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A dataset of ground-based vertical profile observations of aerosol, NO₂, and HCHO from the hyperspectral vertical remote sensing network in China (2019–2023)

Peiyuan Jiao¹, Chengzhi Xing², Yikai Li³, Xiangguang Ji⁴, Wei Tan², Qihua Li⁵, Haoran Liu⁵, and Cheng Liu^{1,2,6,7}

¹Department of Precision Machinery and Precision Instrumentation, University of Science and Technology of China, Hefei 230026, China

²Key Laboratory of Environmental Optics and Technology, Anhui Institute of Optics and Fine Mechanics, Hefei Institutes of Physical Science, Chinese Academy of Sciences, Hefei 230031, China ³School of Environmental Science and Optoelectronic Technology, University of Science and Technology of China, Hefei 230026, China

⁴Information Materials and Intelligent Sensing Laboratory of Anhui Province, Institute of Physical Science and Information Technology, Anhui University, Hefei 230601, China

⁵Institute of Physical Science and Information Technology, Anhui University, Hefei 230601, China ⁶Center for Excellence in Regional Atmospheric Environment, Institute of Urban Environment, Chinese Academy of Sciences, Xiamen 361021, China

⁷Key Laboratory of Precision Scientific Instrumentation of Anhui Higher Education Institutes, University of Science and Technology of China, Hefei 230026, China

Correspondence: Chengzhi Xing (xingcz@aiofm.ac.cn) and Cheng Liu (chliu81@ustc.edu.cn)

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Abstract. Vertical profile observations of atmospheric composition are crucial for understanding the generation, evolution, and transport of regional air pollution. However, existing technological limitations and costs have resulted in a scarcity of vertical profile data. This study introduces a high-time-resolution (approximately 15 min) dataset of vertical profile observations of atmospheric composition (aerosol, NO₂, and HCHO) conducted using passive remote sensing technology across 32 sites in 7 major regions of China from 2019–2023. The study meticulously documents the vertical distribution, seasonal variations, and diurnal pattern of these pollutants, revealing long-term trends in atmospheric composition across various regions of China. This dataset provides essential scientific evidence for regional environmental management and policymaking. Its sharing would facilitate the scientific community in exploring source–receptor relationships, investigating the impacts of atmospheric composition on regional and global climate and feedback mechanisms. It also holds potential for enhancing satellite retrieval methods and advancing the development of regional transport models. The dataset is available for free at Zenodo (https://doi.org/10.5281/zenodo.15211604, Jiao et al., 2024).

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1 Introduction

In recent years, the implementation of policies such as the "Action Plan for Air Pollution Prevention and Control" (2013) and the "Three-Year Action Plan for Winning the Blue Sky War" (2018) has led to significant improvements in air quality across China (Feng et al., 2019; Maji et al., 2020; Jiang et al., 2021). Concentrations of pollutants such as PM_{2.5}, NO₂, and SO₂ have substantially decreased nationwide (Zhang et al., 2020). However, O₃ pollution has shown a continuous upward trend in some regions, emerging as a new environmental challenge (Li et al., 2024; Liu et al., 2022; Fan et al., 2020). As a developing country, China still faces elevated risks of exposure to PM2.5-O3 compound pollution, with its risk levels remaining relatively high on a global scale (He et al., 2024). This exacerbates threats to the ecological environment, climate systems, and human health. Therefore, there is an urgent need to establish a comprehensive monitoring network spanning from the surface to the upper atmosphere to systematically observe and analyze PM_{2.5}, O₃, and their precursors. Such efforts will facilitate the coordinated control of PM2.5-O3 compound pollution and provide critical data to support the development of effective and scientifically grounded mitigation strategies.

Aerosol, as one of the most complex and critical compositions of the atmospheric environment, is closely associated with PM_{2.5}. It serves as both a major constituent and source of PM_{2.5}, significantly impacting regional and global climates through direct and indirect radiative effects while posing multifaceted threats to air quality and human health (Ramanathan et al., 2001; Russell and Brunekreef, 2009). NO2 and volatile organic compounds (VOCs) such as HCHO participate in atmospheric photochemical reactions, acting as key precursors for O₃ and secondary organic aerosol (SOA) (Chan et al., 2023; Liu et al., 2019). Various techniques have been employed to monitor and analyze these pollutants, including in-situ measurements, satellite observations, model simulations, and ground-based remote sensing. The China National Environmental Monitoring Center (CNEMC) has established a nationwide air quality monitoring network based on in-situ measurements, encompassing 2026 sites by 2023, to measure key pollutants such as $PM_{2.5}$, PM_{10} , NO_2 , and SO₂. However, this network lacks VOC monitoring capabilities, preventing the acquisition of data on critical precursors such as HCHO. Additionally, as CNEMC measurements are limited to surface-level concentrations, they cannot adequately characterize pollutant transport within the boundary layer or distributions in elevated atmospheric columns (Huang et al., 2018b). Similarly, the AERONET dataset provides high-temporal-resolution data on aerosol optical depth (AOD) through in-situ measurements (Holben et al., 1998) but lacks insights into the vertical profile of aerosol. NASA's MODIS satellite aerosol products offer global, long-term vertical column density (VCD) data with extensive spatial coverage but are constrained by limited temporal and spatial resolution, making continuous monitoring within specific regions challenging (Justice et al., 2002). Moreover, relying solely on VCD measurements can lead to distorted assessments of local and regional contributions due to the vertical formation, aging, and transport processes occurring above ground level. The CALIPSO satellite offers insights into aerosol vertical profiles (Huang et al., 2007) but has a relatively long revisit cycle (\sim 16 days), limiting its ability to capture processes that change dramatically over time. Chemical transport models (CTMs) and regional climate models (RCMs) provide valuable information on trace gas concentrations and their evolution through atmospheric chemistry simulations (Huang et al., 2018a). However, these models depend heavily on accurate input parameters and initial conditions. Insufficient model representation and parameterization schemes introduce significant uncertainties in characterizing vertical distribution (Huang et al., 2016). To address the shortcomings in vertical distribution monitoring, some studies have employed ground-based remote sensing techniques to obtain the atmospheric vertical profile. Lidar systems provide high-resolution vertical profile data but are limited in the number of detectable pollutants and have a restricted spatial coverage. Airborne and balloon-based platforms offer detailed data on the upper atmospheric composition but are costly to operate and unsuitable for long-term, continuous monitoring. Tower-based observations cover the vertical distribution of pollutants within a few hundred meters above the surface but are unable to capture information from higher altitudes within the boundary layer.

