dozhdemerny	1959-1971	82	111	53
Vstrecha	1951-1997	123	213	60

589

590 Table 5 Mean evapotranspiration (mm) in June – September at different landscapes of KWBS

# site	Landscape	Period	Elevation, m	Slope aspect	Jun	Jul	Aug	Sep	Total*
1	Larch forest	1962-1997	850	n/a	35	37	30	17	119
6	Swampy sparse growth forest	1969-1982	970	North	37	38	30	19	124
7	Dwarf cedar tree bush	1972-1997	1020	n/a	30	33	25	17	104
8	Dwarf cedar tree bush	1976-1997	900	South	47	40	30	14	131
9	Larch forest	1982-1992	669	West	36	39	37	24	136

591 *the sum for warm period

592

593 Table 6 Maximum depth of thawing at the different landscapes

# site	Watershed	Landscape	Period	Elevation, m	Maximum depth of thawing, cm
1	Kontaktovy	Forest	1954-1966	841	150
6	Dozhdemerny	Rocky talus	1960-1965	1048	>240
9	Severny, Ugroza	Rocky talus	1954-1966; 1977-1978	986	168
12	Vstrecha, Severny	Dwarf cedar tree bush at rocky talus	1954-1962; 1966-1968; 1971-1997	866	157
15	Dozhdemerny	Dwarf cedar tree bush at rocky talus	1958-1968; 1970-1982	952	>150
17	Vstrecha	Forest	1960-1965, 1969	914	>124
18 bh7	Kontaktovy	Peat bogs	1959-1960	835	69
18 bh8	Kontaktovy	Peat bogs	1959-1960	835	64

594



595 Table 7 The characteristics of KWBS watersheds

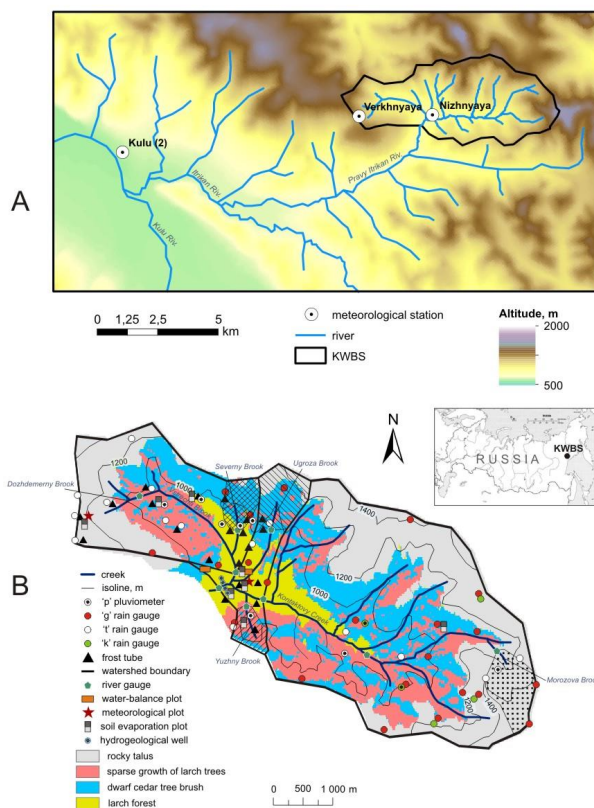
Code	Catchment (creek – outlet)	Period	Area, km ²	X	Y	Stream length, km	Mean watershed width, km	Mean stream slope, ‰	Mean basin slope, ‰	Catchment altitude (max-min, mean), m	Mean annual flow, mm	Maximum observed daily discharge, m ³ s ⁻¹
1104	Yuzhny	1960-1997	0.27	61.84	147.66	0.51	0.35	235	303	1110-917, 985	193	0.14
1107	Severny	1958-1997	0.38	61.85	147.66	0.74	0.38	175	388	1300-880, 1020	227	0.18
1103	Morozova	1968-1996	0.63	61.84	147.75	0.97	0.45	326	649	1700-1100, 1370	453	0.44
1624	Ugroza	1983-1991	0.67	61.86	147.67	0.9	0.74	218	461	1270-914, 1260	354	0.27
1106	Dozhdemerny	1952-1971	1.43	61.86	147.63	0.87	0.99	220	432	1450-950, 1180	208	0.31
1105	Vstrecha	1949-1997	5.35	61.85	147.66	3.4	1.5	92	346	1450-833, 1060	237	3.15
1100	Kontaktovy – Verkhny	1973-1980	5.53	61.84	147.70	2.8	2.1	185	473	1700-909, 1070	317	2.52
1625	Vstrecha – the mouth of Ugroza Cr.	1984-1996	6.57	61.84	147.66	3.6	1.8	76	406	1450-831, 1070	283	2.6
1101	Kontaktovy – Sredny	1948-1997	14.2	61.84	147.67	6.2	2.8	65.2	413	1700-842, 1120	289	7.02
1102	Kontaktovy – Nizhny	1948-1997	21.3	61.85	147.65	7.1	3.7	57.6	413	1700-823, 1070	281	8.15

