



An airborne in-situ dataset of cloud microphysical properties in supercooled large droplet icing conditions

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Abstract. Detailed and comprehensive data sets on microphysical cloud properties in icing conditions are rare. In April 2023, fifteen research flights were performed with the SAFIRE ATR 42 research aircraft during the SENS4ICE-EU airborne measurement campaign over France and adjacent marine regions to measure clouds containing supercooled large droplets (SLD) at altitudes between 2 and 6 km and temperatures of 0 to -18°C . Ten cloud probes were deployed on the aircraft, comprising four imaging probes, two light-scattering probes, and three hotwire probes, in order to characterise natural SLD conditions and serve as reference instruments for novel aircraft icing detection sensors. This work presents a comprehensive cloud dataset derived from the in-situ instruments used during the campaign, which is accessible on the HALO (High Altitude and Long Range Research Aircraft) database (<https://doi.org/10.17616/R39Q0T>, Menekay et al., 2026). The dataset includes measurements of liquid and ice water content, combined particle size distributions, cloud microphysical properties, and meteorological parameters relevant to icing environments. In addition to documenting the dataset structure and processing methods, the paper provides an overview of flight strategies, instrument configurations, and statistical characteristics of the observed cloud properties, including their dependence on temperature and altitude. The dataset is suitable for studies of atmospheric icing conditions, mid-level clouds, sensor development, and model evaluation. It represents a rare collection of in-situ observations of SLD characteristics in icing environments and supports the evaluation of numerical weather prediction models under icing conditions to improve weather forecasts in hazardous conditions.

1 Introduction

Airborne in-situ measurements of cloud properties are crucial for advancing our understanding of atmospheric processes related to clouds. These observations provide valuable data for a wide range of applications, including input for climate and numerical weather prediction models, and validation of satellite retrievals. In particular, measurements of supercooled clouds are of high relevance for characterising aircraft icing environments. However, previous airborne measurement campaigns focusing on supercooled liquid droplets

are limited to confined geographical regions, seasons, or to specific scientific objectives (DiVito et al., 2019; Moser et al., 2023; Sorooshian et al., 2025). A more comprehensive understanding of cloud evolution and radiative effects requires combining observations from multiple campaigns to capture influences such as seasonal variability, geographical differences, and interannual variability. Achieving this goal relies on the availability of well-documented datasets that adhere to established data management principles (Wilkinson et al., 2016).

This article describes the measurement data of the SENS4ICE-EU flight campaign in a way that enables other researchers to understand and reuse the measurement strategy, the derived data products, and the processing steps applied, thereby facilitating comparison with other in-situ cloud datasets. The data evaluation and related uncertainties are discussed in detail, and a general overview of the measurement conditions encountered during the flights is also given.

This work begins with an introduction to the SENS4ICE project and the corresponding European flight campaign. It then outlines the measurement strategy and provides an overview of the flight activities, the general experimental setup, and the instrumentation used in this study. The applied data evaluation procedures and the structure of the resulting dataset are subsequently described, together with the associated uncertainties. Finally, selected results are presented to provide an overview of the cloud microphysical properties encountered during the campaign.

2 The SENS4ICE project

The EU-project SENSors and certifiable hybrid architectures for safer aviation in ICing Environment (SENS4ICE) (Schwarz, 2021, 2023a) aimed to develop sensor technologies that allow to detect and differentiate the icing environments specified in Appendix C (Office of the Federal Register, National Archives and Records Administration, 2014) and Appendix O (Office of the Federal Register, National Archives and Records Administration, 2016) of the European and American certification specifications for large aircraft (CS-25 and 14 CFR Part 25, both are referred to from now on as Part 25). Appendix C specifies the regulations for flights in icing conditions with median volume diameters (MVD) smaller than 50 μm . Appendix O regulates the flight in icing conditions that include SLDs, which are defined as droplets with a diameter larger than 100 μm . Appendix O was added to Part 25 in the mid-2010s and was the consequence of several fatal accidents that occurred in icing conditions where SLDs were likely present (Marwitz et al., 1997; National Transportation Safety Board, 1996, 1998). The SENS4ICE project was initiated as a consequence of the addition of Appendix O to Part 25. It pursued a two-fold approach of detecting and assessing the severity of icing conditions. On the one hand, several companies and research institutes developed and tested sensors for the direct detection of icing and SLD conditions (Pohl, 2022; Roberts et al., 2023; Gonzalez and Frövel, 2022; Schwarz, 2023b). Simultaneously, DLR and partners implemented an indirect approach to detect ice accretion based on flight parameters such as lift and drag (Deiler and Sachs, 2023). Both techniques were combined to test a hybrid ice detection system in flight (Deiler, 2024).

For the testing of the developed sensors and indirect ice detection methodology, two flight campaigns were performed,

one in the Midwestern United States, coordinated by Embraer (Schwarz, 2023a; Lucke et al., 2024) and a second campaign based out of Toulouse, France, with the French ATR 42 environmental research aircraft of the French facility for airborne research (SAFIRE). The second flight campaign is referred to as the European flight campaign and is the subject of this study.

3 European flight campaign

The aircraft used in the European flight campaign was an ATR 42-320, operated by SAFIRE. The scientific instrumentation aboard the aircraft belonged partly to SAFIRE and partly to DLR (Jurkat-Witschas et al., 2023). As part of the SENS4ICE project and prior to the airborne measurements, DLR cloud instruments were used to characterise droplet spectra and total water content in wind tunnels recently enhanced for Appendix O conditions (Lucke et al., 2022a). SAFIRE evaluates and publishes its data from the SENS4ICE campaign on the SAFIRE+ Aeris portal (Bourdon and Schwarz, 2023). These data include flight parameters such as position and airspeed, as well as radiation and humidity measurements, along with the measurements from the cloud probes owned by SAFIRE, which are later described in detail. The evaluation of all DLR instruments is based on the aircraft parameters, like velocity and position, provided by SAFIRE. A list of the aircraft parameters used in the following evaluation, apart from the latitude, longitude, and altitude, is shown in Table 1.

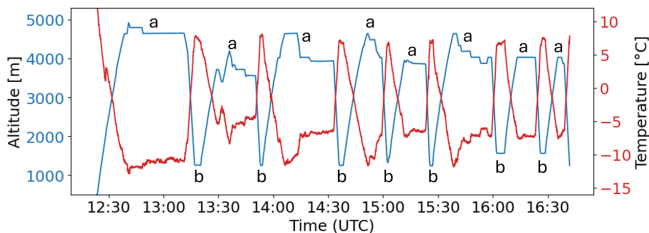
3.1 Measurement strategy

The selection of the campaign location and time was based on extensive climatological analyses of potential icing occurrences in Europe, with a particular focus on SLD formation as discussed in Jurkat-Witschas et al. (2023). In addition to the frequency of potential icing conditions, their severity was also a crucial factor for the aircraft's operational safety requirements. A minimum flight altitude of 8000 ft (2.4 km) above ground level was necessary for icing encounters, along with a warm air layer below the measurement altitude to facilitate deicing. Southern France was chosen as the primary region of interest, as active frontal systems in this area promote moist air masses and atmospheric lifting, creating favourable conditions for icing events. The influx of cleaner marine air with low CCN concentrations from the ocean further increases the likelihood of SLD formation. Conducting the campaign in spring ensured the presence of a warm air layer beneath the clouds, allowing for effective aircraft deicing.