This study, based on the hyperspectral vertical remote sensing network, provides a high-temporal-resolution dataset of vertical profiles of atmospheric composition in China from 2019-2023. To address the lack of vertical profile data for pollutants such as aerosol, NO2, and HCHO across the country, more than 30 hyperspectral vertical remote sensing sites were deployed across 7 major regions in China (North China, East China, South China, Southwest China, Northwest China, Northeast China, and Central China). This established a comprehensive ground-based remote sensing observation network, fulfilling the need for long-term vertical monitoring on a national scale. Section 2 details the distribution of monitoring sites, the MAX-DOAS instrumentation, and the data retrieval algorithms employed. Section 3 focuses on the vertical distribution characteristics of aerosol, NO₂, and HCHO, along with their seasonal and diurnal variations. Sections 4 and 5 address the availability of the dataset and provide a summary of the findings, respectively.

2 Experimental setup

2.1 Description of the monitoring sites

This dataset originates from 32 hyperspectral vertical remote sensing sites, covering 7 major regions in China (9 sites in North China, 13 sites in East China, 3 sites in South China, 2 sites in Southwest China, 2 sites in Northwest China, 2 sites in Northeast China, and 1 site in Central China). Table 1 presents the geographical distribution and detailed information about each site. The sites span typical regions of China, including major metropolitan areas, rural areas, coastal regions, basins, and plateaus. Table S1 in the Supplement offers information on site type and major emission sources in the surrounding areas. Coverage of densely populated urban areas provides data pertinent to anthropogenic emissions, while observations from remote and high-altitude locations offer insights into background atmospheric conditions and long-distance transport processes. The diversity of sites allows for a comprehensive assessment of the vertical distribution characteristics of aerosol, NO_2 , and HCHO across China.

In this study, we set up 32 monitoring sites across 7 regions nationwide, encompassing a range of geographic environments, from coastal plains to inland plateaus. The elevations of these monitoring sites range from near sea level (2 m) to the Tibetan Plateau (3580 m), representing a variety of terrain features, climate conditions, and pollution source distributions. The diversity of these monitoring sites helps provide a complete perspective on the vertical distribution of aerosol, NO₂, and HCHO in China. In North China, monitoring sites are widely distributed, such as the Chinese Academy of Meteorological Sciences (CAMS), the Institute of Atmospheric Physics of the Chinese Academy of Sciences (IAP), and the University of the Chinese Academy of Sciences (UCAS). These sites, situated within the Beijing-Tianjin-Hebei urban cluster, are located at altitudes ranging from 10-120 m, focusing on the impact of urbanization and industrialization on atmospheric composition. Sites like Nancheng (NC) and Wangdu (WD) in suburban and rural areas represent background atmospheric conditions. Additionally, sites in Shijiazhuang (SJZ), Shanxi University (SXU), and Inner Mongolia Normal University (IMNU) extend coverage across diverse terrains such as the Taihang Mountains and Inner Mongolia Plateau, providing essential data for studying pollutant formation and transport between mountainous and plain areas. In East China, sites are primarily distributed across coastal areas and the mid-to-lower reaches of the Yangtze River, such as Xuhui in Shanghai (SH_XH), and Nanjing University of Information Science and Technology (NUIST), which are located in China's economically developed and densely populated areas. Coastal sites like Qingdao (QD), Dongying (DY), and Ningbo (NB) vary in altitude from 2-170 m, significantly influenced by ocean-land interactions. Sites at Mount Tai (TS) and Taian (TA), located at 1500 and 170 m, respectively, offer vertical distribution data across different elevations. Sites like Lin'an (LA) and Huaniao Island (HNI) expand observations into inland mountainous and offshore island areas, supporting the study of trace gas distributions in different terrains and the interface between land and sea. In South China, sites such as the Institute of Urban Environment in Xiamen (IUE), the Guangzhou Institute of Geochemistry (GIG), and Southern University of Science and Technology (SUST) span from coastal cities to inland mountains within a typical subtropical climate zone. These sites, at altitudes ranging from 30-340 m, provide valuable data on the vertical distribution of aerosol, NO2, and HCHO under tropical and subtropical conditions influenced by oceanic monsoons and typhoons. Sites in Southwest China include Shangri-La (SLS), located on the southeastern Tibetan Plateau at 3580 m, which is the highest site in the network, representing background atmospheric conditions at high altitudes and providing unique data on vertical distributions in elevated regions. Chongqing (CQ), located in the Sichuan Basin, is densely populated and industrialized, making it an ideal location to study atmospheric transport between plateaus and basins. Sites in Northwest China include Lanzhou University (LZU) and Xi'an (XA), situated on the Loess Plateau and Hexi Corridor at altitudes of 1600 and 410 m, respectively. These sites represent the typical arid and semi-arid atmospheric environment influenced by dust storms and desert-source aerosol, essential for understanding aerosol and trace gas transport and transformation in arid regions. Northeast China covers the Northeast Plain and Liaodong Peninsula, including Juhua Island (JHI) and the Liaoning University (LNU) site. JHI is located on an offshore island at an elevation of 18 m, monitoring air quality and atmospheric composition changes at the land-sea boundary. LNU is located in Shenyang at an elevation of 46 m, and Shenyang represents an area with both industrial and agricultural activities, aiding in the analysis of aerosol and pollutant profiles in the northeast. Central China includes a site in Luoyang (LY) at 100 m, situated in the middle reaches of the Yellow River, a crucial industrial and agricultural area. The LY site provides critical data on vertical distribution in the central plains, supporting studies on regional transport and accumulation of air pollutants.