596



597

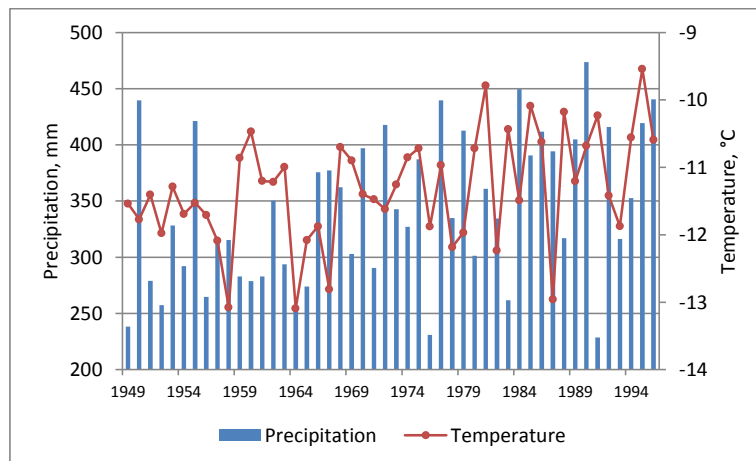
598 Fig. 1 The view of the Kolyma Water Balance Station, August 2016 (photo by O. Makarieva)



599

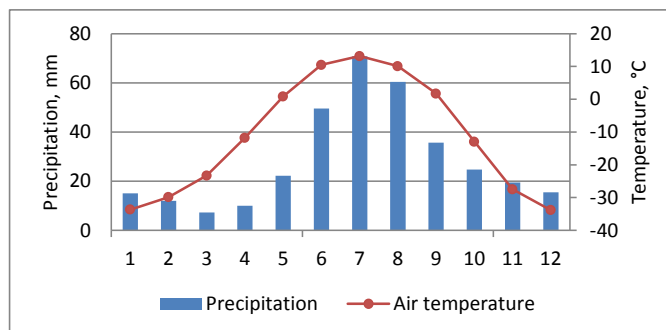
600 Fig. 2 Scheme of the Kolyma Water Balance Station (KWBS) indicating the location of observation sites

601



602

603 Fig. 3 Annual precipitation (mm) and air temperature (°C) at Nizhnyaya weather station, 1949-1996



604

605 Fig. 4 Mean monthly precipitation (mm) and air temperature (°C) at Nizhnyaya weather station, 1949-
606 1996



607

608

Fig. 5 Meteorological observations: rain-gauge plot #2, 1959.



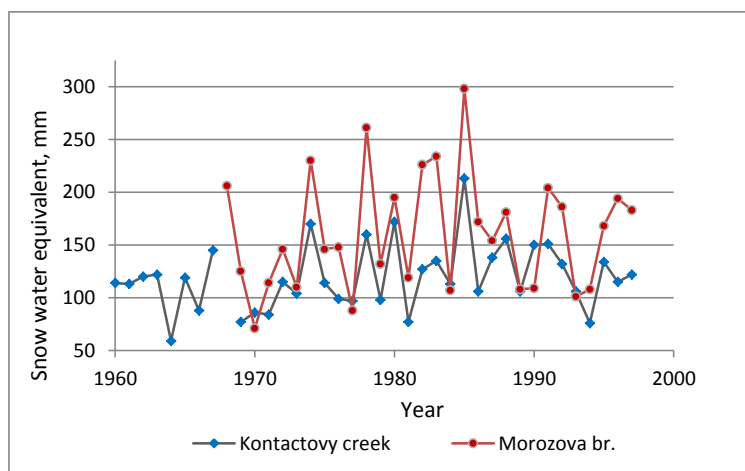
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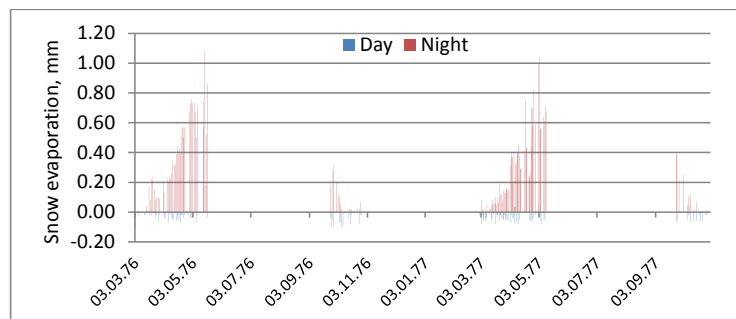
Fig. 6 Snow survey at the Kontakovy creek catchment (A, B) and measurement of snow density (C),
1960



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614

Fig. 7 Snow water equivalent at the Kontaktovy creek and Morozova br.