The flights were conducted as either CER (Contrôle Essais Réception/Temporary Reserved Area) flights or flights on airways. CER flights took place in specifically designated areas reserved for the research aircraft, managed by a dedicated air traffic controller. This setup provided significant flexibil-

Table 1. Aircraft parameters provided by SAFIRE used in the evaluation.

Name	Explanation	Unit
TAS1	True air speed (TAS) from scientific pitot system	m s^{-1}
PRES	Static air pressure corrected for flow distortion	hPa
TEMP2	Air static temperature in the deiced sensor corrected from speed and recovery factor	$^{\circ}\text{C}$
TTEMP2	Total air temperature (TAT) from the deiced sensor, no correction	$^{\circ}\text{C}$

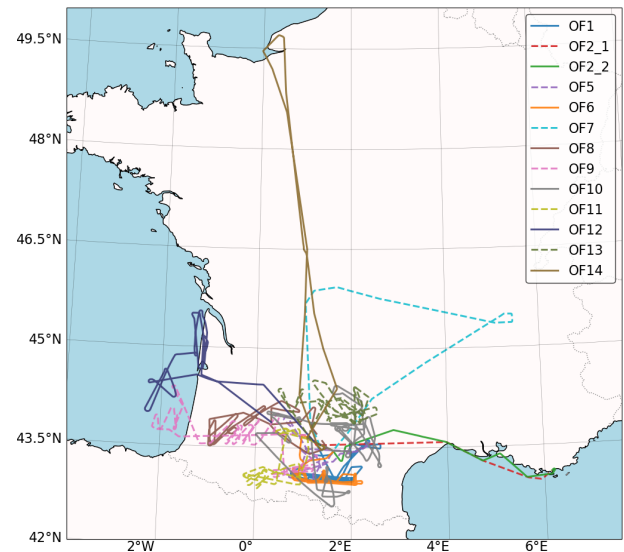
**Figure 1.** Altitude and temperature profile of the CER flight on 24 April 2023 (OF9), illustrating a typical flight strategy used during the campaign. Periods of stable altitude at temperatures below 0°C indicate cloud measurement phases (a), while segments characterised by descends accompanied by temperature increases (b) correspond to aircraft deicing cycles.

ity to adjust the flight plan in real-time. However, CER zones were limited to regions near Toulouse and along the Atlantic coast near Bordeaux. Airways flights were conducted when suitable weather conditions were not forecast for the CER areas. Unlike CER flights, Airways flights followed predefined routes with little to no flexibility for changes in flight path or altitude.

To test both indirect ice detection and accretion-based sensors, the flight included icing segments that promoted ice accretion on the aircraft, followed by warm air segments returning the aircraft to an “ice-free” configuration. The measurement strategy, therefore, involved alternating cycles of ice accretion and deicing by descending to warmer temperatures during the flight. Figure 1 illustrates an example of a flight pattern used during the campaign. Section (a) highlights cloud encounters under icing conditions, while section (b) represents the deicing phases.

3.2 Overview of flight activities

Fifteen scientific flights were conducted and tracks of the evaluated flights are shown in Fig. 2. Additionally, one electromagnetic interference testing flight and two other test flights were carried out but were not evaluated as there was not adequate scientific data collected. Aircraft and instrument issues were encountered during observational flights (OFs) 3 and 4, which were removed from the dataset. Moreover, data from the Nevzorov probe on flights from OF1 to OF8 was unreliable and is therefore not used. For these flights, water content measurements that are used for the as-

**Figure 2.** Flight tracks from the SENS4ICE-EU campaign in France. Flight numbers are shown in the legend. Ocean areas are shaded in light blue and land areas in light grey.

essment of icing conditions are obtained from the LWC-300 and Robust instruments. Table 2 provides an overview of the evaluated flights from the campaign. Determination of the air mass origin is described in Lucke et al. (2025).

4 General setup and instrumentation

A comprehensive understanding of cloud conditions relies on precise measurement instruments and the data they provide. This section focuses on the instruments used to generate the presented dataset, as well as the instrument settings applied. Table 3 presents an overview of the instrument configuration, whereas Fig. 3 presents the position of each instrument onboard the ATR-42. The horizontal and vertical wind components were obtained from the retrieved three-dimensional wind vector. The wind retrieval is based on pressure measurements from the nose-mounted five-hole radome, together with aircraft attitude and navigation data (Bony et al., 2022).

The primary instruments used for particle size, concentration, and water content measurements include the Cloud Droplet Probe (CDP), the Greyscale Cloud Imaging Probe (CIP-GS), and the Precipitation Imaging Probe (PIP). To-

Table 2. Overview of the flights conducted during the SENS4ICE-EU campaign, including flight names used by DLR and SAFIRE conventions, dates and time periods (UTC), flight type, altitude and temperature ranges during cloud measurements, and the dominant air mass origin.

Flight	DLR	SAFIRE	Date	Time (UTC)	Type	Altitude (km)	Temp. (°C)	Air mass origin
OF1	230403a	as230009	03/04/23	06:08–09:38	CER	[2.0, 3.0]	[−8.9, −3.3]	Continental
OF2_1	230404a	as230010	04/04/23	11:39–12:53	Airways	[1.3, 2.4]	[−6.6, 0]	Continental
OF2_2	230404b	as230011	04/04/23	13:12–14:30	Airways	[2.4, 2.4]	[−6.7, −4.7]	Continental
OF5	230415a	as230014	15/04/23	06:04–08:20	CER	[1.8, 2.7]	[−5.8, 0]	Maritime
OF6	230418a	as230015	18/04/23	13:56–17:05	CER	[1.8, 3.0]	[−6.3, 0]	Continental
OF7	230420a	as230016	20/04/23	10:40–13:20	Airways	[1.9, 3.3]	[−8.6, 0.1]	Continental
OF8	230422a	as230017	22/04/23	06:03–08:52	CER	[2.3, 5.1]	[−15.8, 0]	Continental
OF9	230424a	as230018	24/04/23	12:23–16:52	CER	[2.5, 4.8]	[−12.9, 0]	Maritime
OF10	230425a	as230019	25/04/23	11:04–15:45	CER	[2.6, 5.6]	[−16.1, 0]	Maritime
OF11	230426a	as230020	26/04/23	06:31–08:54	CER	[3.8, 4.8]	[−9.5, −3.8]	Maritime
OF12	230426b	as230021	26/04/23	13:34–17:08	CER	[3.1, 6.1]	[−17.7, 0]	Maritime
OF13	230426a	as230022	27/04/23	06:33–09:58	CER	[3.2, 5.0]	[−10.7, 0]	Maritime
OF14	230427b	as230023	27/04/23	12:07–15:46	Airways	[2.9, 4.4]	[−7.9, 0]	Maritime

gether, these instruments cover a wide and complementary cloud particle size range. In addition, measurements from the Nevzorov probe, the LWC-300 sensor, and the Robust hotwire Probe are included to provide reliable estimates of cloud water content. Data from the High-Speed Imager (HSI), the Backscatter Cloud Probe with Polarisation Detection (BCPD), and other probes are used as complementary observations to support the interpretation of the in-situ cloud measurements.