The diversity in geographic location, altitude, and regional characteristics of these monitoring sites enables comprehensive monitoring of various atmospheric environments across China. This extensive coverage captures the vertical distribution features of aerosol, NO₂, and HCHO across urban, rural, coastal, mountainous, and plateau regions, forming a robust monitoring network that reflects the complex air pollution dynamics within China.

2.2 MAX-DOAS measurements

The MAX-DOAS instruments at the 32 sites operated in varying periods from 2019–2023. Each instrument comprises three essential components: two AvaSpec-ULS2048L-USB2 spectrometers (covering the UV range from 296–408 nm and the visible range from 420–565 nm), a telescope with a viewing precision of $< 0.1^{\circ}$ and a field of view of $< 0.3^{\circ}$, and a computer for controlling operations and data storage. The telescope's elevation and azimuth angles are controlled by two stepper motors, allowing measurements of atmospheric

No.	Region	Site (code)	Longitude (° E)	Latitude (° N)	Altitude (m)
1	North China	Chinese Academy of Meteorological Sciences (CAMS)	116.32	39.94	100
2		The Institute of Atmospheric Physics (IAP)	116.37	39.97	100
3		Nancheng (NC)	116.12	39.78	60
4		University of Chinese Academy of Sciences (UCAS)	116.67	40.40	120
5		Wangdu (WD)	115.15	38.17	35
6		Xianghe (XH)	116.97	39.76	10
7		Shijiazhuang (SJZ)	114.60	37.90	70
8		Shanxi University (SXU)	112.58	37.63	780
9		Inner Mongolia Normal University (IMNU)	111.68	40.80	1040
10	East China	Dongying (DY)	118.98	37.76	2
11		Qingdao (QD)	120.67	36.34	10
12		Taishan (TS)	117.10	36.25	1500
13		Tai'an(TA)	117.06	36.20	170
14		Shanghai_Xuhui (SH_XH)	121.43	31.17	3
15		Shanghai_Dianshan Lake (SH_DL)	120.97	31.09	3
16		Nanjing University of Information Science and Technology (NUIST)	118.71	32.20	73
17		Ningbo (NB)	121.89	29.75	10
18		Huaniao Island (HNI)	122.67	30.86	60
19		Lin'an (LA)	119.75	30.30	140
20		Huaibei Normal University (HNU)	116.80	33.98	35
21		Anhui University (AHU)	117.18	31.77	30
22		Changfeng (CF)	117.18	32.21	30
23	South China	Xiamen_Institute of Urban Environment (IUE)	118.05	24.61	340
24		Guangzhou Institute of Geochemistry (GIG)	113.35	23.15	30
25		Southern University of Science and Technology (SUST)	113.99	22.59	40
26	Southwest China	Shangri-La Station (SLS)	99.72	28.00	3580
27		Chongqing (CQ)	106.50	29.60	332
28	Northwest China	Lanzhou University (LZU)	103.85	36.04	1600
29		Xi'an (XA)	109.09	34.52	410
30	Northeast China	Juehua Island (JHI)	120.77	40.47	18
31		Liaoning University (LNU)	123.40	41.81	46
32	Central China	Luoyang (LY)	112.45	34.67	100

Table 1. Geography of the hyperspectral vertical remote sensing network of China. Regional division is based on China's administrative division standards.

composition at different angles. The elevation angles sequence used in this study is set to 1° , 2° , 3° , 4° , 5° , 6° , 8° , 10° , 15° , 30° , and 90° , with each cycle taking approximately 12 min, depending on the intensity of scattered sunlight. During nighttime, dark currents and offsets are recorded and later corrected in the spectral analysis. To minimize stratospheric interference, spectra collected at solar zenith angles (SZAs) greater than 75° are excluded.

For differential slant column density (DSCD) retrievals of oxygen dimer (O₄), NO₂, and HCHO, the QDOAS software developed by BIRA-IASB is utilized, with configuration settings following Xing et al. (2020, 2021b) and detailed in Table 2. In the retrieval process, spectra measured at 90° elevation serve as reference spectra, and the DSCD is calculated by difference with each elevation series spectra, enabling absorption analysis for O₄, NO₂, and HCHO. An example of the DOAS fitting results for these three compositions is shown in Figure 1. To ensure data validity, retrievals with an rms value greater than 5×10^{-4} are filtered out. To account for cloud effects, a color index (CI) threshold is computed using a fifth-order polynomial fit, excluding the lowest 10% of CI data.

