



Supplement of

High-resolution physicochemical dataset of atmospheric aerosols over the Tibetan Plateau and its surroundings

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7 **S1. HR-ToF-AMS operation**

8 The HR-ToF-AMS is one of the most advanced instruments that widely used for
9 the study of atmospheric aerosol chemistry worldwide. The detailed principle of HR-
10 ToF-AMS can be obtained elsewhere (DeCarlo et al., 2006). The HR-ToF-AMS is
11 mainly composed by three different sections that separated by small apertures and
12 differentially pumped, i.e., a particle beam generation section to form a concentrated
13 and narrow particle beam through a critical orifice and a six-stage aerodynamic lens, a
14 particle-sizing chamber to measure the particle aerodynamic sizes through a particle
15 time-of-flight measurement (different velocities and arrival times for size-dependent
16 particles in a known flight distance), and a particle chemical composition detection
17 section to directly vaporize the particle beam at a ~ 600 °C resistively heated surface,
18 ionize the particles into positively charged ion fragments by a 70 eV electron impact,
19 and then detect their chemical composition by a high-resolution mass spectrometer
20 (Jimenez et al., 2003).

21 There are two different operation modes in HR-ToF-AMS, i.e., V-mode (detection
22 limits of about 10 ng m^{-3}) and W-mode ($\sim 5000 \text{ m}/\Delta\text{m}$) with different signal-to-noise
23 ratio (S/N). However, the HR-ToF-AMSs were only operated at the V-mode during
24 almost all the seven field campaigns in consideration of the relatively low aerosol mass
25 concentration level and low S/N ratio over the TP. The mass concentration and size
26 distribution of non-refractory PM_{10} chemical species were obtained by further switching
27 the instrument between mass spectrum (MS) mode and particle time-of-flight (PToF)
28 mode every 15s under the V-mode operation. However, there are no observation of
29 particle sizes during the NamCo and LHG measurements due to the malfunction of the
30 chopper. In addition, the HR-ToF-AMS need to be calibrated for its flow, ionization
31 efficiency (IE), and sizes at the beginning and end of each observation (Jayne et al.,
32 2000). The relative IE (RIE) of ammonium and sulfate were calibrated using the mono-
33 dispersed pure ammonium nitrate and ammonium sulfate particles, respectively, with
34 the selected sizes of 200–300 nm, while the particle size was calibrated using the mono-
35 dispersed ammonium nitrate particles with sizes varied from 60 to 600 nm. Finally,

36 default RIE values were assumed to be 1.1, 1.3, and 1.4 for nitrate, chloride, and OA,
37 respectively, during all the field campaigns, while different RIE values were set for
38 ammonium and sulfate according to their calibration results during each campaign, e.g.,
39 3.9 and 4.2 for ammonium and 1.6 and 1.4 for sulfate based on two calibrations in the
40 QOMS measurement.

41 **S2. HR-ToF-AMS data processing**

42 The HR-ToF-AMS data was processed using the standard data analysis software
43 with SQUIRREL and PIKA toolkits written in Igor Pro (Wavemetrics Inc., Lake
44 Oswego, OR, USA). The SQUIRREL used a fragmentation table to apportion the
45 measured signals at each mass-to-charge ratio (m/z) into different species to quantify
46 the chemical composition of non-refractory PM_{10} species, while the PIKA employed a
47 modified Gaussian fitting algorithm to obtain the ion-speciated high-resolution mass
48 spectra (HRMS) and elemental composition of OA (Allan et al., 2004; DeCarlo et al.,
49 2006). The elemental ratios of OA, i.e., oxygen-to-carbon (O/C), hydrogen-to-carbon
50 (H/C), organic matter-to-organic carbon (OM/OC), and nitrogen-to-carbon (N/C), were
51 determined using the improved method (Canagaratna et al., 2015) during all the seven
52 observation campaigns.

53 In addition, a collection efficiency (CE) was generally introduced to compensate
54 for the incomplete transmission and detection of particles through the aerodynamic lens
55 and bouncing at the vaporizer surface in most AMS studies. Previous study has revealed
56 that the CE is significantly influenced by the relative humidity (RH) in sampling line
57 and the acidity and ammonium nitrate mass fraction (ANMF) in the sampled aerosols,
58 which has been concluded as a build-in composition-dependent CE (CDCE) algorithm
59 in the standard data processing software (Middlebrook et al., 2012). Generally, a high
60 RH, a high aerosol acidity, or a high ANMF often corresponds to a high CE value.
61 However, the RH in the sampling system is always maintained below 40% due to the
62 professional deployments of dryers in the front of the sampling system and the ANMF
63 is basically below 0.4 due to the low contributions of nitrate and ammonium during all
64 the seven observation campaigns, which means the negligible effects of these two

65 parameters on CE in our study. Therefore, default CE value of 0.5 were finally
66 employed during the QOMS, NamCo, Ngari, Waliguan, and Lhasa campaigns in
67 consideration of their overall neutralized or slightly acidic aerosols, whereas the CDCE
68 values were adopted at Motuo, LHG, and Bayanbulak where bulk submicron aerosols
69 were acidic.

70 **S3. OA source apportionment using PMF analysis**

71 Source apportionment of OA during each observation was conducted by the
72 positive matrix factorization (PMF) analysis on organic matrix data using the PMF2.exe
73 algorithm in robust mode (Paatero and Tapper, 1994) and the standard PMF Evaluation
74 Tool (PET, Ulbrich et al., 2009) written in Igor Pro software.

75 The PMF analysis was evaluated thoroughly according to the standard procedures
76 outlined in Zhang et al. (2011) by down-weighting, modifying, or removing some ion
77 fragments in the data and error matrices. Firstly, those ions at $m/z > 120$ and all the
78 isotope ions were generally excluded because of the insufficient ability to resolve the
79 deconvolution due to their low signals. Then, the signals of the four organic ions of O^+ ,
80 HO^+ , H_2O^+ , and CO^+ were scaled to that of CO_2^+ according to the suggested
81 fragmentation table in Aiken et al. (2008) and further down-weighted in PMF analysis.
82 Thirdly, all those “bad” ions ($S/N < 0.2$) were removed from the data matrices, while all
83 the “weak” ions ($0.2 < S/N < 2$) were downweighted by increasing their errors. In
84 addition, some runs and some ions which had obviously huge residual spikes were also
85 removed in order to avoid their unnecessary interference. After the above pre-
86 processing, the PMF solutions were investigated by selecting a certain variation range
87 of factor number and rotational parameter ($fPeak$), e.g., 1–6 factors with $fPeak$ varying
88 from -1 to 1 . Finally, the optimal solution of PMF analysis were determined after a
89 comprehensive evaluation by examining the model residuals at each m/z and each time,
90 comparing the factor mass spectrum with corresponding reference spectrum, comparing
91 the temporal variation of individual factor with external tracers, and analyzing the
92 diurnal variation pattern of each factor. Totally, 2-, 3-, or 4-factor solution were selected
93 during the different field campaigns in this study. The specific high-resolution mass

94 spectrum of each OA factor identified among the eight different field campaigns are
95 shown in Figure S4.