4.1 Cloud Droplet Probe

Two Cloud Droplet Probes (CDP) were installed on the SAFIRE ATR 42 aircraft, integrated into a Cloud Combination Probe (CCP) under the right wing, and the other was incorporated into the same canister as the CDP Robust probe under the left wing. The calibration and binning have been done according to the procedure described in Kleine (2019). The CDP is an open-path forward scattering probe for measuring individual droplets with diameters between 2 and 50 μm (Lance et al., 2010; Lance, 2012; Faber et al., 2018). It uses a laser operating at a wavelength of 658 nm. Light scattered in the forward direction by each droplet is collected within an angular range of 4 and 12°. The droplet size is determined based on the Mie theory, which establishes a relationship between the scattered light intensity and particle diameter (Mie, 1908). For ice particles, the forward-scattered intensity also depends on ice crystal shape and orientation. While the CDP can detect ice crystals, it is generally not suitable for deriving ice crystal sizes (Jang et al., 2022). Due to lower instrument sensitivity for ice crystals and their rapid growth in mixed-phase clouds driven by the Bergeron-Findeisen process (Pruppacher and Klett, 2010), particles detected within the CDP size range are assumed to be liquid. The true air-

speed measured by the SAFIRE aircraft instrumentation was used to correct the CDP number concentration measurements.

4.2 Optical Array Probes

The SAFIRE ATR 42 carried two Cloud Imaging Probes (CIP), one of which forms the second component of the CCP, the second one owned as a standalone instrument. The CIP is an optical array probe (OAP) (Knollenberg, 1970) that measures particle size and shape for particles (Baumgardner et al., 2017). It has been described in various publications, e.g. Braga et al. (2017); O’Shea et al. (2019); De La Torre Castro et al. (2023); Lucke (2024). The DLR-operated probe is a grayscale CIP (CIP-GS) with a resolution of 15 μm , and the SAFIRE-operated probe is a monoscale CIP (CIP-mono), with a resolution of 25 μm . The CIP-GS differentiates between four different degrees of shadowing (unshadowed, more than 25 % shadowed, more than 50 % shadowed, and more than 75 % shadowed). Thus, every pixel is represented by a 2-bit value. The CIP-mono, on the other hand, only differentiates between unshadowed (i.e. less than 50 % shadowed) and shadowed pixels, thus each pixel can be represented by one bit. Both probes contain Korolev tips in order to reduce the effects of ice shattering (Korolev et al., 2013) and have a 658 nm laser. The imaging threshold set in CIP-GS is 50 %, meaning at least one 50 % pixel should exist in the image to be recorded. Additionally, the probe air speed (PAS) (Weigel et al., 2016) value is set to 120 m s^{-1} to avoid wrong recording due to icing or malfunctioning of the pitot tube. This value was deliberately set higher than the actual PAS to ensure that particles could always be imaged with a sufficiently high sampling rate and thus resolution, regardless of the flight conditions. Otherwise, the probe may have

Table 3. Scientific instrumentation aboard the SAFIRE ATR 42 aircraft during the SENS4ICE-EU campaign.

Instrument	Measurement principle	Measured parameters	Range	Owner
CDP	Light scattering	Droplet number and size	2–50 μm	DLR
CDP	Light scattering	Droplet number and size	2–50 μm	SAFIRE
CIP-GS	Imaging (grayscale)	Particle number, size, and shape	15–960 μm	DLR
CIP-Mono	Imaging (monoscale)	Particle number, size, and shape	25–1600 μm	SAFIRE
PIP	Imaging (monoscale)	Particle number, size, and shape	100–6400 μm	DLR
HSI	Imaging (monochromatic)	Particle size and shape	20–2000 μm	DLR
UHSAS	Light scattering	Aerosol number and size	40–1000 nm	SAFIRE
Nevzorov	Hotwire	LWC, TWC	0.03–3 g m^{-3}	DLR
LWC-300	Hotwire	LWC	0.025–3 g m^{-3}	SAFIRE
Robust probe	Hotwire	TWC	0.025–10 g m^{-3}	SAFIRE



Figure 3. Location of cloud instruments on the ATR 42. In total, the aircraft carried 6 underwing cloud probes and 2 fuselage-mounted instruments.

missed important features of the particles if the sampling rate was lower than the flight speed.

For larger particle size ranges, the Precipitation Imaging Probe (PIP) was aboard the aircraft. The PIP measures sizes between 100 and 6400 μm with a resolution of 100 μm . Thus, it allows for measuring larger ice crystals, snowflakes and raindrops compared to the CIP. The wider arm separation of the probe also allows for sampling of higher volumes. The PIP used in the SENS4ICE-EU campaign is a monoscale probe and has been used during several previous flight campaigns (Weigel et al., 2016; Mech et al., 2022; Moser et al., 2023; Voigt et al., 2017; Jurkat-Witschas et al., 2025). The PIP also contains Korolev tips like the CIP and has a 658 nm laser.

4.3 High-Speed Imaging (HSI) Probe

The High-Speed Imaging probe developed by Artium Technologies Inc. was deployed during the campaign to address key limitations of OAPs, particularly issues related to DOF and out-of-focus particles (Esposito et al., 2019). The probe captures high-resolution 2D shadow images of cloud particles, enabling the retrieval of size distribution, thermody-

namic phase (through shape analysis), number concentration, and LWC. In contrast to the single-beam configuration of OAPs, the HSI operates using a multi-beam illumination principle with six lasers at 860 nm, each equipped with a beam expander and collimator, converging at a common focal point that defines the sample volume. The overlapping shadows of particles at this focal point are captured by a CMOS sensor, producing 1624 $\times$ 1240 pixel monochromatic images with 256 grayscale levels and a pixel resolution of 3 μm . From the recorded images, particles are selected based on predefined filter settings and identified as “blobs”, meaning they are isolated as images composed exclusively of the pixels assigned to the respective particle.

The probe was set at a frame rate of 100 Hz to prevent memory buffer overflow during high-concentration cloud encounters. Instrument operation and data evaluation were performed using the manufacturer-provided AIMS (Artium Integrated Management System) software. An adaptive threshold was selected as the image processor type, and the camera mode was used for image acquisition instead of the trigger mode, such that the camera operated continuously at the selected frame rate. This operating mode is recommended by the probe manufacturer to avoid data loss.

4.4 Hotwire Probes

During the campaign, three hot-wire probes were deployed: the Nevzorov probe, the LWC-300 probe, and the Robust probe. The LWC-300 (Droplet Measurement Technologies, 2018) is equipped with a cylindrical collector that measures liquid water content (LWC), as ice crystals are assumed to rebound from the sensor surface after impact and therefore do not significantly contribute to the measurement. In contrast, the Robust probe (Science Engineering Associates, Inc., 2026) uses a single concave collector and measures total water content (TWC).

The Nevzorov probe was equipped with an SN500 sensor head consisting of one LWC sensor, two TWC sensor cones (TWC1: 8 mm and TWC2: 12 mm), and one reference sensor to measure the dry-air term (Lucke et al., 2022b). The LWC sensor has a cylindrical geometry designed to efficiently collect droplets, while the TWC sensors feature concave cones optimised for the collection of both droplets and ice particles. The reference sensor is shielded from particle impacts and therefore measures only convective heat losses. The probe determines cloud water content from the electrical power required to compensate for heat losses caused by evaporation on the collector sensor. A detailed description of the measurement principle is given in Korolev et al. (1998b).

