



# The Coastal Surveillance Through Observation of Ocean Color (COAST $\ell$ OOC) dataset

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**Abstract.** Coastal Surveillance Through Observation of Ocean Color (COAST $\ell$ OOC) oceanographic expeditions were conducted in 1997 and 1998 to examine the relationship between the optical properties of seawater and related biological and chemical properties across the coastal to open-ocean gradient in various European seas. A total of 379 stations were visited along the coasts of the Gulf of Lion in the Mediterranean Sea ( $n = 61$ ), the Adriatic Sea ( $n = 39$ ), the Baltic Sea ( $n = 57$ ), the North Sea ( $n = 99$ ), the English Channel ( $n = 85$ ), and the Atlantic Ocean ( $n = 38$ ). Particular emphasis was placed on the collection of a comprehensive set of apparent and inherent optical properties (AOPs and IOPs) to support the development of ocean color remote-sensing algorithms. The data were collected in situ using traditional ship-based sampling but also from a helicopter, which is a very efficient means for that type of coastal sampling. The dataset collected during the COAST $\ell$ OOC campaigns is unique in that it is fully consistent in terms of operators, protocols, and instrumentation. This rich and historical dataset is still today frequently requested and used by other researchers. Therefore, we present the result of an effort to compile and standardize a dataset which will facilitate its reuse in future development and evaluation of new bio-optical models adapted for optically complex waters. The dataset is available at <https://doi.org/10.17882/93570> (Massicotte et al., 2023).

## 1 Introduction

Since the launch of the Coastal Zone Color Scanner (CZCS) by NASA in 1978, ocean color remote sensing has been used to monitor the state of and changes in global marine ecosystems in both time and space. In open oceans, the main component that affects the variations in the inherent and apparent optical properties (IOPs and AOPs) of seawater is phytoplankton, which is usually represented by the concentration of chlorophyll *a* (Morel and Prieur, 1977). Such open-ocean waters are traditionally termed Case-1 waters. Many simple spectral band-ratio algorithms have been developed to link

changes in remotely sensed ocean color (OC), measured as reflectance, to variations in chlorophyll-*a* concentration (see O’Reilly and Werdell, 2019, for an extensive evaluation of OC band-ratio algorithms). Because these algorithms perform well, a plethora of studies have been conducted based on them, notably about phytoplankton phenology (e.g., Vargas et al., 2009) and phytoplankton primary production (see Carr et al., 2006, and references therein).

Spaceborne monitoring of aquatic ecosystems is more challenging for optically complex Case-2 waters (Morel and Prieur, 1977) often found in coastal areas. In contrast to

Case-1 waters, the optical properties of Case-2 waters are determined by several types of constituents occurring in individually highly variable concentrations. Rivers draining large catchment areas deliver important quantities of optically significant substances such as chromophoric dissolved organic matter (CDOM) and suspended particulate matter (SPM) to coastal waters (Hedges et al., 1997; Cole et al., 2007; Massicotte et al., 2017). In shallow areas, bottom reflectance and resuspension of sediments can also alter the signal of water-leaving reflectance (Lee et al., 1998). Because CDOM and SPM do not necessarily co-vary with chlorophyll *a* and can mask the presence of phytoplankton (Sathyendranath, 2000), bio-optical algorithms developed for Case-1 waters are generally not appropriate for optically complex Case-2 waters (Gordon and Morel, 1983). Hence, concomitant in situ measurements of optically active components (CDOM, SPM, or chlorophyll *a*) and related radiometric quantities (AOPs and IOPs) are, to this day, critically needed to develop and improve bio-optical algorithms for coastal areas.

The objective of the Coastal Surveillance Through Observation of Ocean Color (COAST $\ell$ OOO) oceanographic expeditions was to acquire a comprehensive set of AOPs, IOPs, and concentrations of optically significant components along European coasts to help the development of new remote-sensing algorithms in optically complex waters. Back in 1997 and 1998, the COAST $\ell$ OOO campaigns were among the very first concerted efforts to fill the knowledge gaps in coastal marine optics by establishing a large, consistent, and comprehensive optical dataset for coastal waters. A total of 379 stations were visited along the coasts of the Mediterranean Sea, Adriatic Sea, Baltic Sea, North Sea, English Channel, and Atlantic Ocean, mostly in coastal but also open-ocean waters. While this unique set of data has been used or referenced over the years in numerous peer-reviewed publications (see the list in Appendix A), a reference final dataset, with well-documented quality controls, has never been published. Even though the COAST $\ell$ OOO campaigns were carried out more than 20 years ago, the rich and historical dataset that has been collected still has great potential to contribute to the development and evaluation of new bio-optical models adapted for optically complex waters. The goal of this paper is to formally present a quality-controlled and final reference version of the COAST $\ell$ OOO dataset archived in a public repository where users can download it readily.

## 2 Study area and sampling overview

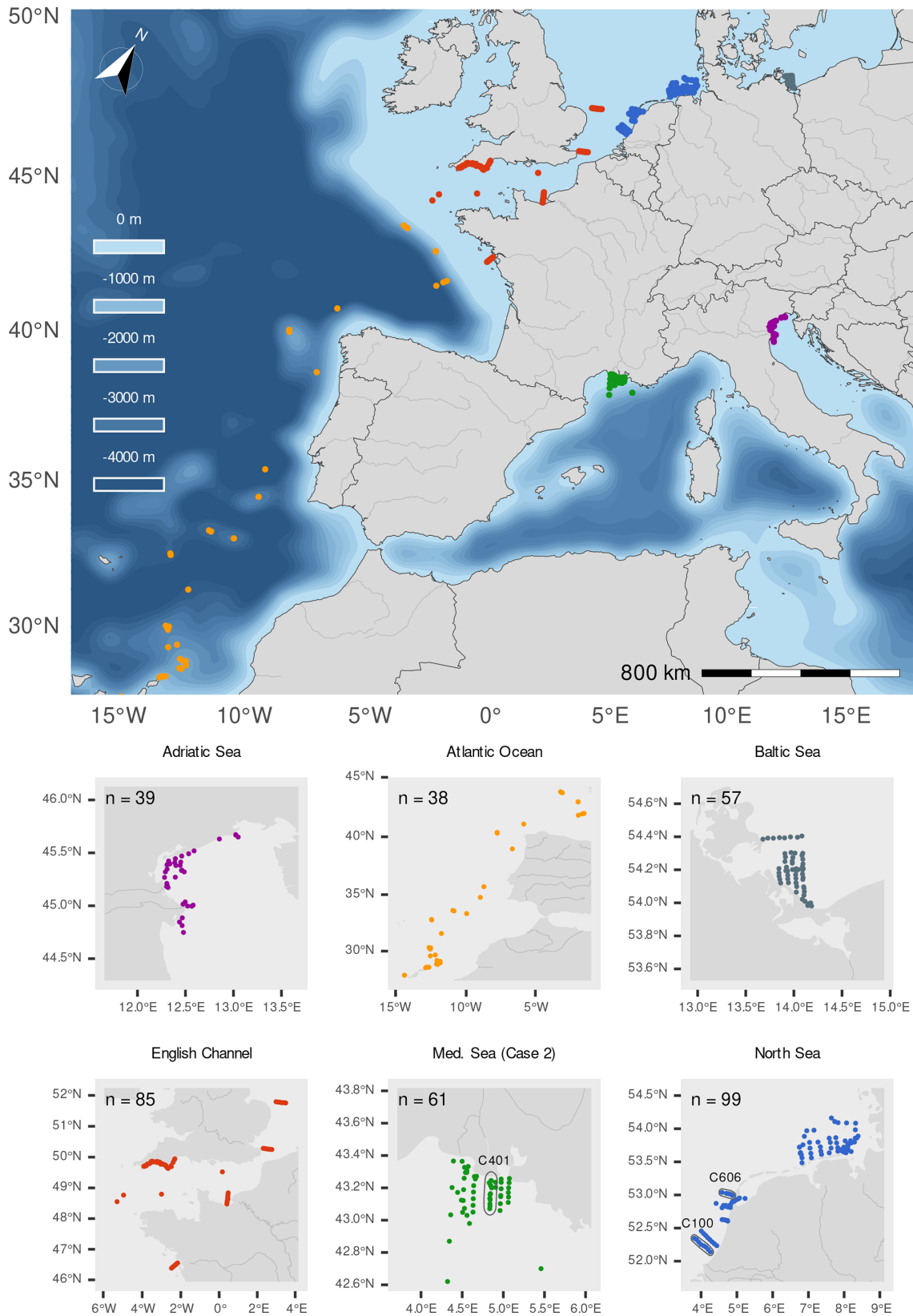
### 2.1 Study area and general sampling strategy

During the six COAST $\ell$ OOO campaigns, further referred to as C1 to C6, a total of 379 locations were visited (Fig. 1). The stations were located along the coasts of the Gulf of Lion in the Mediterranean Sea ( $n = 61$ ), the Adriatic Sea ( $n = 39$ ), the Baltic Sea ( $n = 57$ ), the North Sea ( $n = 99$ ), the