96 **S4. Operation and data processing of other instruments**

97 **1) SMPS**

98 The scanning mobility particle sizer (SMPS) developed by the TSI Inc. is
99 composed by an electrostatic classifier (EC, model 3080) equipped with a long-
100 differential mobility analyzer (long-DMA, model 3081) and a condensed particle
101 counter (CPC, model 3772). Ambient particles are first screened by a particle impactor
102 installed at the front of the DMA and large particles are removed. Ambient particles are
103 measured through an electrical mobility detection technique in this instrument, e.g., a
104 bipolar charger in the EC is utilized to charge the particles to a known charge
105 distribution, then classify them according to their ability to traverse an electrical field
106 in the long-DMA, and finally count those screened monodisperse particles by the CPC.
107 The sample and sheath flow rates are 0.3 and 3.0 L min⁻¹, respectively, at both QOMS
108 and Lhasa which measure particles between 14.6 and 661.2 nm in mobility diameter
109 (D_m), whereas the sample and sheath flow rates are 0.5 and 5.0 L min⁻¹ at LHG and
110 Motuo and sample particles at a size range of 10.9–495.8 nm in D_m . The number
111 concentrations of submicron particles in 107 different size channels are firstly recorded
112 at an initial time resolution of 5 min and then converted to the total number and volume
113 concentrations according to the obtained size distribution of number concentration.

114 **2) PAX**

115 The PAX directly measures the B_{abs} and B_{scat} of aerosol particles at 405 nm by
116 using a modulated diode laser, namely measures the B_{abs} by an in-situ photoacoustic
117 technique while the B_{scat} using a wide-angle integrating reciprocal nephelometer. The
118 B_{ext} is the sum of B_{abs} and B_{scat} while the single scattering albedo (SSA) is calculated as
119 the ratio of B_{scat} to B_{ext} . The BC mass concentration is calculated as the ratio of measured
120 B_{abs} to a fixed BC mass absorption cross-section (MAC) value of 10.19 m²g⁻¹ at 405
121 nm. In addition, the B_{scat} is calibrated using the high-concentration ammonium sulfate
122 particles generated by the aerosol generator, while the B_{abs} is calibrated using the

123 sufficient black smoke from a kerosene lamp before each field campaign according to
124 the operator manual of this instrument.

125 **3) Aethalometer**

126 The Aethalometer (models AE31 or AE33) is used to measure the particle B_{abs} at
127 seven wavelengths, which firstly measures the light attenuation between particle-laden
128 and particle-free sample spots on the filter and finally converts the attenuation to
129 particle B_{abs} in ambient air. Both AE31 and AE33 have seven bands, namely 370, 470,
130 520, 590, 660, 880 and 950 nm, and the concentration of black carbon is mainly
131 measured according to the absorption coefficient at 880 nm. The filter-based loading
132 effect and multiple scattering effect are corrected during all the three observations to
133 eliminate the difference between the light attenuation measured at the filter and the
134 ambient particle B_{abs} .

135 The absorption Ångström exponents (AAE) value is acquired through a power-
136 law fitting of B_{abs} following the typical Beer-Lambert's law, i.e., $AAE =$
137 $\ln(B_{abs,\lambda_1}/B_{abs,\lambda_2})/\ln(\lambda_2/\lambda_1)$. Furthermore, a traditional AAE method was adopted to
138 quantitatively apportion the total B_{abs} into two parts from BC and BrC ($B_{abs,BC}$ and
139 $B_{abs,BrC}$) at 370–660 nm during each campaign. The contribution of BrC to total B_{abs}
140 ($fB_{abs,BrC}$) at a short wavelength λ is calculated as $fB_{abs,BrC,\lambda} = 1 - (B_{abs,880}/B_{abs,\lambda}) \times$
141 $(\lambda/880)^{-AAE_{BC}}$ by assuming its negligible contribution at 880 nm. Detailed information
142 about the data correction and calculation of this instrument can be found in our previous
143 publication (Zhang et al., 2021b).

144 **4) CCN-100**

145 The CCN-100 measures aerosol particles called cloud condensation nuclei that can
146 form into cloud droplets. The instrument supersaturates the sampled aerosol particles
147 in a 50-cm-high column with continuously wetted walls and a longitudinal thermal
148 gradient, so that those particles grow into detectable CCN particles and are measured
149 using an optical particle counter among 20 different size bins. The principle of CCN
150 counters is that diffusion of heat in ambient air is slower than that of water vapor, which
151 diffuses from the warm, moist column walls to the centerline faster than heat in the

152 column. Detail information may refer to Roberts and Nenes (2005). The number
153 concentrations of CCN are measured consecutively at five different *SS* values of 0.2%,
154 0.4%, 0.6%, 0.8%, and 1.0%. The CCN data is recorded every 5 minutes at each *SS* and
155 finally has a time resolution of 30 minutes during a complete measurement cycle.

156 **S5. Calculation and evaluation of the bulk acidity of submicron aerosols**

157 Bulk acidity of submicron aerosols from AMS measurement was generally
158 evaluated following the methods in Zhang et al. (2007) and Schueneman et al. (2021).
159 The mass concentration of ammonium was firstly predicted by assuming to fully
160 neutralize these AMS measured sulfate, nitrate, and chloride using the following
161 equation (S1):

$$162 \quad \text{NH}_4^+_{\text{Predicted}} = 18 \times (2 \times \text{SO}_4^{2-}/96 + \text{NO}_3^-/62 + \text{Cl}^-/35.5) \quad (\text{S1})$$

163 Then, the mass concentration ratio of measured ammonium to predicted
164 ammonium ($\text{NH}_4^+_{\text{Measured}}/\text{NH}_4^+_{\text{Predicted}}$) was further calculated to be a good indicator to
165 evaluate the bulk acidity of submicron aerosols. In this study, linear regression analysis
166 between the mass concentrations of measured and predicted ammonium was performed
167 to evaluate the bulk acidity of submicron aerosols in the different TP regions (Fig. 3).
168 Aerosol particles are generally considered to be “acidic” if the calculated ratio is
169 obviously lower than 1 and to be “more acidic” if the ratio is lower than 0.75, whereas
170 a ratio that roughly near to 1 or larger than 1 indicates the particles are “bulk neutralized”
171 and even there are more excess ammonium that needed to be neutralized. Note that the
172 validity of using this method is based on the assumption that the influence from
173 nitrogen- or sulfur-containing organic ions (e.g., organic acids and organic nitrogen
174 compounds) as well as the mineral and metal ions are negligible (Zhang et al., 2007).