Lucke et al. (2022b) showed that the 12 mm TWC cone exhibits higher droplet collection efficiency for large droplets than the 8 mm cone and is therefore used as the primary sensor. For smaller droplets, the collection efficiency of the TWC sensors is corrected using information from the particle size distribution. A comparison of the Nevzorov probe specifications with the other hot-wire sensors onboard the aircraft is provided in Table 4, while an intercomparison of the measurements from the hot-wire probes is presented in Lucke et al. (2025).

5 Data evaluation

5.1 Data evaluation of hotwire instruments

The three hotwire instruments (LWC-300, Robust probe and Nevzorov probe) were evaluated with a similar procedure. The procedure was established in Lucke (2024), and is summarised here with a focus on aspects relevant for data interpretation and uncertainties.

5.1.1 Removal of convective heat losses

Hotwire instruments determine total power consumption (P_t) needed to maintain a constant sensor temperature. Outside clouds, power is mainly lost to convective cooling by dry air (P_d). Inside clouds, additional power (P_w) is required to heat and evaporate impinging droplets or ice. Thus, to derive the cloud water content (LWC and TWC), P_d must be subtracted from P_t (King et al., 1978).

The magnitude of P_d depends on flight parameters like airspeed, temperature, and pressure (Korolev et al., 1998b). For estimation, the flight data were first separated into in-cloud and out-of-cloud segments based on the variability of the measured power signal, using thresholds adapted to the sampling frequency of each instrument (Lucke, 2024). These segments are grouped into bins by flight conditions (4 m s^{-1} airspeed, 3°C temperature, 20 hPa pressure), averaged, and interpolated across three dimensions. The uncertainty of P_d is evaluated through the standard deviation within each bin. Errors in P_d significantly affect LWC and TWC estimates, especially for sensors like the LWC-300 or the Robust probe that have lower P_w/P_t ratios. A 2 % error in P_d can lead to a 7 % error in LWC for the LWC-300, and 2 %–2.5 % for TWC cones (Lucke et al., 2022a). To address unexplained jumps in P_d , particularly in the Robust probe, a local adjustment method is applied to correct for transient disturbances, such as airflow blockage from ice formation, based on rolling standard deviations and interpolation. A similar technique improves LWC-300 data, despite no ice buildup being expected.

5.1.2 Collision and capture efficiencies

In addition to uncertainties due to convective heat losses, hotwire measurements are influenced by the efficiency with which particles collide with and are retained by the sensor. Collision efficiency depends on particle size, sensor geometry, airspeed, and air properties and primarily affects small droplets (Langmuir and Blodgett, 1946; Korolev et al., 1998b; Strapp et al., 2003; Lucke et al., 2022a; Lucke, 2024; Esposito et al., 2023). Capture efficiency describes the fraction of liquid or ice mass retained by the sensor after impact and becomes relevant for large droplets, particularly under SLD conditions (Schwarzenboeck et al., 2009). Due to remaining uncertainties in published values (Strapp et al., 2003; Lucke et al., 2022a), no corrections for collision or capture efficiency were applied. The published hotwire measurements therefore represent uncorrected values, allowing informed users to apply suitable efficiency corrections if required.

5.2 Data evaluation of imaging probes

The CIP and PIP collect both 1D raw datasets and 2D shadow images of cloud particles. For a detailed evaluation of cloud microphysical properties, 2D images are primarily used. However, this image data requires correction to ensure accurate measurements. In-house code written in Python is used for processing the data (De La Torre Castro, 2024; Lucke, 2024). The cloud data evaluation procedure is described in Lucke et al. (2025) and is also presented here.

Table 4. Comparison of hotwire sensors onboard.

Parameter	LWC 300	Robust Probe	Nevzorov Probe
Sensor temperature (°C)	175	140	110
Sampling frequency (Hz)	1	10	200
Cloud differentiation, power (W), 3 s rolling average	0.1		
Std. dev, current (A), 5 s mean value		0.1	
Std. dev, voltage (V), 5 s mean value			TWC1: 0.002, TWC2: 0.0006

5.2.1 Processing of the image data

A correct representation of the image size is based on the trigger for recording of array frames, which is based on the preset particle air speed. If this deviates from the real speed of air in the sampling volume, the images will be displayed either squeezed or elongated. As explained in the previous section, the particle air speed (PAS) was fixed at 120 m s^{-1} throughout the campaign. To account for deviations from actual flight conditions and ensure accurate sampling volumes for particle concentration measurements, the true airspeed (TAS) data from the SAFIRE database were used. A correction factor was calculated as the ratio of instantaneous TAS to the fixed PAS value. This factor was then used to narrow particle images when TAS was lower than PAS, therefore preserving the accuracy of size estimations.

Once the images were corrected for PAS effects, further processing was required to filter out artefacts and non-physical particles (noise). Several filters were applied to the particle-by-particle data. Particles that were partially out of the image frame were excluded by checking whether 50 % grayscale border pixels were present, as size and phase estimation are only possible with fully imaged particles (Knollenberg, 1970). Stuck bits, which produce long linear artefacts, were filtered by detecting particles with a high aspect ratio at a single grayscale level, typically 25 % or 50 %. Particles with very low grayscale pixel counts, defined as less than 5 % at 25 % level or less than 10 % at 50 % level, were removed to eliminate noise. Extremely small particles, with a width or height of 1 to 2 pixels, were discarded because their physical properties cannot be reliably inferred. Particles with these sizes are instead measured by the CDP. Particles lacking characteristic cloud features, such as those composed entirely of 50 % grayscale pixels or having extreme aspect ratios, were also excluded. Coincident droplets in image frames are detected, separated, and merged back to the dataset. A dead-time correction (Gurganus and Lawson, 2018) was also tested, but it had only a negligible influence on the derived particle concentrations and was therefore not applied.

After the particles had been filtered and non-relevant data removed, the remaining shadow images were processed using the in-house 2D particle image analysis software described in Kirschler (2023). This software provides an extended set of image-derived particle parameters such as the size of the Poisson’s spot, the aspect ratio at each grayscale

level, and the sphericity of the particle shape. The additional parameters support a more detailed discrimination between droplets, ice particles, and image artefacts, which is particularly important for the reliable identification of SLDs in the present study. For spherical particles below $500\text{ }\mu\text{m}$, the Korolev correction (Korolev et al., 1998a) was applied to improve the sizing of droplets affected by diffraction and DOF effects. For the sizing of ice crystals, the maximum diameter enclosing the particle shadow is used. This is consistent with the diameter that is required for the parametrization of the effective diameter (ED) and ice mass according to Baker and Lawson (2006). The maximum diameter is also used in the determination of the DOF. However, this approach may lead to a slight overestimation of the DOF, since particles are not necessarily aligned with their longest axis parallel to the laser beam.