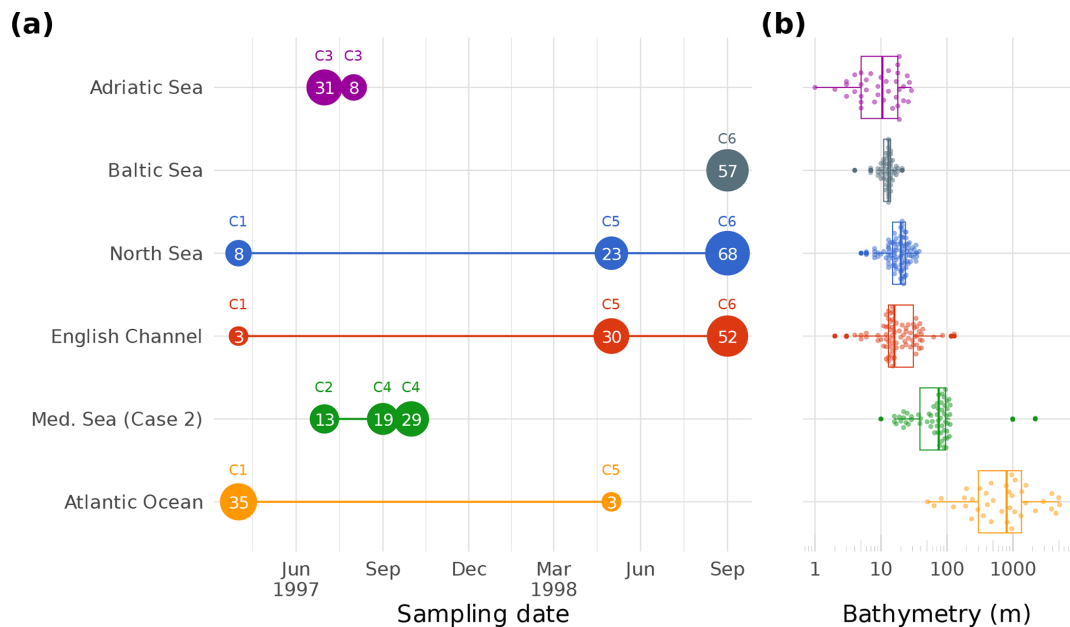
English Channel ( $n = 85$ ), and the Atlantic Ocean ( $n = 38$ ). Within each area, the stations were generally distributed along across-shore or along-shore transects to capture the land-to-sea gradients and document river plumes (Fig. 1b). COAST $\ell$ OOO campaigns were either ship-based (C1, C4, C5) or used a helicopter as a sampling platform (C2, C3, C6) and took place between 2 April 1997 and 25 September 1998 (Fig. 2a). Compared to traditional ship-based sampling, the helicopter platform allowed efficient stations close to the coast in areas difficult to access by boat (some samples were collected in waters as shallow as 1 m, Babin, 2003). Combining both ship and airborne sampling approaches allowed the whole inshore to open-ocean aquatic continuum to be covered. The bathymetry (GEBCO Bathymetric Compilation Group 2021, 2021) varied greatly across the stations, where it averaged ca. 10 m in the Adriatic Sea and ca. 800 m in the Case-1 Atlantic Ocean (Fig. 2b).

### 2.2 Water sampling and processing

On board the ship, water samples were analyzed and/or stored immediately after collection. When using a helicopter as a sampling platform, water samples were kept in polyethylene containers for no longer than 2 h after sampling (see Babin, 2003, for further details). Filtration was conducted at low vacuum onto 25 mm glass-fiber filters (Whatman, GF/F) to collect particles for subsequent analyses as described below. Three separate subsamples (up to 2 L each) were filtered for each sample with the following purposes. To determine the concentration of SPM ( $\text{g m}^{-3}$ ), the particles collected onto a pre-weighed GF/F filter were dried and stored in a freezer at  $-80^\circ\text{C}$  until the dry weight was determined less than 2 months later in the laboratory (Van Der Linde, 1998). For pigment analysis, the filter with collected particles was inserted into a cryotube and kept in liquid nitrogen until analysis less than 3 months later. High-performance liquid chromatography (HPLC) was used as described by Vidussi et al. (1996) to determine liposoluble pigment concentrations. Total chlorophyll *a* is defined here as the sum of chlorophyll *a*, divinyl chlorophyll *a*, chlorophyll-*a* isomer and epimer, chlorophyllids *a*, and pheopigments. The concentrations of particulate organic carbon and nitrogen (POC and PON, respectively) were determined using a Carlo Erba NCS 2500 elemental analyzer as described in Ferrari et al. (2003). These determinations were made on particles collected on a GF/F filter precombusted at  $450^\circ\text{C}$  for 2 h and stored in the freezer at  $-80^\circ\text{C}$  until analysis less than 2 months later. Further details concerning laboratory analyses can be found in Babin (2003), Ferrari (2000), and Ferrari et al. (2003). To determine the concentration of dissolved organic carbon (DOC, M), seawater samples were filtered on a 0.22 mm Millipore membrane. The final filtrate was transferred to a 100 mL amber-glass bottle, stored in the refrigerator, and analyzed within 3 weeks. DOC was determined by high-temperature catalytic oxidation using a Carlo Erba 480



**Figure 1.** Map of the sampling stations. The three black outlines in the bottom panels (Mediterranean Sea and North Sea) identify the stations and the transects used in Fig. 5e (C100) as well as the examples of vertical profiles of  $K_{Ed}(490)$  in Fig. 8 (C401 and C606).



**Figure 2.** (a) Overview of the temporal sampling for the six areas. The numbers in the circles indicate the number of visited stations each month. (b) Boxplot showing the bathymetry at the sampling locations by area. The labels on top of each circle identify the COAST $\ell$ OOO campaigns (C1 to C6).

analyzer. The methodological details can be found in Ferrari (2000).

### 2.3 Optical measurements

At each station, an optical package was deployed in the water column from the surface, down to 1 to 2 m above the ocean bottom in shallow waters, and to maximum depths of ca. 100 m in deep waters. The package included CTD (conductivity, temperature, and depth) and an array of optical instruments to simultaneously measure beam attenuation, absorption and scattering coefficients, as well as irradiance in the water column. The Modular Data and Power System (MODAPS) data-gathering system from Wetlabs Inc., USA, was used to combine the time-stamped data from the different instruments to produce a data matrix with a common depth grid. For further information, please refer to the COAST $\ell$ OOO final report (Aiken et al., 2000).

A specific aspect of COAST $\ell$ OOO is the use of a helicopter as a sampling platform, which has enabled high-frequency visiting of sites in very shallow waters (see Sect. 2.1). An additional advantage is the fact that helicopter-based optical measurements are not influenced by the sampling platform itself, unlike ship-based measurements unavoidably affected by the presence of the ship's hull, most prominently close to the surface. A drawback of this approach is the difficulty in operating a helicopter stationary at a constant height, especially under low or gusty wind conditions: while the former requires substantial flight heights to prevent the helicopter downdraft from impacting the opti-

cal measurements, rendering sampling less efficient, the latter can result in unwanted vertical movements of the instrument package in water on the order of 1–2 ms<sup>-1</sup>. To ensure the availability of data of sufficient quality close to the surface, several consecutive upcasts and downcasts were therefore performed at each “helicopter” site. Among the measured vertical profiles at each site, the most suitable casts for depth merging and surface extrapolation (which are not necessarily the same; see the end of Sect. 2.3.3) were selected using a combination of objective criteria and visual inspection.

#### 2.3.1 Irradiance measurements

The SeaWiFS Profiling Multichannel Radiometer (SPMR, Satlantic Inc., Canada) was used to measure downward ( $E_d$ ) and upward ( $E_u$ ) spectral irradiances ( $\text{W m}^{-2} \mu\text{m}^{-1}$ ) in the water column. Irradiance was measured at 13 wavelengths matching the MEdium Resolution Imaging Spectrometer (MERIS) channels of direct relevance for ocean observations (411, 443, 456, 490, 509, 531, 559, 619, 664, 683, 706, 779, and 866 nm, except for COAST $\ell$ OOO 1 operating at 590 nm instead of 619 nm), ranging from the blue part of the spectrum to the near-infrared at an acquisition rate of 6 Hz. The actual wavelengths differ slightly between upward and downward observations.

The SPMR was fixed to the IOP instrument frame, except for the COAST $\ell$ OOO 5 mission, where it was used in free-falling mode. For the ship-based measurements, concomitant observations of the in-air downward spectral irradiance at the

sea surface were performed to account for changes in the incoming solar radiation during the in situ profiling (e.g., due to clouds). This approach was not achievable for the helicopter-based measurements, since there was no suitable place to mount the radiometer. For these campaigns, measurements of  $E_d$  made by the profiler while in the air (i.e., before and after the in-water measurements) were used to assess  $E_d$  on the sea surface. In order to minimize the time lag between the in-water measurements and the  $E_d$  values on the sea surface, our processing code automatically chose the air measurements closest in time. Under stable atmospheric conditions, no significant error is expected to result from this approach, while it will contribute to increased variability in the retrieved parameters when sampling under varying cloud cover. However, the practical impact of this is limited due to the short duration of helicopter-based sampling, typically lasting between 1 and 3 min per profile.