175 **S6. Estimation of aerosol DRFs using the SBDART and OPAC models.**

176 The aerosol direct radiative forcing (DRF) was modelled by the widely used Santa
177 Barbara DISORT (Discrete Ordinate Radiative Transfer) Atmospheric Radiative
178 Transfer (SBDART) model in the shortwave spectral range of 0.25–4.0 μm . SBDART
179 is a software tool that computes the plane-parallel radiative transfer under both clear

180 and cloudy conditions (Ricchiazzi et al., 1998). Specific simulation of aerosol DRF in
 181 the atmosphere (ATM) can be described by the following equations (2) and (3). In brief,
 182 the net fluxes (ΔF , difference between the downward and upward radiation fluxes) with
 183 and without the investigated variable were calculated twice in this model under cloud-
 184 free conditions at both the earth's surface (SUR) and the top of the atmosphere (TOA).
 185 The differences of net fluxes between the two simulations were then considered as the
 186 DRFs of the specific investigated variable at the SUR and TOA. Finally, the DRF in
 187 the ATM (DRF^{ATM}) was obtained using the DRF at TOA (DRF^{TOA}) minus the DRF at
 188 SUR (DRF^{SUR}). The details of the model description can be found in previous studies
 189 (Xin et al., 2016; Gong et al., 2017).

$$190 \quad \Delta F = F^{\text{downward}} - F^{\text{upward}} \quad (\text{S2})$$

$$191 \quad \text{DRF}^{\text{ATM}} = \text{DRF}^{\text{TOA}} - \text{DRF}^{\text{SUR}} = (\Delta F_{\text{with}}^{\text{TOA}} - \Delta F_{\text{without}}^{\text{TOA}}) - (\Delta F_{\text{with}}^{\text{SUR}} - \Delta F_{\text{without}}^{\text{SUR}}) \quad (\text{S3})$$

192 The principle and specific steps for the estimation of aerosol DRFs using the
 193 SBDART mode can be described briefly as follows:

194 Firstly, the mass concentrations of organic carbon (OC), BC, and water-soluble
 195 ions (WSIs), which were measured from the corresponding off-line filter samples with
 196 one- or two-days time resolutions during each campaign, were the initial measured
 197 parameters used for the estimation of aerosol DRFs.

198 Then, the particle numbers of each species in per cubic meter air (denoted as ρ_n)
 199 were estimated using the above measured mass concentrations of each species divided
 200 by the referred M values in the OPAC model (Hess et al., 1998), which represented the
 201 aerosol mass in per cubic meter air and integrated over the size distribution and
 202 normalized to 1 particle per cubic centimeter of air. Specifically, values of 5.99E-5 was
 203 used for soot/BC particles and 1.34E-3 were used for OC and WSIs.

204 Next, four crucial input optical parameters of aerosol optical depth (AOD), single
 205 scattering albedo (SSA), Ångström exponent (AE), and asymmetry factor (ASY), and
 206 the light absorption and scattering coefficients were estimated for OC, BC, WSIs and
 207 total particles, respectively, using the Optical Properties of Aerosol and Cloud (OPAC)

208 model. Detailed introduction of OPAC model can be found in Hess et al. (1998). Noting
209 that the above estimation in OPAC model were performed six times during each
210 campaign by setting the ρ_n as the above calculated value multiplier 1 to 6, respectively.

211 Then, the entire six datasets of modelled total light absorption and scattering
212 coefficients were further used to evaluate the performance of OPAC model by
213 comparing them with those corresponding measured light absorption and scattering
214 coefficients from PAX and Aethalometer (Figures S6-10). Finally, the optimal ρ_n
215 value was selected when the modelled total light absorption and scattering coefficients
216 were comparable.

217 The performance of the OPAC model was then evaluated and tuned by comparing
218 the modelled total light scattering and absorption coefficients with those
219 correspondingly measured values from online Aethalometer and PAX measurements
220 during the three campaigns, as shown in Figures S5-10. The consistent variation trends
221 were found between them with the coefficient of determination (R^2) varying between
222 0.69 and 0.99. The slightly lower modelled values compared with those measured
223 values was mainly attributed to their inconsistent wavelengths, i.e., the modelled light
224 scattering and absorption coefficient at 550 nm in the OPAC model, whereas the
225 measured light scattering coefficients at 405 nm for PAX and the light absorption
226 coefficients at 520 nm for Aethalometer. Overall, the small difference between the
227 modelled and measured values indicated the reasonable simulation of aerosol optical
228 parameters (e.g., AOD, AE, SSA, and ASY) in the OPAC model in this study. Finally,
229 these four optical parameters belonging to each species were all used in the SBDART
230 model for the simulation of DRFs caused by OC, BC, and WSIs.

231 After the above evaluation, the finally obtained four input parameters of AOD,
232 SSA, AE and ASY belong to each species (OC, BC, WSIs, and total particles) were
233 inputted into the SBDART model to simulation the DRFs of each species, respectively.

234 Table S1. Summary of measurements on ambient aerosols in the TP by different off-line filter sampling and on-line instruments.