5.2.2 Selection of particle rejection mode for CIP

A crucial aspect of interpreting CIP data accurately is the appropriate selection of the depth of field (DOF), which defines the volume from which particles are properly sized. The DOF affects both the precision of the measurements and the number of particles analysed. If the DOF is too large, out-of-focus particles are included, potentially leading to errors in phase discrimination, especially for droplets where sharp edges are needed for the identification. Conversely, a very small DOF may exclude valid particles, reducing statistical robustness. The DOF is calculated based on the coefficient c from the equation in Korolev et al. (1991). Building on this, Lilie et al. (2023) proposed three particle rejection modes corresponding to different c values: Mode 1 accepts particles with at least one pixel at 50 % shadow level ($c = 8.18$), Mode 2 requires at least one pixel at 75 % level ($c = 3.68$), and Mode 3 uses a ratio of $N_{75}/N_{50} > 0.5$, with a tighter DOF ($c = 0.9$). Thus, Mode 1 has the widest DOF and includes more particles (including out-of-focus particles), while Mode 3 is the most restrictive, ensuring particles are well imaged.

Comparative analysis of the CIP modes and the Nevzorov probe’s water content measurements revealed that Modes 1, 2, and 3 yield nearly identical results for particles larger than approximately $100\text{ }\mu\text{m}$. However, Mode 1 aligns better with the CDP at smaller droplet sizes as shown in Fig. 4. Based on this, a hybrid approach was adopted: Mode 1 is used for particles smaller than $90\text{ }\mu\text{m}$, while Mode 3 is applied

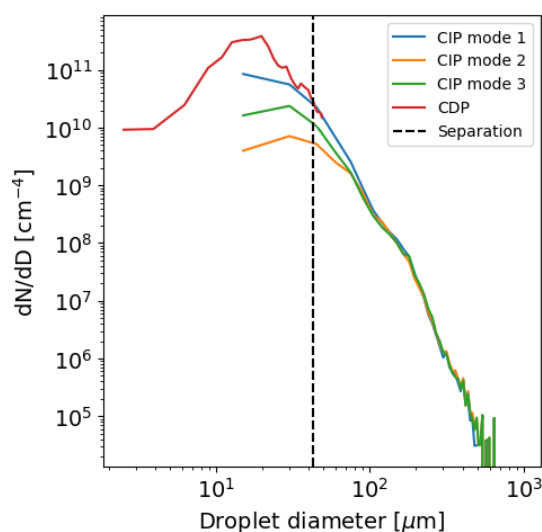


Figure 4. Normalised size distribution derived from three CIP modes and CDP from a cloud encounter. The separation line is located at 43 μm , where the threshold between CDP and CIP is defined.

to larger particles. The 90 μm cutoff aligns with the instrument's 15 μm resolution and reflects the point at which size discrimination and phase identification become more critical. Particles are assumed to be liquid below 90 μm , this threshold should be considered when compared to modelled dataset. For larger particles, Mode 3 ensures the particles are in focus and therefore suitable for detailed phase classification. This approach balances the need for high-quality measurements with statistical significance across the particle size spectrum.

5.3 Combined size distributions

A combined size distribution from the CDP, CIP, and PIP is required for a comprehensive evaluation of particle size distribution between 2 and 6400 μm . Given the overlapping measurement ranges of the probes, thresholds are required to merge the individual size distributions. The threshold between the CDP and CIP was set to 43 μm . This value was chosen because, considering the 15 μm resolution of the CIP, only particles with a minimum size of three pixels, corresponding to 45 μm , were used in the evaluation. This avoids using the lowest two CIP size bins, which have low image resolution and larger uncertainties due to the small DOF. The threshold between the CIP and PIP was set to 600 μm . This value avoids using PIP particle images with less than six pixel widths, while also limiting the use of CIP images for particles larger than 600 μm , for which the requirement of fully imaged particles reduces sampling statistics because the particle size exceeds more than half of the CIP array width. Therefore, the selected thresholds are based on the optical resolution, image quality, depth-of-field uncertainty, and sampling statistics of the individual probes.

After the probe-specific corrections, logarithmic interpolation between bin centres was applied to the combined size distribution to obtain a 1 μm size resolution, following Cober and Isaac (2012). From the resulting distributions, cumulative volume distributions were calculated to derive the cloud microphysical properties. For liquid droplets, the effective diameter is calculated as the ratio of the third to the second moment of the size distribution. This quantity is physically meaningful because droplets are assumed to be spherical, making the third moment proportional to particle volume and therefore mass. The liquid water content is obtained by integrating droplet mass over the probe sample volume using the measured number concentration and particle size information. The median volume diameter is defined as the diameter at which 50 % of the cumulative volume is reached, while the maximum droplet diameter is represented by the 99th percentile of the cumulative volume distribution (VD99) to reduce the influence of outliers.

Ice-phase microphysical properties are derived from the non-spherical particle size distributions measured by the OAPs for particles larger than 90 μm . The effective diameter is defined as the maximum diameter as explained previously. Ice water content is determined by summing particle mass derived from least squares fit with respect to the projected area, which is an area-to-mass parametrization described in Baker and Lawson (2006).

Together with the 1 Hz data, the combined data of the water content and size is averaged in sequential 15 s intervals, corresponding to a horizontal length scale of 1.8 ± 0.15 km assuming an average aircraft speed of 120 m s^{-1} . Time averaging ensures statistically robust size distributions. 15 s of averaging time is chosen because it represents a short averaging scale compared to the cloud extension and provides sufficient measurement data for statistical significance (Cober et al., 2003). Computed values represent the running average in the $-7/+7$ s intervals to provide continuous values of the parameters.

5.4 Discussion on the uncertainties

Uncertainties in particle number concentration and sizing primarily arise from counting statistics, probe sampling characteristics, and particle imaging limitations. For the CDP, scattering intensity generally increases with droplet diameter. However, the relationship is not strictly monotonic due to Mie ambiguities and exhibits oscillations, which lead to increased sizing uncertainties for droplets smaller than 20 μm (Lance et al., 2010). A comprehensive analysis of these uncertainties is provided by Faber et al. (2018). Additionally, Baumgardner et al. (2017) estimate uncertainties between 10 % and 30 % for counting and between 10 % and 50 % for sizing, with relative uncertainties generally decreasing toward larger particle sizes.

For OAPs, including the CIP and PIP, uncertainty estimates depend on particle size, phase, concentration, parti-

cle shape, image quality, and the applied processing method. Baumgardner et al. (2017) estimate uncertainty ranges between 10 % and 100 % for both particle counting and sizing, reflecting the complexity of OAP error estimation. The uncertainties reported here are therefore interpreted as representative relative uncertainties for the processed dataset, rather than as absolute errors or size-resolved uncertainty estimates. The applied corrections and the use of the particle rejection mode in the CIP processing are expected to reduce uncertainties in both sizing and concentration measurements. For droplets, sizing uncertainties are reduced by the Korolev correction, whereas larger uncertainties remain for ice particles due to their irregular shapes. Additional sizing uncertainties arise from DOF effects, pixel-resolution discretisation, and assumptions related to particle shape (Baumgardner et al., 2017; Vaillant de Guélis et al., 2019). The relative sizing uncertainty generally decreases with increasing particle size and increases for smaller particles. Combining measurements from the CDP, CIP, and PIP further reduces the influence of uncertain size ranges by excluding those with the highest uncertainty from the merged particle size distribution.