Vertical profiles of aerosol and trace gases (NO₂ and HCHO) are retrieved using an Optimal Estimation Method (OEM), with the VLIDORT radiative transfer model serving as the forward model. The retrieval process minimizes the following cost function, χ^2 , to estimate the posterior state vector x:

$$\chi^{2} = (\mathbf{y} - \mathbf{F}(\mathbf{x}, \mathbf{b}))^{\mathrm{T}} \mathbf{S}_{\varepsilon}^{-1} (\mathbf{y} - \mathbf{F}(\mathbf{x}, \mathbf{b})) + (\mathbf{x} - \mathbf{x}_{\mathrm{a}})^{\mathrm{T}} \mathbf{S}_{\mathrm{a}}^{-1} (\mathbf{x} - \mathbf{x}_{\mathrm{a}}) , \qquad (1)$$

Here, $\mathbf{F}(\mathbf{x}, \mathbf{b})$ represents the forward model, describing the measurement vector \mathbf{y} (DSCDs at different elevation angles) as a function of the state vector \mathbf{x} (pollutant profiles) and actual atmospheric parameters (temperature and pressure profiles, surface albedo, and aerosol phase function). \mathbf{x}_a denotes a priori state vector, while \mathbf{S}_e and \mathbf{S}_a are the covariance matrices for \mathbf{y} and \mathbf{x}_a , respectively. The retrieval uses a two-step Gaussian–Newton (GN) approach, first deriving aerosol profiles from O₄ DSCDs and subsequently inputting them as forward model parameters to obtain NO₂ and HCHO profiles. Profile data with a degree of freedom (DOF) below 1.0, χ^2 above 200, or relative retrieval error above 50 % are excluded.

Parameter	Data source	Fitting intervals (nm)		
		O ₄	NO ₂	НСНО
Wavelength range		338-370	338-370	322.5-358
NO ₂	298 K, IO-corrected*, (Vandaele et al., 1998)	\checkmark	\checkmark	\checkmark
NO ₂	220 K, I0-corrected*, (Vandaele et al., 1998)	\checkmark	\checkmark	\checkmark
O ₃	223 K, I0-corrected*, (Serdyuchenko et al., 2014)	\checkmark	\checkmark	\checkmark
O ₃	243 K, I0-corrected*, (Serdyuchenko et al., 2014)	\checkmark	\checkmark	\checkmark
O_4	293 K, (Thalman and Volkamer, 2013)	\checkmark	\checkmark	\checkmark
НСНО	298 K, (Meller and Moortgat, 2000)	\checkmark	\checkmark	\checkmark
H ₂ O	HITEMP (Rothman et al., 2009)	\checkmark	\checkmark	×
BrO	223 K, (Fleischmann et al., 2004)	\checkmark	\checkmark	\checkmark
Ring	Calculated with QDOAS	\checkmark	\checkmark	\checkmark
Polynomial degree		Order 5	Order 5	Order 5
Intensity offset		Constant	Constant	Constant

Table 2. Detailed retrieval settings of O₄, NO₂, and HCHO.

* Solar I0 correction, Aliwell et al. (2002).



Figure 1. Examples of DOAS fits for (a) O_4 , (b) NO_2 , and (c) HCHO.

3 Vertical profile observations of atmospheric composition

Figure 2 shows the monthly data integrity for the aerosol, NO₂, and HCHO vertical profiles at each site during 2019-2023, with full continuity indicated by 100 %. The observation period varied among sites due to differing setup dates. Of the 32 sites, 81.3 % recorded over 1 year of profile data, and 59.4 % provided more than 2 years. Long-term sites, such as CAMS, UCAS, WD, and NC in North China; AHU, NUIST, and HNI in East China; SUST and GIG in South China; CQ in Southwest China; LZU in Northwest China; and LY in Central China achieved data coverage spanning 2-5 years. These sites reflect the impact of natural and anthropogenic factors on atmospheric composition over extended time scales in different regions of China. Some sites, like LNU and IUE, operated for shorter periods due to experimental scheduling and equipment availability but still captured more than 4 months of data. For regional representativeness, these sites provide references for variations in atmospheric composition under diverse geographic and climatic conditions. Among all sites, 90.6% recorded more than 6 months of continuous data, and 68.8 % achieved over 1 year of continuous observations. This continuity enhances data reliability and provides robust insights into long-term trends and patterns of aerosol and trace gases. Some sites experienced data gaps due to factors such as power or equipment failures, adverse weather conditions (e.g., rain or heavy clouds), or high retrieval error and low DOF, which occasionally led to retrieval failures. Given the heterogeneity of atmospheric composition, the vertical profile resolution was set to 0.1 km for the 0-1.0 km range and 0.2 km for the 1.0-3.0 km range (Xing et al., 2021a). As the daytime boundary layer height typically ranges from 0.1-2 km (Guo et al., 2016), we used three altitude layers of 0-100, 500-600, and 9001000 m to represent the lower, middle, and upper boundary layers, respectively. The time series of the aerosol extinction coefficient, NO₂ concentration, and HCHO concentration for the three altitude layers during 2019–2023 are shown in Figures S1–S3, with the corresponding correlation analysis provided in Section S2.