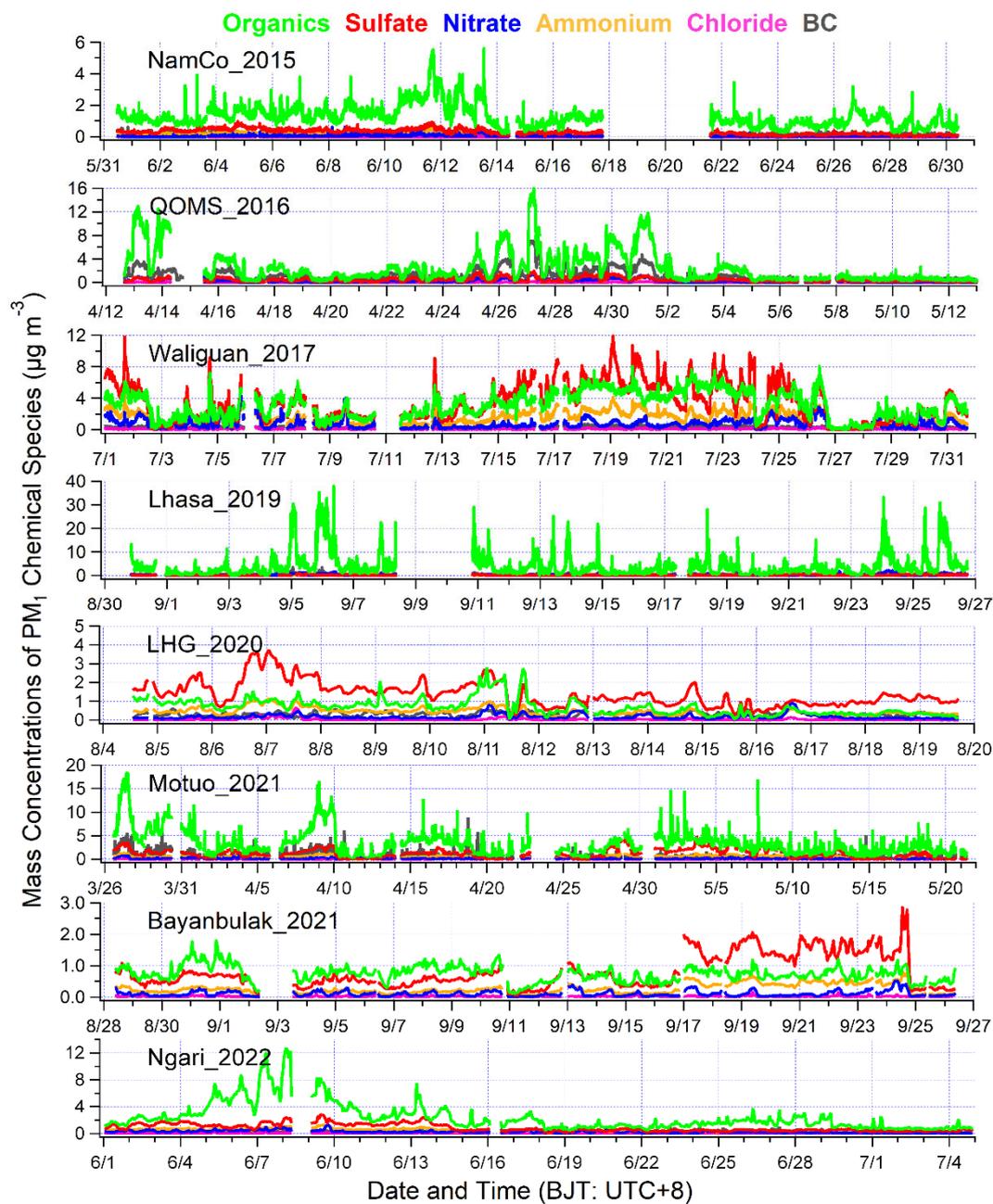
	Measurement content	Period	Site	Instrument/Method	Resolutions	References	
	Off-line filter sample analysis for aerosol concentration and chemical composition: CAs (BC/EC/OC/WSOC/BrC), WSIs, element, etc.						
1.	BC and OC dataset over the Third Pole	Different periods during different campaigns	19 stations from APCC network	TSP samples	off-line, 24-48 h	Kang et al. (2022)	
2.	CAs characteristics on the Third Pole			TSP samples	off-line, 24-48 h	Chen et al. (2019a)	
3.	CAs in the TP: An investigative review			Lhasa, Lulang, Nam Co, QOMS, etc.	PM _{2.5} and TSP samples	off-line	Li et al. (2021a)
4.	BC sources reaching the TP			7 stations across the Himalayas	TSP samples	off-line, 24-48 h	Li et al. (2016a)
5.	Concentrations and light absorption of CAs	2013/5-2014/3	Lhasa, urban city	PM ₁₀ and PM _{2.5} samples	off-line, 12-24h/week	Li et al. (2016b)	
6.	CAs (OC and BC)	2006/7-2009/12	Nam Co, inland TP	TSP samples	off-line, 4-14 days	Zhao et al. (2013a)	
7.	Mass, elements, CAs, and WSIs	2008/7/16-2009/7/26	Lulang, southeastern TP	TSP samples	off-line, 6 days	Zhao et al. (2013b)	
8.	Concentrations, seasonality and sources of CAs	2009/8-2010/7	QOMS, southern TP	TSP samples	off-line, 24 h/week	Cong et al. (2015)	
9.	Size distribution of CAs	2012	Nam Co, inland TP	Size-segregated samples	off-line, 72 h	Wan et al. (2015)	
10.	Size-segregated characteristics of CAs	2014/8-9, 11-12	Lhasa, urban city	Six-stage cascade impactors	off-line	Wei et al. (2019a)	
11.	Near surface PM _{2.5} concentration	2016/12/20-12/26	Lhasa, urban city	PM _{2.5} samples	off-line, 12 h	Li et al. (2019)	
12.	BC concentrations	2003/12/5-2006/2/17	Muztagh Ata Mountain, northern TP	TSP samples	off-line, 1 week	Zhou et al. (2018)	
13.	Elemental concentrations	2005/11-2007/11	Nam Co Station, inland TP	TSP samples	off-line, weekly	Kang et al. (2016)	
14.	Elemental composition	2019/12-2020/11	Yaze, southeastern TP	TSP samples	off-line, 15 days	Xu et al. (2022b)	
15.	Elements and WSIs	2009/3-5, 7-8, 11-12	Lijiang, southeastern TP	TSP samples	off-line, 1-2 days	Zhang et al. (2016)	
16.	Characteristics of WSIs	2010/7/16-2011/7/28	Qilian Shan, northeastern TP	PM _{2.5} samples	off-line, weekly	Xu et al. (2014)	
17.	Seasonal variations and size distributions of WSIs	2012/10-2013/9	Shigatse, southern TP	Eight-stage impactor	off-line, weekly	Yang et al. (2016)	
18.	Characteristics of size distributions and sources of WSIs	2014/8-9, 11-12	Lhasa, urban city	Six-stage cascade impactors	off-line	Wei et al. (2019b)	
19.	OC, EC, WSOC and HULIS	2014/12/21-2015/2/2 2015/8/3-9/7	Nam Co, inland TP	TSP samples	off-line, 12 h	Wu et al. (2018)	
20.	Overestimation of BC concentration caused by high OC	2018/7-2019/7; 2020/4- 2021/3; 2017/1-2017/12	Yaze, southeastern TP; QOMS, southern TP; Nam Co, inland TP	TSP samples	off-line, 2-20 days	Hu et al. (2023)	
21.	Chemical composition of PM _{2.5} and TSP	2010/6-2010/9	Qinghai Lake, northeastern TP	PM _{2.5} and TSP samples	off-line, 3 days	Zhang et al. (2014)	
22.	Chemical characterization and sources of PM _{2.5}	2015/5-2016/5	Mt. Gongga, eastern TP	PM _{2.5} samples	off-line, weekly	Meng et al. (2020)	
23.	Chemical composition and size distribution of PM _{2.5}	2012/7/11-9/6	Qilian Shan, northeastern TP	PM _{2.5} samples	off-line, 3 days	Xu et al. (2015)	
24.	Chemical compositions in PM _{2.1}	2012/4-2014/12	Mt. Gongga, eastern TP	Nine-stage Anderson sampler	off-line, 48 h	Su et al. (2018)	
25.	Chemical composition of size-segregated aerosols	2013/3-2014/2	Lhasa, urban city	Size-segregated samples	off-line, 2 weeks	Wan et al. (2016)	
26.	Chemical composition and optical properties	2016/4/12-5/12	QOMS, southern TP	PM _{2.5} samples	off-line, 48 h	Xu et al. (2020)	
27.	Molecular composition and optical properties of WS-BrC	2017/7/1-7/31	Waliguan, northeastern TP			Xu et al. (2022a)	
28.	Characteristics and sources of PM _{2.5}	2018/3/11-5/13	Gaomeigu, southeastern TP	PM _{2.5} samples	off-line, 24 h	Zhao et al. (2019b)	
29.	Abundance, composition and source of PM _{2.5}	2010/7/3-8/26	Qinghai Lake, northeastern TP	PM _{2.5} samples	off-line, 24 h	Li et al. (2013)	
30.	Seasonal/diurnal variation of molecular tracers for OAs	2014/12/21-2015/2/1 2015/8/4-9/6	Nam Co, inland TP	TSP samples	off-line, 12 h	Wan et al. (2023)	