### 2.3.2 Water-leaving reflectance

To calculate the water-leaving reflectance, accurate measurements of both  $E_d$  and  $E_u$  just below the water surface ( $z = 0^-$ ) are needed. In case no spectrally matching in-air  $E_d$  reference data are available (as is the case for the helicopter-based campaigns), this was done by extrapolating both  $E_d(z)$  and  $E_u(z)$  vertical profiles toward the sea surface by fitting an exponential model for the depth dependence of the measured spectral irradiance values (Eq. 1):

$$\Gamma(\lambda, z) = \Gamma(\lambda, 0^-) e^{-K_\Gamma(\lambda)z}, \quad (1)$$

where  $\Gamma(\lambda, z)$  is either  $E_d$  or  $E_u$  measured at wavelength  $\lambda$  and at depth  $z$ ,  $\Gamma(\lambda, 0^-)$  is  $E_d$  or  $E_u$  estimated just below the air–water interface, and  $K_\Gamma$  is the estimated diffuse attenuation coefficient for downward ( $K_{E_d}$ ,  $\text{m}^{-1}$ ) or upward ( $K_{E_u}$ ,  $\text{m}^{-1}$ ) irradiance. Note that a correction to account for instrument self-shading of the upwelling irradiance was not applied. We assume that the impact of this omission is relatively small except for highly absorbing waters (depending on constituent concentration and/or observation wavelength), where it may lead to an underestimation of near-surface upwelling irradiances.

In case concomitant spectrally matching in-air  $E_d$  ( $0^+$ ) reference data are available (as is the case for most ship-based campaigns),  $E_d(0^-)$  is calculated from  $E_d(0^+)$  by applying a constant factor of 0.943 to account for losses of the downwelling irradiance (direct + diffuse) due to Fresnel reflection at the rough sea surface. The derived  $E_u(0^-)$  and  $E_d(0^-)$  values were thereafter used to calculate the water reflectance  $R(0^+)$  just below the sea surface:

$$R(0^+) = \frac{E_u(0^-)}{E_d(0^-)}. \quad (2)$$

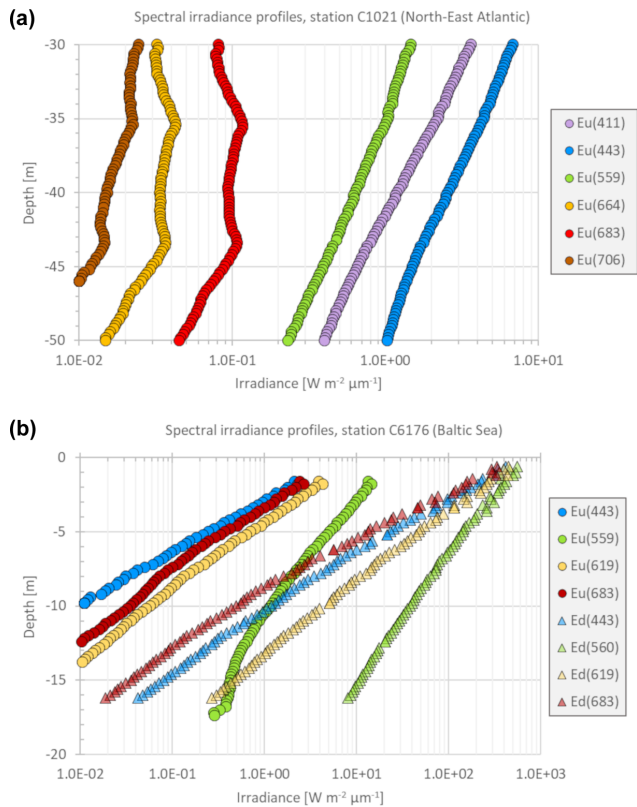
Vertical irradiance profiles in the red and near-infrared channels at wavelengths beyond 600 nm were fitted using a

sum of two exponential functions, of which one represents the diffuse attenuation due to absorption and elastic scattering for the channel under consideration, while the other accounts for Raman scattering from the corresponding (shorter) excitation wavelengths significantly contributing to the upward light field (e.g., Sugihara et al., 1984). Not doing so, i.e., using just a single exponential function, leads to substantial underestimation of the extrapolated subsurface values due to significant Raman contributions from excitation wavelengths with significantly lower  $K$  values (see Appendix A). Figure A1 showing the near-surface depth dependence of  $E_u$  at 779 nm demonstrates why the two-function approach is required. Below a depth of ca. 1.5 m, the signal is strongly dominated by Raman scattering and is therefore characterized by a diffuse attenuation coefficient very close to that of the corresponding excitation wavelengths.

### 2.3.3 Irradiance depth merging

Depth-merged irradiance profiles were derived for all the sites by applying a standardized approach. The underlying strategy for depth merging was to provide users with the information they need to do their analyses while modifying the original data as little as possible. Processing was therefore limited to procedures that could be applied equally to all COAST $\ell$ OOC campaigns, whether helicopter- or ship-based. For example, variations of the incoming solar radiation were not considered in the depth-merging process, since such observations were not available for the helicopter-based campaigns. However, if measured while operating from research vessels, such information is included in the depth-merged profile data to allow users to do their own subsequent analyses.

Typically, several upcast and downcast vertical profiles were performed, especially when using the helicopter as a sampling platform, to increase the potential availability of high-quality observations. The actual merging process was executed as follows. First, any observation not meeting both of the following two conditions was discarded: (1) SPMR instrument tilt smaller than  $20^\circ$ ; (2) SPMR vertical speed between 0.1 and  $2.0 \text{ ms}^{-1}$ . To automatically identify the best-suited cast for profile generation, the observations fulfilling the above two conditions were then distributed cast-wise over 0.2 m wide vertical bins, starting at zero depth. Finally, the cast with the highest number of bins containing at least one valid observation was selected for depth profile generation if it fulfilled two further conditions: (3) at least 2 m vertical distance between the highest and lowest bins and (4) at least 50 % of the bins containing at least one observation. No depth-merged data are provided for stations where there is no cast fulfilling all of the above four conditions. Combining several casts taken at one station has not proven successful due to potentially large discontinuities in the aggregated values at depth levels where the number of available casts



**Figure 3.** (a) Vertical profile of the upwelling irradiance taken on 11 April 1997 at station C1021 located in the northeastern Atlantic offshore of northern Portugal. (b) Vertical profiles of the upwelling and downwelling irradiances taken on 25 September 1998 at station C6176 located in the Baltic Sea offshore of Usedom in northeastern Germany. See the text for further details.

changes, especially under unstable atmospheric conditions with the corresponding short-term irradiance fluctuations.

Actual depth merging was done by taking the median of all the observations of a particular irradiance parameter within each depth bin. Spectral irradiance observations were only considered for values above  $0.01 \text{ W m}^{-2} \mu\text{m}^{-1}$  to reduce the impact of residual dark-current and radiometric noise. Finally, an attempt was made to identify and mask out remaining outliers in the depth-merged irradiance profiles. This was done by applying Grubb's test at a 5% significance level (Guthrie, 2012) to the residual of a second-order polynomial fit of the logarithmic-transformed spectral irradiance observations for a 2 m deep window sliding down in steps of 1 m. While this procedure has successfully identified individual spikes, some obvious outliers extending over several depth bins remain in the merged dataset, and users will need to make their own judgment on how to identify and remove those.

Examples of depth-merged irradiance profiles are shown in Fig. 3 for one clear-water site in the northeastern Atlantic (Fig. 3a) and one CDOM-rich site in the Baltic Sea (Fig. 3b).

Figure 3a features a double deep chlorophyll maximum with the corresponding irradiance increase due to chlorophyll fluorescence at water depths between ca. 35 and 45 m, most prominently visible at 683 nm but also affecting the neighboring channels at 664 and 706 nm. Diffuse attenuation in the blue to green wavelengths is correspondingly enhanced, as reflected by the change in slope at a depth of ca. 34 m. Figure 3b features Baltic Sea waters rich in CDOM, resulting in strong absorption in the UV and blue parts of the spectrum. The reduced decrease in the upwelling irradiance at 559 nm toward larger depths is likely caused by the reflection of downwelling irradiance from the ocean bottom, while chlorophyll fluorescence is likely the reason for the relatively enhanced (change in slope) upwelling irradiance at 683 nm for depths below ca. 7 m.

Note that extrapolated surface values and depth-merged profiles for a specific site are not necessarily derived from the same observations. This is especially true for helicopter-based operations typically comprising several upcasts and downcasts close to the water surface to increase the likelihood of having sufficient observations for the surface extrapolation, while usually only one or two deep-water casts were made to obtain observations at greater depth. The best-suited observations for surface extrapolation, therefore, do not necessarily correspond to the cast used for depth merging.

#### 2.3.4 Profiles of the diffuse attenuation coefficient

Downward ( $K_{E_d}$ ,  $\text{m}^{-1}$ ) and upward ( $K_{E_u}$ ,  $\text{m}^{-1}$ ) spectral diffuse attenuation coefficients have been derived from the depth-merged irradiance data by fitting an exponential function through the irradiance observations within a 2 m deep window, sliding down in steps of 0.2 m. The calculation was only performed if at least five observations were available for fitting within the depth window. The choice of window width was based on a compromise between two conflicting requirements. On the one hand, it should be small enough to allow resolution of relatively narrow vertical features, such as a deep chlorophyll maximum. On the other hand, it should be large enough to allow for an accurate estimation of  $K$ . The window width of 2 m chosen herein favors vertical resolution over accuracy: while even relatively narrow vertical features remain distinguishable in their own derived  $K$  profiles, these are on the other hand affected by remaining outliers, instrument tilt variations, or instrumental noise. Obviously, users may use depth-merged irradiance profiles to derive  $K$  profiles or depth-integrated  $K$  values that better meet individual requirements.