3.1 Aerosol

The box plot of the total AOD for the 32 sites during the period from 2019–2023 is shown in Figure 3. The median of the box plot represents the typical or representative AOD value at each site during the observation period, while the length of the box reflects the range of AOD variation over the observation period. AOD values in East China, South China, Southwest China, and Central China are generally higher than those in North China, Northwest China, and Northeast China, showing an overall trend of a higher AOD in the south and east, and lower in the north and west. This pattern is consistent with the findings of Filonchyk et al. (2019). The median of AOD in Central China is 0.73, the highest among all regions, followed by South China at 0.69. These values are 181 % and 165 % higher than the lowest median of AOD in Northeast China, respectively. It is important to note that Central China has only one site (LY) located in the industrial city of Luoyang, which may result in an overestimation of the representative AOD for the region. In addition, the SLS site in Southwest China is located in a high-altitude, clean environment, with a median of AOD of 0.11, which is close to 0. In contrast, the CQ site, located in the Sichuan Basin, has a median of AOD of 0.89. Sites in North China, such as XH, NC, and IAP, show relatively high upper quartile values of AOD (0.97, 0.88, and 0.93, respectively), indicating more frequent high AOD events in this region. This is likely associated with the higher level of industrialization and pollution accumulation during the winter heating season in North China (Xiao et al., 2015; Cai et al., 2020). Notably, the AOD upper quartile for the SJZ site is 1.18, significantly higher than other North China sites, reflecting the heavier aerosol load in this area. In East China, the AOD range varies from 0.39–0.94, with the lower bound attributed to the TS site, which is located on the summit of Mount Tai, far from pollution sources, in a relatively clean environment. The upper bound occurs at the SH_DL site, which is located near a lake. The higher air humidity at this site leads to hygroscopic growth of aerosol particles (Lee et al., 2008; Wang et al., 2021), resulting in consistently high AOD values, with the upper whisker reaching 3.67. The AOD range in South China is between 0.46 and 1.09, while in Northwest China, it ranges from 0.40-0.81. Northeast China exhibits the lowest AOD values, with an average range of 0.12-0.55. This may be due to the colder temperatures and lower solar radiation in Northeast China, which result in more active cold air movements and stronger winds that accelerate the diffusion of pollutants (Zhao et al., 2013).

Figure 4 shows the total averaged vertical profiles of aerosol extinction from 2019-2023. The averaged profiles for almost all sites exhibit a decreasing trend, with the highest aerosol extinction occurring near the surface, rapidly decreasing with the altitude. In North China, the near-surface aerosol extinction coefficients (AECs) at Beijing sites average 0.67, which is 16% lower than that of other cities in the region. This may be attributed to the strict environmental regulations and denser environmental monitoring in the capital. Overall, the SH_DL site shows a significantly higher surface AEC of 1.96, which can be attributed to its proximity to a lake-wetland environment, where aerosol particles condensate under high air humidity, leading to an increase in aerosol content (Wang et al., 2021). Similarly, the CQ site, located in a subtropical humid climate, experiences higher air humidity, and the basin topography restricts the diffusion of pollutants, resulting in elevated surface aerosol concentrations, with a coefficient reaching 1.48 (Yuan et al., 2023). The maximum extinction coefficient of the total averaged profile at the SUST site occurs at an altitude of 200 m above the ground, possibly due to the regional transport of pollutants from the north (Wu et al., 2022). The seasonal averaged vertical profiles of AEC are shown in Figures S4–S7. The ratios of the middle to lower boundary layer AECs in the total averaged, spring averaged, summer averaged, autumn averaged, and winter averaged profiles are 63.71%, 64.58%, 68.50%, 59.37%, and 52.76%, respectively. The upper-layer ratios decrease to 32.93 %, 33.85 %, 35.90 %, 30.45 %, and 26.08 %. Near-surface aerosol content exhibits seasonal variation. The averaged near-surface AECs in spring, summer, autumn, and winter account for 23.60 %, 24.63 %, 24.69 %, and 27.08 % of the total averaged values of the four seasons, respectively. The peak in aerosol concentration occurs in winter, which is closely linked to the largescale coal combustion emissions from heating in northern regions and the formation of thermal inversion layers in winter (Bai et al., 2022; Cai et al., 2020). In spring, sandstorms are frequent in North China and Northwest China, and the accumulation of dust aerosol in the upper atmosphere significantly increases the extinction coefficient (Gui et al., 2022; Filonchyk et al., 2018). This results in seasonal increases in aerosol concentrations at sites such as SXU, IMNU, and LZU during spring. Some southeastern coastal sites, such as SUST and GIG, exhibit lower AECs across the entire vertical profile during the summer, likely due to the frequent rainfall and monsoon influences. This reduction in aerosol concentration may be attributed to the efficient removal of aerosol particles through wet deposition (Liu et al., 2021). Some southern sites exhibit moderate pollution accumulation in winter, which may be associated with long-range transport of pollutants from the north (Xiao et al., 2015; Lee and Hills, 2003). The SLS site in the southwestern plateau is located in a remote area, far from industrial and urban pollution sources, resulting in generally low AECs. At the CQ site, aerosol concentrations peak during the winter, which may be associ-



Figure 2. Monthly data integrity for (a) aerosol, (b) NO₂, and (c) HCHO vertical profiles during 2019–2023.



Figure 3. Total AOD box plot for each site in 2019–2023, and the color of each box represents the region where the site is located.

ated with topography and adverse meteorological conditions, such as low boundary layer height, which impede the dispersion of pollutants (Wang and Cai, 2023; Zhao et al., 2019).