5.5 Assessment of icing environments

To identify icing conditions during the flights, the recorded in situ measurements were analysed to detect cloud encounters corresponding to Appendix O icing environments containing SLD. The classification is based on a set of criteria involving ambient temperature, liquid water content, ice crystal concentration, droplet size characteristics, and the relative contribution of large droplets to the total water content. A detailed description and justification of the applied criteria are provided by Lucke et al. (2025). In brief, the filters ensure that the detected encounters occur at sub-freezing temperatures, contain sufficient liquid water, are not dominated by ice crystals, and include a significant fraction of SLD. All parameters except the static ambient temperature (SAT) were evaluated using 15 s averages, and each encounter was required to persist for at least 5 s to avoid rapidly varying cloud conditions. Encounters that do not meet the Appendix O criteria but exhibit liquid water content exceeding 0.025 g m^{-3} at temperatures below 0°C are classified as Appendix C icing conditions.

Figure 5 summarises the flight time spent in different cloud environments for each flight, including freezing drizzle, small-droplet icing, and mixed-phase conditions. Freezing drizzle corresponds to the Appendix O icing encounters identified using the criteria mentioned above, as freezing rain was not observed during the campaign. Small-droplet icing refers to encounters at sub-freezing temperatures where droplet sizes are characterised by VD99 values below $100 \mu\text{m}$. Mixed-phase conditions are identified when liquid water is present together with elevated ice crystal concentrations exceeding 1 L^{-1} .

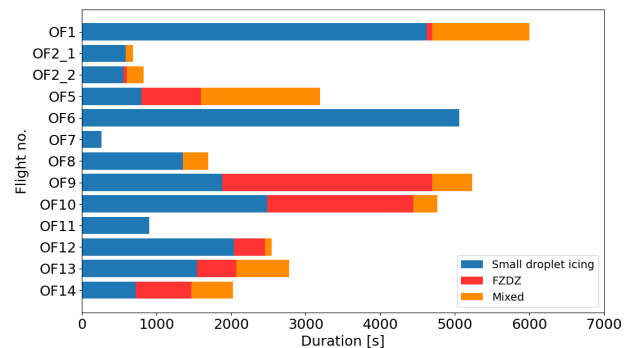


Figure 5. Duration of measurements (in s) in small droplet icing, freezing drizzle (FZDZ), and mixed-phase conditions in each flight of the campaign. Small droplet icing conditions were encountered during all flights and were sampled at the highest frequency. The figure is adapted from Lucke et al. (2024).

6 Overview of the dataset

For each flight of the SENS4ICE-EU campaign, the data are archived in the HALO Database (HALO-DB) (Menekay et al., 2026) as a set of associated datasets linked to the corresponding flight entry. These associated datasets comprise a merged cloud dataset, particle size distribution (PSD) datasets, and particle-by-particle image property data from the HSI instrument.

The primary product for each flight is the merged cloud dataset, provided in NASA Ames format and additionally in NetCDF3 format. These datasets are identified by the filename **SENS4ICE EU Cloud Dataset [flight number] [HALO-DB name]**. They contain 1 Hz in-situ measurements of cloud microphysical properties, including cloud particle number concentrations, liquid and ice water content estimates, particle size metrics (MVD and ED), and icing indicator flags. Each data record is time-stamped and accompanied by aircraft and atmospheric data such as static air temperature, true airspeed, altitude, geographic position (latitude and longitude), and three-dimensional wind components. A detailed overview of the variables included in the cloud dataset, grouped by instrument and data product, is provided in Table 5.

Cloud Particle size distribution data are provided as two separate associated datasets in NASA Ames format. The dataset **SENS4ICE-EU PSD all [flight number] [HALO-DB name]** contains total particle size distributions, including both liquid and ice particles, expressed as number concentration per unit size (dN/dD) across 118 size bins covering particle diameters from 2 to $6400 \mu\text{m}$ at a temporal resolution of 1 Hz. The dataset **SENS4ICE-EU PSD liquid [flight number] [HALO-DB name]** provides the particle size distribution of only spherical particles, thus mainly liquid.

Particle-by-particle measurements from the High-Speed Imager are provided as a separate associated dataset, **SENS4ICE-EU HSI [datetime of the start of recording]**

Table 5. Structure of the dataset and parameters. Time-averaged parameters are not shown in this table to save space.

Group	Parameter	Unit	Description
Flight and position data	Datetime	YYYYMMDD. HHMMSS	Timestamp of measurement
	SAT	°C	Static air temperature
	TAT	°C	Total Air Temperature
	TAS	m s^{-1}	True airspeed
	Altitude	m	Aircraft altitude above sea level
	Pressure	hPa	Static air pressure
	Latitude/Longitude	°	Geographic position of the aircraft
Particle number concentration	N	m^{-3}	Total particle concentration from CDP, CIP-GS, and PIP
	SLD N	m^{-3}	Concentration of SLD from CIP-GS and PIP
	LAS N	m^{-3}	Concentration of large aspherical ice particles from CIP-GS and PIP
	LIQ N	m^{-3}	Liquid particle concentration from CDP, CIP-GS, and PIP
	ICE N	m^{-3}	Ice particle concentration from CIP-GS and PIP
Particle size metrics	MVD	m	Median volume diameter from CDP, CIP-GS, and PIP for liquid particles
	CDP MVD	m	MVD from CDP only
	SLD MVD	m	MVD of SLD from CIP and PIP
	ED	m	Effective diameter from particles in the size range 2 to 6000 μm .
	ICE ED	m	Effective diameter of ice particles from particles in the size range 15 to 6000 μm .
	VD99/VD90	m	99th/90th percentile volume-based diameters in the size range 15 to 6000 μm .
	CDP VD99/VD90	m	Diameter of the 90th volume percentile from the CDP data only
	MVD filtered VD99/VD90 filtered	m m	MVD if LAS N is smaller than 1000 m^{-3} else CDP MVD VD99/VD90 if LAS N is smaller than 1000 m^{-3} else CDP VD99/VD90
Cloud water content measurements	LWC	kg m^{-3}	Liquid water content from CDP, CIP-GS, and PIP
	IWC	kg m^{-3}	Ice water content from CIP and PIP (non-liquid particles)
	SLD LWC	kg m^{-3}	LWC of SLD from CIP and PIP
	LWC 2–15...100up	kg m^{-3}	Size-binned LWC from CDP, CIP-GS, and PIP across specified diameter ranges (2–15, 15–30, 30–45, 45–60, 60–75, 75–100, and 100up μm)
	Hotwire LWC	kg m^{-3}	Liquid water content from the LWC-300 sensor
	Hotwire IWC	kg m^{-3}	Ice water content
	Hotwire TWC	kg m^{-3}	Total (liquid and ice) water content from Robust
	Nevz LWC	kg m^{-3}	Liquid water content from Nevzorov LWC sensor
	Nevz TWC1	kg m^{-3}	Total water content from 8 mm cone
	Nevz TWC2	kg m^{-3}	Total water content from 12 mm cone
	Robust TWC	kg m^{-3}	Total water content from the Robust probe
Meteorological parameters	Absolute humidity	kg m^{-3}	Absolute humidity calculated with WVSS2 volumetric concentration SAT and pressure
	Upward wind	m s^{-1}	Upward wind component
	Eastward wind	m s^{-1}	Eastward wind component
	Northward wind	m s^{-1}	Northward wind component
Flags (indicators)	Icing flag	0 or 1	1 if Hotwire LWC > 0.025 g m^{-3} and SAT < 0°C
	Appendix O flag	0 or 1	1 when Appendix O conditions are encountered, else 0
	Icing encounter flag	0 or 1	Similar to the icing flag, but smoothed to avoid very short encounters.