#### 2.3.5 Chromophoric dissolved organic matter (laboratory measurements)

Measurements of CDOM absorption ( $a_{\text{CDOM}}$ ,  $\text{m}^{-1}$ ) were performed in the laboratory using a spectrophotometer (Perkin Elmer, Lambda 12) on water samples (10 cm cu-

vette path length) filtered on a 0.22  $\mu\text{m}$  Millipore membrane pre-rinsed with 50 mL of Milli-Q water (Babin, 2003).  $a_{\text{CDOM}}$  spectra were measured between either 300–750 or 350–750 nm at a 1 nm increment.

### 2.3.6 Particulate, non-algal, and phytoplankton absorption (laboratory measurements)

Water samples were filtered onto 25 mm glass-fiber filters (Whatman, GF/F) at a low vacuum before absorption measurement (Babin, 2003). Total particulate absorption ( $a_{\text{P}}$ ,  $\text{m}^{-1}$ ) was measured on particles retained on the filters between 380 and 750 nm at a 1 nm increment using a spectrophotometer equipped with a 60 mm integrating sphere (Perkin Elmer, Lambda 19) and following the transmittance–reflectance (TR) method (Tassan and Ferrari, 1995, 1998). Afterward, pigments were removed from the particles with sodium hypochlorite to measure non-algal (or non-pigmented) absorption ( $a_{\text{NAP}}$ ,  $\text{m}^{-1}$ ). Finally, phytoplankton absorption ( $a_{\text{P}}$ ,  $\text{m}^{-1}$ ) was retrieved by subtracting  $a_{\text{NAP}}$  from  $a_{\text{P}}$ .  $a_{\text{CDOM}}$  and  $a_{\text{NAP}}$  absorption spectra were fitted according to the following equation (Jerlov, 1968; Bricaud et al., 1981):

$$a_{\tau}(\lambda) = a_{\tau}(\lambda_0)e^{-S_{\tau}(\lambda-\lambda_0)} + k, \quad (3)$$

where  $a_{\tau}$  is the absorption coefficient, either  $a_{\text{CDOM}}$  or  $a_{\text{NAP}}$  ( $\text{m}^{-1}$ ),  $\lambda$  is the wavelength (nm),  $\lambda_0$  is a reference wavelength (443 nm),  $S_{\tau}$  is the spectral slope ( $\text{nm}^{-1}$ ) that describes the approximate exponential decrease in absorption with increasing wavelength, and  $k$  is a background constant ( $\text{m}^{-1}$ ) accounting for scatter in the cuvette and drift of the instrument.  $a_{\text{CDOM}}$  spectra were fitted using observations below 700 nm, whereas  $a_{\text{NAP}}$  fits were performed between 380 and 730 nm, excluding spectral regions between 400–480 and 620–710 nm to avoid possible residual pigment absorption (Babin, 2003).  $a_{\text{CDOM}}$  and  $a_{\text{NAP}}$  were baseline-corrected by subtracting the background parameter ( $k$ ) derived from the following equation:

$$\hat{a}_{\tau}(\lambda) = a_{\tau}(\lambda) - k. \quad (4)$$

Finally, the spectral background of  $a_{\text{P}}$  spectra (i.e., the average absorption measured between 746 and 750 nm) was added to  $a_{\text{NAP}}$ . The underlying rationale is that absorption by phytoplankton in the near-infrared (NIR) is null and that absorption measured in the NIR for total particles using the TR method is real (Tassan and Ferrari, 2003) and belongs exclusively to non-algal particles. The absorption signal measured for total particles retained on untreated filters is assumed to be more reliable than that measurement on bleached filters.

### 2.3.7 Total absorption and scattering (in situ measurements)

In situ vertical profiles of absorption ( $a$ ,  $\text{m}^{-1}$ ) and beam attenuation ( $c$ ,  $\text{m}^{-1}$ ) were acquired at nine wavelengths (412,

440, 488, 510, 555, 630, 532 or 650 depending on the instrument configuration, 676, and 715 nm) using a flow-through in situ absorbance–attenuance meter (AC9, Wetlabs). As the measurements are referenced to pure (Milli-Q) water, the obtained absorption and attenuation coefficients exclude the contribution of water. To correct absorption measurements for incomplete recovery of scattered light,  $a(715)$  was subtracted from  $a(\lambda \leq 715)$ . Scattering coefficients ( $b$ ,  $\text{m}^{-1}$ ) were calculated by subtracting absorption from attenuation. As the contribution of molecular (water) scattering is excluded (see above), this coefficient essentially corresponds to particle scattering and is hereafter denoted by  $b_{\text{p}}$  ( $\text{m}^{-1}$ ).

The derived values for  $a$ ,  $b_{\text{p}}$ , and  $c$  derived in this way were subsequently averaged over the first attenuation length ( $1/K_{\text{Ed}}$ ) to provide values representative of the surface layer. Due to practical difficulties when operating the AC9 from a helicopter (e.g., trapped air in the instrument), the availability of  $a$ ,  $b_{\text{p}}$ , and  $c$  values is limited for helicopter-based campaigns, especially for COAST $\ell$ OOOC 3 (Adriatic Sea). Surface-to-bottom depth merging of the full AC9 profiles has not been attempted so far, but this is envisaged for potential future reprocessing of the COAST $\ell$ OOOC dataset (see Sect. 5).

## 3 Data quality control and data processing

Different general quality-control procedures were adopted to ensure the integrity of the data. First, the raw data were visualized and screened to eliminate both errors originating from the measurement devices, including sensors (systematic or random), and errors inherent to measurement procedures and methods. Statistical metrics such as average, standard deviation, and range were computed to detect and remove anomalous values in the data. Then, data were checked for duplicates and remaining outliers. The complete list of variables is presented in Table 1.

## 4 Data description (an overview)

### 4.1 Spatial variability along the coastal–ocean gradient

The COAST $\ell$ OOOC sampling strategy was primarily designed to capture the bio-optical gradient across the sampled ecosystems and along transects from the coast toward the open ocean within each area (Fig. 1). The following sections present an overview of a few selected variables measured in the different areas. An extensive and detailed explanatory visualization analysis is presented in the final report of the COAST $\ell$ OOOC project (Aiken et al., 2000).

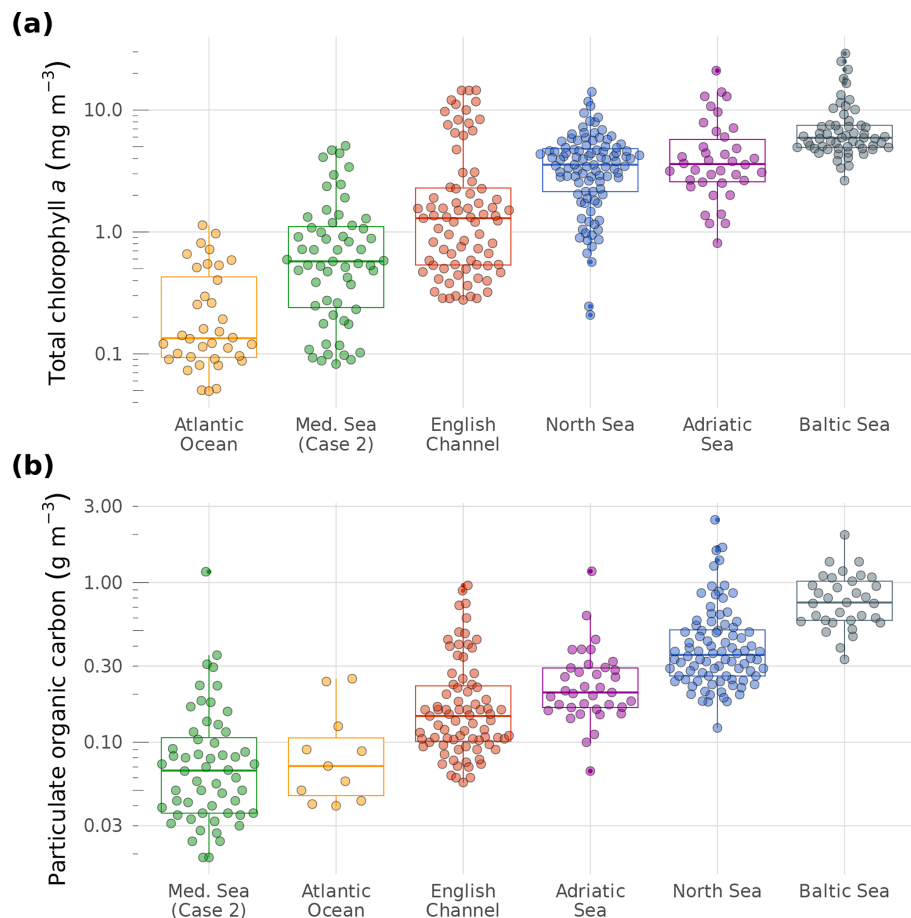
#### 4.1.1 Chlorophyll *a* and particulate organic carbon

As observed in Fig. 4, both total chlorophyll-*a* and POC concentrations varied markedly across the sampled areas, reflecting the natural gradient captured by the sampling

Table 1. List of measured parameters.