Off-line filter sample analysis for aerosol light absorption of CAs (BC/EC/OC/WSOC/BrC/HULIS)						
31.	Light absorption of CAs	2014/8-2015/8	QOMS and Lulang	TSP samples	off-line, 2 samples/month	Li et al. (2016c)
32.	Seasonal variation and light absorption property of CAs	2014/12-2016/12	Mt. Yulong and Ganhaizi basin	TSP samples	off-line, 6 days	Niu et al. (2017)
33.	Light absorption by WSOC	2015/6/2-7/1	Nam Co Station, inland TP	PM _{2.5} samples	off-line, 24 h	Zhang et al. (2017)
34.	Sources and light absorption characteristics of WSOC	2016/11-2017/11	Nam Co, inland TP	PM _{2.5} and TSP samples	off-line	Li et al. (2021b)
35.	Multi-wavelength light absorption of BC and BrC	2008/7/16-2009/7/26	Lulang, southeastern TP	TSP samples	off-line, 72 h	Zhao et al. (2019a)
36.	Light absorption, fluorescence properties and sources of BrC	2013/8-2014/1	Lulang, southeastern TP	TSP samples	off-line, 11.5 h	Wu et al. (2020)
37.	Light absorption properties of BrC	2015/11-2016/11	Lulang, southeastern TP	TSP samples	off-line, weekly	Zhu et al. (2018)
38.	Light absorption of CAs from the Sichuan Basin to TP	2018/12/21-2019/12/18	Six sites from SCB to TP	PM ₁ samples	off-line, day/night	Zhao et al. (2022a)
39.	Molecular compositions and light absorption of HULIS	2015/3/22-2015/4/14	Mt. Yulong, southeastern TP	PM _{2.5} samples	off-line, 48 h	Wang et al. (2019b)
Off-line filter sample analysis for PAHs, Individual particle analysis by TEM, Isotope Analysis, and others						
40.	Concentrations of PAHs	2008/8/5-2009/7/13	Lhasa, urban city	TSP samples	off-line, 24 h/week	Ma et al. (2013)
41.	Source apportionment and risk assessment of atmospheric PAH	2013/4-2014/3	Lhasa, urban city	TSP samples	off-line, 24 h	Chen et al. (2018a)
42.	n-Alkanes and PAHs	2015/11-2016/11	Lulang, southeastern TP	TSP samples	off-line, weekly	Zhu et al. (2020)
43.	Individual particle analysis (Mixing state and sources)	2013/9/15-10/15	Menyuan, northeastern TP	TEM samples	off-line	Li et al. (2015)
44.	Individual particle analysis (Mixing State and Fractal Dimension)	2016/5/26-6/2	Lulang, southeastern TP	TEM samples	off-line	Yuan et al. (2019)
45.	Individual particle analysis (composition and sources)	2013/2/2-3/8	Lhasa, urban city	TEM samples	off-line	Duo et al. (2015)
46.	Radiocarbon and stable isotope ¹³ C of OC and EC	2018/9-2019/8	Qinghai Lake, northeastern TP	TSP samples	off-line, 1 week	Ni et al. (2023)
47.	Δ ¹⁴ C and δ ¹³ C of TC and IPC	2020/7-2021/6	Hongyuan, eastern TP	PM _{2.5} and TSP samples	off-line, every 2 weeks	Li et al. (2022)
48.	Stable isotope (¹⁵ N) analysis of nitrogen compounds	2017/3/23-2018/3/19	QOMS, southern TP	TSP samples	off-line	Bhattarai et al. (2023)
49.	Nitrogen and oxygen isotopic compositions	2017/3-2018/2	QOMS, southern TP	TSP samples	off-line, 1 week	Lin et al. (2021)
50.	S-isotope characteristic	2021/3-2021/6	QOMS, southern TP	PM _{2.5} samples	off-line, 96 h	Dasari et al. (2023)
51.	Primary biological aerosol particles (PBAPs)	2018/8-2019/9; 2018/9-2019/7; 2019/7-2020/7	Qinghai lake, Beiluhe, Ngari	TSP samples	off-line	Zhu et al. (2022)
52.	Oxalic acid and related SOA	2010/7-2010/8	Qinghai Lake, northeastern TP	PM _{2.5} samples	off-line, 24 h	Meng et al. (2013)
On-line measurement of BC concentration and light absorption of BC and BrC by Aethalometers and other optical instruments						
53.	Concentration, temporal variation, and sources of BC	2015/5/15-2017/5/31	QOMS, southern TP	AE-33 Aethalometer	on-line, 1-min	Chen et al. (2018b)
54.	BC concentration	2009/5-2011/3	Qilian Shan, northeastern TP	AE-31 Aethalometer	on-line, 5-min	Zhao et al. (2012)
55.	BC concentration	2008/11-2009/1	Linzi, southeastern TP	AE-16 Aethalometer	on-line, 5-min	Cao et al. (2011)
56.	Aerosol light absorption	Autumn 2015	Lhasa and Lulang	AE-33 Aethalometer	on-line, 1-min	Zhu et al. (2017)
57.	Light absorption contributions of BC, BrC _{pri} , and BrC _{sec}	2018/8-2019/9; 2018/9-2019/7; 2019/7-2020/7	Qinghai lake, Beiluhe, Ngari	AE-33 Aethalometer	on-line	Zhu et al. (2021)
58.	Secondary BrC absorption	2018/3/14-5/13	Gaomeigu, southeastern TP	AE-33 Aethalometer	on-line, 1-min	Wang et al. (2019a)
59.	Source contribution of fossil fuels and biomass-burning BC	2018/3/14-5/13	Lijiang, southeastern TP	AE33 Aethalometer/PAX	on-line	Liu et al. (2021)
60.	Mass concentrations, size distributions, mixing states, and light absorption properties of refractory BC	2015/9/17-10/31	Lulang, southeastern TP	SP2 and PAX	on-line, hourly	Wang et al. (2018)