Figure 5 illustrates the diurnal variation of the total averaged aerosol extinction vertical profiles. High-concentration aerosols with extinction coefficients exceeding $1.0 \,\mathrm{km}^{-1}$ are primarily distributed below 600 m, while aerosols with extinction coefficients greater than 0.6 km⁻¹ are concentrated below 1000 m. At many sites, the AEC exhibits a bimodal pattern, with one peak occurring before 12:00 Beijing time (BJT) and the other between 16:00 and 18:00 BJT. The afternoon decline in aerosol concentration can be attributed to increased atmospheric mixing driven by higher temperatures, while the evening increase may result from the formation of a stable boundary layer that inhibits vertical dispersion of pollutants (Liu et al., 2024). This bimodal pattern is more pronounced at urban sites such as CAMS, IAP (located in central Beijing) and SH XH (located in central Shanghai), likely due to the significant contribution of traffic emissions during the morning and evening rush hours. In contrast, the SH_DL site, located near a lake and subject to high humidity, exhibits persistently elevated aerosol concentrations without distinct peaks. At urban sites in southeastern cities such as SH_XH (Shanghai), GIG (Guangzhou), and SUST (Shenzhen), a concentration peak around midday is observed. This can be attributed to intensified photochemical reactions driven by elevated surface temperatures, leading to secondary aerosol formation (Chang and Lee, 2007; Wang et al., 2016). The bimodal pattern is less distinct in regions such as East China and Southwest China. This difference could stem from variations in aerosol sources. In northern regions, a greater proportion of aerosol originates from sources like dust and coal combustion, which are prone to accumulation and peak formation. Conversely, finer aerosol prevalent in southeastern regions have shorter atmospheric residence times, resulting in less pronounced concentration changes (Zhang et al., 2022; Cheng et al., 2024). Most sites exhibit lower AECs between 14:00 and 16:00 BJT, likely due to the enhanced solar radiation and elevated boundary layer height during this period, which facilitate vertical dispersion of aerosol and lead to reduced extinction coefficients (Gu et al., 2020; Huang et al., 2023). The diurnal variation of vertical profiles across different seasons is shown in Figures S8-S11. During winter, the bimodal aerosol extinction pattern is most prominent, closely linked to coal-burning heating activities concentrated in the morning and evening in northern rural regions (Cai et al., 2020). The frequent occurrence of temperature inversions during these times further exacerbates the bimodal pattern by trapping aerosol from rush-hour traffic and industrial emissions near the surface (Bai et al., 2022). In summer, the maximum aerosol dispersion height is observed, with high extinction coefficients at some sites appearing in 200-500 m. This phenomenon may be associated with active atmospheric circulation facilitating regional pollutant transport during the summer months.

3.2 Nitrogen dioxide (NO₂)

Figure 6 presents the total averaged vertical profile of NO₂ from 2019-2023. NO2 concentrations exhibit distinct variations across the lower, middle, and upper atmospheric layers, with the highest levels concentrated near the surface. At most sites, NO₂ concentrations are elevated near the surface and decrease sharply with the altitude, a pattern closely linked to ground-based emission sources such as transportation and industrial activities (Liu et al., 2017; Jiang et al., 2020). However, some sites, such as SXU, show an atypical vertical profile where elevated NO₂ concentrations are observed at approximately 100 m above ground level. This anomaly may result from variations in the height of local emission sources, atmospheric temperature inversions, and local wind field effects that facilitate the accumulation of pollutants at this altitude (Bai et al., 2022; Li et al., 2023). At urban sites, such as IAP and AHU, located in the central areas of Beijing and Hefei, high NO₂ concentrations near the surface are primarily attributed to high population density, rapid industrialization, and heavy vehicular emissions (He et al., 2023; Lamsal et al., 2013). In contrast, suburban and rural sites, such as UCAS and CF near Beijing and Hefei respectively, exhibit lower NO₂ concentrations due to fewer emission sources and more favorable dispersion conditions. Specifically, the averaged near-surface NO₂ concentration at the IAP site (Beijing urban center) is 13.96 ppb, 125.16 % higher than the UCAS site (Beijing suburb), which has an averaged concentration of 6.20 ppb. Similarly, the AHU site (Hefei urban center) records an averaged near-surface NO2 concentration of 9.39 ppb, 103.69 % higher than the CF site (Hefei rural area), which averages 4.61 ppb. At remote locations, such as the SLS site in Southwest China and the TS site atop Mount Tai, where anthropogenic emission sources are minimal, and atmospheric convection in the mountainous areas facilitates vertical dispersion (Triantafyllou and Kassomenos, 2002), NO₂ concentrations are generally lower, with minimal vertical gradients. Seasonal averaged vertical profiles of NO2 are illustrated in Figures S12–S15. The ratios of the middle to lower boundary layer NO2 concentrations in the total averaged, spring averaged, summer averaged, autumn averaged, and winter averaged profiles are 33.50 %, 36.46 %, 36.85 %, 34.04 %, and 30.69 %, respectively. The proportions in the upper layer decrease to 15.92 %, 17.57 %, 18.19 %, 16.20 %, and 14.26%, respectively. The averaged near-surface NO₂ concentrations in spring, summer, autumn, and winter account for 23.06 %, 16.57 %, 25.74 %, and 34.63 % of the total averaged values of the four seasons, respectively. During the winter heating season, northern sites experience significant NO₂ emissions from coal combustion (Wang et al., 2023; Meng et al., 2018), coupled with temperature inversions that inhibit vertical diffusion, leading to a vertical profile where concentrations decline rapidly with the altitude (Bai et al., 2022; Cai et al., 2020). In summer, stronger atmospheric convection and more favorable dispersion conditions result in



Figure 4. Total averaged aerosol extinction vertical profiles during 2019–2023.