615



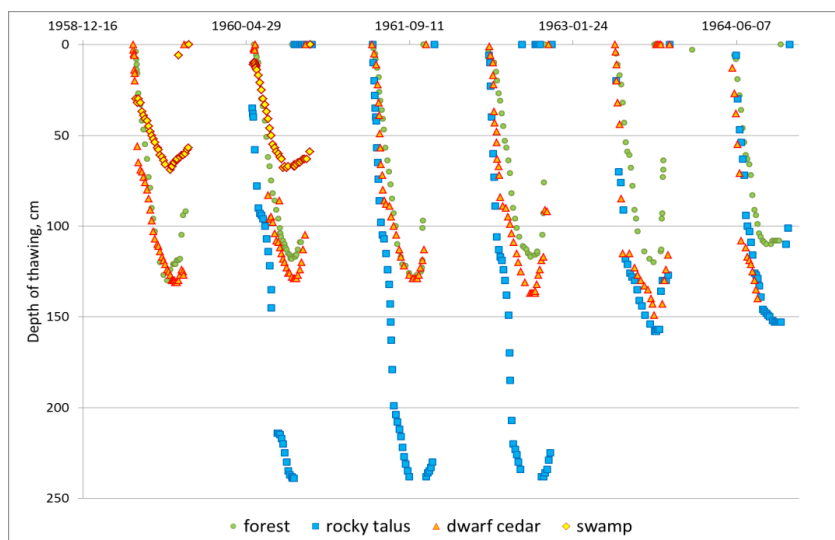
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Fig. 8 Snow evaporation (mm) during day and night period, 1976-1977

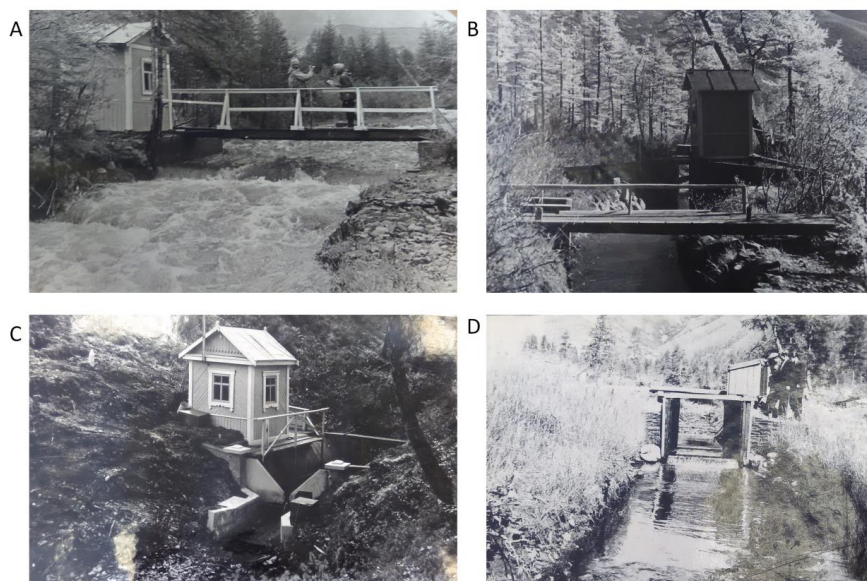
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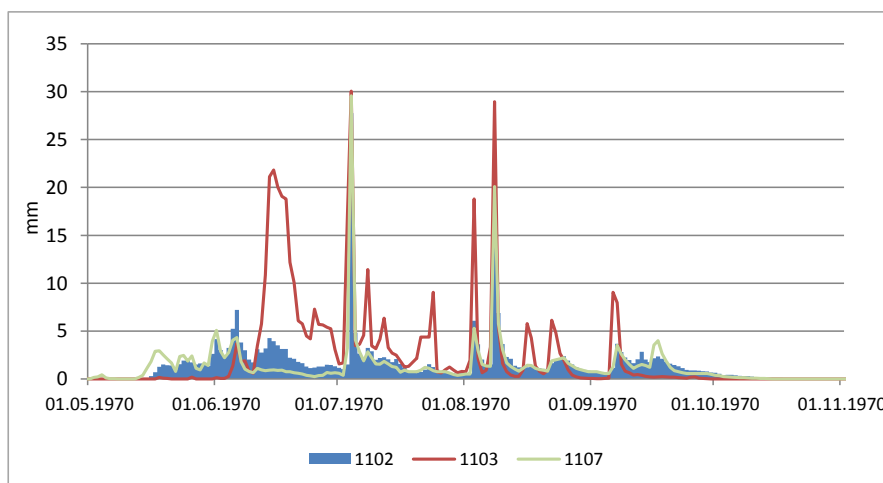
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Fig. 6 Depth of ground thawing at the different landscapes of KWBS



623
 624
 625
 626

Fig. 7 Runoff observations: A – runoff gauge at the Kontaktovy creek, 1979; B – runoff gauge at the Dozhdemerny creek, 1959; C – runoff gauge at the Yuzhniy creek, 1960; D – runoff gauge at the Vstrecha creek, 1953.



627

628 Fig. 11 Flow depth (mm) at the Kontaktovy creek – Nizhny (1102), Severny br. (1107) – south-facing
629 slope with cedar dwarf bush landscape and Morozova br. (1103) – rocky talus landscape at watershed
630 divides, 1970