61.	Light absorption properties of BC and BrC	2015/5/31-7/1 2016/4/12-5/12 2017/7/1-7/31	Nam Co, inland TP QOMS, southern TP Waliguan, northeastern TP	Aethalometer PAX/MAAP HR-ToF-AMS	on-line, 5-min on-line, hourly	Zhang et al. (2021b)
62.	Optical properties, size distributions, chemical compositions	2019/7/8-8/2	Shiquanhe, southwestern TP	Nephelometer/PAX/APS	on-line, 10s-5 min	Zhang et al. (2021a)
63.	Optical characteristics of aerosol	2020/8/5-9/11	Nam Co, inland TP	APS/MAPP	on-line, 5/1-min	Wang et al. (2023)
On-line measurement of aerosol mass and number concentrations by TEOM RP1400, SMPS, and other instruments						
64.	Concentrations of BC, PM _{2.5} , PM ₁₀ , and CAs and WSIs	2004/4/7-5/24	Tengchong, southeastern TP	TEOM RP1400/Aethalometer PM _{2.5} and PM ₁₀ samples	on-line, 5-min off-line, 24-48 h	Engling et al. (2011)
65.	Characteristics of aerosol mass loadings	2011–2013	Ngari, QOMS, Nam Co, SET	TEOM RP1400 PM _{2.5} samples	on-line, 5-min off-line, 3 days	Liu et al. (2017)
66.	Spatial and temporal variations of particulate pollutants	2016/6-2017/5	Lhasa, Ngari, Qamdo, Nyingchi, Nagchu, and Shigatse	CNEMC download data	on-line, hourly	Chen et al. (2019b)
67.	New particle formation	2019/4/26-5/22 2019/6/15-6/25	Nam Co, inland TP	SMPSs AE-33 Aethalometer	on-line	Tang et al. (2022)
68.	Vertical profiling of particle size distributions	2020/8/8-8/28	Lhasa, urban city	POPS and GRIMM 11-C attached to a tethered balloon	on-line	Ran et al. (2022)
69.	Aerosol number size distribution	2016/8/2-8/20	Lhasa, urban city	Optical particle size spectrometer	on-line, 10-min	Cui et al. (2018)
70.	First simultaneous measurements of PAN and O ₃	summer in 2011/2012	Nam Co, inland TP	PAN Analyzer	on-line, 10-min	Xu et al. (2018b)
71.	Measurements of O ₃ and PAN as well as their precursors	2019/5/1-7/31	Nam Co, inland TP	GC-ECD analyzer	on-line	Xu et al. (2023)
On-line measurement of ambient aerosol chemical composition by aerosol mass spectrometers						
72.	Chemical characterization of PM ₁	2013/9/5-10/15	Menyuan, northeastern TP	ACSM	on-line, 15-min	Du et al. (2015)
73.	Chemical composition of PM ₁	2015/3/22-4/14	Mt. Yulong, southeastern TP	HR-ToF-AMS	on-line, 5-min	Zheng et al. (2017)
74.	Chemical characterization of rBC	2015/5/30-6/30	Nam Co, inland TP	SP-AMS/HR-ToF-AMS MAAP	on-line, 5-min	Wang et al. (2017)
75.	Chemical characteristics and sources of PM ₁	2015/5/31-7/1	Nam Co, inland TP	HR-ToF-AMS/SMPS/MAAP etc.	on-line, 5-min	Xu et al. (2018a)
76.	Chemical characteristics and sources of PM ₁	2016/4/12-5/12	QOMS, southern TP	HR-ToF-AMS/SMPS/PAX	on-line, 5-min	Zhang et al. (2018)
77.	Chemical characteristics and sources of PM ₁	2017/7/1-7/31	Waliguan, northeastern TP	HR-ToF-AMS/PAX/CCN-100	on-line, 10-min	Zhang et al. (2019)
78.	Chemical characteristics and sources of PM ₁ during a festival	2017/7/5-7/6	Waliguan, northeastern TP	HR-ToF-AMS/PAX/CCN-100	on-line, 10-min	Zhang et al. (2020)
79.	Chemical characteristics and sources of PM ₁	2019/8/31-9/26	Lhasa, urban city	HR-ToF-AMS/SMPS/PAX	on-line, 5-min	Zhao et al. (2022b)
80.	Regional difference of aerosol chemical compositions and sources in the TP and its surroundings	Different periods during 2015-2022	QOMS, Motuo, NamCo, Ngari, Waliguan, LHG, Bayanbulak, and Lhasa	HR-ToF-AMS SMPS/ /CCN-100 PAX/Aethalometer	on-line, hourly	This study

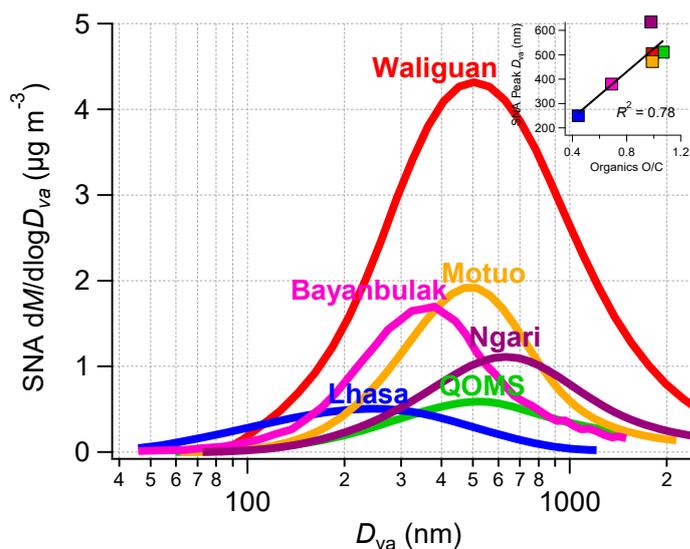
- 235 **Abbreviation:** BC: black carbon; OC: organic carbon; EC: elemental carbon; CAs: carbonaceous aerosols; WSOC: water soluble organic carbon; BrC: brown carbon; WSIs:
236 water soluble ionic species; HULIS: Humic-Like Substances; SOA: secondary organic aerosol; TEM: Transmission electron microscopy; TC: total carbon; IPC: water-insoluble
237 particulate carbon; PAH: polycyclic aromatic hydrocarbons; PAN: peroxyacetyl nitrate; APCC: Atmospheric Pollution and Cryospheric Change; CNEMC: China National
238 Environmental Monitoring Center;

239 **Figures**



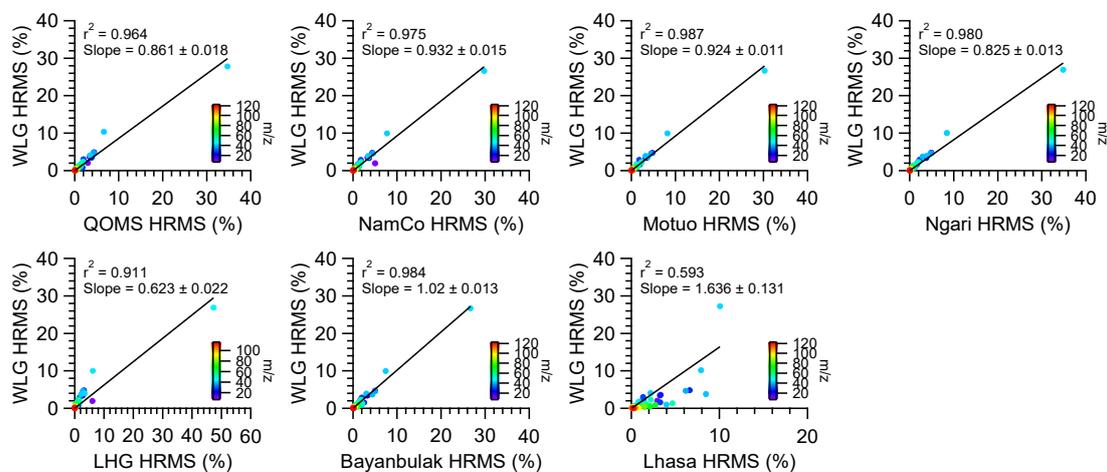
240

241 **Fig. S1** High-time-resolution temporal variations of the mass concentrations of PM₁ chemical
 242 compositions during the eight online aerosol field measurement campaigns over the Tibetan Plateau
 243 and its surroundings.