Figure 5. Diurnal variation of the total averaged aerosol extinction vertical profiles during 2019–2023.

smoother vertical profiles and lower overall NO_2 concentrations, attributable to the scavenging effects of precipitation and strong winds (Shen et al., 2021; Harkey et al., 2015).

Figure 7 displays the diurnal variation of the total averaged NO₂ vertical profiles. At most sites, peak NO₂ concentrations occur during the morning and evening rush hours, with morning peaks generally higher than evening peaks due to nighttime emissions and accumulation (Yang et al., 2019; Wen et al., 2020; Shen et al., 2023). NO₂ levels rise sharply during the morning peak from 07:00-09:00 BJT, followed by a decline in the afternoon due to enhanced photochemical reactions and vertical mixing driven by the rising boundary layer (Xie et al., 2016). In the evening, NO₂ concentrations increase again as solar radiation weakens and the boundary layer contracts (Shen et al., 2023), forming a pronounced diurnal variation. Seasonal diurnal variations of vertical profiles are shown in Figures S16-S19, with NO₂ concentrations exhibiting a consistent bimodal pattern across all four seasons, peaking at 08:00–10:00 BJT in the morning and 16:00– 18:00 BJT in the evening.

3.3 Formaldehyde (HCHO)

HCHO is one of the most significant VOCs in the atmosphere (Lei et al., 2009) and it acts as a critical precursor driving O₃ formation through atmospheric photochemical reactions (Hassan et al., 2018; Luecken et al., 2012). HCHO originates from diverse sources, including fossil fuel combustion (Ho et al., 2012; Schauer et al., 2002), biomass burning (Carlier et al., 1986; Lee et al., 1997), transportation, and industrial activities (Buzcu Guven and Olaguer, 2011). Additionally, it can form in the atmosphere through the oxidation of biogenic VOCs, further amplifying its concentration variability (Altshuller, 1993; Carter, 1995; Seinfeld et al., 1998). Figure 8 presents the total averaged vertical profiles of HCHO from 2019-2023. Similar to NO₂, the vertical profiles of HCHO at some sites exhibit an exponential decay pattern, with the highest concentrations in the lower boundary layer. At certain locations, however, HCHO peaks appear at altitudes of 100–200 m. The GIG site in South China experiences elevated HCHO concentrations, driven by high temperature and humidity levels combined with intense solar radiation, which enhance the photolysis of biogenic VOCs (Cheng et al., 2014; Ho et al., 2002). Figures S20-S23 illustrate the seasonal averaged vertical profiles of HCHO. The ratios of the middle to lower boundary layer HCHO concentrations of the total, spring, summer, autumn, and winter averaged profiles are 77.33 %, 83.32 %, 76.65 %, 74.75 %, and 69.28 %, respectively. At upper layers, the ratios decrease to 54.34 %, 59.59%, 53.23%, 52.66%, and 48.60%, respectively. The near-surface HCHO concentrations during spring, summer, autumn, and winter account for 21.52 %, 35.34 %, 24.30 %, and 18.84 % of the seasonal total averages, with peak concentrations observed in summer. In North China, low temperatures and weak solar radiation in winter limit photochemical reactions, resulting in HCHO production being dominated by primary emissions (Cerón et al., 2007). The presence of temperature inversions in winter further inhibits vertical dispersion, leading to typical vertical profiles with concentrations near the surface. In contrast, summer observes enhanced solar radiation and active photochemical reactions, driving elevated HCHO concentrations (Huang et al., 2008).

Figure 9 illustrates the diurnal variation of the total averaged HCHO vertical profiles. Starting at 08:00 BJT, HCHO concentrations increase and reach the first peak between 11:00 and 13:00 BJT. In the afternoon, the boundary layer rises and vertical mixing intensifies, leading to a decline in HCHO concentrations. From 17:00 BJT, HCHO concentrations begin to rise again, peaking at 18:00 BJT, likely due to the lowering of the boundary layer (Franco et al., 2016). At certain suburban and rural sites, high vegetation coverage leads to substantial emissions of VOCs (Cao et al., 2022), which undergo photochemical reactions to produce HCHO, resulting in its accumulation over time (Nussbaumer et al., 2021). In the evening, the weakening of HCHO photolysis reduces its consumption, combined with the decrease in the boundary layer height that limits dispersion, further enhancing the concentration peak (Biswas et al., 2020). At the SLS site, elevated HCHO levels are observed between 08:00 and 09:00 BJT. This may be attributed to strong morning solar radiation on the plateau, coupled with stable nighttime atmospheric conditions and morning inversion layers at high altitudes, which promote the accumulation of HCHO and its precursors near the surface and initiate early photochemical reactions, resulting in a concentration peak. Figures S24-S27 present the seasonal averaged diurnal variation of HCHO vertical profiles. During summer, most sites exhibit Gaussian-shaped diurnal profiles, with concentrations peaking in the high noon, driven by active photochemical reactions under intense solar radiation (Hong et al., 2021; Gao et al., 2023). In winter, some sites show elevated HCHO concentrations in the morning, likely due to the accumulation of HCHO precursors near the surface under temperature inversion conditions. Upon exposure to solar radiation, rapid photochemical reactions lead to a sharp increase in HCHO concentrations (Wei et al., 2023).