Variable	Units	Source file	Description
Wavelength	nm	100303.csv	
a_p_m1	m <sup>-1</sup>	100303.csv	Total particulate absorption
a_nap_m1	m <sup>-1</sup>	100303.csv	Non-algal absorption
a_nap_adjusted_m1	m <sup>-1</sup>	100303.csv	Non-algal absorption adjusted so that the baseline background is equal to that of particulate absorption
a_cdom_m1	m <sup>-1</sup>	100303.csv	Chromophoric dissolved organic matter absorption
a_cdom_adjusted_m1	m <sup>-1</sup>	100303.csv	Chromophoric dissolved organic matter absorption with the background baseline removed
a_phy_m1	m <sup>-1</sup>	100303.csv	Phytoplankton absorption
background_a_p_m1	m <sup>-1</sup>	100303.csv	Baseline background of total particulate absorption
background_a_cdom_m1	m <sup>-1</sup>	100303.csv	Baseline background of chromophoric dissolved organic matter absorption
background_a_nap_m1	m <sup>-1</sup>	100303.csv	Baseline background of non-algal absorption
a_m1	m <sup>-1</sup>	100304.csv	Total non-water absorption coefficient
c_m1	m <sup>-1</sup>	100304.csv	Total non-water attenuation coefficient
bp_m1	m <sup>-1</sup>	100304.csv	Particle-scattering coefficient
Longitude	Degree decimal	100305.csv	Longitude of the pixel used to extract the bathymetry
Latitude	Degree decimal	100305.csv	Latitude of the pixel used to extract the bathymetry
bathymetry_m	m	100305.csv	Bathymetry depth at the sampled stations
eu_w_m2_um	W m <sup>-2</sup> μm <sup>-1</sup>	100307.csv	Upward irradiance just beneath the water surface (Eu0-)
ed_w_m2_um	W m <sup>-2</sup> μm <sup>-1</sup>	100307.csv	Downward irradiance just beneath the water surface (Ed0-)
k_eu_m1	m <sup>-1</sup>	100307.csv	Attenuation coefficient for upward irradiance (just beneath the water surface)
k_ed_m1	m <sup>-1</sup>	100307.csv	Attenuation coefficient for downward irradiance (just beneath the water surface)
measured_reflectance_percent	%	100309.csv	Surface water reflectance
s_cdom_nm1	nm <sup>-1</sup>	100310.csv	Spectral slope that describes the approximate exponential decrease in a <sub>CDOM</sub>
s_nap_nm1	nm <sup>-1</sup>	100310.csv	Spectral slope that describes the approximate exponential decrease in a <sub>NP</sub>
a_cdom443_m1	m <sup>-1</sup>	100310.csv	Chromophoric dissolved organic matter absorption at 443 nm
a_nap443_m1	m <sup>-1</sup>	100310.csv	Non-algal absorption at 443 nm
Station	Free text	100311.csv	Unique ID of the sampled station: this can be used as a unique key to merge the data across other files.
Date	YYYY-MM-DD	100311.csv	Date on which the measurement was made
depth_m	m	100311.csv	Depth at which the measurement was made
Longitude	Degree decimal	100311.csv	Longitude of the sampling station
Latitude	Degree decimal	100311.csv	Latitude of the sampling station
Area	Free text	100311.csv	Region where the measurement was made: one of (1) North Sea, (2) English Channel, (3) Atlantic Ocean, (4) Med. Sea (Case 2), (5) Adriatic Sea, or (6) Baltic Sea
System	Free text	100311.csv	Dominant river system influencing the site
gmt_time	Hours decimal	100311.csv	Representative time of observations at the site in Greenwich mean time (UTC)
solar_zenith_angle	Degree	100311.csv	Solar zenith angle
chlorophyll_a_mg_m3	mg m <sup>-3</sup>	100308.csv	Chlorophyll <i>a</i>
chlorophyll_b_mg_m3	mg m <sup>-3</sup>	100308.csv	Chlorophyll <i>b</i>
chlorophyll_c_mg_m3	mg m <sup>-3</sup>	100308.csv	Chlorophyll <i>c</i>
pheopigment_mg_m3	mg m <sup>-3</sup>	100308.csv	Pheopigment
fucoxanthin_mg_m3	mg m <sup>-3</sup>	100308.csv	Fucoxanthin
hexanoyloxyfucoxanthin_19_mg_m3	mg m <sup>-3</sup>	100308.csv	Hexanoyloxyfucoxanthin-19
butanoyloxyfucoxanthin_19_mg_m3	mg m <sup>-3</sup>	100308.csv	Butanoyloxyfucoxanthin-19
alloxanthin_mg_m3	mg m <sup>-3</sup>	100308.csv	Alloxanthin
zeaxanthin_mg_m3	mg m <sup>-3</sup>	100308.csv	Zeaxanthin
prasixanthin_mg_m3	mg m <sup>-3</sup>	100308.csv	Prasixanthin
neoxanthin_mg_m3	mg m <sup>-3</sup>	100308.csv	Neoxanthin
violaxanthin_mg_m3	mg m <sup>-3</sup>	100308.csv	Violaxanthin
diatoxanthin_mg_m3	mg m <sup>-3</sup>	100308.csv	Diatoxanthin
diadinoxanthin_mg_m3	mg m <sup>-3</sup>	100308.csv	Diadinoxanthin
peridinin_mg_m3	mg m <sup>-3</sup>	100308.csv	Peridinin
carotene_mg_m3	mg m <sup>-3</sup>	100308.csv	Carotene
lutein_mg_m3	mg m <sup>-3</sup>	100308.csv	Lutein
suspended_particulate_matter_g_m3	g m <sup>-3</sup>	100306.csv	Suspended particulate matter
particulate_organic_nitrogen_g_m3	g m <sup>-3</sup>	100306.csv	Particulate organic nitrogen
total_particulate_carbon_g_m3	g m <sup>-3</sup>	100306.csv	Total particulate carbon
particulate_organic_carbon_g_m3	g m <sup>-3</sup>	100306.csv	Particulate organic carbon
dissolved_organic_carbon_g_m3	g m <sup>-3</sup>	100306.csv	Dissolved organic carbon
Cast	l	100312.csv	Processed cast number
Depth	m	100312.csv	Depth of the vertical bin, e.g., -1.00 representing the depth bin [-0.90, -1.10 m]
TmpWat	°C	100312.csv	Water temperature
Cond	ms cm <sup>-1</sup>	100312.csv	Conductivity
Salin	PSU	100312.csv	Salinity
SigmaT	l	100312.csv	Quantity (no units) to specify the density of seawater
TiProf	Degree	100312.csv	Tilt of the profiling radiometer
TiRef	Degree	100312.csv	Tilt of the reference radiometer
VSpeed	ms <sup>-1</sup>	100312.csv	Vertical speed
Altim	m	100312.csv	Altimeter sounding of the distance from the ocean floor
N_OBS	l	100312.csv	Number of observations within the depth bin
EU <sub>nnn</sub>	W m <sup>-2</sup> μm <sup>-1</sup>	100312.csv	In-water upwelling irradiance at wavelength nnn
ED <sub>nnn</sub>	W m <sup>-2</sup> μm <sup>-1</sup>	100312.csv	In-water downwelling irradiance at wavelength nnn
ER <sub>nnn</sub>	W m <sup>-2</sup> μm <sup>-1</sup>	100312.csv	In-air downwelling irradiance at wavelength nnn
KU <sub>nnn</sub>	m <sup>-1</sup>	100312.csv	Diffuse attenuation at wavelength nnn calculated from the upwelling irradiance
KD <sub>nnn</sub>	m <sup>-1</sup>	100312.csv	Diffuse attenuation at wavelength nnn calculated from the downwelling irradiance
PAR_ABS	μmol m <sup>-2</sup> s <sup>-1</sup>	100312.csv	Photosynthetically active radiation (PAR)
PAR%SRF	%	100312.csv	PAR at depth <i>z</i> relative to PAR on the sea surface
K_PAR	m <sup>-1</sup>	100312.csv	Diffuse attenuation for PAR