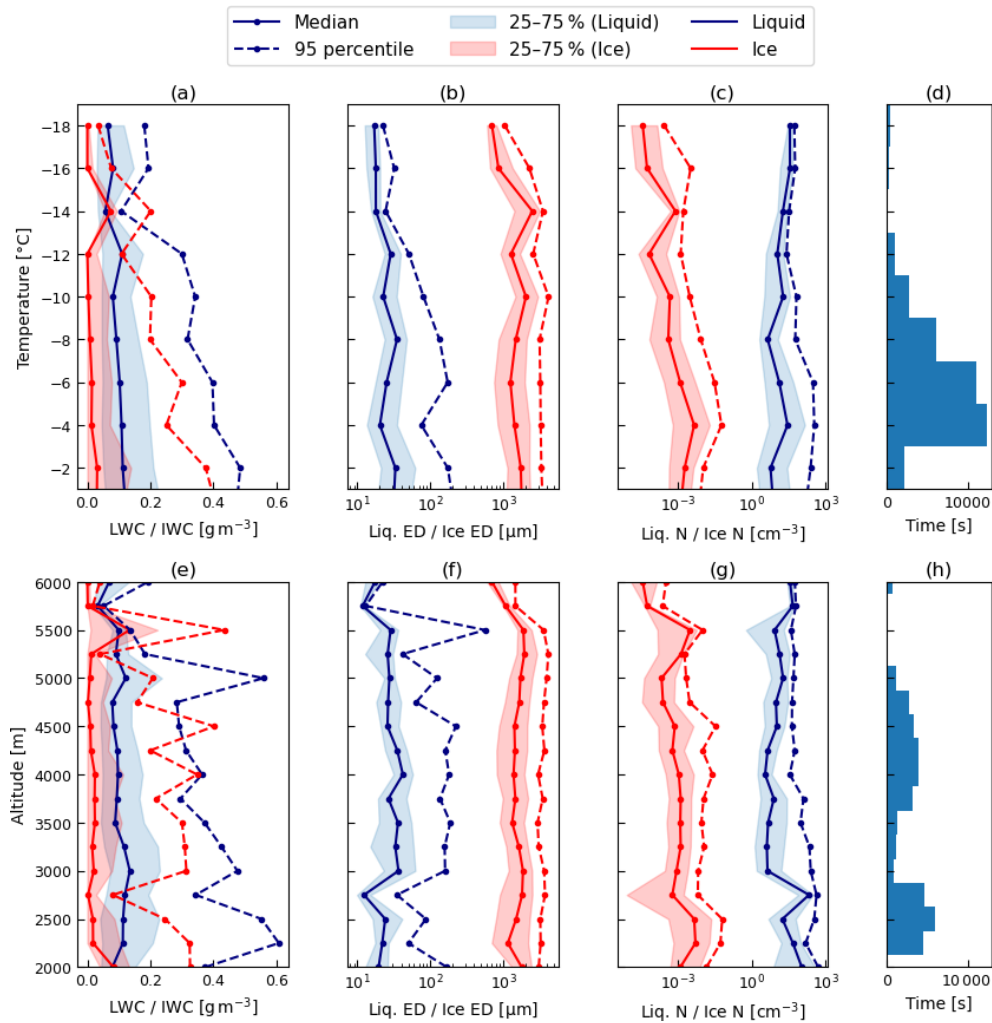


Figure 6. Distribution of (a, e) LWC in blue and IWC in red, (b, f) ED, and (c, g) N observations, sorted into 2 °C temperature and 250 m altitude bins. The histogram in (d, h) shows the total frequency of observations per temperature bin for each group. Thicker solid lines represent the median values per temperature bin, while dashed lines indicate the 95th percentile. Shaded areas highlight the interquartile range, covering the 25th and 75th percentiles of the measurements.

Valid Blobs Table]. This dataset contains all valid (in-focus) particles and includes properties such as detection time, projected area, perimeter, aspect ratio, focus quality, and particle size metrics. Only particles meeting the applied quality criteria ($\text{IntensityMean} \leq 150$, $\text{GradientMean} \geq 200$, and $\text{MeanDiameter} \geq 10 \mu\text{m}$) are included. A complete list of variables provided in the HSI dataset is given in Table 6.

Figures 6–8 provide an overview of the cloud microphysical properties observed during the SENS4ICE-EU campaign and illustrate key characteristics of the dataset. Figure 6 presents the temperature and altitude dependent distributions of cloud microphysical parameters derived from all cloud encounters during the campaign. Blue profiles represent liquid-phase properties, while red profiles indicate ice-phase measurements. Panels (a)–(c) show liquid and ice water content (LWC and IWC), effective diameter (ED), and number con-

centration (N) as functions of temperature, while panels (e–g) present the same parameters as functions of altitude. The corresponding histograms in panels (d) and (h) indicate the sampling duration within the respective temperature and altitude bins. The results show that most observations were obtained within the temperature range from approximately -2 to -12 °C, indicating that statistical distributions are most robust within this interval.

Figure 6 shows cumulative mass distributions calculated from the mean particle size distributions for different cloud thermodynamic regimes encountered during the campaign, including supercooled large droplets, supercooled small droplets (SSD), mixed-phase, and liquid-only encounters. The corresponding MVD of the mean distributions are $45.63 \mu\text{m}$ for SLD, $23.39 \mu\text{m}$ for SSD, $36.77 \mu\text{m}$ for mixed-phase, and $33.96 \mu\text{m}$ for all liquid-only encounters, illus-

Table 6. Structure of the dataset and parameters in the HSI files.

Parameter	Unit	Description
Frame	–	Frame number that the particle is pictured
Time	s	Time from the start of data acquisition moment
Area	μm^2	Area of the particle image (blob)
Perimeter	μm	Perimeter around the blob
Mean diameter	μm	Mean diameter of the blob
Circularity	–	Circularity of the blob, the value of 1 indicates a perfect circle
Mean intensity	pixel	The average intensity of all pixels within the blob, indicating the focus quality
X	pixel	The number of pixels in the longest continuous row along the x -axis of the blob
Y	pixel	The number of pixels in the longest continuous row along the y -axis of the blob
Gradient mean	–	The average quantity for the gradient around the edge of the blob, indicating the focus quality
Aspect ratio	–	The ratio of the length over the width of the blob
Intensity SD	pixel	The standard deviation of intensity for all pixels in the blob
Relative intensity	–	The local background intensity relative to the particle shadow intensity

trating the larger characteristic droplet sizes associated with SLD conditions.

Figure 8 illustrates joint distributions between selected cloud microphysical parameters derived from all cloud encounters. Panel (a) shows the relationship between effective diameter (ED) and droplet number concentration (N), while panel (b) presents liquid water content (LWC) as a function of median volume diameter (MVD). The color shading represents the number of observations within logarithmically spaced bins, highlighting the typical ranges and variability of these parameters in the dataset.

Air motions are also relevant for understanding the microphysical variability and the occurrence of SLD conditions. Figure 9 provides dynamical context for the observed cloud microphysical variability by showing the distribution of vertical wind speed derived from the five-hole radome measurements. The distribution is centred close to 0 m s^{-1} , with a median of 0.25 m s^{-1} , indicating that most cloud encounters occurred in weak vertical motion. However, the positive tail shows that some cloud regions were associated with enhanced upward motion, which can promote condensational growth and support the maintenance of supercooled liquid water. The Appendix O subset follows a similar distribution, suggesting that SLD conditions were not limited to the strongest updrafts but occurred over a range of dynamical conditions.