3.4 Validations

This study employs two methods to validate the dataset. The first method involves comparing surface AECs and NO₂ concentrations measured at the hyperspectral sites with ground-based data from the CNEMC stations for the period 2019–2023. We calculate the distances between each hyperspectral site and nearby CNEMC station, excluding sites based on the O₄ effective optical path thresholds (Song et al., 2023). The relevant distances are listed in Table S4. As shown in Figure 10a and b, there is a strong correlation between NO₂ concentrations from the hyperspectral sites and CNEMC measurements, with a Pearson correlation coeffi-



Figure 6. Total averaged NO₂ vertical profiles during 2019–2023.

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Figure 7. Diurnal variation of the total averaged NO₂ vertical profiles during 2019–2023.



Figure 8. Total averaged HCHO vertical profiles during 2019–2023.



Figure 9. Diurnal variation of the total averaged HCHO vertical profiles during 2019–2023.



Figure 10. Correlation analysis (a) of $PM_{2.5}$ measured by CNEMC and surface AEC (0–100 m) measured by MAX-DOAS, (b) of NO_2 concentration measured by CNEMC and MAX-DOAS, (c) of the tropospheric HCHO VCD measured by TROPOMI and MAX-DOAS, and (d) of the tropospheric NO₂ VCD measured by TROPOMI and MAX-DOAS.

cient (R) of 0.71. The correlation between PM_{2.5} concentrations and AEC is 0.53, which is influenced by the relative humidity. Figures S28 and S29 show the individual correlations of aerosol and NO2 between each hyperspectral site and CNEMC station. For sites in regions such as North China and Northwest China, where the air humidity is relatively low, the correlation between PM2.5 concentrations and AEC is generally stronger than that observed at sites in the southwest, coastal, or lakeside areas. The second validation method uses TROPOMI satellite data, comparing the VCD of HCHO and NO₂ from the hyperspectral sites with the corresponding measurements from TROPOMI. The data from the hyperspectral sites are averaged within a 30 min window (from 13:30–14:00 BJT) around the satellite overpass time. TROPOMI data are spatially averaged within a $7 \text{ km} \times 7 \text{ km}$ grid centered on the hyperspectral site, as TROPOMI has a spatial resolution of $7 \text{ km} \times 7 \text{ km}$. The linear relationships between hyperspectral sites and TROPOMI measurements are shown in Figure 10c and d. A strong correlation was found between hyperspectral sites and TROPOMI observations, with R of 0.68 and 0.60 for NO₂ and HCHO, respectively. Figures S30 and S31 show the individual correlations of HCHO and NO₂ between each site and TROPOMI.

4 Data availability

The vertical profiles of atmospheric composition over the major regions of China presented here are available in .xlsx format at Zenodo (https://doi.org/10.5281/zenodo. 15211604; Jiao et al., 2024).

5 Summary

High-temporal-resolution monitoring and analysis methods that include vertical profiles and VOC gases are currently scarce but critically important. This study provides a dataset of high-resolution vertical profiles of aerosol, NO₂, and HCHO, observed over periods ranging from 5 months to 5 years at 32 sites across China between 2019 and 2023. The dataset captures the vertical distribution, diurnal pattern, and seasonal variations of these compositions. It has been validated against TROPOMI satellite observations and groundbased CNEMC measurements, showing good correlations. The results highlight the spatial and temporal complexities of these atmospheric compositions, with the following key findings: (1) AOD levels are generally higher in East China, South China, and Southwest China compared to North China and Northwest China, exhibiting a south-high-north-low and east-high-west-low distribution. The vertical profile of AEC decreases with the altitude, with near-surface AEC showing significant seasonal variation influenced by human activities (e.g., winter heating) and meteorological conditions, peaking prominently in winter. (2) In northern regions, NO₂ concentrations exhibit a wintertime emission peak, while southern regions demonstrate stronger cleaning and dispersion capacity in summer. Diurnal patterns show significant influence from traffic emissions during the morning and evening rush hours. (3) HCHO concentrations are strongly affected by photochemical reactions, peaking during summer and around noon. In humid and warm regions such as South China and Southwest China, a secondary concentration peak around 200 m altitude is observed.

The dataset has the following limitations: (1) Under conditions such as heavy fog or extensive cloud cover, the MAX-DOAS instrument is affected by photon scattering noise and reduced spectral signal intensity, which typically increase retrieval errors. To ensure data reliability, retrievals with relative errors exceeding 50% are excluded from the dataset. (2) MAX-DOAS operates using passive sunlight scattering, which limits retrievals of aerosol, NO₂, and HCHO vertical profiles to daytime observations. To avoid interference from strong stratospheric absorption, spectral data collected when the solar zenith angle exceeds 75° are filtered out. (3) Due to hardware and algorithmic limitations, the current maximum vertical resolution is 100 m. While this represents a significant advancement, further improvements are needed in the future. (4) In regions such as Central China and Southwest China, the number of monitoring sites is relatively limited. This uneven distribution may potentially affect the representativeness of the results for these areas. Despite these limitations, the dataset provides higher temporal resolution and broader temporal coverage, addressing critical gaps in vertical profiling. It helps mitigate the underestimation of nearsurface concentrations in satellite retrievals and reduces uncertainties in the vertical direction for CTMs and RCMs. Potential applications of this dataset include the following: (1) As a valuable complement to national monitoring efforts, the dataset offers detailed regional atmospheric composition data, providing scientific support for governmental and research institutions in air pollution control and policymaking. (2) By leveraging high-resolution data, the dataset facilitates source apportionment analyses for aerosol, NO₂, and HCHO, revealing inter-regional interactions and pollutant transport pathways. (3) The long-term monitoring data can inform RCMs, improving the assessment of climate change impacts and enabling more robust evaluations of mitigation and adaptation strategies.

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