**Figure 4.** (a) Total chlorophyll *a* and (b) particulate organic carbon across the sampled areas.

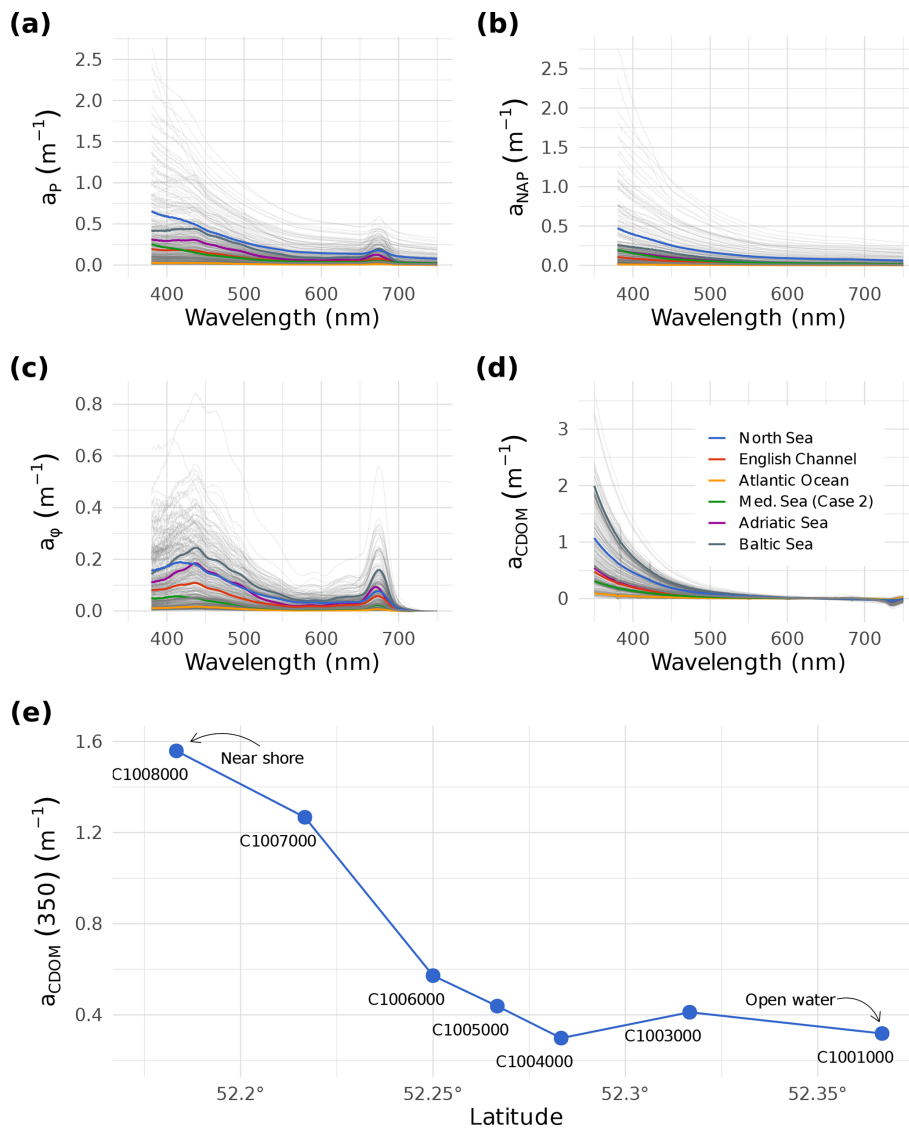
strategy. Across all the stations, total chlorophyll *a* ranged between 0.05 and 29 mg m<sup>-3</sup>. The median chlorophyll *a* ranged between ca. 0.15 mg m<sup>-3</sup> in the Atlantic Ocean and 5.9 mg m<sup>-3</sup> in the Baltic Sea (Fig. 4a). Based on the definition proposed by Antoine et al. (1996), the sampled stations were representative of a wide range of trophic statuses: oligotrophic ( $n = 17$ , chlorophyll *a*  $\leq 0.1$  mg m<sup>-3</sup>), mesotrophic ( $n = 106$ ,  $0.1 \leq$  chlorophyll *a*  $\leq 1$  mg m<sup>-3</sup>), and eutrophic ( $n = 245$ , chlorophyll *a*  $> 1$  mg m<sup>-3</sup>). A similar shore-to-sea gradient pattern could be observed for POC, with median values varying between ca. 0.07 g m<sup>-3</sup> and 0.8 g m<sup>-3</sup> in the Mediterranean Sea (Case 2) and the Baltic Sea, respectively (Fig. 4b, range between 0.02 and 2.5 g m<sup>-3</sup>).

#### 4.1.2 IOPs

Measured absorption spectra in each area for total particulates ( $a_p$ ), non-algal particles ( $a_{NAP}$ ), phytoplankton ( $a_\phi$ ), and chromophoric dissolved organic matter ( $a_{CDOM}$ ) are presented in Fig. 5. Phytoplankton maximum absorption peaks at around 440 and 675 nm are easily distinguishable in Fig. 5c and were found to be highly correlated with the concentration of total chlorophyll *a* (Pearson's  $r > 0.90$ ). Across

all the areas,  $a_{CDOM}(350)$ , a proxy for dissolved carbon concentration, varied between 0.03 and 3.66 m<sup>-1</sup>. These values fall within the ranges of the mean values reported globally in the ocean (0.14 m<sup>-1</sup>), coastal (1.82 m<sup>-1</sup>), and estuarine (4.11 m<sup>-1</sup>) ecosystems (Massicotte et al., 2017). Overall, the highest absorption by dissolved organic matter was observed in the Baltic Sea (Fig. 5d). Sampled stations in this area were located west of the Oder River plume, which drains important quantities of humic substance from its catchment area. In contrast, the lowest CDOM absorption was observed in the Atlantic Ocean, where stations were located away from land and thus less influenced by terrestrial inputs (Fig. 1). The spatial variability of CDOM is further emphasized in Fig. 5e, where  $a_{CDOM}(350)$  decreased by a factor of 4 between the two most distant points (approximately 40 km) of the westernmost transect sampled in the North Sea (see Fig. 1b).

Light scattering by suspended particles in the water column is a driver of reflectance variability and is often used by remote-sensing applications to discriminate between Case-1 and Case-2 waters (Sathyendranath, 2000; Morel and Bélanger, 2006). The particulate scattering coefficient at 440 nm,  $b_p(440)$ , ranged between 0.05 and 35.8 m<sup>-1</sup>



**Figure 5.** (a) Average total particulate ( $a_p$ ), (b) non-algal ( $a_{\text{NAP}}$ ), (c) phytoplankton ( $a_\phi$ ), and (d) chromophoric dissolved organic matter ( $a_{\text{CDOM}}$ ) absorption spectra in each area. Light gray lines in panels (a–d) show individual absorption spectra measured across the sampled areas. (e)  $a_{\text{CDOM}}(350)$  along the westernmost transect in the North Sea (see Fig. 1b).

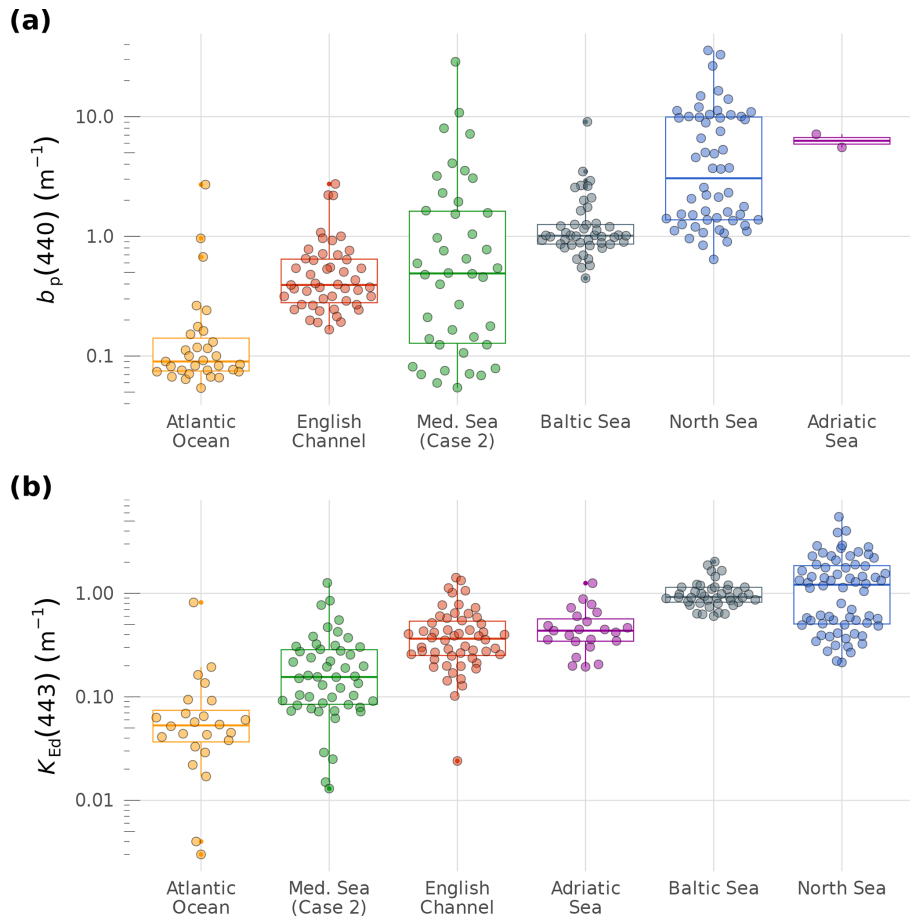
(Fig. 6a). The median values varied by almost 2 orders of magnitude between the Atlantic Ocean and the Adriatic Sea (Fig. 6a). Furthermore,  $b_p(440)$  showed more or less the same spatial pattern as the downward irradiance attenuation coefficient at 443 nm,  $K_{E_d}(443)$ , with median  $b_p(440)$  values varying between 0.05 and  $1.21 \text{ m}^{-1}$  in the Atlantic Ocean and the North Sea, respectively (Fig. 6b, Pearson's  $r = 0.76$ ).

#### 4.1.3 AOPs

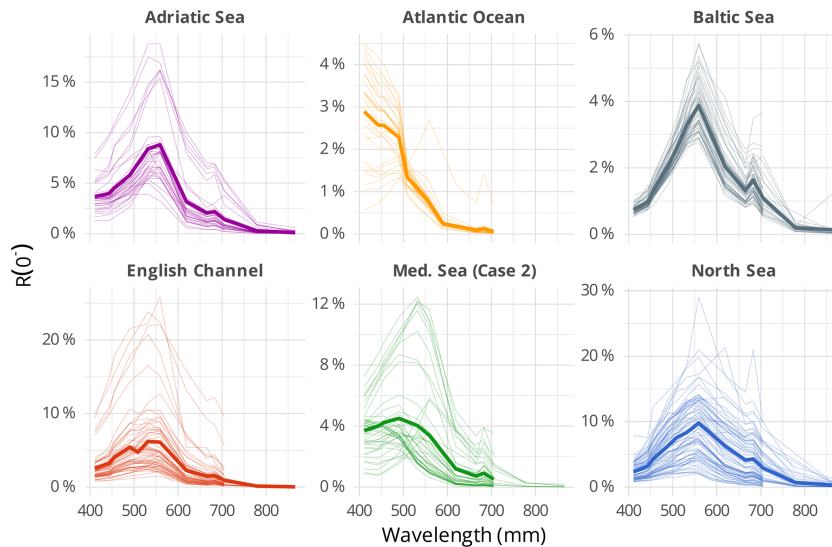
Subsurface reflectance  $R(0^-)$  for all COAST $\ell$ OOC sites is shown in Fig. 7, reflecting the fundamental differences in the bio-optical properties of the different areas visited. For example, due to enhanced CDOM absorption, reflectance in

the blue part of the spectrum is generally low at the Baltic Sea sites, while it is higher in the Case-1 waters encountered in the Atlantic Ocean or the Mediterranean Sea. In the red and NIR, very high reflectance values are observed in the sediment-rich Case-2 waters of the North Sea and the English Channel.