7 Data availability

The datasets for this study are available at the HALO database (Menekay et al., 2026, <https://halo-db.pa.op.dlr.de/mission/146>, last access: 3 August 2026) under the following DOI: <https://doi.org/10.17616/R39Q0T> (Mission: SENS4ICE-EU, mission no. 146; Menekay et al., 2026). The cloud datasets are provided in both NASA Ames and NetCDF formats, with corresponding dataset IDs #11170–#11182 and #11274–#11286, respectively.

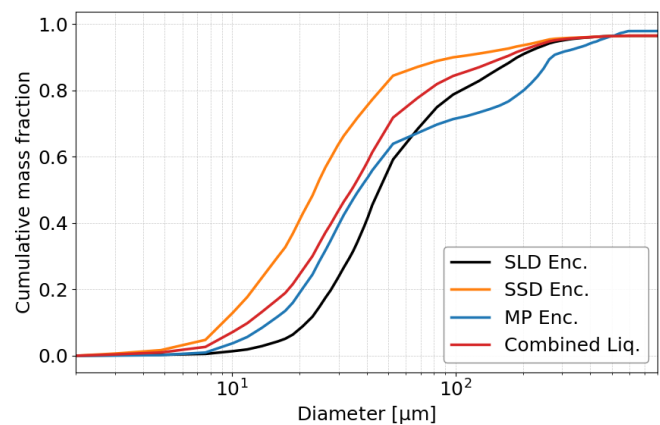


Figure 7. Cumulative mass distributions for the mean of supercooled large droplets (SLD), supercooled small droplets (SSD), mixed-phase encounters, and liquid (supercooled) encounters.

The remaining datasets are provided in NASA Ames format: HSI datasets #11183–#11195, cloud PSDs #11196–#11208, and droplet PSDs #11210–#11221 and #11288. The cloud probe image dataset is available from the corresponding author upon request. Supporting aircraft and meteorological measurements from SAFIRE are available through the SAFIRE+/AERIS data centre (Bourdon and Schwarz, 2023). These include data from the CDP (<https://doi.org/10.25326/477>, Schwarz, 2023c), UHSAS (<https://doi.org/10.25326/476>, Schwarz, 2023d), CIP (<https://doi.org/10.25326/481>), ice accretion measurements (<https://doi.org/10.25326/473>, Schwarz, 2023e), LWC measurements (<https://doi.org/10.25326/474>, Schwarz, 2023f), radiation measurements (<https://doi.org/10.25326/475>, Schwarz, 2023g), thermodynamic and dynamic core parameters (<https://doi.org/10.25326/478>, Schwarz, 2023h), and aircraft navigation and platform characteristics at 10

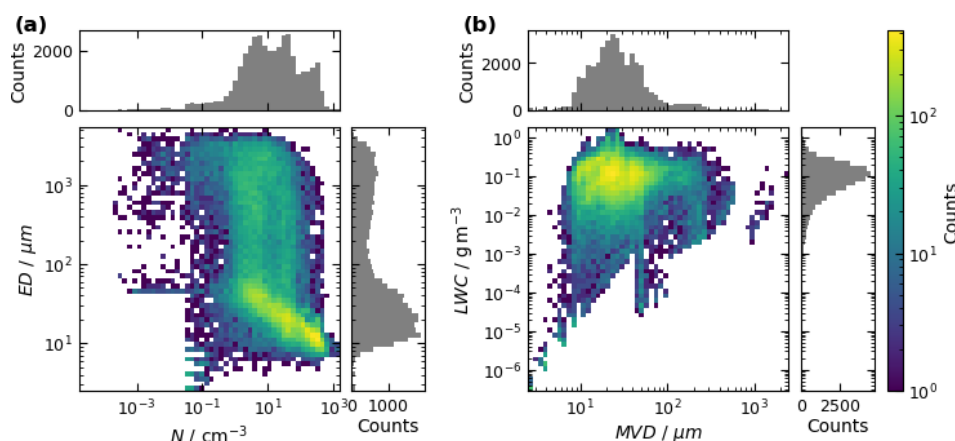


Figure 8. Joint distributions of ED versus N (a) and LWC vs MVD (b) from the cloud encounters during the campaign, color coded by their occurrence.

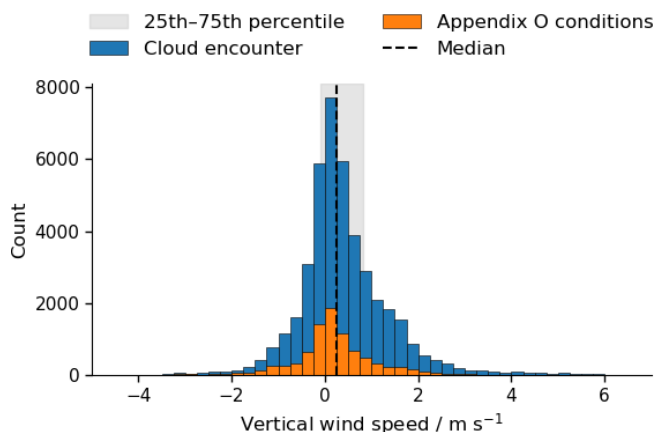


Figure 9. Distribution of vertical wind speed derived from the five-hole radome measurements during cloud encounters. Blue bars show the cloud-encounter dataset, while orange bars indicate the subset fulfilling Appendix O icing conditions.

and 1 Hz (<https://doi.org/10.25326/479>, Schwarz, 2023i; <https://doi.org/10.25326/480>, Schwarz, 2023j).

8 Summary

This paper presents a cloud microphysical dataset derived from in-situ measurements collected by in-situ cloud probes during the SENS4ICE-EU flight campaign conducted in April 2023, based in Toulouse. The instruments were primarily operated as reference sensors for the validation of aircraft icing conditions and the development of new icing detection technologies. Beyond this primary objective, the observations provide a valuable resource for a broad range of scientific applications, including weather and climate research, atmospheric cloud modelling, and improved understanding of the occurrence and characteristics of in-flight icing conditions. The dataset mainly provides properties of mid-level

icing clouds of both continental and marine origin above Europe at altitudes between 2 and 6 km and temperatures between 0 and -18°C . The resulting dataset is intended to support future investigations of aircraft icing environments and cloud microphysical processes, as well as the evaluation and improvement of numerical weather prediction and climate models, particularly with respect to the representation and prediction of SLD icing conditions.

Author contributions. TJW conceived the project, led and organised the flight campaign, coordinated the scientific activities, and supervised the data analysis and interpretation. JL prepared the cloud microphysics instrumentation for the campaign, calibrated the instruments, developed the data evaluation protocols, evaluated the hotwire probe data, and generated the combined size distributions. DM contributed to data collection during the campaign, evaluated the imaging probe data, and prepared the dataset for public release. SK supported the data evaluation procedures and assisted with data collection. CV supervised the project and provided scientific guidance and feedback on the manuscript. AB contributed to the organisation of the flight campaign, provided the aircraft platform and SAFIRE data. All authors contributed to the manuscript.

Competing interests. The contact author has declared that none of the authors has any competing interests.

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