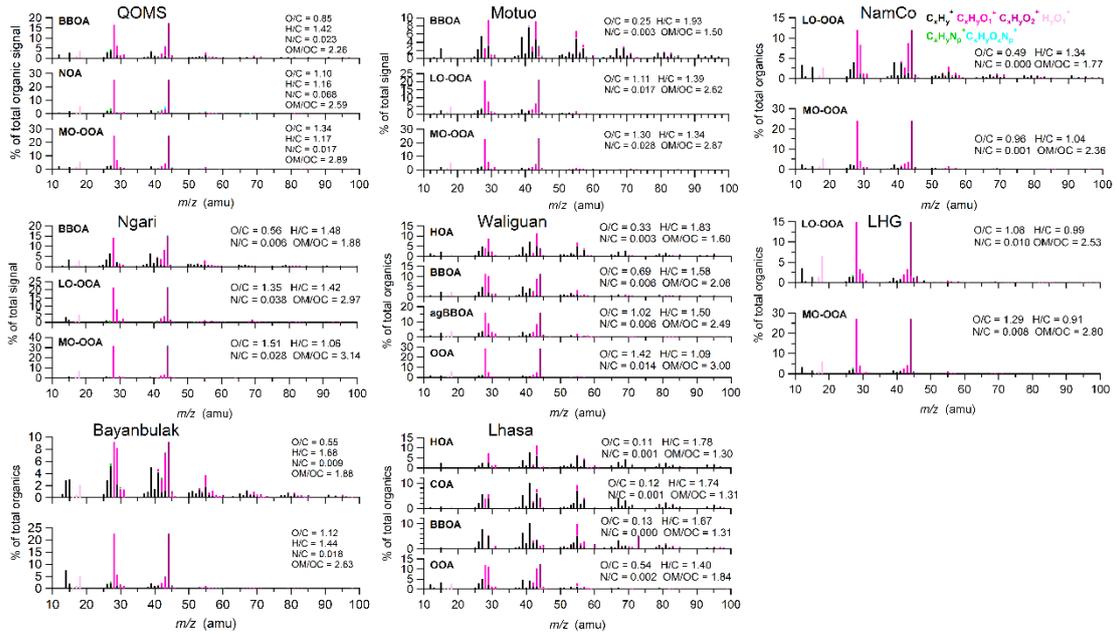
244

245 **Fig. S2** Average size distributions of SNA mass concentrations during six field measurement
 246 campaigns in the Tibetan Plateau and its surroundings. Insert graph is the scatter plot of peak
 247 diameters in these size distributions versus the average O/C ratio of organics.



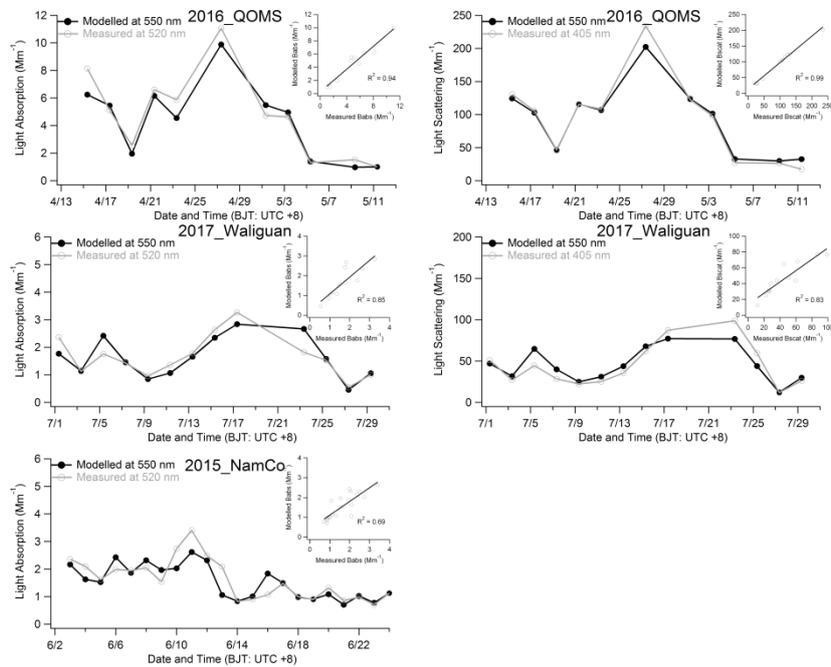
248

249 **Fig. S3** Comparison of OA HRMS between Waliguan and the other seven sites in this study.



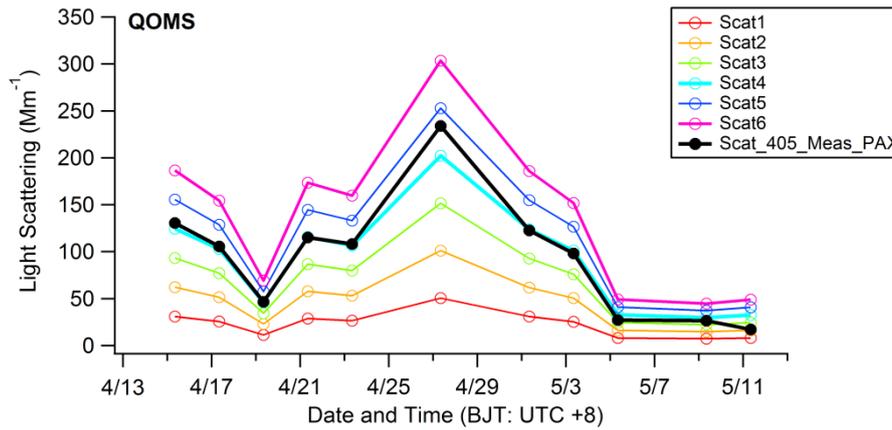
250

251 **Fig. S4** The high-resolution mass spectra (HRMS) of different OA factors identified by the PMF
 252 source apportionment among the eight field campaigns in our study. All the HRMSs are colored by
 253 six ion categories at $m/z < 120$.