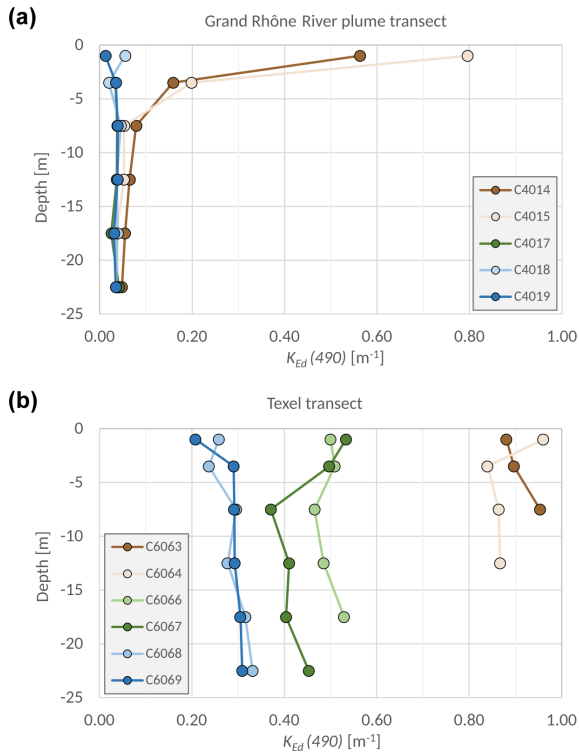
Sample profiles of the diffuse attenuation coefficient  $K_{E_d}(490)$  at a vertical resolution between 2.0 m close to the surface and 5.0 m underneath are shown in Fig. 8 for two transition zones ranging from turbid waters close to the coast to clearer waters offshore. Figure 8a shows  $K_{E_d}(490)$  values along a transect taken in the Gulf of Lion on 30 September 1997 from site C4014 located ca. 5 km to the west of the mouth of the main Rhône branch (Grand Rhône) to site



**Figure 6.** (a) Particulate scattering coefficient at 440 nm ( $b_p(440)$ ) and (b) attenuation coefficient for downward irradiance at 443 nm ( $K_{Ed}(443)$ ) across the sampled areas.



**Figure 7.** Subsurface reflectance  $R(0^-)$  across the sampled areas. The thick lines represent the regional averages. Note that  $R(0^-)$  at wavelengths above 705 nm could only be derived from helicopter-based measurements. Please also note that the y axes are adjusted to the data presented within each subplot.



**Figure 8.** (a)  $K_{Ed}(490)$  profiles taken in the Gulf of Lion on 30 September 1997 along a transect extending from site C4014 located ca. 5 km to the west of the mouth of the main Rhône branch (Grand Rhône) to site C4019 located some 20 km further south in clear Mediterranean waters. (b)  $K_{Ed}(490)$  profiles taken in the North Sea on 12 September 1998 along a transect extending from site C6063 close to the western shore of the island of Texel to site C6069 located ca. 22 km further west in the open North Sea.

C4019 some 20 km further south in clear Mediterranean waters. While  $K_{Ed}(490)$  at the offshore sites C4017 to C4019 adopts low values between ca. 0.02 and 0.05  $\text{m}^{-1}$  reflecting clear-water conditions, sites C4014 and C4015 are strongly influenced by the Rhône River plume, resulting in  $K_{Ed}(490)$  reaching values of up to 0.8  $\text{m}^{-1}$  in the surface layer. The turbid Rhône freshwater floats on top of denser and clearer Mediterranean seawater, resulting in a strong stratification. Figure 8b shows a similar transect taken in the North Sea on 12 September 1998 extending from site C6063 close to the western shore of Texel to site C6069 located ca. 22 km further west in the open North Sea. Again, there is a distance-to-coast-dependent decrease in  $K_{Ed}(490)$  values likely related to reduced turbidity, from ca. 0.8 to 1.0  $\text{m}^{-1}$  close to the coast down to values of ca. 0.2 to 0.3  $\text{m}^{-1}$  offshore. In contrast to the transect in the Gulf of Lion, no obvious stratification is observed:  $K_{Ed}(490)$  values are rather homogeneously distributed over the entire water column, pointing to effective mixing.

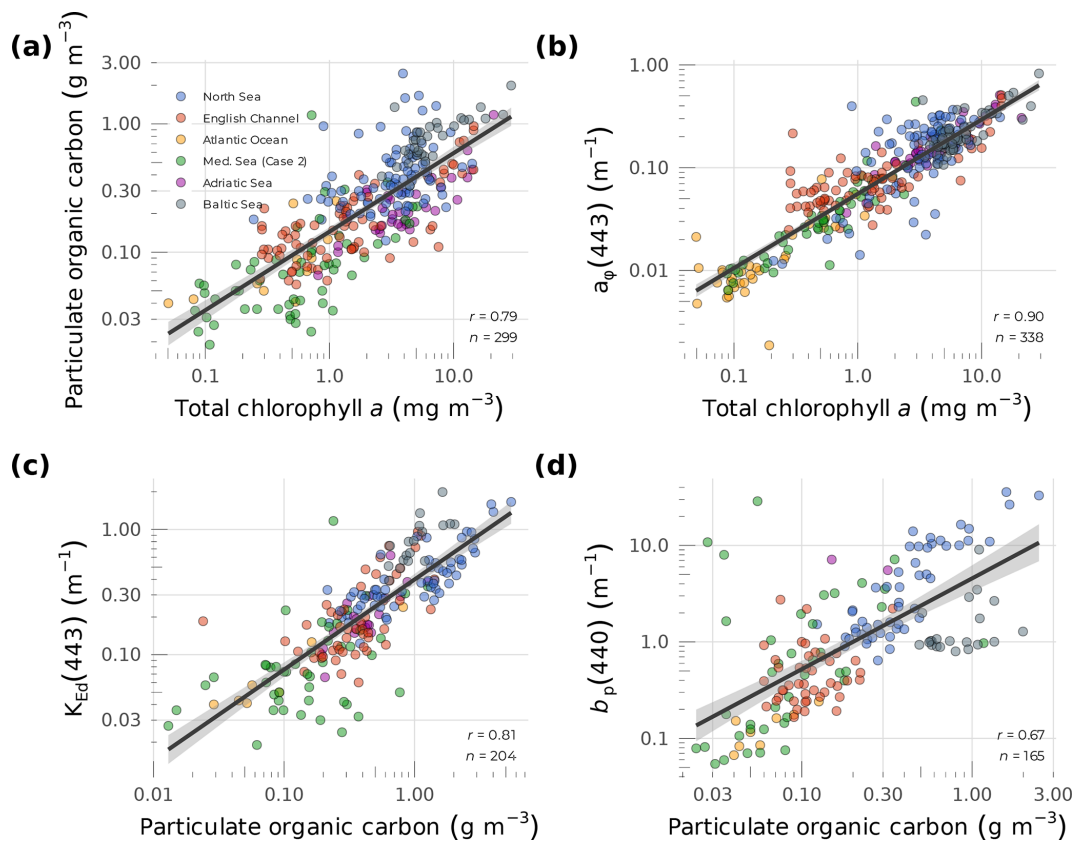
## 4.2 Co-variability across bio-optical measurements

Few research papers have previously shown how optically significant components and AOPs and IOPs measured during the COAST $\ell$ OOO missions co-varied (Ferrari, 2000; Ferrari et al., 2003; Babin, 2003; Babin et al., 2003a). In this section, a brief overview of selected pairwise relationships is presented. Total chlorophyll *a* co-varies with POC (Fig. 4), and the relationship was highly variable (Fig. 9a). While the global Pearson correlation is 0.79 ( $n = 299$ ), it ranged between 0.36 in the North Sea ( $n = 88$ ) and 0.78 in the Adriatic Sea ( $n = 37$ ), showing that optically significant components do not necessarily co-vary altogether. Likewise, the global Pearson correlation between  $a_{\phi}$  and total chlorophyll *a* (Fig. 9b) was relatively high ( $r = 0.90$ ,  $n = 338$ ). The lowest and highest correlations were observed in the North Sea ( $r = 0.69$ ,  $n = 88$ ) and the Adriatic Sea ( $r = 0.91$ ,  $n = 38$ ), respectively. POC concentration is well known to be an important driver of both IOPs and AOPs in aquatic ecosystems (Stramski et al., 2008; Cetinić et al., 2012). Unsurprisingly, positive correlations were observed between POC and  $K_{Ed}(443)$  (Fig. 9c,  $r = 0.81$ ,  $n = 204$ ) and  $b_p(440)$  (Fig. 9d,  $r = 0.67$ ,  $n = 165$ ).

Based on the size–reactivity continuum model proposed by Benner and Amon (2015), physicochemical and photochemical processes shape the size distribution of organic matter along the aquatic continuum and determine the contrasted intrinsic nature of the particles for each type of ecosystem. This is likely one key reason explaining why the observed relationships are quite variable across the different sampled ecosystems (Fig. 9). For example, in coastal areas (e.g., the Baltic Sea, Fig. 1), the POC content is generally influenced by large humic organic particles drained by rivers from the surrounding watershed (Babin et al., 2003a). Consequently, in these areas, POC concentration (Fig. 4b), absorption (Fig. 5), and downward attenuation coefficients (Fig. 9c) were higher than in other sampled areas. In contrast, the stations that were farther away from the coast (such as the Atlantic Sea, Fig. 1) had lower POC concentrations (Fig. 4b) and absorption (Fig. 5).

## 5 Code and data availability

The COAST $\ell$ OOO data are provided as a collection of comma-separated value (.csv) files that regroup measurements associated with each measurement. The processed and tidied version of the data is hosted at SEANOE (SEA scieNtific Open data Edition) under the CC-BY license (<https://doi.org/10.17882/93570>, Massicotte et al., 2023). To aid the user in merging these files, there is a lookup table file designated station list (<https://www.seanoe.org/data/00824/93570/data/100311.csv>, Massicotte, 2023c) that can serve as a table to join the data together based on the station's unique identifier. Table 1 shows the complete list of available measurements. Please note that the actual filenames



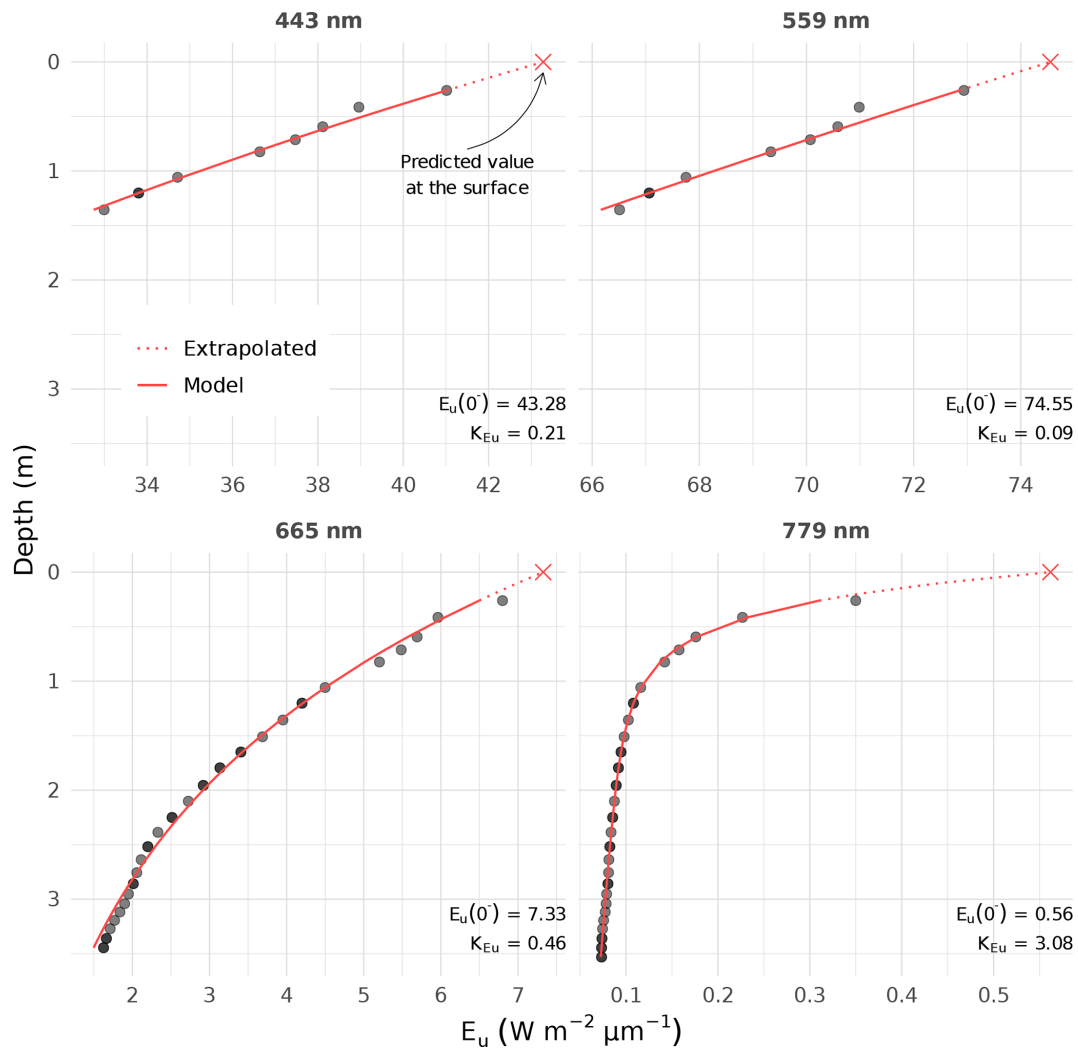
**Figure 9.** Scatterplots showing relationships among different selected variables. **(a)** Particulate organic carbon (POC) and **(b)** phytoplankton absorption at 443 nm ( $a_{\phi}(443)$ ) against total chlorophyll  $a$ . **(c)** Downward irradiance at 443 nm and **(d)** particulate scattering at 440 nm against particulate organic carbon. The red lines show the linear relationships between the variables. The shaded gray areas represent the 95 % confidence intervals around the fitted models.

(e.g., /100311.csv) are assigned by SEANOE during the upload process and cannot be altered. All statistical analyses were performed in R version 4.3.0 (R Core Team, 2023). The code used to produce the figures and the analysis presented in this paper is available under the GNU GPLv3 license (<https://doi.org/10.5281/zenodo.8091717>, Massicotte, 2023b). Irradiance depth merging was done with an Interactive Data Language (IDL) code available under the Creative Commons Attribution 3.0 Germany license (<https://doi.org/10.5281/zenodo.8096682>, Fell, 2023). Note that there are plans to update and further improve the current dataset. More data will likely be recovered from the backup archives, most prominently AC9 measurements to derive in situ vertical profiles of  $a$ ,  $b_p$ , and  $c$ . Intended improvements to radiometer processing comprise (not exhaustively) a better dark-current correction, the application of a self-shading correction, a more accurate treatment of the air–sea transmission of downward irradiance ( $E_d$ ), and a closer look at uncertainties. Any update of the COAST $\ell$ OOC dataset will be made publicly available through SEANOE.

## 6 Conclusions

The consolidated and quality-controlled data collected during the COAST $\ell$ OOC oceanographic expeditions offer, even if collected more than 20 years ago, many possibilities for better understanding the bio-optical dynamics along the land-to-sea gradient where optically complex water drained from watersheds mixes with seawater. To date, nearly 40 peer-reviewed studies have made use of the COAST $\ell$ OOC dataset to develop remote-sensing algorithms and to study the optical properties of seawater across the coastal to open-ocean gradient (see Table A1). This now consolidated and easily accessible dataset will facilitate future development and evaluation of new bio-optical models adapted for optically complex waters. In this paper, only a subset of variables has been presented. The reader is referred to Table 1 for a complete list and description of variables collected during the COAST $\ell$ OOC expeditions.

## Appendix A



**Figure A1.** Example of extrapolation of the upward irradiance toward the sea surface (station C3004000). The irradiance values just below the sea surface (red crosses,  $E_u(0^-)$ ) and the corresponding diffuse attenuation coefficient for the topmost water layer are specified in the figures. The full lines represent the interpolated values, whereas the dashed lines show extrapolation values up to the sea surface.

**Table A1.** Scientific articles in peer-reviewed journals using or referencing COAST $\ell$ OOO.

Publications
Babin et al. (2023b), Begouen Demeaux and Boss (2022), Belanger et al. (2008), Beltrán-Abaunza et al. (2014), Blix et al. (2019), Caillault et al. (2021), Chami and Platel (2007), Claustre et al. (2000), D'Alimonte et al. (2014), Defoin-Platel and Chami (2007), Doerffer and Schiller (2007), Doron et al. (2007), Doron et al. (2011), Dransfeld et al. (2005), Ferrari (2000), Groom et al. (2009), Jamet et al. (2012), Kratzer and Moore (2018), Loisel et al. (2001), Loisel et al. (2017), Loisel et al. (2018), Matsuoka et al. (2011), Matsuoka et al. (2014), Morel and Bélanger (2006), Neukermans et al. (2012), Oubelkheir et al. (2007), Schroeder et al. (2007), Schroeder et al. (2022), Shahraiyni et al. (2007), Tassan and Ferrari (2002), Tassan et al. (2000), Wei et al. (2021), Yu et al. (2016, 2019), Zhang et al. (2003), Zhang and Fell (2004, 2007), Zheng and Stramski (2013).

**Author contributions.** Conceptualization: MB. Formal analysis: FF, DD, VFS, MB. Investigation: MB, FF, VFS. Software: DD (AC9), FF (radiometry), VFS (data acquisition and preprocessing). Writing: PM with support from all the co-authors.

**Competing interests.** The contact author has declared that none of the authors has any competing interests.

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