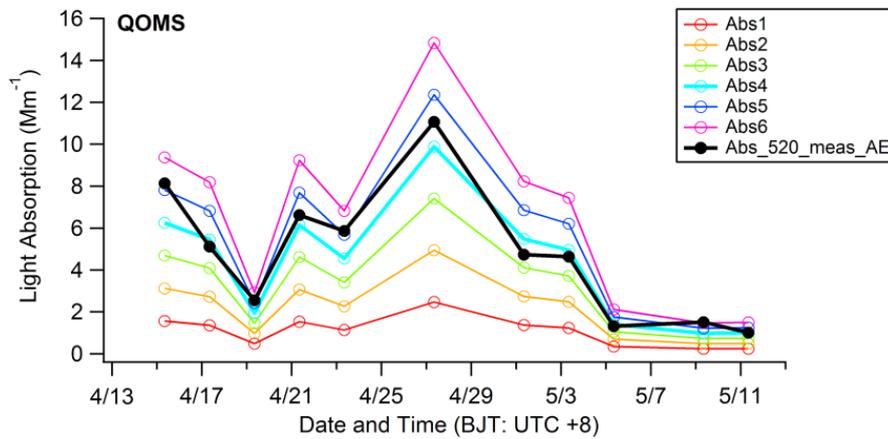
254

255 **Fig. S5** Comparisons of the particle light scattering and absorption coefficients between the
 256 modelled values (final selected) from Optical Properties of Aerosol and Cloud (OPAC) model and
 257 measured values from the photoacoustic extincitometer and aethalometer during the three
 258 campaigns.



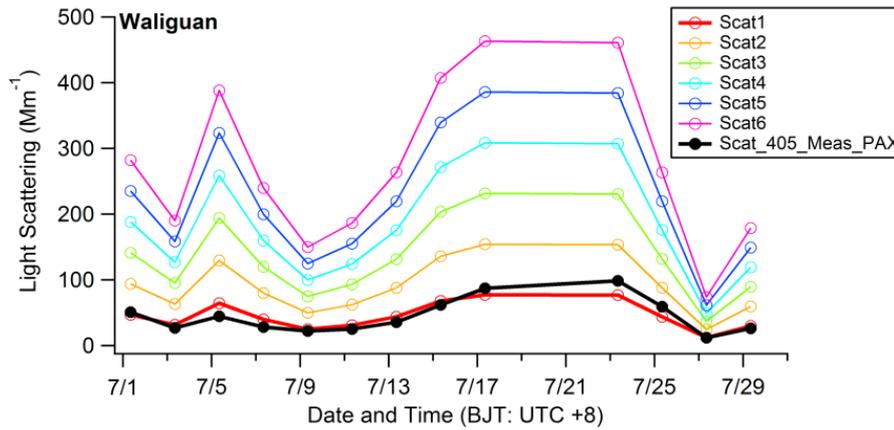
259

260 **Fig. S6** Time series of the particle light scattering coefficients at 405 nm measured by the PAX and
 261 those modelled light scattering coefficients at 550 nm under different particle numbers in per cubic
 262 meter air (the calculated value multiplier 1 to 6, respectively) from the OPAC model during the
 263 QOMS campaign.



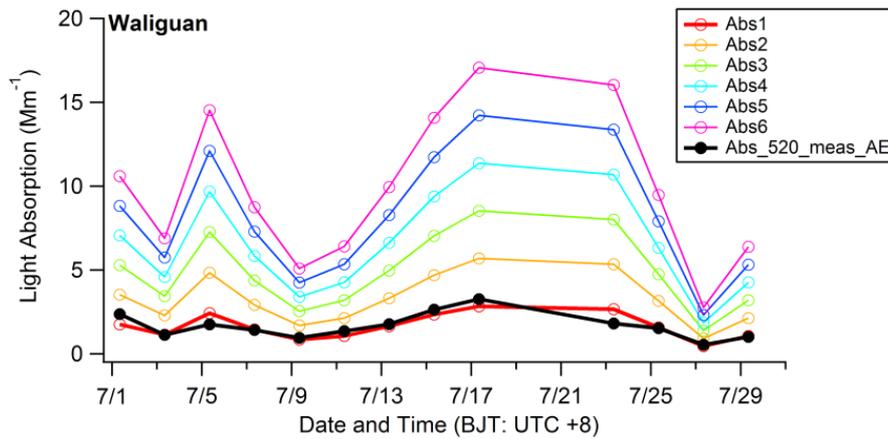
264

265 **Fig. S7** Time series of the particle light absorption coefficients at 520 nm measured by the
 266 Aethalometer and those modelled light absorption coefficients at 550 nm under different particle
 267 numbers in per cubic meter air (the calculated value multiplier 1 to 6, respectively) from the OPAC
 268 model during the QOMS campaign.



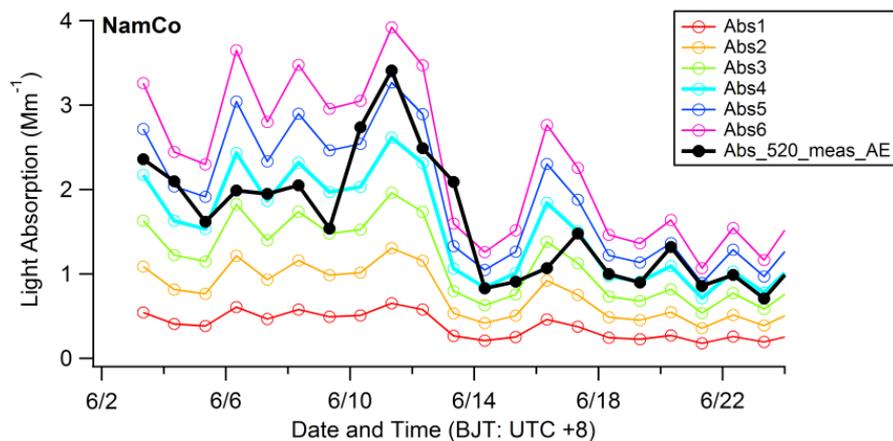
269

270 **Fig. S8** Time series of the particle light scattering coefficients at 405 nm measured by the PAX and
 271 those modelled light scattering coefficients at 550 nm under different particle numbers in per cubic
 272 meter air (the calculated value multiplier 1 to 6, respectively) from the OPAC model during the
 273 Waliguan campaign.



274

275 **Fig. S9** Time series of the particle light absorption coefficients at 520 nm measured by the
 276 Aethalometer and those modelled light absorption coefficients at 550 nm under different particle
 277 numbers in per cubic meter air (the calculated value multiplier 1 to 6, respectively) from the OPAC
 278 model during the Waliguan campaign.



279

280 **Fig. S10** Time series of the particle light absorption coefficients at 520 nm measured by the
 281 Aethalometer and those modelled light absorption coefficients at 550 nm under different particle
 282 numbers in per cubic meter air (the calculated value multiplier 1 to 6, respectively) from the OPAC
 283 model during the NamCo campaign.

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